

PECHORA-KARA SEAS COASTS HYDROMETEOROLOGICAL STRESS EVOLUTION AND INTENSIFICATION IN RECENT 35 YEARS Natalia Shabanova, Stanislav Ogorodov Moscow State University, Geoecology of the North Laboratory

I. Arctic coasts retreat and quite rapidly





Arctic coasts retreat rate is about 0.5 - 2 m/year. If the coastal cliff is composed of ice-rich material or contains (ice bodies (ice wedges or chunks, massive ground ice) the coastal destruction may reach 5 – 10 m/year.

II. In recent years some coasts displayed accelerated erosion

Reference	Region	Method	Erosion ra periods)	ites, m/yec	ar (averaged	for the indi	cated
Jones,	Alyaska,	Satellite		1955-79	1979-2002		2002-2007
2009	Beaufort Sea	images		6.8	8.7		13.6
Vasiliev, 2011	Western Yamal, Kara Sea	Direct field observations		1978-81 ~1.2	1988-1992 1.7-3.5	1997-2002 ~1	2006-2010 2.5
Grigoriev, 2006	Eastern Siberian Seas	Literature sources and others	~1935-45 ~6-7	~1955-65 ~3	~1985-1995 ~5-6	The beg. of 2000s ~2	

All the authors indicate the increased retreat rates of coastal erosion for different Arctic regions. Some authors indicate also that there were periods of heightened values in about 1985 -1995 and after 2002 and low values in about 1995 - 2002

III. Arctic coastal erosion depends mainly on hydrometeorological factors





Kara Sea stations year mean temperature with 11year running mean



DO CHANGING WAVES, ICE AND TEMPERATURE COOPERATE IN COASTAL RETREAT ACCELERATION OR COMPENSATE EACH OTHER ?

IV. Climate changes and drives coastal dynamics

V. Methods

The observation data from Varandey and Marresalya polar stations are used

The amount of heat added to or extracted from the ground and the permafrost during the warm and cold periods is assessed by **air** thawing and air freezing index $(I_{at}$ and I_{af} respectively) showing the number of positive/negative °C·days per vear (Andersland and Ladanyi, 2004)

Residual mass curve method for periods detection

$$X_N_i = \frac{X_i / \overline{X} - 1}{C_v} = \frac{X_i - \overline{X}}{\delta}$$
$$X_R M_i = \sum_{i=1}^{i} X_N_i$$

High-value period $X_i > \overline{X} \rightarrow X_N_i > 0 \rightarrow X_RM_i$ ascends \checkmark

Low-value period $X_i < \overline{X} \rightarrow X_N_i < 0 \rightarrow X_RM_i$ decends

Popov-Sovershaev method for wave energy assessment (Popov, Sovershaev, 1982)

Deep water conditions

 $E = 3 \times 10^{-6} V^3 x \times p \times n,$

where V – wind velocity, m/sec;

x – wave fetch, km,

p - wave-dangerous winds frequency, n – ice-free period duration, days.

Shallow water conditions

$$E = 2 \times 10^{-6} \left(\frac{gH}{V^2}\right)^{1.4} V^3 x \times p \times n$$

References

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Grigoriev, M.N., Razumov, S.O., Kunitzky, V.V. and Spektor, V.B., 2006. Dinamika Beregov Vostochnykh Arkticheskikh Morei Rossii: Osnovnye Faktory, Zakonomernosti i Tendencii (Dynamics of the Russian East Arctic Sea coast: major factors, regularities and tendencies), Kriosfera Zemli (Earth's Cryosphere) 10(4), 74-94.

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VI. Two periods of increase: 1979 - 1995, 2003 - 2011(...)

Hydrometeorological factors evolution. Varandey



VII. Wave energy and ice-free period

Varandey. Wind wave energy flux got 72% higher due to both atmospheric circulation change and 16% due to ice-free period extension: energetic stress has increased by ~90%

Marresalya. Wind wave energy flux could get 62% higher due to ice-free period extension in case of stable atmosphere conditions. But due to AC change the increase amounts only 50%.









Contacts

Natalia Shabanova Nat.Volobuyeva@gmail.com



Stanislav Ogorodov Ogorodov@aha.ru

