



Volcanic and ENSO Effects on Climate in China in the last millennium

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Summary

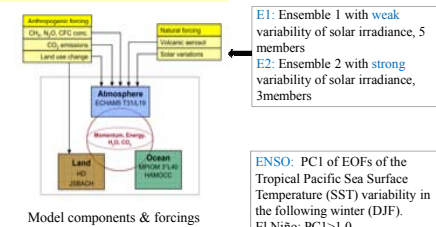
The effects of **volcanic eruptions** in two ensemble (8 simulations) by the COSMOS ASOB earth system model are analyzed. Following results are found:

■ Volcanic eruptions with neutral SOI cool Asia, dry NE China and Tibetan Plateau, and cause an **EI Niño like warming** in the tropical Pacific. The global **time scale** of temperature decay is assessed by an exponential function fitting.

■ The **favor of EI Niño** events following eruptions in previous studies is **not found**. ENSO overcome volcanoes with disturbed patterns of T2M and SPI.

■ A case study of Tambora in 1815 is compared with reconstructions.

Model and Method



21 volcanic eruptions with at least -2.0 W/m² decrease in net top solar irradiation

| No. | Year | Name | VEI* |
|-----|------|-----------------|-----------------|
| 1 | 842 | Unknown | *Newhall C. G. |
| 2 | 854 | Unknown | S. Steve (1982) |
| 3 | 897 | Unknown | |
| 4 | 971 | Unknown | |
| 5 | 1193 | Unknown | |
| 6 | 1228 | Unknown | |
| 7 | 1258 | Unknown | |
| 8 | 1286 | Unknown | |
| 9 | 1442 | Unknown | |
| 10 | 1456 | Kuonene/Vamutu? | 6 |
| 11 | 1600 | Huaynaputina | 6 |
| 12 | 1641 | Parker | 6 |
| 13 | 1673 | Capelo | |
| 14 | 1694 | Senia | |
| 15 | 1809 | St-Helen? | |
| 16 | 1815 | Tambora | 7 |
| 17 | 1832 | Babuyan Claro | 4? |
| 18 | 1835 | Cosiguina | |
| 19 | 1884 | | 6 |
| 20 | 1903 | Grimsvotn | 4 |
| 21 | 1992 | Pinatubo | 6 |

Temperature Decay

The decay of the temperature anomaly is fitted by an exponential function:

$$\Delta T = \Delta T_1 * \exp(-(t-1)/\tau)$$

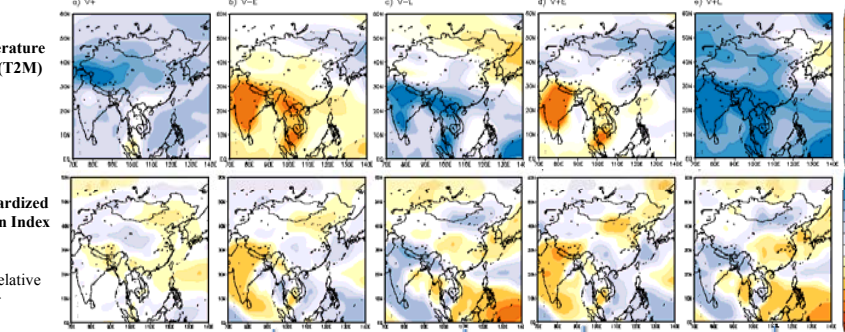
2 fitting parameters:
 ΔT_1 temperature anomaly at year 1
 τ time scale
(1 ≤ t ≤ 10, time after event, years)

(1) Volcanoes EI Niño La Niña Vol/EI Niño Vol/La Niña

Fig 1 Temperature at 2 meters (T2M) anomaly

Fig 2 Standardized Precipitation Index (SPI, JJA)

at year one relative to event year



◆ **T2M:** Volcanoes (Fig1a) cool Asia with West China up to 1°C.
◆ **SPI:** Drought (Fig2b) in China: NE, south, east coast, and Tibetan Plateau. Wetness in Central and West China.

