



## Out-of-Sequence Thrust in the Higher Himalaya- a Review & Possible Genesis

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An out-of-sequence thrust (OOST) has been established inside the Higher Himalaya by previous workers more frequently from Nepal- and Bhutan Himalaya. The OOST lies between the South Tibetan Detachment System-Upper (STDSU) and the South Tibetan Detachment System-Lower (STDLSL). The thrust has been recognized as the Kakhtang Thrust in Bhutan (Grujic et al., 2002 and references therein); Khumbu Thrust (Searle, 1999), Modi Khola Shear Zone (Hodges et al., 1996), Kalopani Shear Zone (Vannay and Hodges, 1999), Toijem Shear Zone in Nepal (Carosi et al., 2007), Chaura Thrust (Jain et al., 2000)- also designated as the Sarahan Thrust (Chambers et al., 2008) in the western Indian Himalaya in Sutlej section, Zimithang Thrust in the eastern Indian Himalaya (Yin et al., 2006), as 'physiographic transition' in Marsyandi valley, Nepal (Burbank et al., 2003). We note that considering the upper strand of the Main Central Thrust (the MCTU) as the lower boundary of the Higher Himalaya, the physiographic transition has also been referred to lie in the Lesser Himalaya. The period of activity of the OOST was 22.5-18.5 Ma (Hodges et al., 1996), 14-10 Ma (Grujic et al., 2002), 4.9-1.5 Ma (Jain et al., 2000), and from Late Pliocene to even Holocene Period (Burbank, 2005). The out-of-sequence thrusting was followed after the initiation of channel flow at  $\sim 15$  Ma in the Higher Himalaya with a maximum delay of  $\sim 13$  Ma. However, in the Bhutan Himalaya, the thrusting continued along with the extensional ductile shearing in the STDSU at 11-10 Ma (Hollister and Grujic, 2006). The OOST in the Higher Himalaya lies inside the zone of the top-to-SW sense of ductile shearing. The OOST, at Kakhtang, Toijem, and Chaura are ductile shear zones with a top-to-SW sense of shearing. The OOST merges with the MCT and the Main Himalayan Thrust (MHT) at a depth of 30 km or more and either runs 200-300 km beneath the Tibetan plateau (Grujic et al., 2002; Hollister and Grujic, 2006). The hanging wall side of the OOST is more dominant with migmatites and leucogranites (Searle, 1999; Yin et al., 2006; Carosi et al., 2007; Grujic et al., 2002; Hollister and Grujic, 2006), but the footwall side does contain these rocks (Hodges et al., 1996; Chambers et al., 2008). The thickness of the OOST are 50 m (Carosi et al., 2007),  $>150$  m (Yin et al., 2006), 3-6 km (Searle, 1999) and  $\sim 1.5$  km (Vannay and Hodges, 1996). A number of hypotheses have been put forward to explain the genesis of the OOST. These are (i) a disparity in erosion rates triggered mainly by a spatial variation in the intensity of rainfall (Wobus et al., 2005). (ii) The lower boundary of the channel flow extrusion defined the OOST (Hollister and Grujic, 2006). (iii) As a result of a heterogeneous velocity profile of channel flow extrusion across lithologic discontinuity (Carosi et al., 2007). The granitic melt at depth in some way led to this thrusting (Swapp and Hollister, 1991). Had channel flow been the extrusion mechanism of the Higher Himalaya, the genesis of the OOST might somehow be related to this extrusion. In this work, a channel flow box was prepared and polydimethylsiloxane was used as the model material. A channel flow was generated in the horizontal channel and was allowed to extrude through an inclined channel similar to the Higher Himalaya (Mukherjee, 2007). In different considerations, the walls of the Higher Himalaya are parallel and diverging-up. A late formed blind thrust plane forms at the corner joining the inclined and the horizontal wall and crops to the surface much later to the initiation of channel flow. On the basis of its late arrival to the surface than the channel flow and its relative position in the model Higher Himalaya, the thrust is comparable with the OOST. This means that (i) climatic factors nor lithologic discontinuity were a trigger to the OOST; and (ii) the OOST is a delayed product of channel flow that initiated at a sub-horizontal channel below the Tibetan plateau and extrude the Higher Himalaya.

