

INTRODUCTION

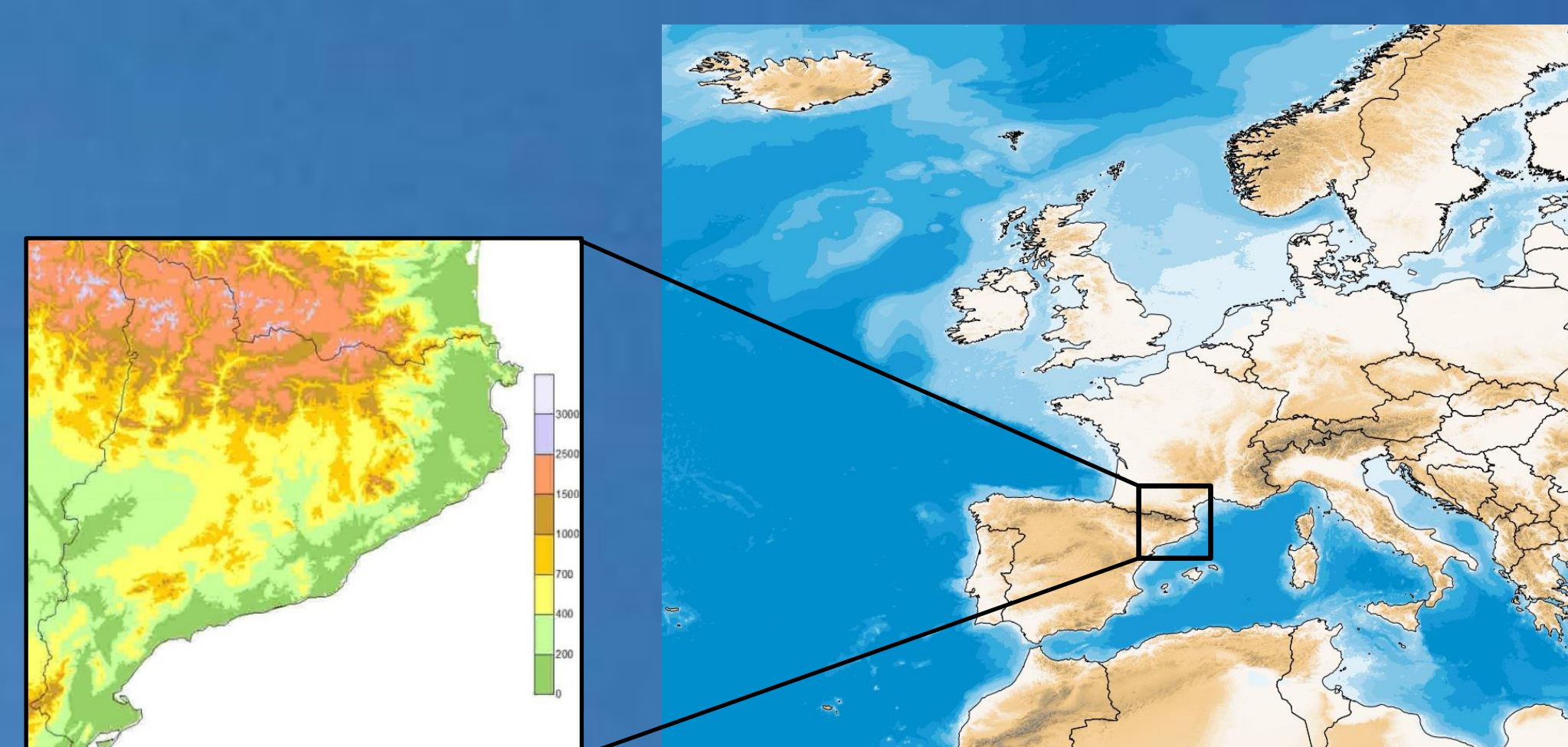
MAIN GOAL

Improve the calculation of the snow level line (SL) and determine which method works better in the different weather situations..

