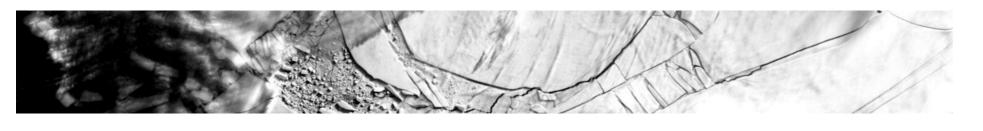
Fracture-induced softening for large-scale ice dynamics





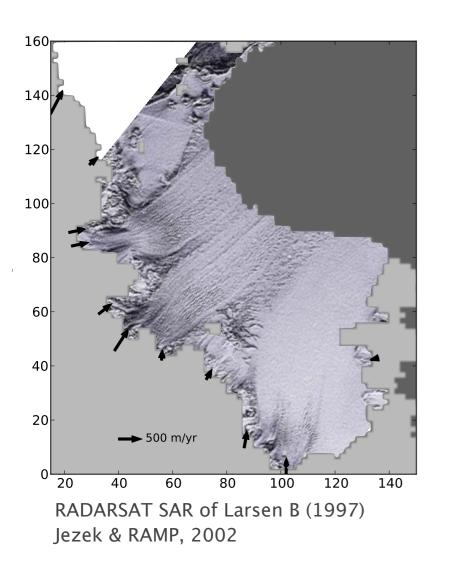
Potsdam Institute for Climate Impact Research Torsten Albrecht Anders Levermann







Motivation



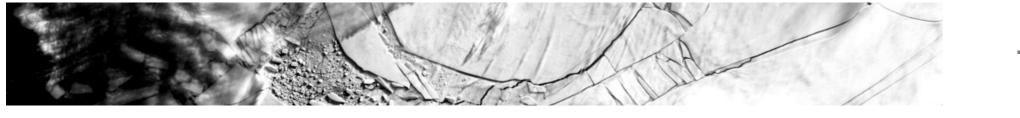
fractures \rightarrow ice flow

• ice fractures have a significant effect on flow dynamics, iceberg calving and on stability

ice flow \rightarrow fractures

- formation along shear margins, in the wake of topographic features and where flow units merge
- fractures elongated in band structures of certain frequency
- simple macroscopic parameterization needed for large-scale ice flow models (e.g. PISM) to evaluate location and feedback of fracture processes







The concept of the fracture density field

define $0 \le \phi \le 1$ as measure of fracture density

advective transport

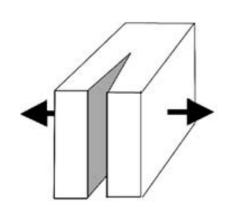
$$\frac{\partial \phi}{\partial t} + \boldsymbol{v} \cdot \nabla \phi = f_s$$

defining source

$$f_{\rm s} = \gamma \cdot \dot{\varepsilon}_1 \cdot (1 - \phi) \cdot \Theta(\sigma_{\rm t} - \sigma_{\rm cr})$$

fracture growth rate (maximal spreading) fracture interaction (upper bound)

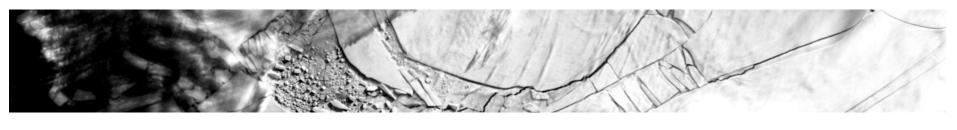
von Mises criterion for fracture initiation (max. octahedral shear stress)



