

Late Holocene environmental trends in the center of the Russian Plain

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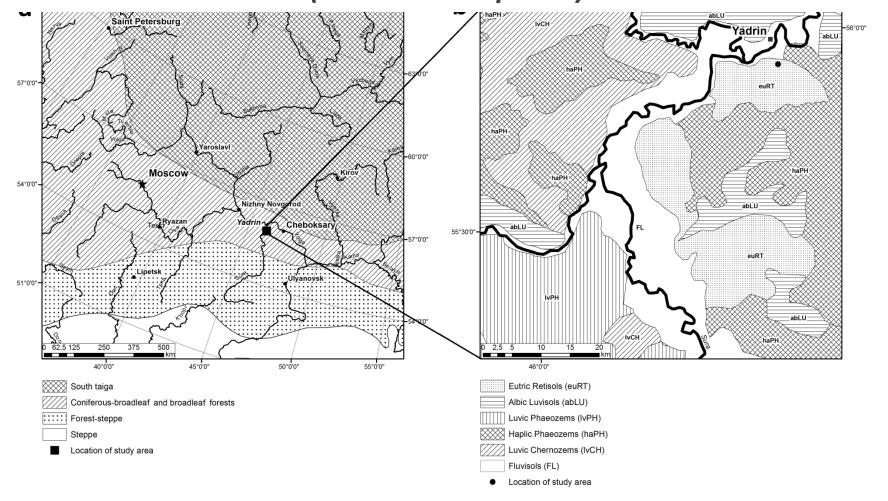


Central Forest-Steppe,

Lipetsk region

Starting from Dokuchaev time scientists suggested multiple shifts of steppe and forest landscapes based on intricate borders between Chernozems, Phaeozems and Luvisols

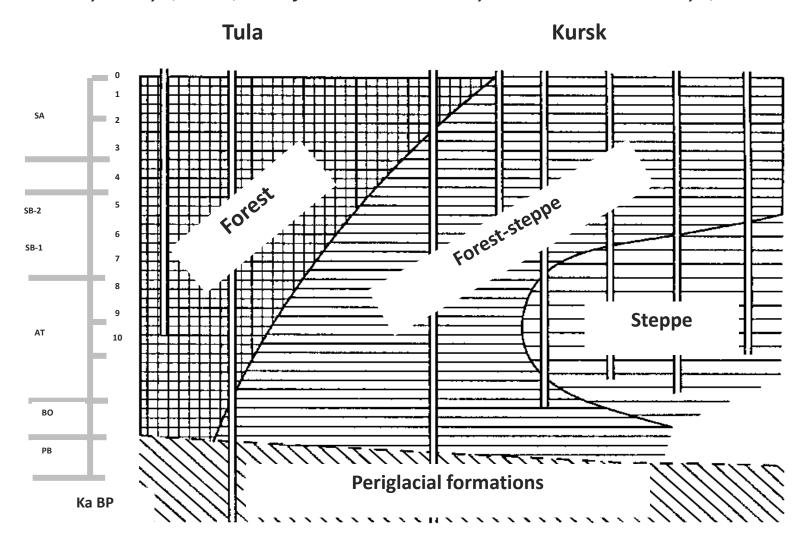
Complicated intersection of Retisols, Luvisols, Phaeozems and Chernozems at the Southern fringe of the Forest zone (Chuvashia Republic)

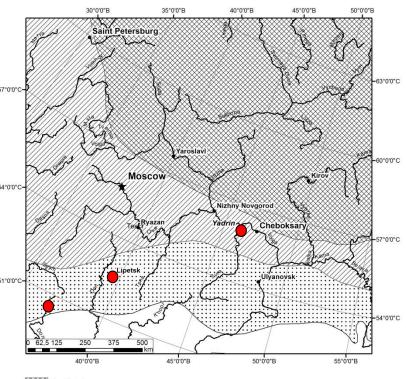


Holocene landscape shifts

based on palynological assemblages in peat of raised bogs,

Serebryannaya, 1992; modified: Alexandrovskiy and Alexandrovsakaya, 2005





South taiga
Coniferous-broadleaved and broadleaved forests
Forest-steppe
Steppe
Location of study area

