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# Detection of hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>) in the Martian atmosphere with MEX / PFS

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### **Abstract**

We first derived the long-term averaged abundance of hydrogen peroxide  $(H_2O_2)$  in the Martian atmosphere with data sets of Planetary Fourier Spectroscopy (PFS) onboard Mars Express (MEX). The total averaged amounts of  $H_2O_2$  at three Martian years were  $45 \pm 21$  ppb and  $25 \pm 18$  ppb in the forward/reverse pendulum direction, respectively. It could not explain the observed short lifetime of  $CH_4$  in the Martian atmosphere.

#### 1. Introduction

Recently, a non-negligible amount of CH<sub>4</sub> was detecteded in the Martian atmosphere [1]. This discovery is remarkable because possible sources include biological or geological activities. Recent observations show fast temporal and large spatial variation of CH<sub>4</sub> [2,3]. These results imply that the lifetimes of CH<sub>4</sub> is the order of weeks or months which, in turn, imply a strong sink, about 600 times faster than predicted by standard photochemistry processes [4]. It would suggest the existence of strong oxidants. Hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>) is possibly one of the most important tracer elements for possible oxidants in the Martian atmosphere. However, only ground-based investigations [5] are currently available. The long term characteristics for the abundance, variability, and spatial distribution of H<sub>2</sub>O<sub>2</sub> are uncertain.

# 2. Instrument and objective

The detection of  $H_2O_2$  has not been reported yet by instruments onboard Martian orbiters. It requires higher spectral resolution in order to distinguish the weak absorption of  $H_2O_2$  from the stronger absorptions by major constituents. At the moment, the Planetary Fourier Spectroscopy (PFS) aboard

Mars Express (MEX) is the best instrument for this objective, which has the highest spectral resolution in the mid-IR range ( $\sim$ 1.3 cm<sup>-1</sup>) in previous orbiters. It has been operated for seven years on orbit. In this study, we derive the long-term averaged amount of  $H_2O_2$  from this data set, and try to evaluate the possible loss scenario of  $CH_4$  in the Martian atmosphere by oxidants in the atmosphere.

## 3. Data analysis and Results

We derived the long-term averaged abundance of H<sub>2</sub>O<sub>2</sub> in the Martian atmosphere among three Martian years (from January 2004 to December 2009). By careful rejections of weak instrumental noises and averaging several thousands of measurements, we got the mixing ratio of Martian H<sub>2</sub>O<sub>2</sub> with the accuracy of several tens ppb in the spectral range from 360 cm<sup>-</sup> to 385 cm<sup>-1</sup>. The total averaged H<sub>2</sub>O<sub>2</sub> mixing ratios were  $45 \pm 21$  ppb in the forward pendulum direction, and  $25 \pm 18$  ppb in the reverse pendulum direction, respectively. The annual averaged H<sub>2</sub>O<sub>2</sub> mixing ratios were  $50 \pm 27$  ppb and  $43 \pm 26$  ppb at the MY 27,  $0 \pm 89$  ppb and  $0 \pm 74$  ppb at the MY 28, and 46  $\pm$  21 ppb and 28  $\pm$  20 ppb at the MY 29, respectively. The seasonal variation of H<sub>2</sub>O<sub>2</sub> mixing ratio was 0 -120 ppb during the observational period (see Fig. 1). We are now analyzing the origin of the discrepancy in forward/reverse data sets, but the amount of both is converged to similar values.

### 4. Discussions and conclusion

Our results are slightly larger than past ground-based and space-born observations and the photochemical models. This discrepancy can possibly be due to (1) artificial - the criteria used for our data selection, i.e., surface temperature between 250K and 270K and local time between 10 and 16; or (2) natural - unexpected transient enhancement might be included

(by electrostatic discharge appear during dust storms and dust devils). In both cases, this little discrepancy are not sufficient to explain the short lifetime of CH<sub>4</sub> recently observed [2,3]. By the assumption of photochemical steady state, we conclude that the atmospheric oxidation loss by OH cannot promote the observed fast variation of CH<sub>4</sub> (see Fig. 2 and Table. 1). We suggest that atmospheric oxidation loss by O(<sup>1</sup>D) and O may play an important role. On the other hand, no correlation between the variation of H<sub>2</sub>O<sub>2</sub> and CH<sub>4</sub> has been found. This also suggests that the atmospheric electrochemical loss do not play major role in the fast variation for CH<sub>4</sub>.

Table 1: The molecular number density assumed in photochemical steady state for H<sub>2</sub>O<sub>2</sub> and OH.

Mol.	Number density (cm <sup>-3</sup> )	Ref.
H <sub>2</sub> O <sub>2</sub>	1×10 <sup>9</sup> - 1×10 <sup>10</sup>	this study
$H_2O$	$1 \times 10^{13}$	this study
CO	$1 \times 10^{14}$	[6]
O	$5\times10^7$	[7]
$O(^{1}D)$	0.9	[7]
$O_3$	5×10 <sup>9</sup>	[7]
H	$5 \times 10^4$	[7]
$H_2$	5×10 <sup>8</sup>	[7]

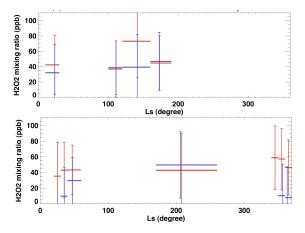


Figure 1: The seasonal variation of the mixing ratio of  $\mathrm{H_2O_2}$  derived by the PFS/LW in the MY 27 (top), and in the MY 29 (bottom). The vertical bars of each data point are the standard deviation and the horizontal bars are the averaged duration. The red lines are the FWD data sets and blue lines are REV data sets.

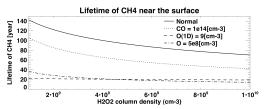


Figure 2: The expected calculated photochemical lifetime of  $CH_4$  affected by the abundance of  $H_2O_2$  near the surface. Solid line shows the lifetime with the input normal parameters listed in Table 1. The dotted line shows the lifetime with the  $5\times10^8$  cm<sup>-3</sup> of CO. The dashed lines show the values with the 10 times larger abundance of  $O(^1D)$  and O than the lifetime predicted by [7].

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