PREVIOUS WORK

- Asses the usage of the Wet Bulb Temperature (WBT) as an indicator to differentiate the atmospheric air masses leading to different types of precipitation.
- Evaluate three different methods aiming to obtain an accurate and improved snow level (SL) forecast

AREA OF STUDY



Altitude of Catalonia (m.a.s.l)

Fig. 1: Catalonia location

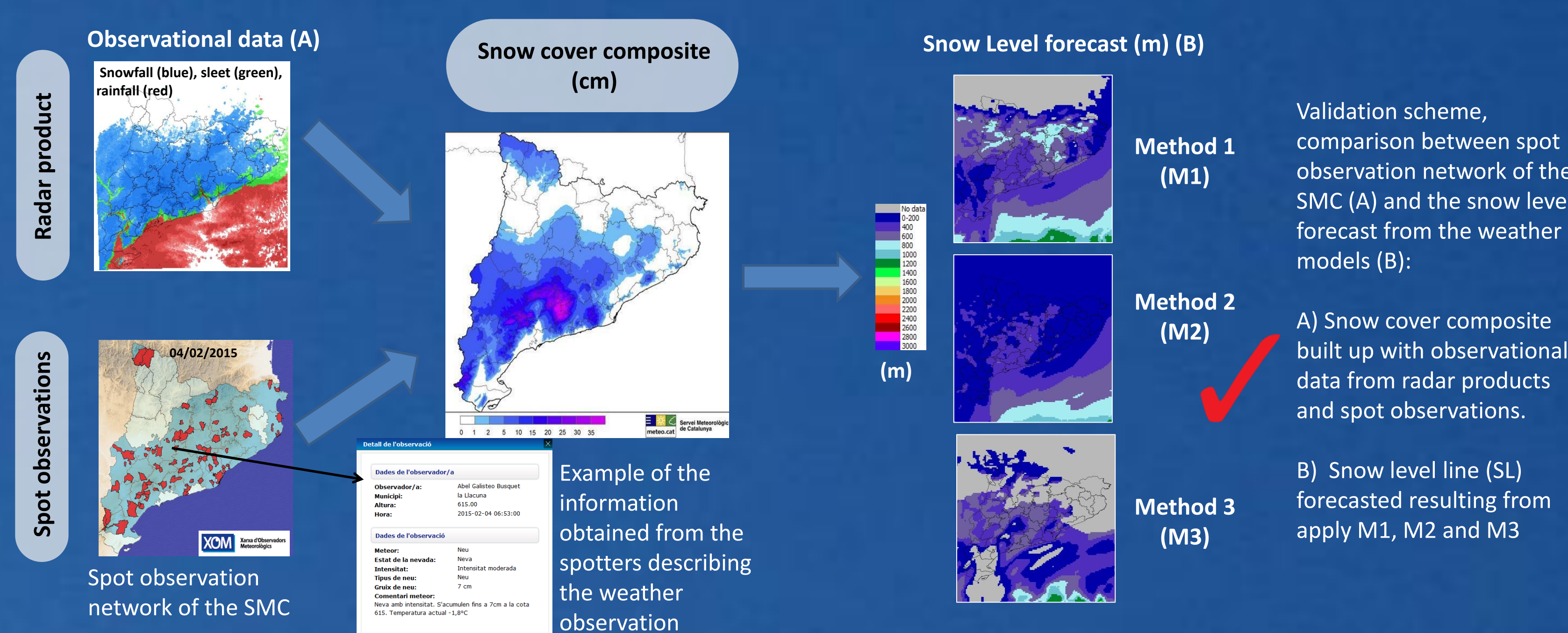
METHODOLOGY

3 methods are used to compared them in different synoptic situations.

Method 1 (M1): SL is defined using isozero altitude (m), and dropping 200m if HR>60%; and dropping 400m to the Wet Bulb Zero Level (WBZL) if HR<60%.

Method 2 (M2): SL is determined using the altitude (m) of the 1.5°C Wet Bulb Temperature (WBT1,5).

Method 3 (M3): SL is obtained from Lumb Curve (cooling by melting process is taken into account).