Vaughan, 1993 Pralong & Funk, 2005

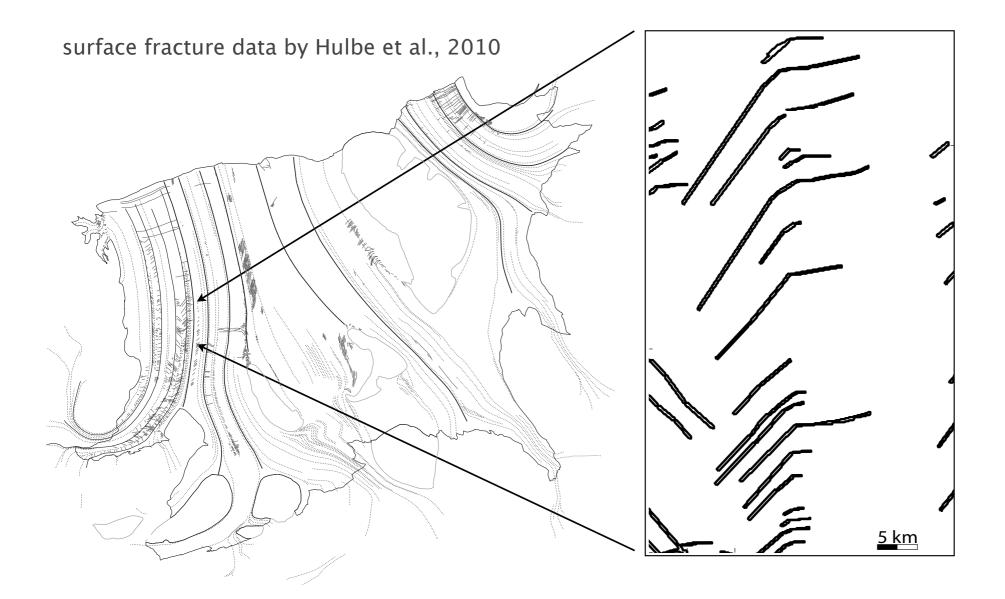




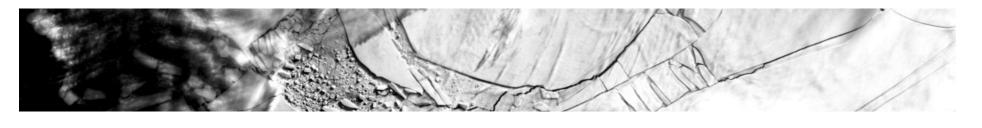




Observations of surface features (Filchner-Ronne / Antarctica)

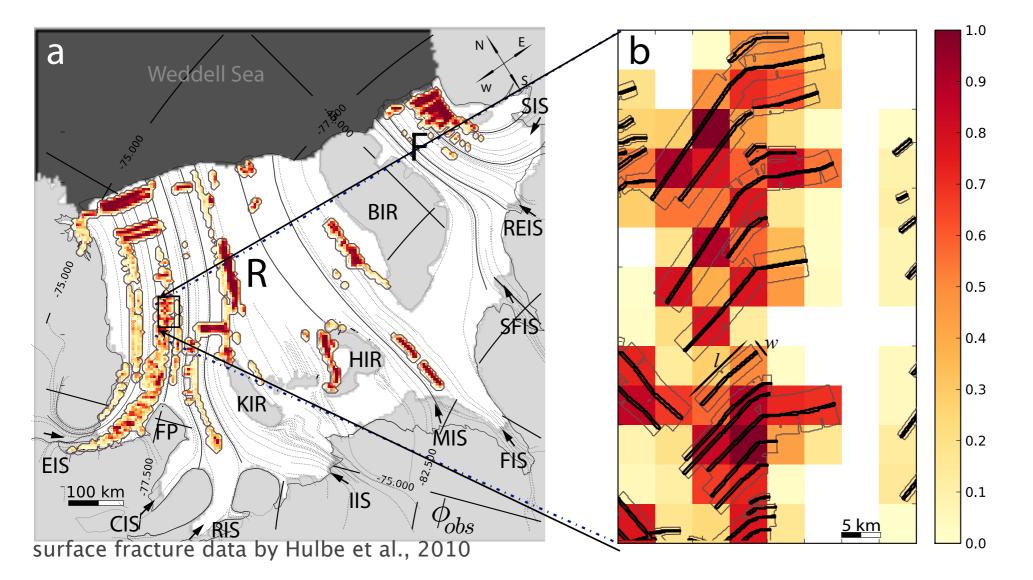








Validation against surface observations (Filchner-Ronne)