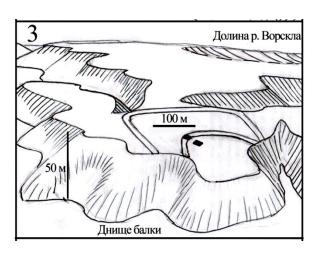


The Early Iron Age ~2500 yr BP



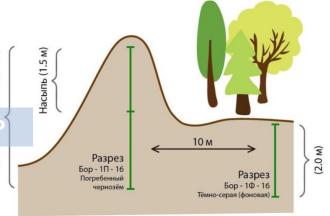
Sedentary tribes introduced fortifications with earth walls

Typical position – promontories of uplands between ancient gullies



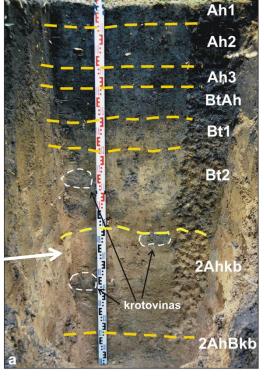
Buried surface 14C 2540±100 years Call Обороните ¹⁴C 6750±1**2**0 years CalBP Ahkb ABthkb1 ABthkb2 Btkb 2Ahkb 2AhBkb krotovinas

Southern Forest steppe, Belgorod region





Cambic Cryosol
Calcaric

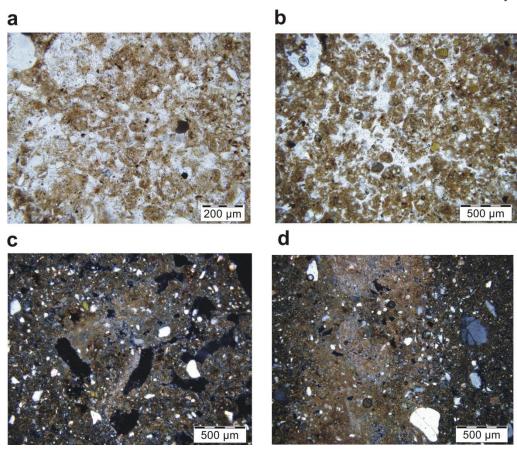


Buried soil – Chernic Luvic Phaeozem

Surface soil – Chernic Greyzemic Luvic Phaeozem

Pedogenetic stages recorded in soils of Southern Forest-Steppe Stage 1. Cryo-arid pedogenesis.

Late Pleistocene, MIS2 (?)



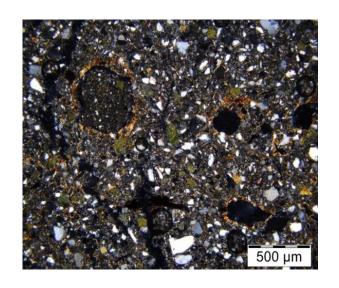
a, b - fine granular microfabric with rounded aggregates common for Upper Pleistocene paleosols

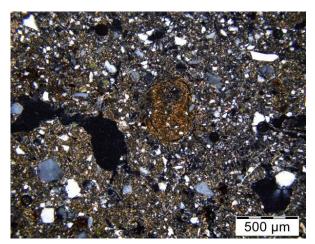
c, d- carbonate neoformations and crushed grains

Buried Cambic Cryosol in the layer of carbonate loess

(recorded both in surface and buried soils ~100 cm below the former surface)

Pedogenetic stages recorded in soils of Southern Forest-Steppe Stage 2. Forest pedogenesis. Early to mid-Holocene (?)





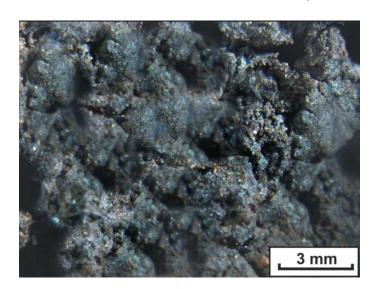
Argic horizon:

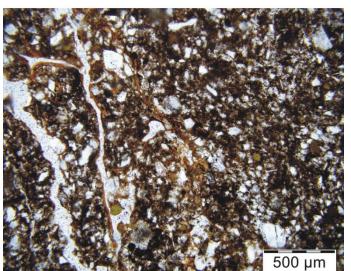
Subangular blocky peds

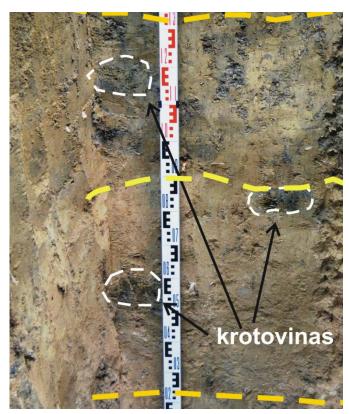
Multi-layered clay cutans in a sequence of the Bt horizons.

