



Oxygen isotope exchange and isotopic fractionation during N₂O production by denitrification

Dominika Lewicka-Szczebak (1), Reinhard Well (1), and Jan Kaiser (2)

(1) Thünen Institute of Climate-Smart Agriculture, Braunschweig, Germany (dominika.lewicka-szczebak@vti.bund.de), (2) Centre for Ocean and Atmospheric Sciences, School of Environmental Sciences, University of East Anglia, Norwich, UK

Stable isotopic analyses of N₂O may help in quantification of soil denitrification processes. N₂O reduction to N₂ during denitrification is associated with significant isotopic fractionation. Theoretically, this would allow the product ratio (N₂O/(N₂+N₂O)) to be calculated based on the isotopic signature of the remaining N₂O if the isotopic signature of the produced N₂O were precisely predictable. Oxygen isotopes are especially useful, in particular the ¹⁷O excess, Δ(¹⁷O), to quantify any oxygen isotope exchange between soil water and intermediate products (NO₂⁻, NO). This significantly influences the δ(¹⁸O) values of the N₂O produced. Previous studies showed this exchange to be nearly complete, up to 90%^{1,2}. However, there are very few studies on the associated oxygen isotopic fractionation and its potential coupling to the magnitude of oxygen isotope exchange with soil water^{2,3}. We hypothesize that for high oxygen exchange, the oxygen isotopic fractionation will show little variability.

In previous experiments, an ¹⁸O-enriched tracer was applied to quantify the magnitude of oxygen exchange^{1,2}. Such an approach does not permit determining any oxygen isotope fractionation. Here, we applied two novel experimental approaches: (1) waters of two different δ(¹⁸O) values within the natural range (-1.5 and -14.8 ‰ relative to VSMOW) were used for rewetting the incubated soils and the oxygen exchange was calculated from comparing the relative isotope ratio difference between N₂O and H₂O for these two water treatments; (2) soils were amended with Chile Saltpeter characterized by a high ¹⁷O excess (20.2 ‰) and the ¹⁷O excess of the N₂O product was measured. Both approaches were applied simultaneously on the same soil samples, which allowed quantifying the oxygen isotope exchange with two independent methods at the same time. The N₂O reduction was inhibited with acetylene method and δ(¹⁸O) values of the N₂O produced were measured to determine the oxygen fractionation during N₂O production.

Both methods showed nearly complete (95-99%) oxygen exchange with soil water for all soil types (loamy sand and silt loam), water contents (50 to 80% water-filled pore space), temperatures (8 and 22 °C), and N₂O production rates. The oxygen isotope fractionation ε(¹⁸O, N₂O-H₂O), defined as the relative isotope ratio difference between product (N₂O) and substrate (soil water), was quite stable between 16 and 20 ‰. This confirms that δ(¹⁸O) values may be useful in determination of the product ratio. Moreover, the very low ¹⁷O excess found in N₂O indicates that the hypothesis of soil denitrification contributing to the oxygen isotope excess of atmospheric N₂O can be rejected^{4,5}.

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