



HYDROGEOLOGICAL SETTING

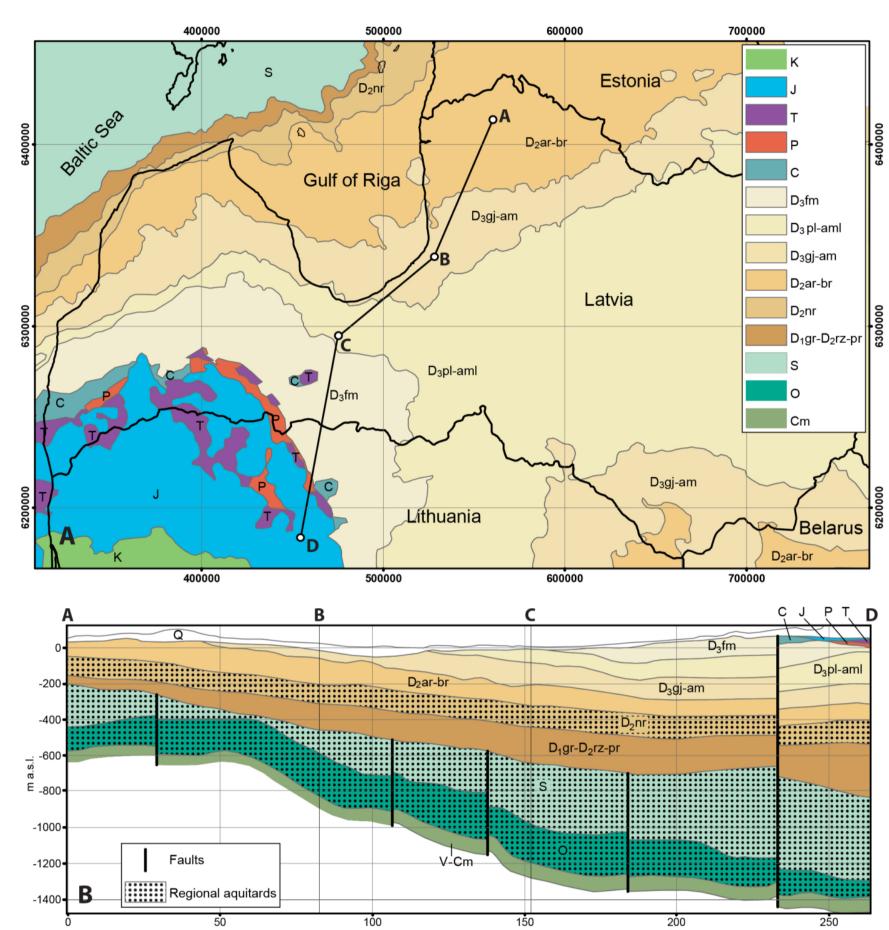


Figure 1. Geological map and geological cross-section of the study region without Quaternary cover (modified after Popovs et al. in print; Virbulis et al. 2013).

Location of cross-section indicated in A. V-Cm - Ediacaran- Cambrian sequence O-Ordovician sequence S- Silurian sequence D₁gr-D₂rz-pr- lower Devonian Gargzdu Fm to middle Devonian Parnu Fm; D₂nr- middle Devonian Narva formation; D₂ar-br- middle Devonian Burnieki Fm to Arukila Fm; D_3 gj-am upper Devonian Gauja Fm to Amata Fm; D₃fm- upper Devonian Famena Fm; C- Carboniferoussequence, P- Permian sequence, T-Triassic sequence J- Jurassic sequence K- Cretaceous sequence

The study area covers the central part of the Baltic Artesian Basin. The thickness of the sedimentary cover varies from about 500m in northern part to more than 2000m in southwestern part of Latvia.

zones of groundwater are traditionally identified within study area (Figure 1):

- stagnation zone- Ediacaran- Cambrian aquifer complex with brines;
- passive (slow) water exchange zone lower and middle Devonian aquifer complex with brackish groundwater;
- exchange active water zonefreshwater aquifers above Narva regional aquitard.