Pedogenetic stages recorded in soils of Southern Forest-Steppe Stage 3. Steppe pedogenesis.

(Holocene climatic optimum - ¹⁴C 6110±100 years CalBP)

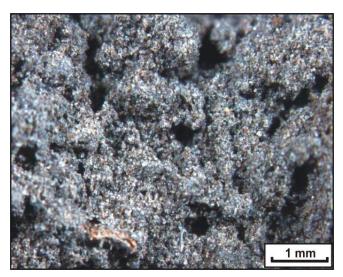


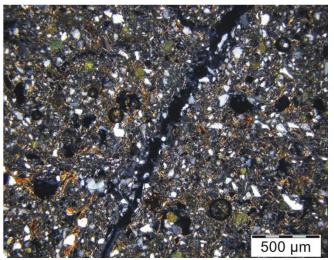


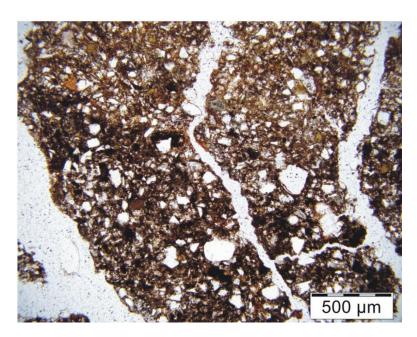


- Dark color, fine granular structure and high humus content
- Groundmass impregnated with dark humus
- Humus enriched in Humic acids $(C_{HA}/C_{FA} < 2)$
- Krotovinas

Pedogenetic stages recorded in soils of Southern Forest-Steppe Stage 4. Pedogenesis under broadleaf forests (Early Iron Age till present)





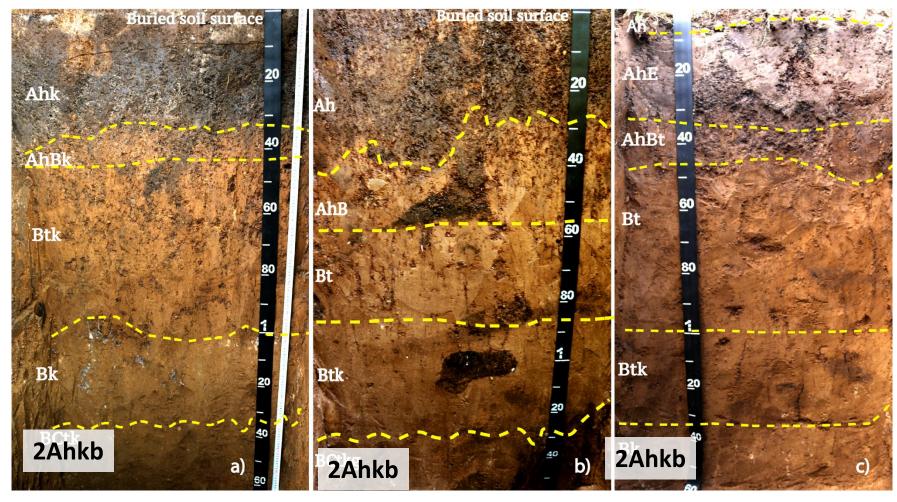


Degradation of the upper part of Chernic horizon

Greyzemic features and Albeluvic glossae in Ah horizon

Recorded in surface soil

Central Forest-Steppe, Lipetsk region

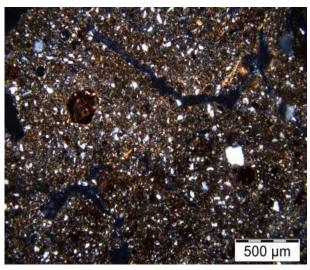


Buried soil
Greyzemic Luvic Phaeozem,
V-VI centuries BC

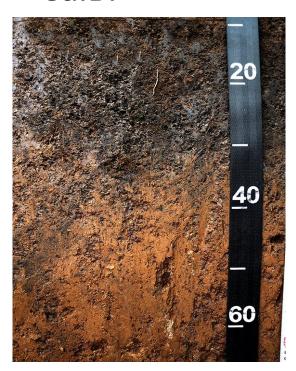
Buried soil Luvic Chernozem, V century AD Surface soil, Greyzemic Luvic Phaeozem