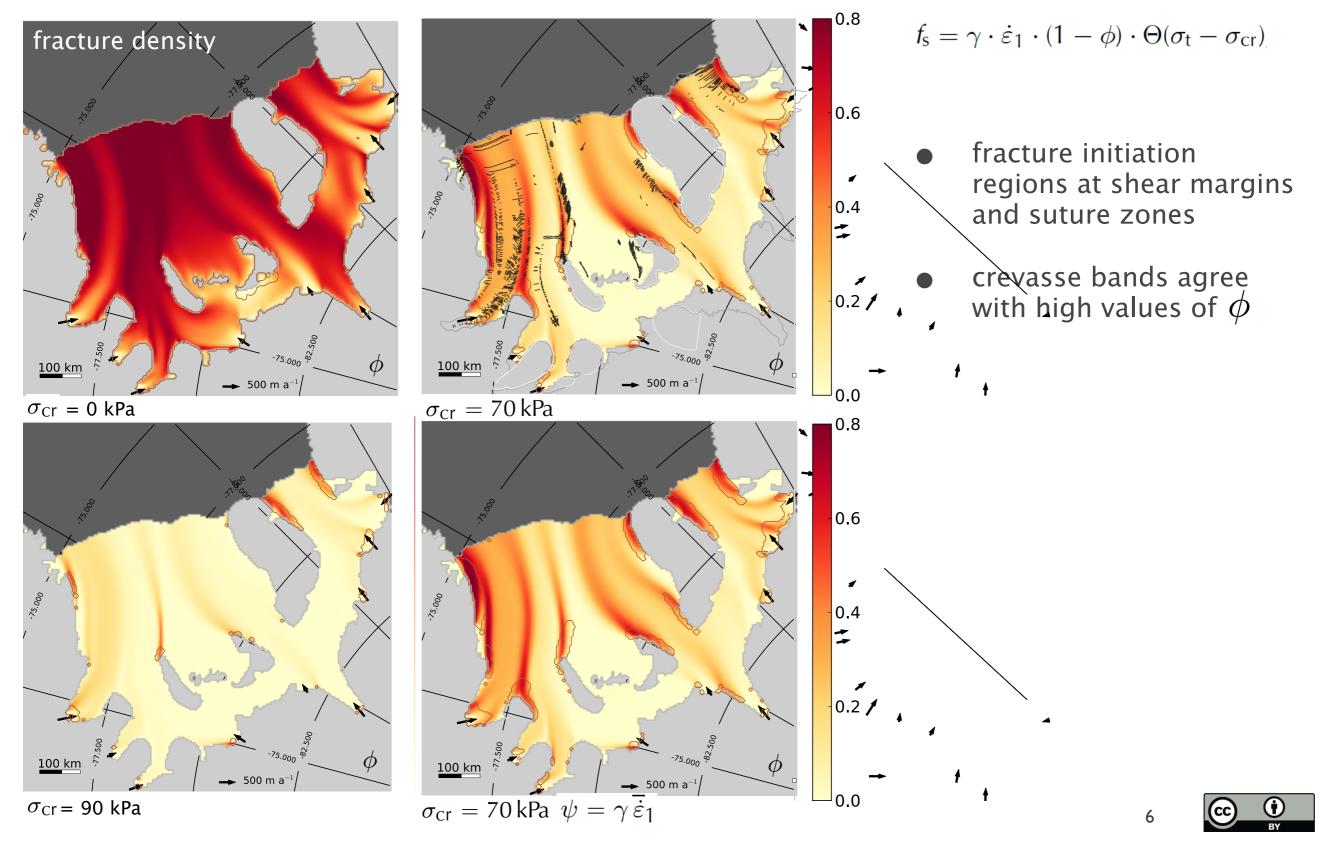


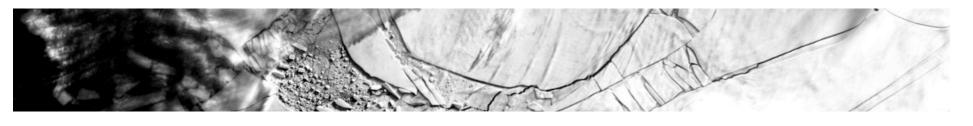
- simplified zone of influence around (visible) surface fractures
- fraction of covered grid cell (here 5km) determines $\phi_{
 m obs}$





