

Detection of shallow marine cumulus convection with airborne and spaceborne lidar systems over the tropical North Atlantic Ocean

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Motivation

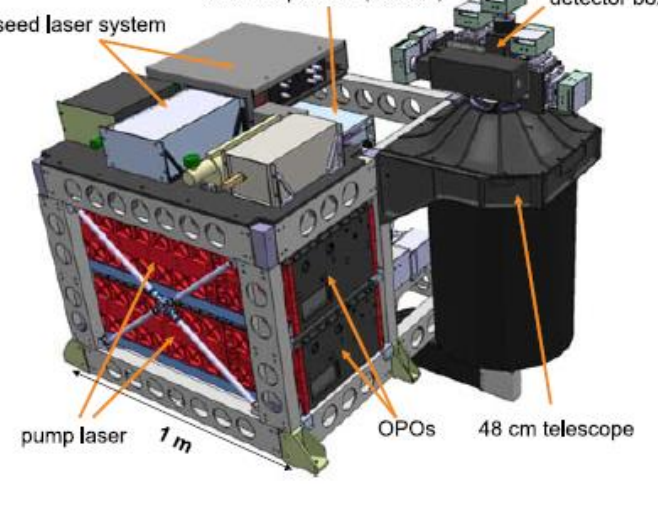
- Differences in the representation of shallow marine convection (SMC) in global climate models lead to large differences in climate sensitivity estimates (Bony and Dufresne (2005))
- Cumulus clouds are often of smaller extent than grid spacing of General Circulation Models → parametrizations needed

→ Study small scale shallow marine clouds with the advanced airborne lidar system WALES

Scientific questions:

- What are typical distributions of vertical and horizontal cloud extent in the subtropical Atlantic dry season and wet season?
- How accurate can CALIOP, a spaceborne lidar system, resolve small scale shallow cloud regimes in contrast to WALES?


Method and Instrumentation



WALES

- Combined DIAL and HSRL system onboard the German HALO aircraft
- Resolution: 15 m (vertical), ~40 m (horizontal)

Wirth et al. (2009), Esselborn et al. (2008)