MATERIALS AND METHODS

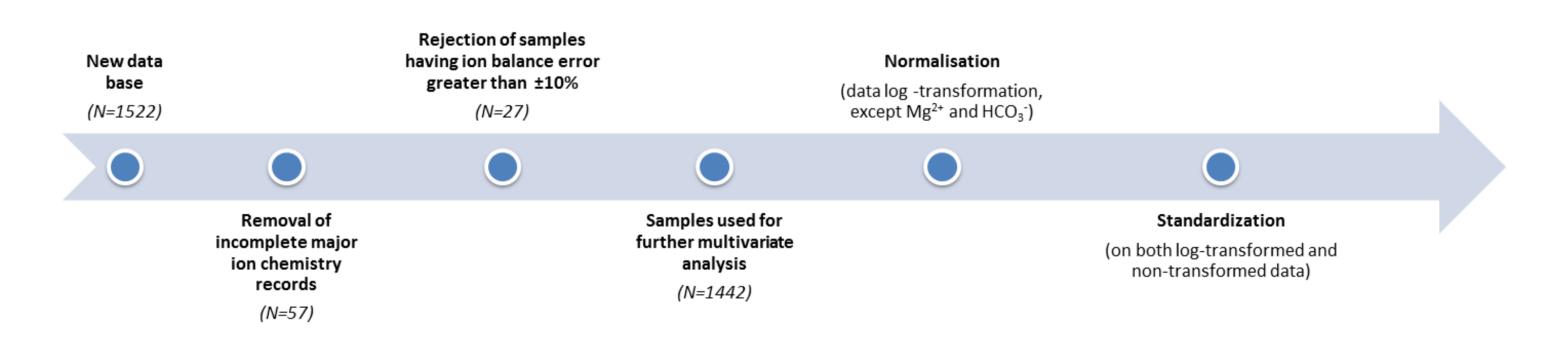


Figure 2. Data preparation for multivariate statistical analysis

Groundwater hydrochemical groups were defined using hierarchical cluster analysis (HCA) and principal component analysis (PCA). For HCA Euclidean distance as a similarity measure and Ward's method as a linkage method were used. For PCA Varimax rotation was used.

Data pre-treatment, PCA and HCA were performed using SPSS Statistics 22. Electrical balances and saturation indices of calcite, dolomite, gypsum and halite minerals were calculated using software PHREEQC, version 3.





