



Changes of fluid flow regimes in a complex calcite vein network (Natih Formation, Oman Mountains): Insights from Stable Isotope Analysis

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We measured $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ compositions of calcite veins and their immediate limestone host-rock from an intensely veined outcrop at the top of the middle Cretaceous (Turonian) Natih A Formation in the Central Oman Mountains (Virgo and Arndt, 2010). The $\delta^{18}\text{O}$ composition of the limestone host-rock in the studied pavement ranges from $22.5\text{\textperthousand}$ to $23.7\text{\textperthousand}$. The $\delta^{13}\text{C}$ composition ranges from $1.1\text{\textperthousand}$ to $1.9\text{\textperthousand}$. This range of compositions is depleted in ^{18}O relative to unaltered Cretaceous marine limestones ($24.7\text{--}28.8\text{\textperthousand}$ after Veizer and Hoefs, 1976). However, in a regional isotopic survey of the limestone sequence, Wagner (1990) has shown that the $\delta^{18}\text{O}$ composition of the Natih A Formation can range from $23.3\text{\textperthousand}$ to $26.3\text{\textperthousand}$. The depleted C/O isotopic compositions are results of meteoric diagenesis during subaerial exposure (Wagner, 1990; Grelaud et al., 2006). The $\delta^{18}\text{O}$ compositions of vein calcite vary from $22.5\text{\textperthousand}$ to $26.2\text{\textperthousand}$, while $\delta^{13}\text{C}$ compositions range from $-0.8\text{\textperthousand}$ to $2.2\text{\textperthousand}$. Two compositional trends are apparent for vein calcite data. In trend A there is a spread in $\delta^{13}\text{C}$ values from host rock compositions to values nearly $1.3\text{\textperthousand}$ lower than the immediate host rock, while $\delta^{18}\text{O}$ remains constant. Microstructural observations have shown high contrasts of $\delta^{13}\text{C}$ within the same sample, indicating episodic fluid flow. We don't observe reaction haloes. In the second composition range (trend B) a number of vein calcite samples have $\delta^{18}\text{O}$ values up to $3.3\text{\textperthousand}$ higher than the immediate host rock range, whereas the $\delta^{13}\text{C}$ compositions are similar to the host-rock values. The majority of the trend B samples are from a late, E-W trending fault vein that cross cuts any other extension vein of the network and has a normal displacement. Episodic fluid flow is indicated by high contrast of $\delta^{18}\text{O}$ values within the same sample. By combining our observations with existing literature we propose that (1) meteoric diagenesis has altered the top of Natih A during meteoric diagenesis. (2) After burial a complex and dense network of crack-seal extension veins formed promoting vertical fluid flow (bringing in lower $\delta^{13}\text{C}$ values) in terms of meters and lateral fluid flow in terms of 10s of meters (rock buffered veins). (3) The change in fluid flow is reflected by trend B of enriched $\delta^{18}\text{O}$ values constraint to a later fault vein. The fault vein has tapped a fluid reservoir at a deeper stratigraphic level with high $\delta^{18}\text{O}$ values that have a typical Cretaceous marine limestone composition ($26.2\text{\textperthousand}$).

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