

Late Holocene Palaeotsunami Events Archived along the Gujarat Coast, Western India

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extered and mobilized these boulders to their inland final position. Using optical dating technique, the age of deposition of the dune on which those boulders to their inland. Using various physical parameters and mobilized these boulders to the dune on which those boulders to the dune on which those boulders were lying was estimated to be 3.4 to 23 ka. the suggests the tsunami event took place sometime during the last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of about 1.5 m was dug at an elevation of 2 m from high tide line. This sequence shows a typical tidal flat sedimentation comprising silty – clayey layers (unit - 1 to unit 6). However unit 6 and unit 4 were sandy in nature and supported their deposition in form of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of about 1.5 m was dug at an elevation of 2 m from high tide line. This sequence shows a typical tidal flat sedimentation comprising silty – clayey layers (unit - 1 to unit 6). However unit 6 and unit 4 were sandy in nature and supported their deposition in form of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similarly at the Mundra coastline of Kachchh, a shallow trench of a high energy marine last 3.4 ka. Similar at the Mundra coastline of Kachchh, a mudclasts, broken shell fragments and layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of lore that a tsunami has been recorded historically around 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait of Hormoz, Iran. This sand layer (Unit-5) most likely is geological signature of the 1008 AD in Strait Hormoz Tsunami.



For Saurashtra coast

- We investigated a total of 210 boulders and recorded their physical dimensions and shapes along with their distance from high tide line.
- The boulders were present in scattered, imbricated and embedded manner along the entire southwestern Saurashtra coastline.
- On the basis of imbrication directions of these boulders, it is suggested that the source of tsunami wave lies in southwestern direction, most likely the OFZ (Prizomwala et al.2015).
- Numerical stimulations using Pignatelli et al. (2009) suggested that 'to dislodge all the boulders from their offshore position and to place them to their present landward position' would require a storm with a minimum wave height of 17.1 m.
- Whereas the same boulders can all be placed to their present position by a tsunami wave of 3.5 m.
- We compared the wave heights calculated from our data with strongest storm to hit the west coast of India as well as with strongest probable modelled storm in the Arabian Sea.
- Strongest storm along Fig 2a) west coast of India (May 2001) Strongest probable storm ---ibajnia et al., 2010) Chorwad Adri -20 -10 0 10 20 30 40 50 60 Distance from the high tide line (m) Chorwad Adri 30180280380480Distance from the high tide line (m)
- Fig 2 a) is the plot of the wave height at the final position of boulders with the distance and comparison of it with highest storm that the coast of Saurashtra has experienced along with strongest probable storm in Arabian Sea and b) a tsunami wave computed from these boulders. • An OSL age of 3.4 ± 0.3 ka was estimated from dune on which these boulders were scattered which implies that this event took place sometime during last 3.4 ka period i.e. Late Holocene Period.

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Results

For Kachchh coast

•The presence of mudballs, assorted nature of fragments and an abrastional contact with both suggests the origin of unit-5 as tsunamigenic, co which has landward dipping foreset ripples, mo account of ebb storm surge wave.

•Geochemistry of this sequence showed that the s and Unit-5) shows depletion in major oxides an concentrations. The concentration of CaCO₃ shows and a simultaneous decrease in Unit-3. This is resul to presence of broken shell fragments in Unit-5 and Unit-3.

•We presume deposition of Unit-5 to be on account wave, which might have eroded shell/bioclastic ric Kachchh and deposited it on landward side.

•Similar to CaCO₃ concentrations the Unit-3 shows l and Cr compared to Unit-5. The higher amount of testifies its origin from offshore sands which are no elements compared to fluvial sands of Kachchh. Si shows enrichment in Sr concentrations compared concentrations in Unit-5 is comparable to Sr conce silt units, which points at their alike provenance.

•The Mollusc shells dated from bottom of the unitserve as older age bracket to our sequence.

