



SIMS oxygen isotope analyses of white mica from the Larderello geothermal field reveal a complex magmatic – hydrothermal petrogenesis

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The Larderello geothermal field (LGF) in southern Tuscany is underlain by a series of shallow intrusions of isotopically and geochemically distinct two-mica granites belonging to the Miocene-Pleistocene Tuscan Magmatic Province (1). Muscovite is ubiquitous in the granites, the contact metamorphic aureole and the deep pre-Alpine metasedimentary units that constitute the lower reservoir of the LGF. Texturally and chemically distinct generations of muscovite record stages of magmatic crystallization, thermometamorphic/hydrothermal replacement and fluid/rock interaction (2,3).

Existing O isotope data of whole rock and mineral separates indicates that granite and contact metamorphic rocks have been gradually depleted in ^{18}O ($\delta^{18}\text{O}$ from 5-12‰ [4]). *Although some secondary minerals, such as muscovite, proved highly retentive and retained their pre- alteration $\delta^{18}\text{O}$ signature (8-10‰), biotite in the alteration zone can have $\delta^{18}\text{O}$ values as low as 1.5‰ (3).*

In-situ Secondary Ion Mass Spectrometry (SIMS) O isotope analysis of white mica is used to resolve the fluid/mineral interaction on microscale in the dynamic environment of the LGF. Data of white mica from granite, contact metamorphic micaschist and phyllite show a large and variable spectrum of $\delta^{18}\text{O}$ values ranging from 14‰ to as low as 2‰. The highest $\delta^{18}\text{O}$ data ($\leq 14‰$) are found in chemically homogeneous white mica from a phyllite sample, which are close to values from unaltered micaschist from the Northern Appennines (4). White mica from a contact metamorphic micaschist shows the largest spread of $\delta^{18}\text{O}$ values, from 12‰ in homogeneous domains in larger grains to 2‰ in a more heterogeneous small grain. A granite and a high-temperature magmatic/hydrothermal vein contain white mica with a magmatic $\delta^{18}\text{O}$ of $\sim 9‰$ and higher values of $\sim 13‰$ respectively. In both samples, white mica show a pronounced *intra-grain* $\delta^{18}\text{O}$ variability of up to 5‰ either as core-rim zoning, or as ^{18}O -depletion in Fe-Mg-enriched domains (~ 2.5 wt.% higher in Fe-Mg) that surround inclusions of chlorite.

The highest measured $\delta^{18}\text{O}$ values from white mica in metasediments are either related to a pre-intrusive, syn-tectonic metamorphic overprint, or were caused by recrystallization from a more recent ^{18}O -rich fluid, as proposed for other secondary minerals (2). Large homogeneous domains in the igneous white mica retained their magmatic $\delta^{18}\text{O}$ values, whereas the selective ^{18}O -depletion in the Fe-Mg-enriched, recrystallized domains results from hydrothermal alteration by meteoric water that also affected some metasedimentary white mica (2,4). The new SIMS $\delta^{18}\text{O}$ data of white mica, therefore, reveal, for the first time, their multifaceted petrogenesis on a microscale, and may prove a valuable tool in further describing the fluid dynamics of complex magmatic-hydrothermal-metamorphic systems.

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