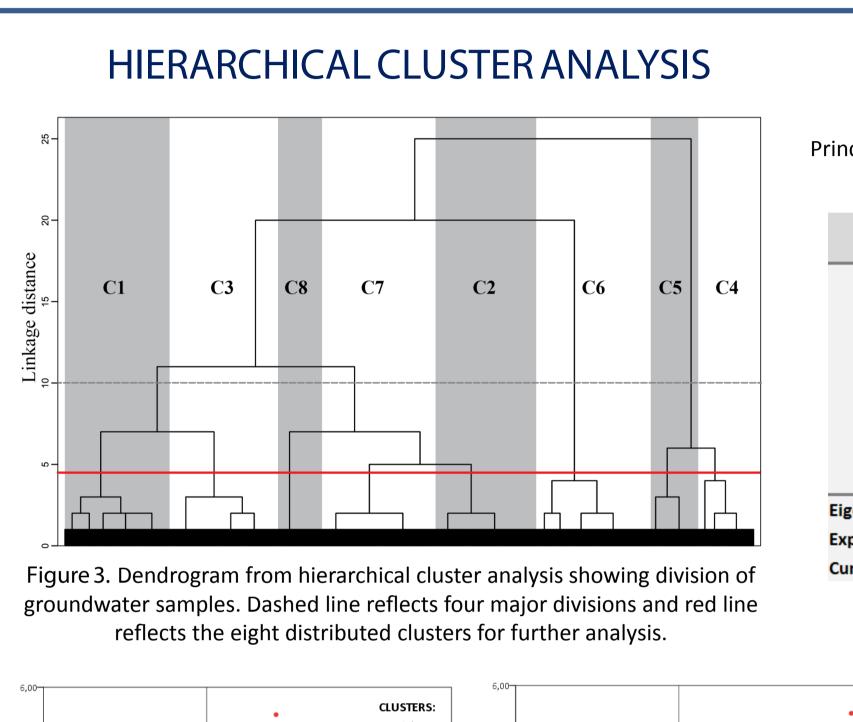


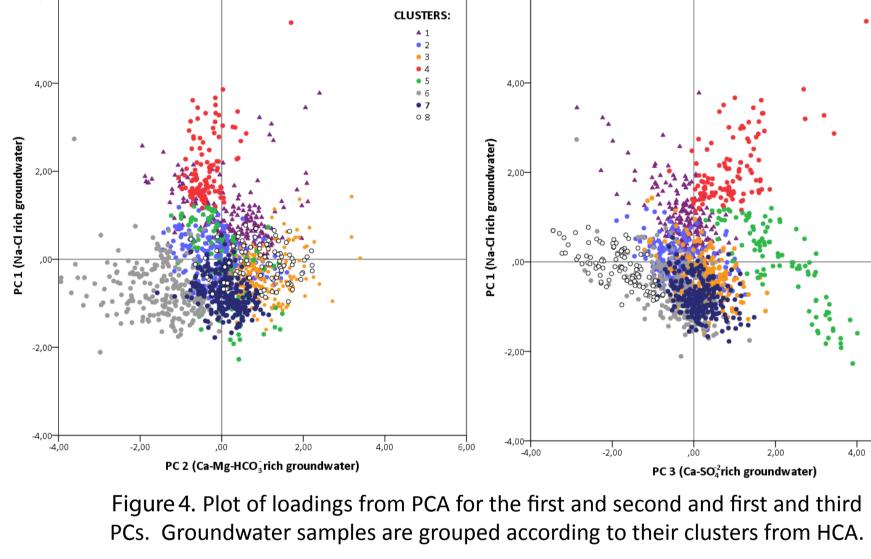


Hydrogeochemical characteristics of groundwater in Latvia using multivariate statistical analysis Inga Retike ¹, Andis Kalvans ², Janis Bikse ¹, Konrads Popovs ¹ and Alise Babre ¹

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Three hydrodynamical and hydrochemical





Eight geochemically distinct groundwater groups (C1-C8) can be observed characterised by particularly elevated or depressed major ion, trace elements and NO_3^- and NH_4^+ concentrations:

- C5 corresponds to gypsum dissolution in the active water exchange zone. concentrations in case of C2.
- result.

The results show that although trace elements and nitrogen compounds were not included in multivariate statistical analysis, their variance in groundwater can be observed by analyzing their composition within each of the subdivided groups based on major ion chemistry.

Acknowledgement

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RESULTS

| PRINCIPAL COMPONENT ANALYSIS Table pal component loadings and explained variance for three components with Varimax rotation. | | | | | |
|--|------|-------|-------|--|--|
| Parameter | PC 1 | PC 2 | PC 3 | | |
| Ca ²⁺ | 0.17 | 0.65 | 0.67 | | |
| Mg ²⁺ | 0.52 | 0.63 | 0.43 | | |
| Na⁺ | 0.92 | 0.10 | 0.18 | | |
| K⁺ | 0.82 | 0.14 | 0.12 | | |
| HCO3 ⁻ | 0.01 | 0.93 | -0.14 | | |
| CI | 0.78 | 0.00 | 0.38 | | |
| SO 4 ²⁻ | 0.33 | -0.10 | 0.88 | | |
| nvalue | 3.7 | 1.4 | 0.8 | | |
| ained variance (%) | 52.7 | 20.1 | 11.3 | | |
| ulative % of variance | 52.7 | 72.7 | 84.0 | | |

PC 1 reflects the Na-Cl rich groundwater. This type of water is generally found starting from middle-lower Devonian aquifers till Cambrian aquifer. Na-Cl rich groundwater also can be a result of freshwater mixing with saline water in fracture zones and with seawater close to the coastline.

PC 2 is defined as Ca-Mg-HCO₃ water type and is the most common groundwater type in active water exchange zone.

PC 3 account for gypsum dissolution process and reflects Ca-SO₄ water type. This type of water is characteristic to areas where gypsum is encountered in upper and lower Devonian aquifers, however, is also found in other parts of the active water exchange

The main objective was to examine characteristic trace elements in each of the distributed groundwater groups and to propose an insight in major geochemical processes responsible for evolution of each group.

| | Main geochemical and hydrog | | | | |
|--------------------------|-----------------------------|-------------------------------|-----------------------|---------------------|--|
| Cluster (sample size) | | Positive PC scores | Average depth, (m) | Aquifer n | |
| C1 (N=218) | | PC1 and PC2 | 4-100 | sandston dolomit | |
| C2 (N=213) | ullet | PC2; PC1 or PC3 | 4-90 | sandston dolor | |
| C3 (N=223) | ightarrow | PC2 and PC3 or PC2 and PC1 | 2-7 | till, sand, o | |
| C4 (N=115) | | PC1 and PC3 | 65-170 | sands | |
| C5 (N=98) | ullet | PC3 | 15-100 | sandstone, gyps | |
| C6 (N=242) | igle | none or PC3 | 3-15 | sand, sar | |
| C7 (N=240) | ullet | PC2 or PC2 and PC3 | 3–50 | sand, san dolon | |
| C8 (N=93) | 0 | PC2 or PC2 and PC1 | 25-75 | sandstone, | |
| | | | | | |

CONCLUSIONS

C6 is interpreted as recharge water not yet equilibrated with most of the sediment forming minerals.