- Stage 1. Cryo-arid pedogenesis.
 - Late Pleistocene, MIS2 (?), 140 cm
- Stage 2. Forest pedogenesis Argic horizon
 - Early to mid-Holocene (?)





- Stage 3. Steppe pedogenesis dark Ah, krotovinas
 - Holocene climatic optimum ¹⁴C 5530±80 years
 CalBP

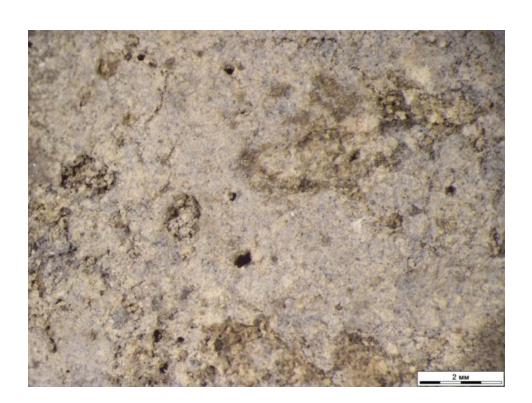




- Stage 4. Pedogenesis under broadleaf forests degradation of the upper part of Chernic horizon, Greyzemic features and Albeluvic glossae in Ah horizon
 - Since the Early Iron Age

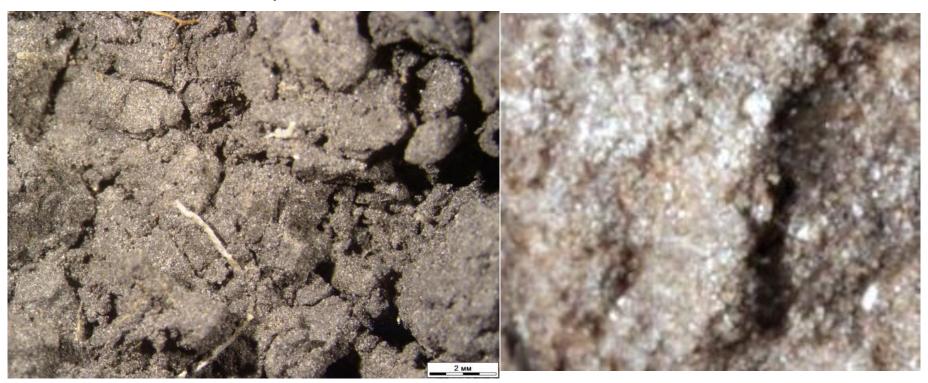


- Stage 5. Arid pedogenesis dark Ah, carbonate impregnation, carbonate films above clay cutans
 - V century AD