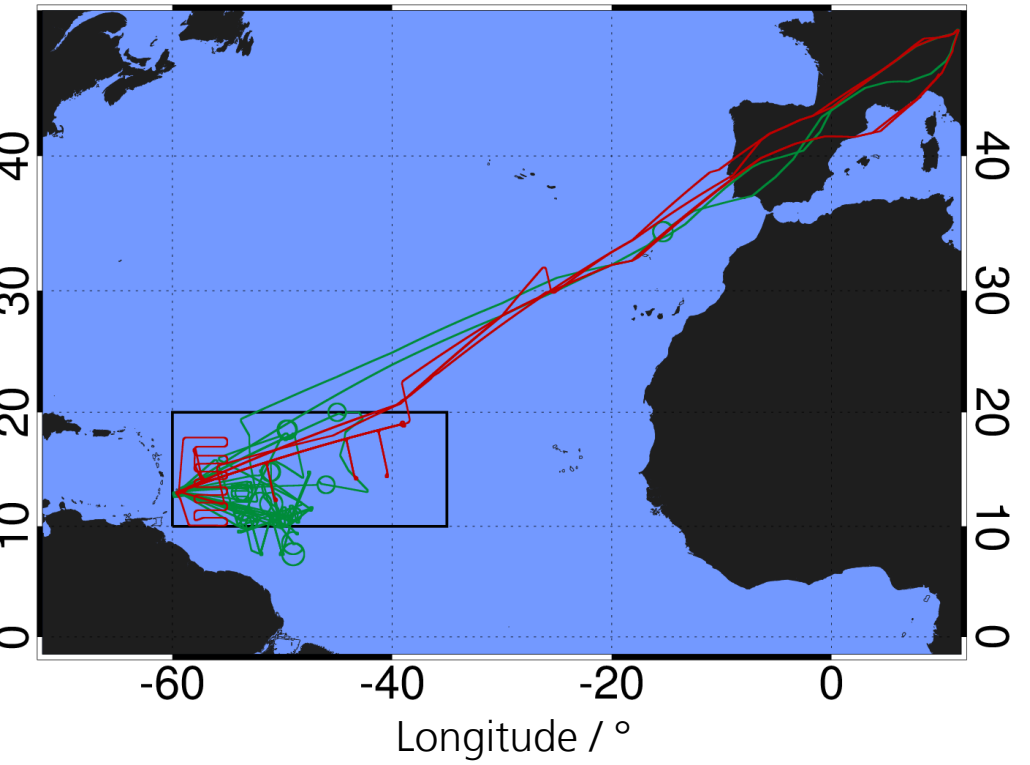
CALIOP

- Backscatter Lidar on-board CALIPSO
- Resolution: 30 m (vertical), 330 m (horizontal)

Winker et al. (2007)


This study: WALES HSRL measurements (532 nm)
CALIOP Backscatter Lidar measurements (532 nm)

NARVAL campaigns



NARVAL I

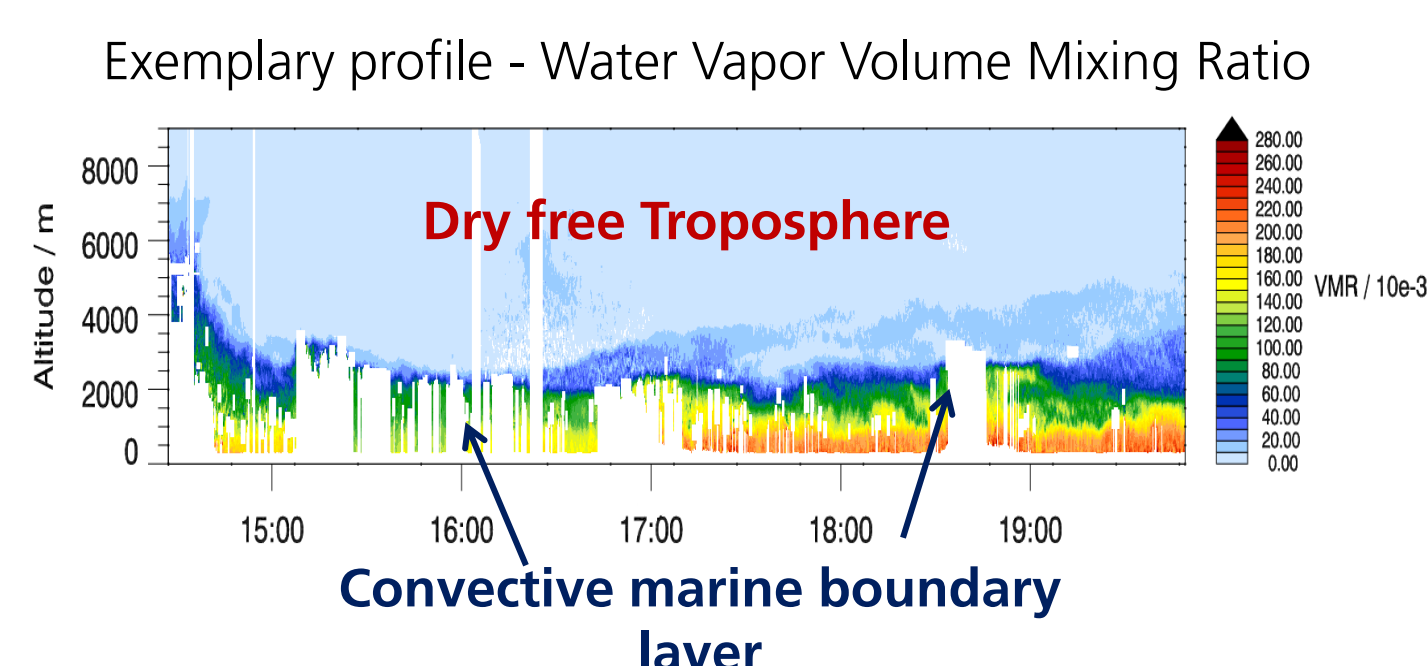
- 10 to 20 Dec 2013 - **dry season**
- 65 flight hours within 8 research flights
- 6 CALIOP underflights over the subtropical North Atlantic Ocean