C3 is interpreted as groundwater form water table aquifers affected by diffuse agricultural pollution.

Groundwater in C4 reflects brine or seawater mixing with fresh bicarbonate groundwater.

C7 and C2 belong to typical bicarbonate groundwater resulting from calcite and dolomite weathering with slightly elevated K⁺

Extremely low Cl⁻ and SO₄²⁻ are observed in C8 and interpreted as pre-industrial groundwater or solely carbonate weathering

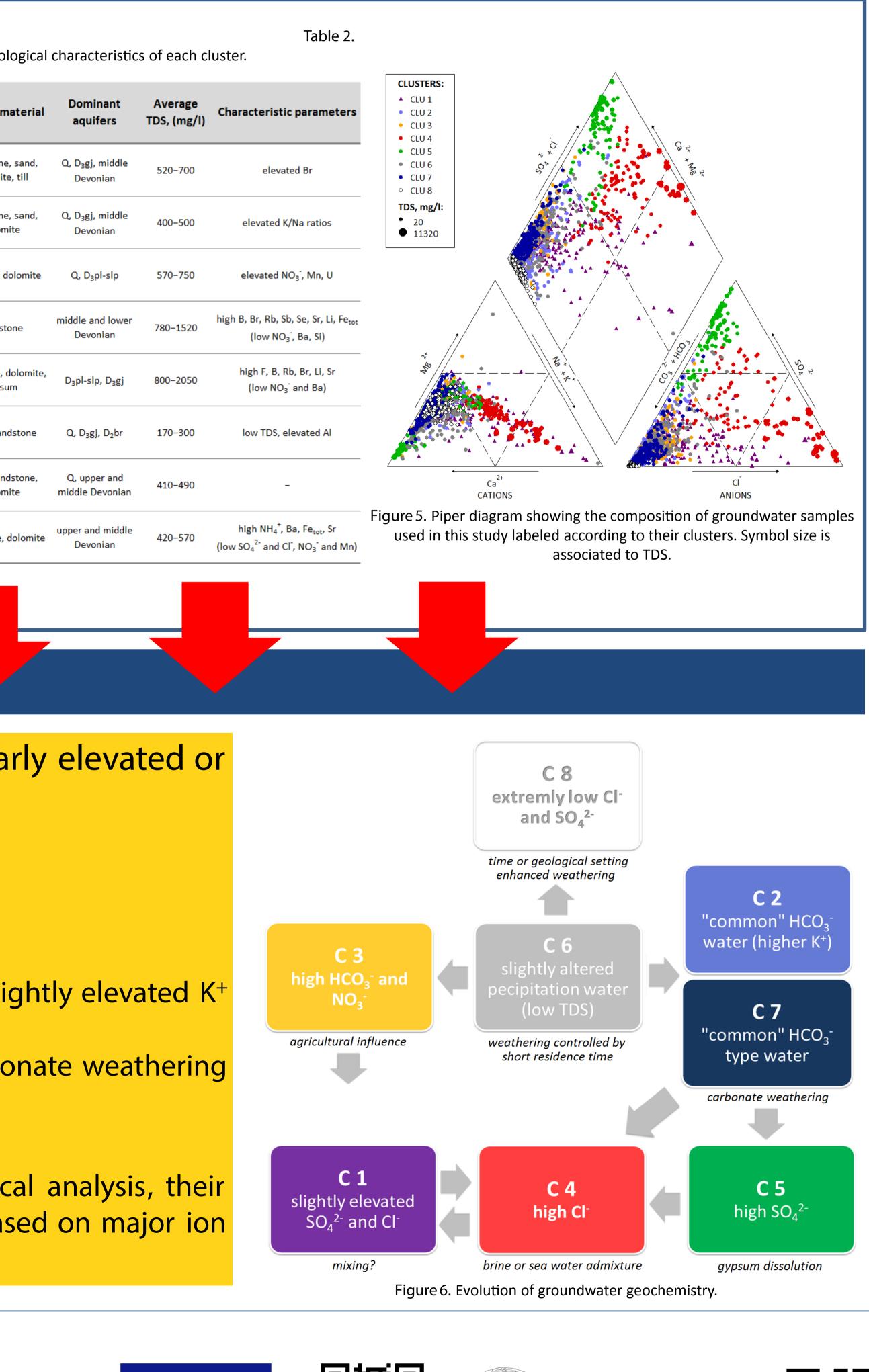
C1 seems to be a poorly definite subgroup resulting from mixing between other groups.



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