WARM FRONT

SYNOPTIC ANALYSIS

SPOT OBSERVATIONS

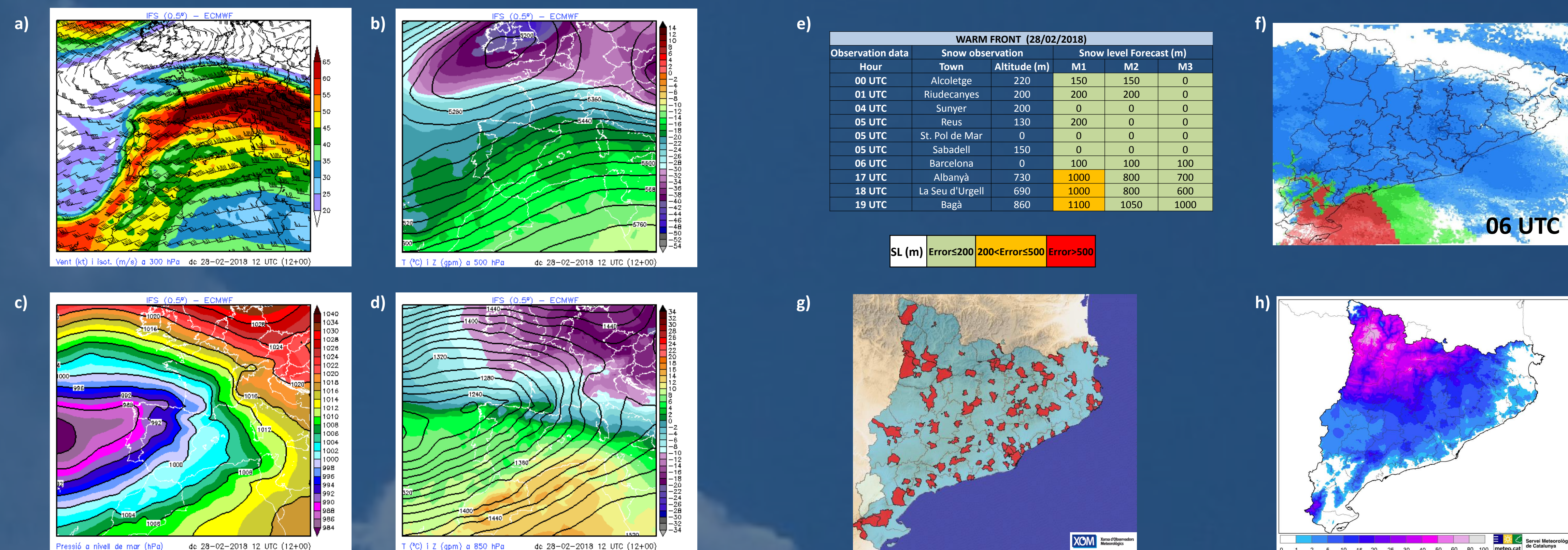


Figure. 2. a) Jetstream at 300hPa, b) T and geopotential at 500hPa, c) surface pressure, d) T and geopotential at 850 hPa.

Figure. 3. e) Table with the error of each method compared with observations (only showed 10 of 309 spot observation), f) radar product is shown with an estimation of rain, sleet or snow; g) spot observations; h) snow depth.