NARVAL II

- 8 to 30 Aug 2016 - **wet season**
- 85 flight hours within 10 research flights
- 5 CALIOP underflights over the subtropical North Atlantic Ocean

NARVAL I – dry season



- Clean and very dry free troposphere in winter months
- Sharp moisture gradient at trade wind inversion (2000–2500 m)

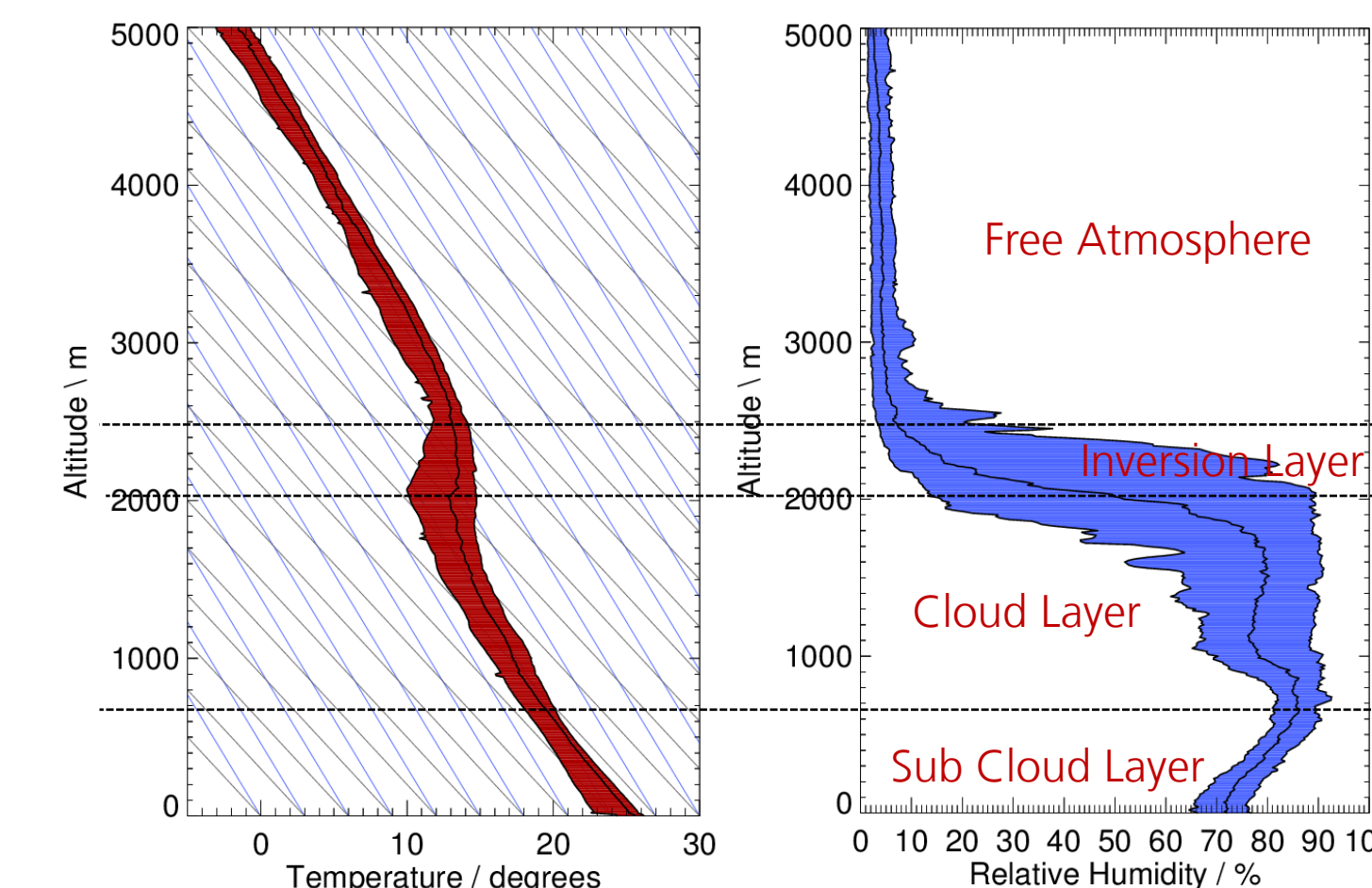
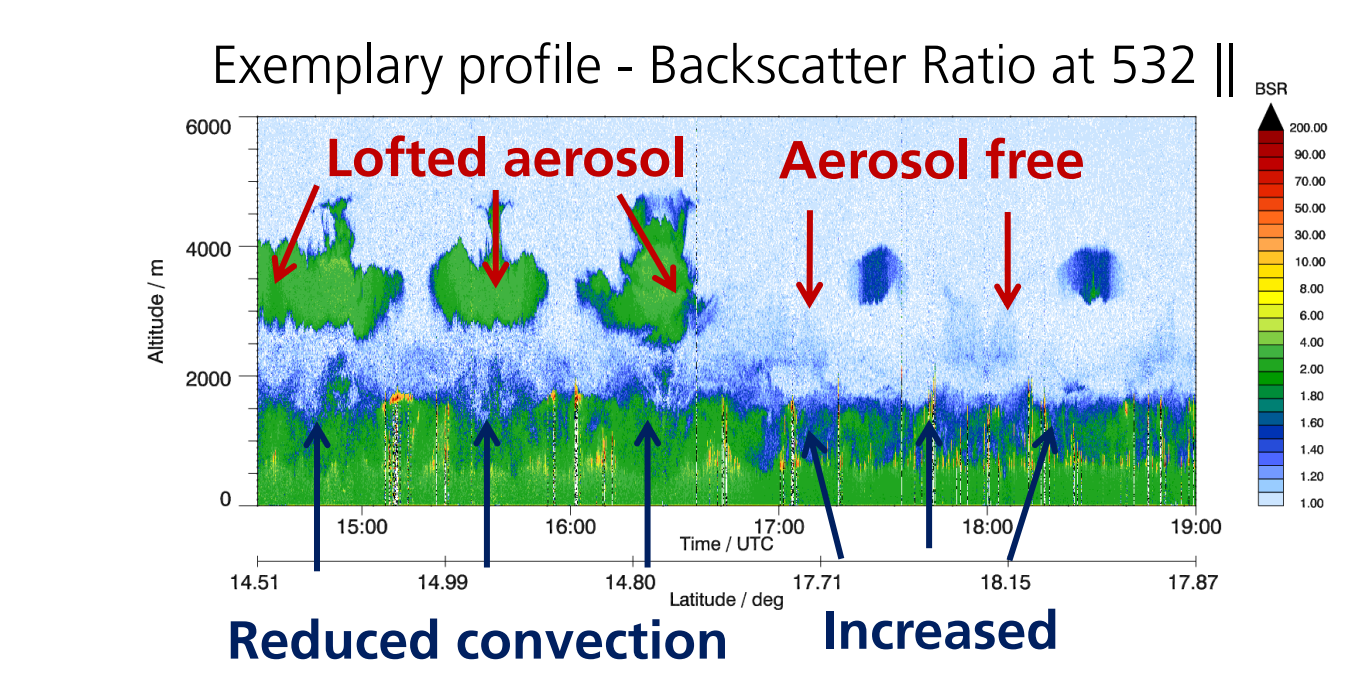


Figure: Temperature and relative humidity profiles from all dropped radiosondes during NARVAL I (shaded areas: Interquartile Range).