◆ **T2M:** EI Niño (Fig1b) warms Asia except NE China. La Niña (Fig1c) effect antisymmetric.
◆ **SPI:** dipole pattern (Fig2b), NE China moderate drought, India severe drought and SE China wet. La Niña (Fig2c) effect antisymmetric.

◆ **T2M:** Cooling weakened/intensified by EI Niño/La Niña (Fig1d/Fig1e) descending from SW to NE. Similar to ENSO effects (Fig1b & c).
◆ **SPI:** ENSO effects overcome volcanoes by disturbed pattern of drought/wetness in North China Plain (Fig2d/Fig2e). EI Niño effects after eruptions slightly stronger than EI Niño (Fig2b).

Table 1: T2M and precipitation anomaly at year one relative to event year

| Variable | Ensemble | Volcanoes only | Volcanoes+ EI Niño | Volcanoes+ La Niña |
|-------------------|----------|----------------|--------------------|--------------------|
| T2M /C | E1 | -0.54 | -0.12 | -0.73 |
| | E2 | -0.38 | -0.31 | -0.68 |
| Precipitation /mm | E1 | -56 | -64 | 29 |
| | E2 | -43 | -68 | |

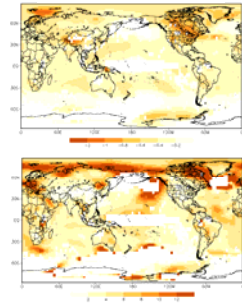
◆ Simultaneous chances of volcanic eruptions and ENSO is **44 % (38%)** in E1 (E2). **No favor to EI Niño** following volcanic eruptions (**half and half** for EI Niño and La Niña).

◆ **Dynamic causes:** The Monsoon Hadley Circulation index (MHI) and East Asia Monsoon Index (EAMI) no significant changes by V+, by V+E, **significant phase-locking** of MHI at year one & **significant enhancement** at year two, by V+L, the **enhancement** comes at year one.

(2) Global temperature decay

Fig 3 a) Temperature anomaly ΔT_1

Fig 3 b) Decay time scale τ



◆ **Big ΔT_1** (Fig3a) and **big time scale τ** (Fig3b) in NH.
◆ **Small ΔT_1** in SH (Fig3a) with small τ or no-finite-positive values (Fig3b blank) through fitting due to the **small values** of temperature decrease of ocean to the volcanic eruptions, or **even positive values**. Same regions have intense **long term memory** found by Fraedrich and Blender (2003).
◆ Similar pattern between fitting ΔT_1 and simulated T2M at year 1 (not shown) **with an EI Niño like warming** up at tropical Pacific).

(3) Tambora in 1815: a case study

◆ **Cooling:** -1.2 for E1, -1.3°C for E2, in reconstructions/historic documents (Fig4a). Droughts only in SE China in reconstructions but not in all 8 simulations, too strong ENSO signal.
◆ **ENSO effects:** Ocean conditions differ strongly in 8 simulations, 4 typical cases (Fig4b colored-dots, EI Niño, La Niña, Neutral and best-guess which matches reconstructed ENSO by Quinn et al., 1998) are selected for T2M and SPI (Fig5).

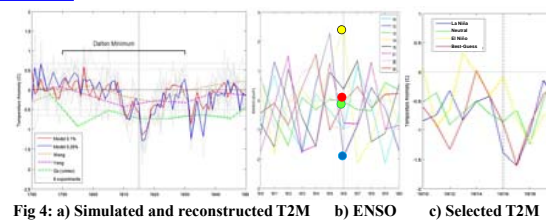


Fig 4: a) Simulated and reconstructed T2M b) ENSO c) Selected T2M

◆ **T2M:** La Niña enhances temperature reduction while EI Niño compensates it at year one (Fig4c) with difference of 1.5 °C.

◆ **SPI:** Droughts with neutral SOI (Fig5b, right); La Niña case is not 'standard' due to strong EI Niño signal in 1815 (Fig5a & Fig4b). EI Niño has less 'power' in North China Plain (section 1, hint: ENSO at year 0 needs to be considered as well).

Fig 5: simulated T2M (left) and SPI (JJA, right) in 1816