OCCCLUDED FRONT

SYNOPTIC ANALYSIS

SPOT OBSERVATIONS

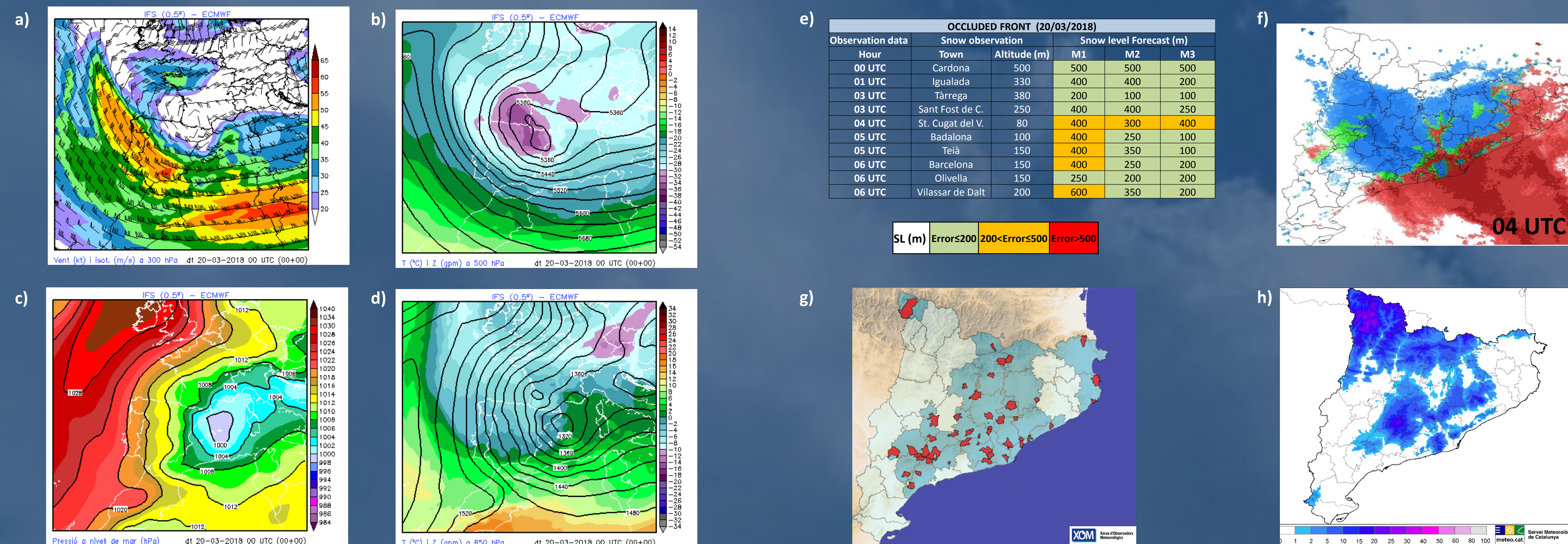


Figure. 6. a) Jetstream at 300hPa, b) T and geopotential at 500hPa, c) surface pressure, d) T and geopotential at 850 hPa.

Figure. 7. e) Table with the error of each method compared with observations (only showed 10 of 83 spot observation), f) radar product is shown with an estimation of rain, sleet or snow; g) spot observations; h) snow depth.

CONCLUSIONS

This work evaluates three different powerful methods to forecast the snow level line (SL) comparing to observational data in different synoptic weather situations.

- ✓ In warm front, models have a good response in general. However, models have some troubles in the valleys oriented to south when warm air arrives because of topography. In this event, M2 and M3 offer a better forecast. Independently, according to Kain et al. (2000) in cases with strong low-level temperature advection, M3 could drop more the snow level line than the observed.
- ✓ There are no significant differences between the methods in a cold front. All of them estimate precisely the snow level line, but there is a tendency to underestimate the SL in the leeward side of the Pyrenees, especially M1 and M3 if precipitation jumps to the south slope. This situations usually occurs when the Jetstream is perpendicular to the mountain range.
- ✓ The case of the occluded front is complex. The big challenge to this situations is interpret the isotherm layer created by the melting process, especially in the north-west quadrant of the cyclone. This episodes seem to be well-considered by M3 and M2, specially if precipitation is well forecasted.
- ✓ There are no doubts that in convection weather, the cold air is fallen down to lower altitudes. Numerical models have some difficulties to describe the cooling due to melting process. In this case, this phenomenon is well forecasted by M3, and by M2 if the method is complemented with the radiosonde forecasted (in order to see if an isotherm stratum exists).

