

## **Petrology and deformation style of lithospheric mantle beneath the Heldburg Dike swarm (Central Germany) subset of Central European Volcanic Province**

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The Heldburg Dike swarm is a set of Cenozoic alkali basalt dikes occurring in the central part of Germany at the border between Thuringia and Bavaria. We studied xenoliths from Strauf, Feldstein, Bramberg and from the active quarry in Zeilberg. The peridotites from Strauf, Feldstein and Bramberg have the composition of spinel lherzolite (15), spinel harzburgite (9) and dunite (3). They vary in size from 1.5 cm (Strauf) up to 20 cm (Zeilberg).

We distinguish groups (A, A- and B) of peridotites based on different forsterite content in olivine. Group A consists of olivine (89.6 – 91.8 Fo), orthopyroxene (Mg# 0.90–0.93, Al 0.05–0.18 a pfu), clinopyroxene (Mg# 0.87–0.95, Al 0.06–0.26 a pfu) and spinel (Cr# 0.13–0.65, Mg# 0.54–0.78). Clinopyroxene rare earth elements (REE) patterns are S-shaped (Feldstein, Bramberg) or U-shaped (Strauf); spoon-shaped patterns occur occasionally. Trace element (TE) patterns show negative Nb, Ta, Zr, Hf, Ti and positive Th, U anomalies. The most magnesian clinopyroxene (xenolith 3140, Feldstein) is strongly aluminous and LREE depleted with weak anomalies in TE patterns. Group A- contains olivine (88.9–89.5 Fo), orthopyroxene (Mg# 0.89–0.90, Al 0.10–0.13 a pfu) and clinopyroxene (Mg# 0.90–0.92, Al 0.10–0.17 a pfu). Clinopyroxene is increasingly enriched in REEs from Lu to La. TE patterns are similar to those of group A but with less pronounced anomalies. Group B (3 xenoliths only) consists of olivine Fo 86.7–88.9, orthopyroxene (Mg# 0.88–0.89, Al 0.07–0.19 a pfu), clinopyroxene (Mg# 0.88–0.90, Al 0.10–0.26 a pfu). Clinopyroxene is enriched in LREE, concave upward in Pr. TE patterns are similar to those in group A. One of group B harzburgites contains grains (up to 0.5 mm) of Ca-Mg carbonate located in interstices.

The clinopyroxene chemical composition plots away from the melting trend in the MgO-Al<sub>2</sub>O<sub>3</sub> diagram of Upton et al. (2011), suggesting a later addition of the clinopyroxene. The composition of orthopyroxene corresponds to ca. 15–30 % of melting of primitive mantle, which was overprinted by silicate and/or carbonatite metasomatism. The xenolith 3140 seems not to be affected by metasomatic overprint.

Based on the EBSD analyses of 15 xenoliths, olivine grains are characterized by relatively strong CPO (crystal preferred orientation) with J indices 4.4 – 13.3, and they have orthorhombic (8 xenoliths) or [100]-fiber CPO (6 xenoliths) symmetries except for one [010]-fiber symmetry observed in group B (Tommasi et al., 1999). Pyroxenes have weaker CPO and the distribution of their crystallographic axes is inconsistent with their coeval deformation with olivine. We propose that their CPO postdates that of olivine, hence strongly support a later origin for pyroxenes.

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