



## Quantification and kinetics of H<sub>2</sub> generation during hydrothermal serpentinisation experiments

Teddy Castelain, Colin Fauguerolles, Johan Villeneuve, and Michel Pichavant  
ISTO, Orléans, France (teddy.castelain@cnrs-orleans.fr)

H<sub>2</sub>-rich hydrothermal fluids generated by serpentinisation of mantle rocks at slow-spreading ridges have been revealed by recent studies [1, 2]. Fluxes and the future of the H<sub>2</sub> produced by this process are poorly constrained [1, 3]. In this study, we aim to quantitatively evaluate the H<sub>2</sub> production fluxes associated with these hydrothermal systems and to document the kinetics of the hydrogen-producing reaction.

For this matter, hydrothermal serpentinisation experiments are being undertaken on mixtures composed of a natural peridotite from the Pindus ophiolite and olivine crystals from San Carlos. The experiments are conducted at a temperature of  $\sim 300^{\circ}\text{C}$  and a pressure of 450-500 bars in large-volume Dickson-Seyfried bombs for periods of  $\geq 1$  month. Starting materials are powders between 1 - 100  $\mu\text{m}$  for the peridotites and individual grains ranging from 1 - 2 mm for the San Carlos olivine. They are reacted with a homemade artificial seawater in such proportion that water-rock ratio = 1.8. The reactants are loaded in a modified Ti cell fitted with a semi-permeable Au-Pd membrane simultaneously allowing direct sampling of the hydrothermal fluid and *in situ* monitoring of the pH<sub>2</sub> during the advancement of the reaction. The gas fraction of the fluid sampled is then analyzed by gas chromatography (GC).

The pH<sub>2</sub> readings show traces of H<sub>2</sub> to be present from the second day of experiment. The increase of the pH<sub>2</sub> reaches a maximum after  $\sim 6$  days and the pH<sub>2</sub> finally stabilizes after  $\sim 16$  days at  $\sim 12.5$  bars, which corresponds to a local fO<sub>2</sub> of about NNO-4. The GC measurements, performed after 30, 43, 51 and 65 days, yield respectively, H<sub>2</sub> concentrations of 82.4, 89.7, 90.3 and 101 mmol.kg<sup>-1</sup> of water, in reasonable agreement with results from previous studies [4-6]. Further experiments are being undertaken in order to: duplicate observations, especially the pH<sub>2</sub> readings, more closely link the GC measurements and the *in situ* pH<sub>2</sub> readings, especially during the first 15 days of experiment, and relate H<sub>2</sub> production with the mineralogical composition of products of the serpentinisation reaction. The possible influence of the oxidation of the Ti cell on the H<sub>2</sub> production will be also checked by using a Au bag instead of a Ti cell. However, from our results, it appears that H<sub>2</sub> generation via serpentinisation is surprisingly rapid.

[1] J.-L. Charlou et al., *Chem. Geol.*, **191**, 2002. [2] C. Mével, *C.R. Geosc.*, **335**, 2003. [3] M. Cannat et al., *Geophys. Mono. Series*, **188**, 2010. [4] D.G. Allen, and W.E. Jr Seyfried, *Geochim. Cosmochim. Acta* **67** (8), 2003. [5] M.E. Berndt, et al., *Geology* **24** (4), 1996. [6] W.E. Seyfried, et al., *Geochim. Cosmochim. Acta* **71**, 2007.