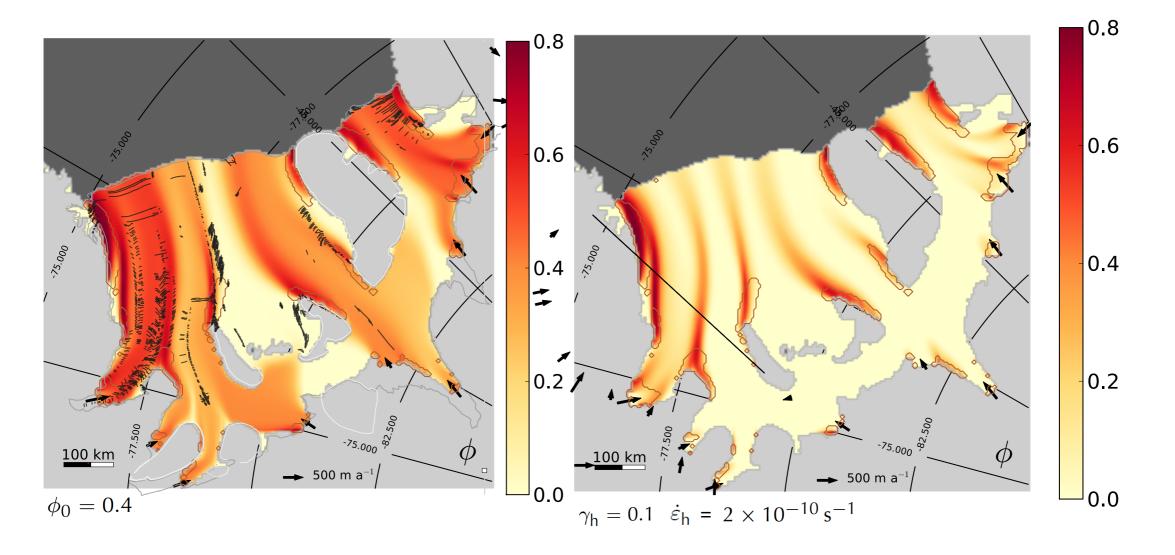
Application in diagnostic simulation of Filchner-Ronne





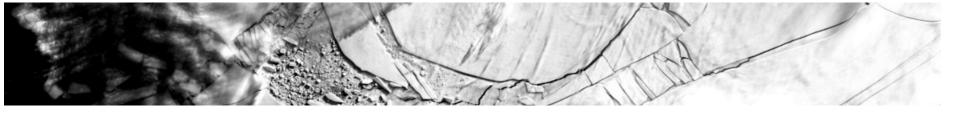


Boundary conditions and Healing



- ullet Boundary conditions (fractures from inlets, tidal tilt data Green et al.) ϕ_0
- Healing rate $f_{\rm h} = \Theta \left(-\psi_{\rm h}\right) \cdot \psi_{\rm h}$, with $\psi_{\rm h} = \gamma_{\rm h} \left(\dot{\varepsilon}_1 \dot{\varepsilon}_{\rm h}\right)$



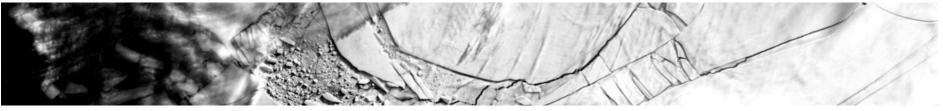




Fracture-induced softening (Larsen A + B) 700

C = 0.928C = 0.920C = 0.829600 0.8 500 speed (m/a) 400 300 0.6 200 100 80 100 60 120 0.4 observations (1997-2000) 0.7 diagnostic E = 1 0.0 0.5 0.4 0.3 0.2 diagnostic E = 5prognostic E = 3 prog. soft. $E = 3 + 10\phi$ 0.2 20 km _ 500 m/yr (D)0.1 0.0 0.0 80 90 100 60 70 110 120 $\sigma_{\rm Cr}$ = 50 kPa $\gamma_{\rm h}$ = 0.5 $\dot{\varepsilon}_{\rm h}$ = 6 imes 10⁻¹⁰ s⁻¹

• macroscopic softening as enhancement factor E in Glen's law $\dot{\boldsymbol{\varepsilon}} = EA au_{ ext{e}}^{n-1} \boldsymbol{ au}$