### References.

Burbank, D.W., 2005. Cracking the Himalaya. *Nature* 434, 963-964.

Burbank, D.W., Blythe, A.E., Putkonen, J., Pratt-Sitaula, B., Gabet, E., Oskin, M., Barros, A., Ojha, T.P., 2003.

Decoupling of erosion and precipitation in the Himalayas. *Nature* 426, 652-654.

Carosi, R., Montomili, C., Visonà, D., 2007. A structural transect in the lower Dolpo: insights in the tectonic evolution of Western Nepal. *Journal of Asian Earth Sciences* 29, 407-423.

Chambers J.A., Argles, T.W., Horstwood, M.S.A., Harris, N.B.W., Parrish, R.R., Ahmad, T., 2008. Tectonic implications of Palaeoproterozoic anatexis and Late Miocene metamorphism in the Lesser Himalayan Sequence, Sutlej valley, NW India. *Journal of the Geological Society, London* 165, 725-737.

Godin, L., Grujic, D., Law, R.D. and Searle, M.P., 2006. Channel flow, extrusion and extrusion in continental collision zones: an introduction. In: R.D. Law and M.P. Searle (Editors) *Channel Flow, Extrusion and Extrusion in Continental Collision Zones*. Geological Society of London Special Publication 268, 1-23.

Grujic, D., Casey, M., Davidson, C., 1996. Ductile extrusion of the Higher Himalayan Crystalline in Bhutan: evidence from quartz microfabrics. *Tectonophysics* 260, 21-43.

Grujic, D., Hollister, L.S., Parrish, R.R., 2002. Himalayan metamorphic sequence as an orogenic channel: insight from Bhutan. *Earth Planetary Science Letters* 198, 177-191.

Harris, N., 2007. Channel flow and the Himalayan-Tibetan orogen: a critical review. *Journal of Geological Society, London* 164, 511-523.

Hollister, L.S. and Grujic, D., 2006. Himalaya Tiber Plateau. Pulsed channel flow in Bhutan. In: R.D. Law, M.P. Searle and L. Godin (Editors). *Channel flow, Ductile Extrusion and Extrusion in Continental Collision Zones*. Geological Society of London Special Publication 268, pp. 415-423.

Jain, A.K., Kumar, D., Singh, S., Kumar, A., Lal, N., 2000. Timing, quantification and tectonic modelling of Pliocene-Quaternary movements in the NW Himalaya: evidences from fission track dating. *Earth Planetary Science Letters* 179, 437-451.

Mukherjee, S. 2007. Geodynamics, deformation and mathematical analysis of metamorphic belts of the NW Himalaya. Unpublished Ph.D. thesis. Indian Institute of Technology Roorkee. pp. 1-267.

Searle, M.P., 1999. Extensional and compressional faults in the Everest-Lhotse massif, Khumbu Himalaya, Nepal. *Journal of Geological Society, London*, 156, 227-240.

Swapp, S.M., Hollister, L.S., 1991. Inverted metamorphism within the Tibetan slab of Bhutan: evidence for a tectonically transported heat source. *Canadian Mineralogist* 29, 1019-1041.

Vannay, J-C., Hodges, K.V., 1996. Tectonomorphic evolution of the Himalayan metamorphic core between the Annapurna and Dhaulagiri, central Nepal. *Journal of Metamorphic Geology* 14, 635-656.

Wobus, C., Heimsath, A., Whipple, K., Hodges, K., 2005. Active out-of-sequence thrust faulting in the central Nepalese Himalaya. *Nature* 434, 1008-1011.

Yin, A., Dubey, C.S., Kelty, T.K., Gehrels, G.E., Chou, C.Y., Grove, M., Lovera, O., 2006. Structural evolution of the Arunachal Himalaya and implications for asymmetric development Himalayan orogen. *Current Science* 90, 195-206.