	METHODOLOGY	
abian Sea which hosts multiple Subduction Zone (MSZ), Owen ridge. An earthquake with generate tsunami which may	 For Boulder Deposits, Saurashtra coastline Measurements of the boulder dimension namely the long (a), intermediate (b) and shorter (c) axis were recorded in addition to their distance from the high tide line using Bosch GLM 250VF Professional distometer. 	 For Sand Deposits, Kachchh coastline Geomorphological maps of various segments of Kachchh and Saurashtra coast were prepared using Survey of India topographic sheets on 1:50,000 scale and high resolution satellite images.
Iowever the knowledge of past is ' <i>Terra Incognita</i> '. I both sandy (Kachchh) as well each in Kachchh coast (Fig 1b) dri shows presence of a jointed imestone, which acts as source (Fig 1c). Shows scattered, imbricated as (Fig 1d) along the coastline gnated Kachchh region, which as a segment with higher dflat along Mundra region we ched between the mud layers	 Samples were collected for estimating the bulk density in laboratory following equation ρ_{bi} = ρ_{sw}* W_a/(W_a-W_f) Where, ρ_{bi} is the bulk density of boulder; ρ_{sw} is the density of sea water = 1.02 g/cm3; W_a and W_f are the weights of the sample in air and in sea water respectively We used equations suggested by Pignatelli et al. (2009) for joint bounded boulder transport for storm and tsunami waves to estimate the minimum wave height required to dislodge and transport these boulders H_i ≥ [0.5c (ρ_s - ρ_w/ρ_w)]/CL H_i ≥ [2c(ρ_s - ρ_w/ρ_w)]/CL Where, H_i is the tsunami wave height and H_s is the storm wave height at breaking point; a, b and c are the boulder long, intermediate and short axis length; ρ_w is the sea water density=1.02; ρ_s is the boulder's density=2.75; CL is the coefficient of lift=0.178. Simultaneously samples were collected from the sand dune on which these boulders were scattered for estimating the age of the event using OSL chronology technique. 	 We then designated various landforms along the coastline as per their response to the tsunami waves: For example (i)river mouths, creeks, tidal channels etc. act as 'conveyor' which allow the tsunami waves inland and cause wide destruction up to considerable distance (ii) sandy beaches and mangrove swamps absorbs the energy and velocity of tsunami waves. The extent to which such absorber landform would reduce the energy of a tsunami wave depends on height and width of the landform (iii) palaeomudflats, swales, back swamps etc. act as 'accommodators' which accommodate most of tsunami wave energy and material carried by it. Accommodator has a high preservational potential for tsunami deposits (iv)landforms like beach ridges, cliffs and coastal dunes act as a 'barrier' to the tsunami wave by obstructing it inland. The capacity of such barrier will depend upon height of tsunami surge and elevation and width of the barrier. The remote sensing study was followed by field investigation near Mundra region of Kachchh in which we dug a trench of about 1.5 m depth in a palaeomudflats environment which is 3m high from the present sea level. The sedimentary section was studied in detail for its sedimentological attributes and sampled at 2 cm interval. The samples packed in polythene bags and brought to laboratory where they were subjected to geochemical analysis (major and trace elements).
		Conclusion
sand, broken shell trom mud horizon ompared to unit-3 ostly deposited on sand layers (Unit-3 nd trace elements s increase in Unit-5 lted most likely due hlack of the same in th sea bed of Gulf of U_{11} and $U_$		 The coastal archives of Gujarat provides a fascinating scenario to study and decouple the geological signatures of palaeotsunami and palaeostorm deposits. Our present study suggest a tsunami wave of 3.5 m wave height at the coastline of Saurashtra dislodged boulders from there offshore position and deposited them inlands. The event took place sometime during the last 3.4 ka. Based on boulder imbrications, the Owen Fracture Zone is believed to be source of this event. Another tsunami event took place around 997-1107 AD. This is archived in form of a sand layer (Unit-5) sandwiched between mud layers of Mundra coast, Kachchh. We also for the first time report evidences of a palaeostorm deposit in form of Unit-3 which was deposited on account of high flux of sediment during ebb-tide
lesser amount of Zr Zr and Cr in Unit-5		References
rmally rich in these imilarly, the Unit-5 d to Unit-3. The Sr entrations in clayey 5 unit gave an AMS ¹⁴ C age of AD	$\frac{1}{150} \int_{150}^{1} \int_{150}^{1} \int_{150}^{1} \int_{100}^{1} \int_{100}^{1} \int_{100}^{100} \int_{200}^{200} \int_{100}^{1} \int_{100}^{1} \int_{200}^{1} \int_{100}^{1} \int_{100}^{1} \int_{200}^{1} \int_{100}^{1} \int_{10}^{1} \int_{100}^{1} \int_{10}^{1} \int_{10}^{1} \int_{10}^{1} \int_{10}$	 Pignatelli C, Sanso P, Mastronuzzi G (2009) Evaluation of tsunami flooding using geomorphologic evidence. Mar 331 Geol 260(1-4):6-18. 332. Prizomwala S.P., Gandhi D., Ukey V.M., Bhatt N. & Rastogi B.K. (2015) Coastal boulders as evidences of high-energy marine events from Diu Island, west coast of India: storm or palaeotsunami? Natural Hazards, 75(2):1187-1203. Ramasamy S.M., Kumanan C.J., Saravanavel J. & Selvakumar R. (2006) Geosystem responses to December 26, 2004, Tsunami and mitigation strategies for Cuddalore Naganattinam coast Tamil Nadu Jour Geol Soc of India 68:967-983





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