1.0

0.9

0.8

0.7

0.6

0.5

0.4

0.3

0.2

0.1

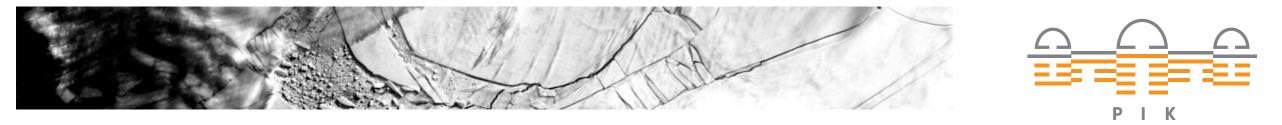
0.0

Application to grounded ice streams (Pine Island, Thwaites)

5000 ice speed in m/a fracture density -75.000 2000 1000 500 200 100 104.000 -104.000 50 40 km 40 km

softening and larger gradients along shearing margins





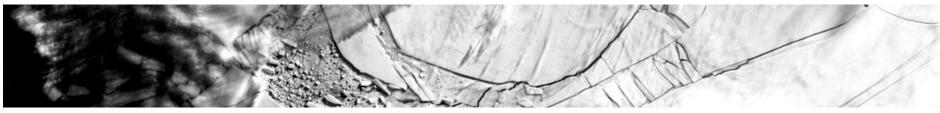
Conclusions

- first-order parameterization of fracture formation and its softening effect for coarse-scale ice sheet/shelf simulations
- reproduces observed surface-fracture pattern in ice shelves
- more realistic representation of flow dynamics

next steps

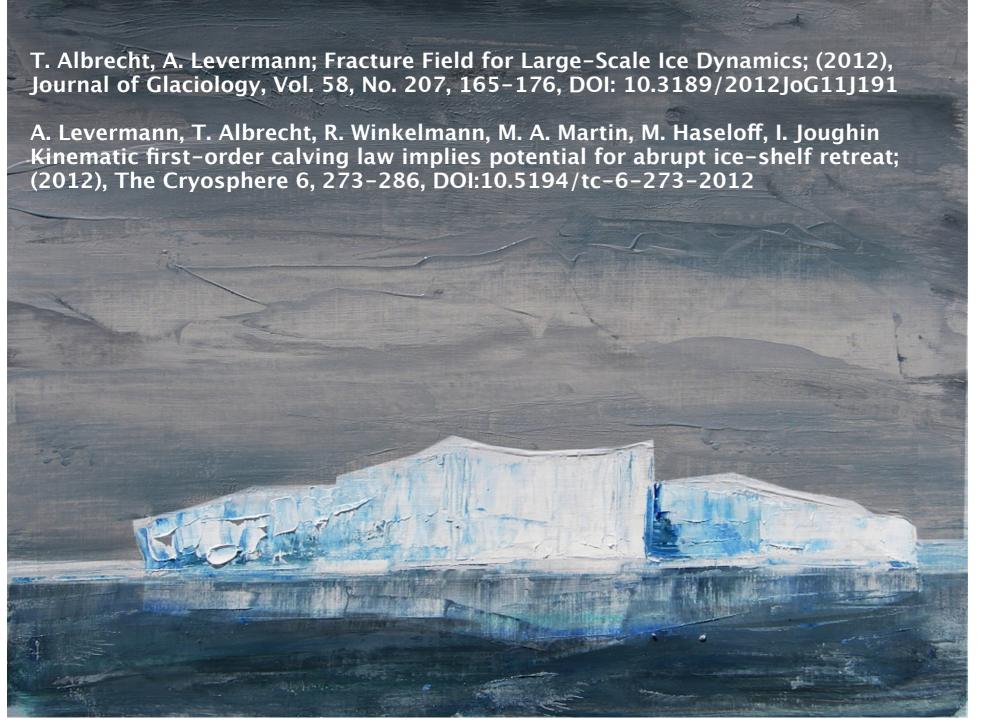
- consider fracture depth (parameterize hydro fracturing and refreezing processes)
- expand "eigencalving" parameterization (material dependent $C = K(\phi) \cdot \det(\dot{\epsilon})$)
- need fracture-data for validation!







Thank you!



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