COLD FRONT

SYNOPTIC ANALYSIS

SPOT OBSERVATIONS

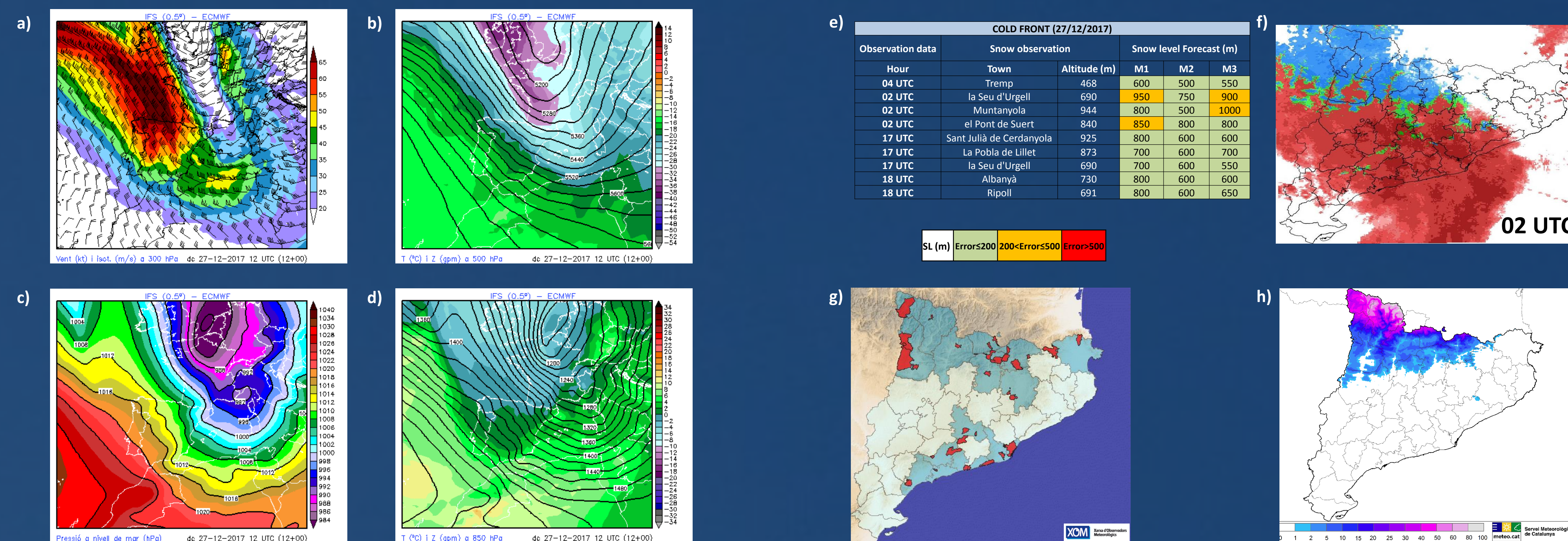


Figure. 4. a) Jetstream at 300hPa, b) T and geopotential at 500hPa, c) surface pressure, d) T and geopotential at 850 hPa

Figure. 5. e) Table with the error of each method compared with observations (only showed 9 of 53 spot observation), f) radar product is shown with an estimation of rain, sleet or snow; g) spot observations; h) snow depth.