NARVAL II – wet season



- Moist and often aerosol loaded lower free atmosphere
- Weakly pronounced moisture gradient at trade wind inversion (1500 – 1800 m)

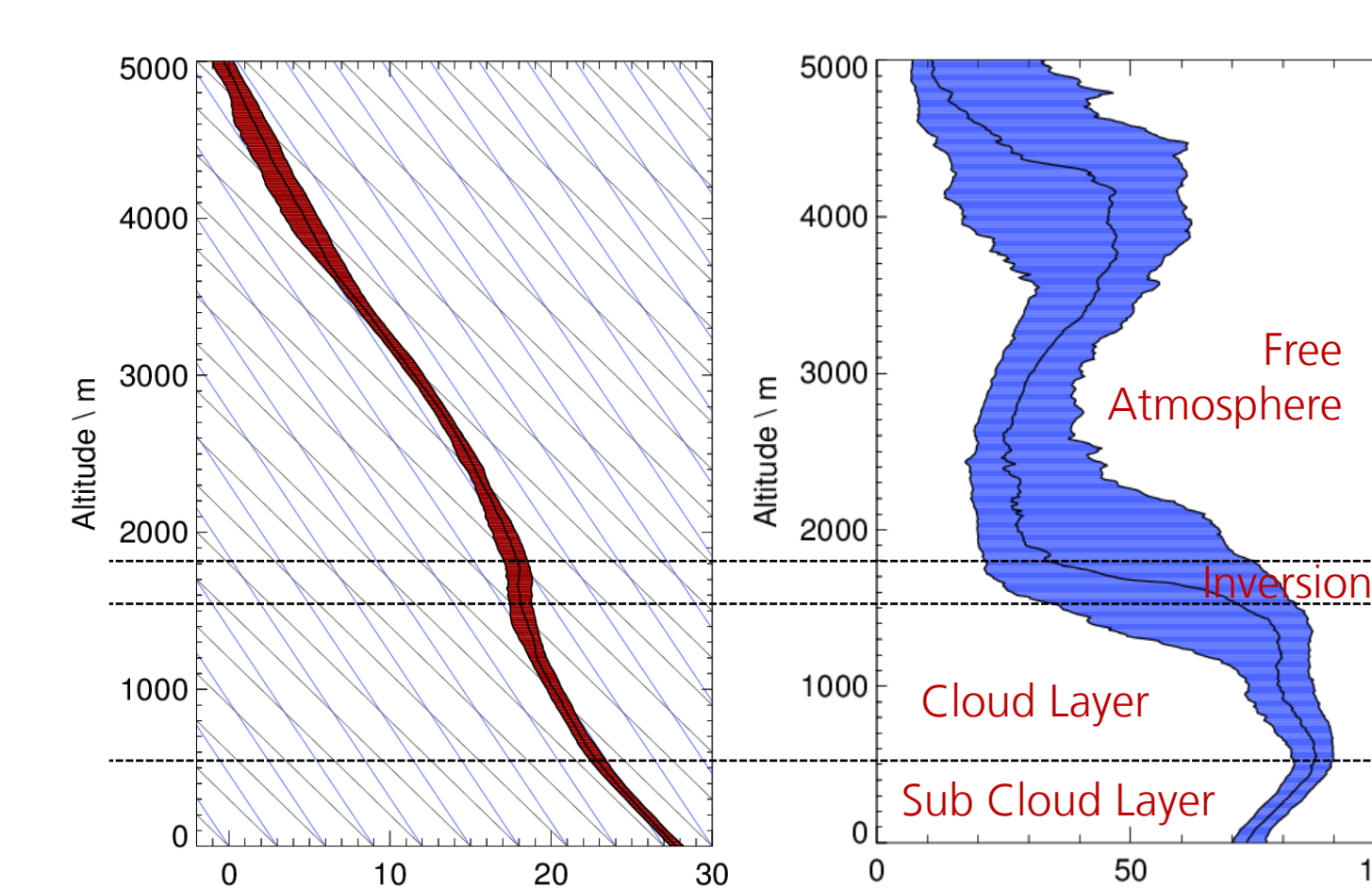


Figure: Temperature and relative humidity profiles from all dropped radiosondes during NARVAL II (shaded areas: Interquartile Range).