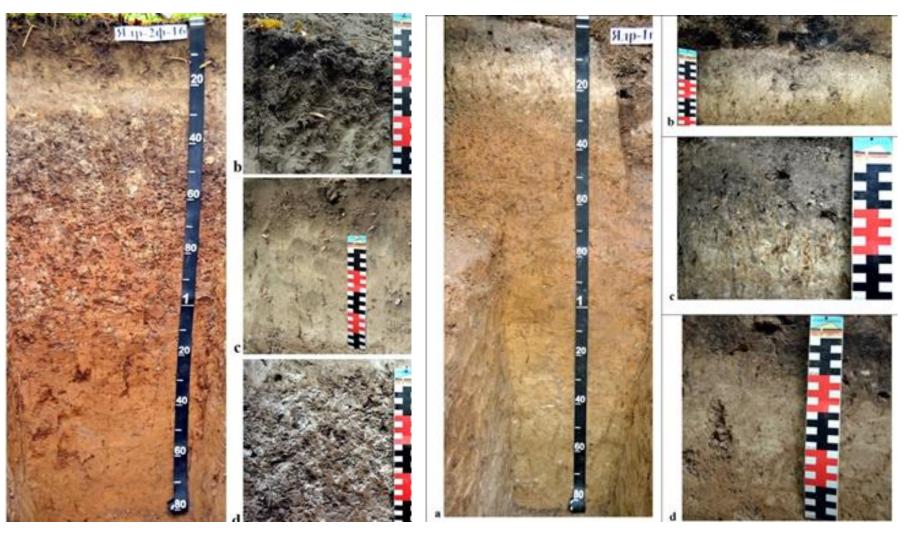


Carbonate films over clay cutans in the Bt horizon

- Stage 6. Pedogenesis under broadleaf forests degradation of the upper part of Chernic horizon, Greyzemic features and Albeluvic glossae in Ah horizon
 - The last 1500 years



Broadleaf forest, Chuvashia Republic Albic Retisols, left – surface soil; right – buried soil (14C 2068 CalBP)



Surface soil

Buried soil

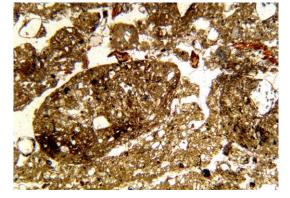
Pedogenetic stages recorded in soils of broadleaf forest Stage 1. Cryo-arid pedogenesis.

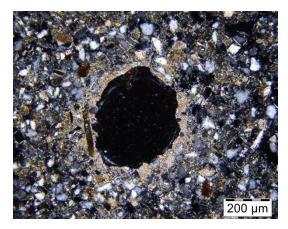
Late Pleistocene, MIS2 (?)



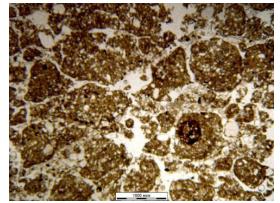
1 mm

Fine granular microfabric





Carbonate neoformations, crushed



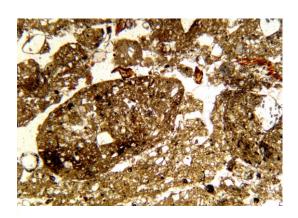
Cambic Turbic Cryosol, Central Yakutia

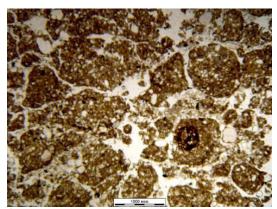
Buried Ca

(recorded both in surface and buried soils ~100 cm below the former surface)

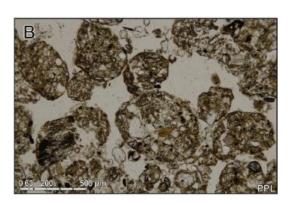
Pedogenetic stages recorded in soils of broadleaf forest Stage 1. Cryo-arid pedogenesis.

Late Pleistocene, MIS2 (?)





Cambic Turbic Cryosol, Central Yakutia





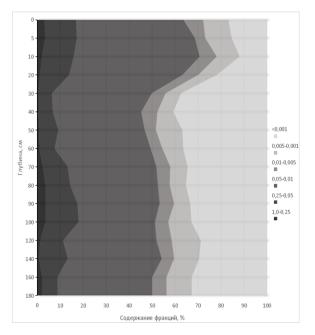
Microstructure of Stilfrid B paleosol (MIS3),

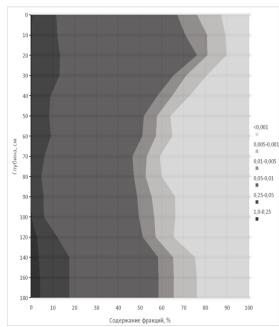
Terhorst et al., 2013

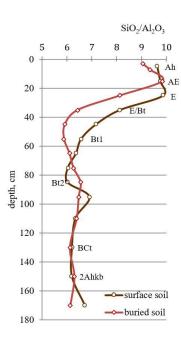
Pedogenetic stages recorded in soils of broadleaf forest

Stage 2. Forest pedogenesis (clay cutans, Retic properties, Albeluvic glossae)