CONVECTION

SYNOPTIC ANALYSIS

SPOT OBSERVATIONS

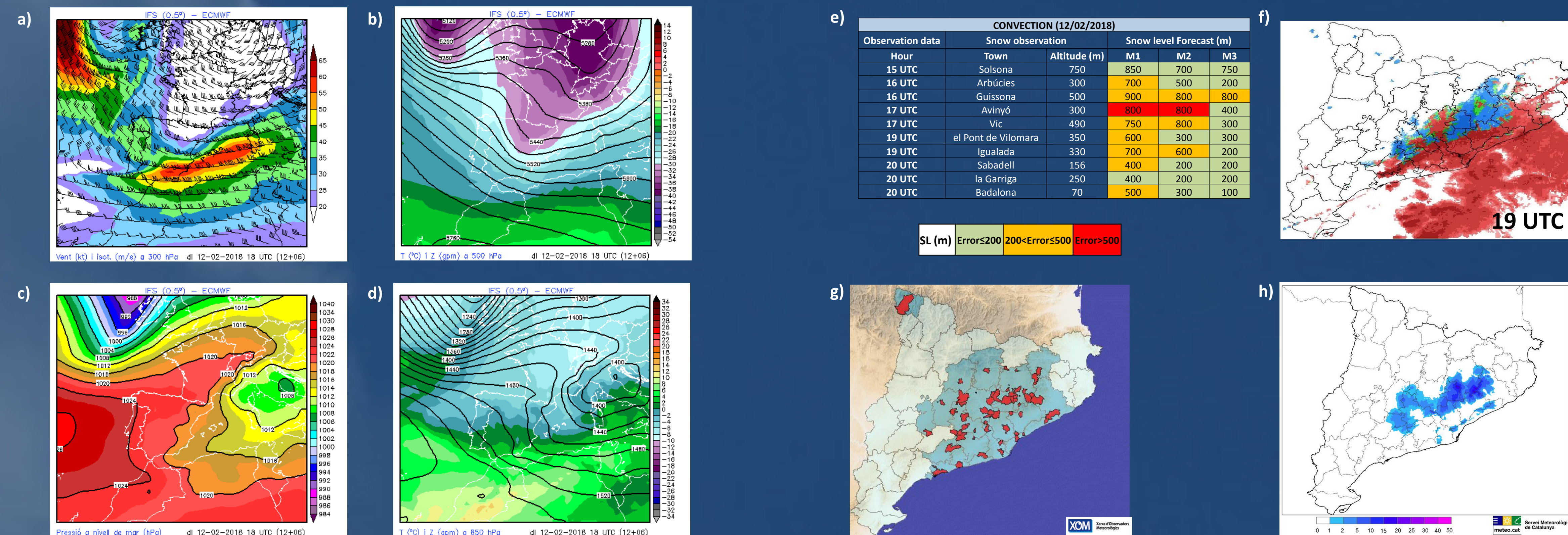


Figure. 8. a) Jetstream at 300hPa, b) T and geopotential at 500hPa, c) surface pressure, d) T and geopotential at 850 hPa

Figure. 9. e) Table with the error of each method compared with observations (only showed 10 of 95 spot observation), f) radar product is shown with an estimation of rain, sleet or snow; g) spot observations; h) snow depth.

FUTURE WORK

- Going on building a bigger data base in order to obtain a precise method to estimate snow level line.
- Further research on the different meteorological parameters characterizing the dry and wet snowfall to achieve a better forecast for the snow accumulation rate.
- Try to implement a procedure that would take into account the best method in each situation.

REFERENCES

- Matsuo, T. and Sasyo, Y, 1980a: Melting of snowflakes below freezing level in the atmosphere. Journal of the Meteorological Society of Japan, 59, No. 1, 10-25.
- Matsuo, T. and Sasyo, Y, 1980b: Non-melting phenomena of snowflakes observed in subsaturated air below freezing level'. Journal of the Meteorological Society of Japan. 59, No. 1, 26-32.
- Kain, J., 2000: The melting effect as a factor in precipitation-type forecasting. Weather and forecasting. No. 15, 700-714.
- Lumb, F. E.: 1961, 'The problem of forecasting: the downward penetration of snow', Meteor. Mag. 90, 310-319.
- K.K. Szeto and R.R. Stewart, 1997: Effects of Melting on Frontogenesis, Volume 54, No. 6, 689-702
- Mr W.S. Pike, 2001: Meteorologist's profile – Frank Emerson Lumb. Weather. 42-48

ACKNOWLEDGEMENTS

Paul Gibson (MetService New Zealand); Dr. Will Land (MetOffice); Sylvain Boutot (Environnement Canada); Lionel Peyraud (MétéoSuisse); Marina Ealo (BSC); Different departments of Servei Meteorològic de Catalunya (SMC), especially Jordi Moré, Alex Serra and Bel Martínez; and all the weather spotters members of the Meteorological Observations Network (XOM) of SMC.