WALES vs. CALIOP

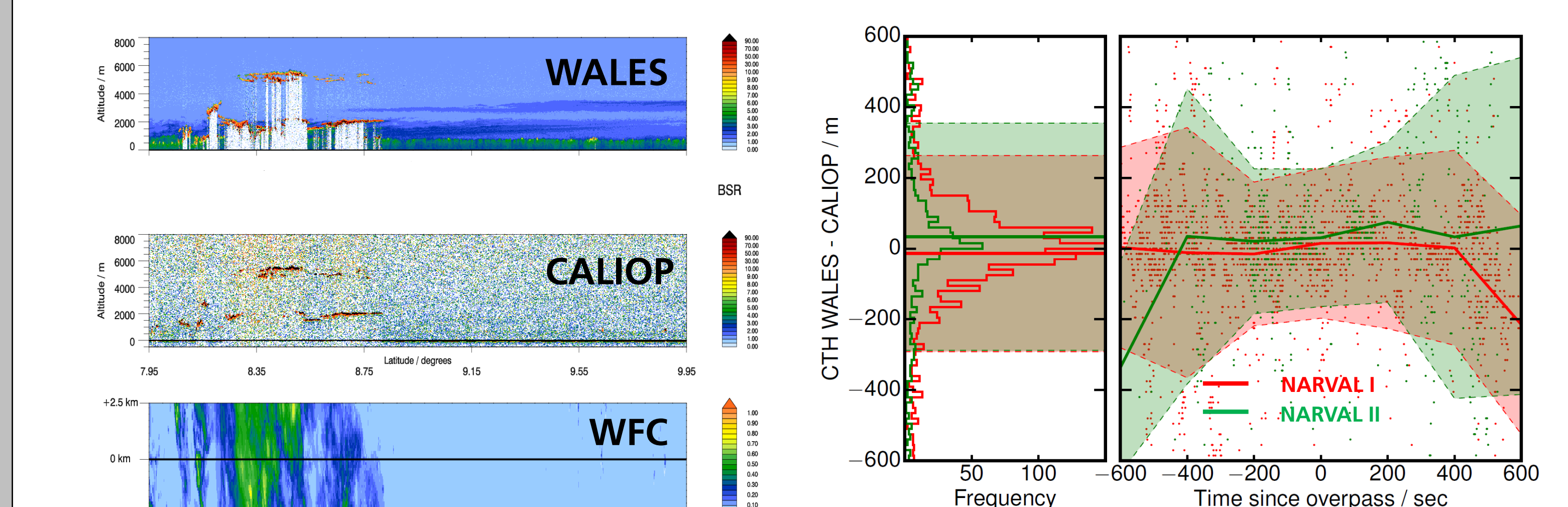


Figure: Exemplary underflight on 15 Aug 2016 (WFC: Wide Field Camera onboard CALIPSO)

Figure: Distributions of differences in detected cloud top heights (CTH) from CALIOP and WALES from all underflights (shaded: 1σ).

- Good agreement of cloud structure and cloud top frequency in WALES and CALIOP underflight measurements
- Differences in detected cloud top heights of CALIOP and WALES are smallest, when the temporal separation between the two instruments is small.

