



## **Equilibrium Line Altitudes and paleotemperature reconstructions from Nevado Hualcán (9°S) and Nevado Coropuna (15°S), Tropical Andes (Peru).**

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We have reconstructed the Equilibrium Line Altitude (ELA) in seven valleys on the SW slope of Nevado Hualcán (9°S, 77°W; 6122 m asl) and on the SE slope of Nevado Coropuna (15°S, 72°W; 6377masl) using the Area x Altitude Balance Ratio method (Osmaston 2005). We have also deduced the paleotemperatures using the following equation:  $\Delta T = \text{ATLR} \cdot \Delta \text{ELA}$ ; being  $\Delta T$  (°C) the paleotemperature depression; ATLR (°C/m) the Atmospheric Temperature Lapse Rate; and  $\Delta \text{ELA}$  (m) the ELA depression. The ATLR for Coropuna was deduced through the use of data loggers. For Hualcán we used the value  $\text{ATLR} = 0.0065^\circ\text{C}/\text{m}$ , valid for the tropics (Kaser and Osmaston, 2002).

We obtained the following results:

- 1) Hualcán: a) ELAs: 5124m (2003); 5018m (1962); 4994m during the Little Ice Age (LIA); and 4652m during the last studied maximum advance considered to be the Younger Dryas (YD, ~13-11ka) by correlation with nearby mountains (Glasser et al., 2009). b)  $\Delta \text{ELAs}$ : 106m (1962); 130m (LIA); and 199m (YD). c)  $\Delta T$ :  $-0.69^\circ\text{C}$  (1962);  $-0.85^\circ\text{C}$  (LIA);  $-3.07^\circ\text{C}$  (YD).
- 2) Coropuna: a) ELAs: 5862m (2007); 5853m (1986); 5787m (1955); 5776 (LIA); and 4951m in the 13-1136Cl ka phase (Úbeda, 2011). b)  $\Delta \text{ELA}$ : 9m (1986); 66m (1955); 86m (LIA); and 911m in 13-1136Cl ka. C)  $\Delta T$ :  $-0.20^\circ\text{C}$  (1986);  $-0.71^\circ\text{C}$  (1955); and  $-7.65^\circ\text{C}$  (13-1136Cl ka).

The values  $\Delta T$  during LIA in Hualcán and Coropuna ( $0.85$  and  $0.72^\circ\text{C}$ ) are consistent with the global warming considered to be  $0.74^\circ\text{C}$  between 1906 and 2005 (IPCC, 2007). During the mid XXth century and the LIA,  $\Delta T$  is higher in Hualcán ( $0.69^\circ\text{C}$  and  $0.85^\circ\text{C}$ ) than in Coropuna ( $0.55^\circ\text{C}$  and  $0.72^\circ\text{C}$ ), with a regional gradient of  $-0.02^\circ\text{C}$  per degree of latitude ( $^\circ\text{C}/^\circ$ ). However, during the YD (13-1136Cl ka),  $\Delta T$  was higher in Coropuna ( $7.65^\circ\text{C}$ ) than in Hualcán ( $3.07^\circ\text{C}$ ), with a gradient of  $0.76^\circ\text{C}/^\circ$ . Although other evidences exist of a pantropical cooling of  $>5^\circ\text{C}$  during the last glaciation, in Coropuna this cooling was strengthened by the retro-alimentation of its glacial system which in 13-1136Cl ka had a surface of  $>400 \text{ km}^2$  (Úbeda, 2011).

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