



## A large scale hydrological model combining Budyko hypothesis and stochastic soil moisture model

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Based on the Budyko hypothesis, the actual evapotranspiration,  $E$ , is controlled by the water conditions and the energy conditions, which are represented by the amount of annual precipitation,  $P$  and potential evaporation,  $E_0$ , respectively. Some theoretical or empirical equations have been proposed to represent the Budyko curve. We here select Choudhury's equation to describe the Budyko curve (Mezentsev, 1954; Choudhury, 1999; Yang et al., 2008; Roderick and Farquhar, 2011).

$$\varepsilon = (1 + \varphi^{-\alpha})^{-1/\alpha}, \varepsilon = \frac{E}{P}, \varphi = \frac{E_0}{P}$$

Rodriguez-Iturbe et al. (1999) proposed a stochastic soil moisture model based on a Poisson distributed rainfall assumption. Porporato et al. (2004) described the average water balance based on stochastic soil moisture model as following,

$$\varepsilon = 1 - \frac{\varphi \cdot \gamma^{\frac{\gamma}{\varphi} - 1} \cdot e^{-\gamma}}{\Gamma\left(\frac{\gamma}{\varphi}\right) - \Gamma\left(\frac{\gamma}{\varphi}, \gamma\right)}, \gamma = \frac{Zr}{h}$$

where,  $h$  means the average rainfall depth,  $Zr$  means basin water storage ability. Combining these two equations, we can get the relation between  $\alpha$  and  $\gamma$ . Then we develop a large scale hydrological model to estimate annual runoff from  $P$ ,  $E_0$ ,  $h$  and  $Zr$ .

$$R = (1 - \varepsilon) P, \varepsilon = (1 + \varphi^{-\alpha})^{-1/\alpha}, a = 0.7078\gamma^{0.5946}, \gamma = \frac{Zr}{h}$$

This method has good performance when it is applied to estimate annual runoff in the Yellow River Basin and the Yangtze River Basin. The impacts of climate changes ( $P$ ,  $E_0$  and  $h$ ) and human activities ( $Zr$ ) are also discussed with this method.