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## Si-F complexing in aqueous fluids: experimental study and implications for transport of immobile elements

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Intepretation of fluid-mineral interaction mechanisms and hydrothermal fluxes requires knowledge of predominant solubility and speciation reactions and their thermodynamic properties. Fluorine represents a hard electron donor, capable of complexing and transporting high-field strength elements, which are traditionally considered to be immobile. Reactions responsible for element mobility have general form  $MO_x + y$  HF (aq) + x - y H<sub>2</sub>O =  $M(OH)_{2x-y}F_y$  (aq), and their extent and transport efficiency relies on hydrogen fluoride activity. In natural fluids, a[HF] is controlled by various fluorination equilibria including neutralization of silicates with consequent formation of silicohydroxyfluoride complexes. Quartz solubility in HF-H<sub>2</sub>O fluids was experimentally determined at 400-800 °C and 100-200 MPa using rapid-quench cold-seal pressure vessels and the mineral weight-loss method. Quartz solubility significantly increases in the presence of hydrogen fluoride: at 400 °C and 100 MPa, dissolved SiO<sub>2</sub> ranges from 0.18 wt. % in pure H<sub>2</sub>O to 12.2 wt. % at 8.3 wt. % F in the fluid, whereas at 800 °C and 200 MPa it rises from 1.51 wt. % in pure H<sub>2</sub>O to 15.3 wt. % at 8.0 wt. % F in the fluid. The isobaric solubilities of quartz appear to be temperature-independent, i.e., effects of temperature vs. fluid density on the solubility are counteracting. The experimental data are described by the density model:  $\log m[SiO_2] = a + b\log \rho + c\log m[F] + dT$ , where  $a = -1.049 \text{ mol kg}^{-1}$ ,  $b = 0.816 \text{ mol cm}^{-3}$ ,  $c = 0.802 \text{ and } d = 1.256 \cdot 10^{-3} \text{ mol kg}^{-1} \text{ K}^{-1}$ . Solubility isotherms have similar  $d(\log m[SiO_2])/d(\log m[F])$  slopes over the entire range of conditions indicating that  $Si(OH)_2F_2$  is the major aqueous species. Several factors promote breakdown of silicohydroxyfluoride complexes and precipitation of silica solute: (i) decreasing temperature and pressure, i.e., fluid ascent and cooling and/or (ii) neutralization and increase in the alkali/H ratio of fluids during alteration reactions or removal of hydrogen halides by fluid boiling. Thermodynamic analysis of mineral equilibria in the system SiO<sub>2</sub>-Al<sub>2</sub>O<sub>3</sub>-FeO-MgO-CaO-Na<sub>2</sub>O-K<sub>2</sub>O-H<sub>2</sub>O-F<sub>2</sub>O<sub>-1</sub> indicates that cryolite, topaz, fluorite and sellaite represent fluoride buffers with decreasing chemical potential of F<sub>2</sub>O<sub>-1</sub> or a[HF], in a sequence from peralkaline to peraluminous silicic, intermediate to progressively Ca-rich mafic and, finally, ultramafic environments. Corresponding a[HF] decrease from  $10^{0.2}$  to  $10^{-1}$  and from  $10^{-1.6}$  to  $10^{-3.0}$  mol kg<sup>-1</sup> at 800 and 400 °C, respectively, and 100 MPa. These results imply that: (i) silicohydroxyfluoride and aluminumhydroxyfluoride complexes transport Si and Al in quantities appreciably greater than SiO<sub>2</sub> (aq) and aluminate species in peraluminous granite and greisen environments only, and (ii) significant transport (10-100 ppm) of high-field strength (e.g., Ti, Zr) and rare earth elements in aqueous fluids is predicted when formation constants of metal-fluoride complexes exceed  $10^{1-2}$  under hydrothermal conditions. This study concludes that in fluorine-bearing environments the transport of Si and Al remains little affected, but HFSE and REE are largely mobile.