Through the whole Holocene till present







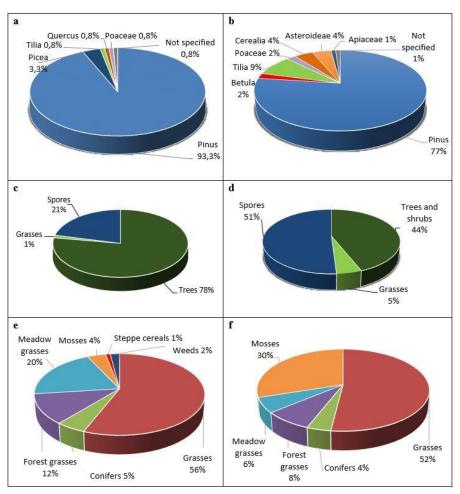
Buried soil Surface soil Grain size distribution pattern

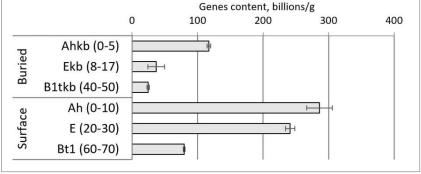
Micromorphology of Albic Retisols,

broadleaf forest, Chuvashia Republic

Surface soil Buried soil **EBt** Bt

Broadleaf forest, Chuvashia Republic

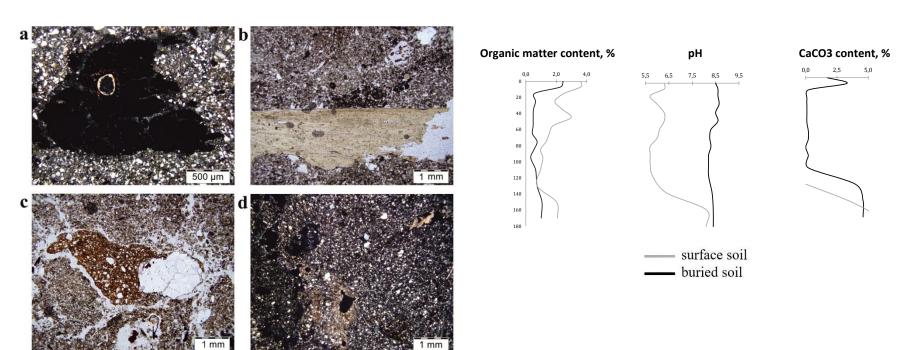




Bacterial 16S rRNA genes content in buried and surface soils estimated by qPCR.

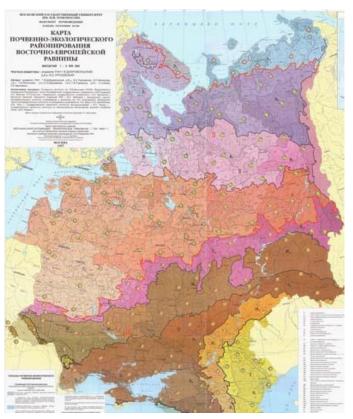
Microbiomorphic data for surface (left) and buried (right) soils. a, b – ratio of taxa in the spore-pollen spectrum; c, d – ratio of major groups in the spore-pollen spectrum; e, f – composition of the phytoliths complex.

Anthropogenic impact and diagenesis in the buried Retisol, broad-leaved forest, Chuvash Republic



- a Akhb horizon. Charcoal with the ring of diagenetic carbonates. 10II;
- b Ahkb horizon. Bone fragment. 4II;
- c AEkb horizon. Bone fragments, diagenetic carbonates and admixture of Bt horizon. 4X;
- d Ekb horizon. Diagenetic carbonates. 4X.

Shallow buried Late Pleistocene paleosols are widespread within the profile of surface soils



- Due to shallow depth of the upper loess layer on the uplands buried soils are included within the profile of surface soils
- The lower horizons (BC, C) are former Ah horizons of buried soils
- Calcaric Cryosols indicate cold and arid environment of the Late Pleistocene
- Within the steppe forest ecotone buried soils show striking similarities indicating simplified periglacial zonality (hyper-zonality of A. Velichko, 1973)

Dokuchaev bioclimatic zonal soil sequence



Conclusions



- Due to a combination of more dynamic and more stable features soils of the forest-steppe areas are polygenetic and show features of both forest and steppe pedogenesis
- The Retisols at the southern fringe of the forest zone show landscape stability in a studied time/space range
- Dynamic soil porperties (humus and calcareous profile, Greyzemic features) are proxy of multi-directional landscape shifts
- Clay cutans are more stable and indicate one-way soil evolution: once appeared they are then inherited by subsequent stages