



## **Role of mesostructural inheritances in inversion-related fracturing. Example form the Mataporquera Anticline, Western Pyrenees (Spain).**

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Since the pioneering work of Stearns (1968), several studies have focused on the description and interpretation of fracture patterns in reservoir-scale thrust-related anticlines (e.g. Srivastava and Engelder, 1990; Cooper, 1992; Lemiszki et al. 1994; Tavani et al., 2006; Savage et al., In press). It is widely recognised that, during thrust-related folding, fractures development results from the interplay of different factors, including the mechanical stratigraphy (e.g. Corbett et al., 1987, Woodward and Rutherford, 1989; Gross, 1995; Couzens and Wiltschko, 1996; Fischer and Jackson, 1999; Tavani et al., 2008), the environmental conditions (e.g. Chester et al., 1991; Jamison, 1992; Lemiszki et al., 1994), and the folding process (e.g. Thorbjornsen and Dunne, 1997; Tavani et al., 2006). More recently, many works have also documented how previously developed mesostructures exert a strong control on fracturing (e.g. Silliphant et al., 2002; Bergbauer and Pollard, 2004; Bellahsen et al., 2006; Lash and Engelder, 2009). Investigating the role that mesostructural inheritances exert on fracture development is particularly important in inversion-related anticlines, where the widespread presence of sin-extensional mesostructures can strongly influence the sin-inversion fracturing.

In this work we present data from the Mataporquera Anticline (Western Pyrenees, Spain). This WNW-ESE striking and ESE plunging anticline formed during the Cenozoic, when previously developed, mostly Early Cretaceous, extensional structures were reworked. Seismic cross-sections across the structure allow to constrain both Mesozoic and Cenozoic evolution of the shallower portion of the anticline. On the other hand, the same cross-sections indicate that the tectonic style was fully thick-skinned, during both Mesozoic and Cenozoic.

The mesostructural pattern observed in the Mataporquera Anticline includes extensional faults, joints and veins striking parallel and perpendicular to the Lower Cretaceous extensional master faults (ENE-WSW). These structures have been tilted together with beds during folding and are interpreted as sin-extensional (i.e. Mesozoic). Both faults and joints have been, at places, re-worked as faults. This reworking was accompanied by the development of other structures (E-W striking pressure solution cleavages and N-S striking extensional structures) allowing to infer a Cenozoic stress field consistent with the reworking of the sin-extensional structures.

Cenozoic mesostructures are mostly observed in areas poorly affected by Mesozoic extensional deformation and in the vicinity of major re-worked map-scale faults. On the other hand, in other sites, the inversion is testified only by the reworking of previously developed mesostructures. These observations allow us to conclude that, in the Mataporquera Anticline, the network of sin-extensional Mesozoic discontinuities inhibited the development of sin-inversion mesostructures.

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