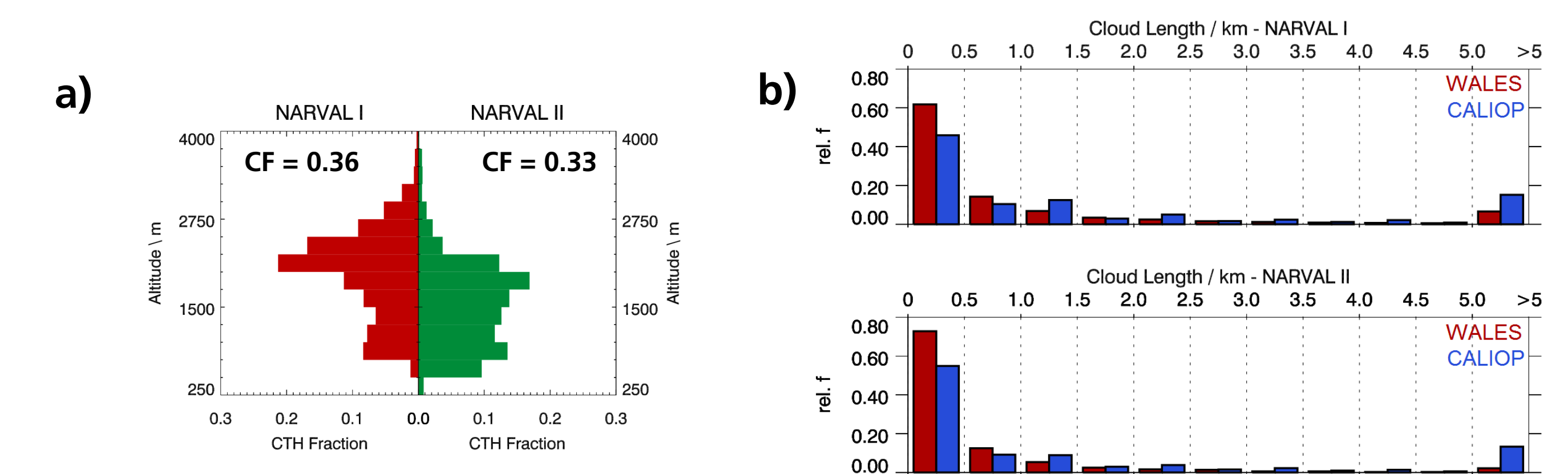


Figure: Distributions of cloud top heights deduced from CALIOP measurements (a); Cloud length distributions derived from WALES and CALIOP measurements (b).

- CALIOP measurements underestimate small cloud amount. This may be caused by missing information on the inter-profile region between two consecutive CALIOP lidar shots.

Cloud Statistics: NARVAL I vs. NARVAL II

- Threshold of $BSR = 90$ is defined for single layer cloud detection
- Decrease in cloud top height in wet season compared to dry season

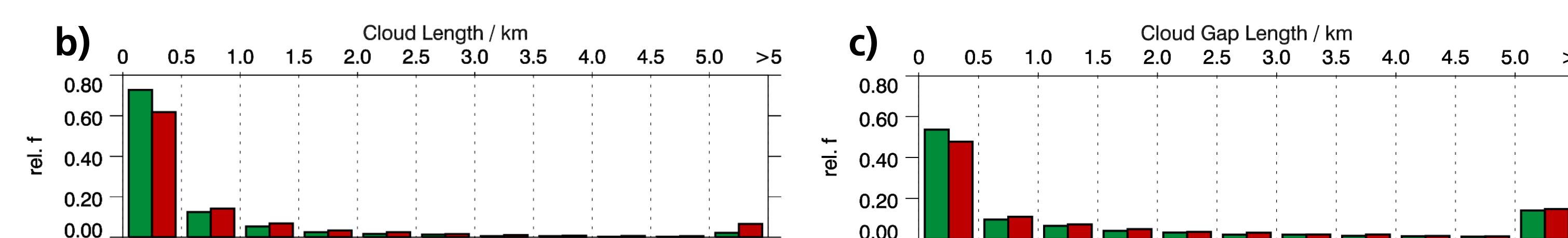


Figure: Distributions of cloud top heights (a), cloud lengths (b) and cloud gap lengths (c) deduced from WALES measurements during NARVAL I (red) and NARVAL II (green).

References

Bony and Dufresne, 2005: Marine boundary layer clouds at the heart of tropical cloud feedback uncertainties in climate models, *Geophysical Research Letters*, 32 | Esselborn et al., 2008: Airborne high spectral resolution lidar for measuring aerosol extinction and backscatter coefficients. *Appl. Opt.*, 47 (3) | Wirth et al., 2009: The airborne multi-wavelength water vapor differential absorption lidar WALES: system design and performance. *Applied Physics B*, 96 (1) | Winker et al., 2007: Initial performance assessment of CALIOP. *Geophys. Res. Lett.*, 34 (19).