



# Mesoclimate Regulation Induced by Landscape Restoration and Water Harvesting in Agroecosystems of the Horn of Africa



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#waterharvesting

# Agroecosystems

**Agroecosystems:** communities of plants and animals interacting with their physical and chemical environments that have been **modified by people to produce food, fiber, fuel and other products for human consumption and processing** (Altieri, 2002).

Agroecosystems offer a **myriad of possibilities** for the implementation of new practices and management techniques, **larger than other ecosystems**. Agroecosystems management can be shifted on agricultural production AND Ecosystems services provision with **relatively small changes** (DeClerck et al., 2016).



# Landscape Restoration/Water Harvesting (LRWH)

**Water Harvesting** is the process of concentrating precipitation through runoff and storing it for beneficial use (Critchley et al., 1991)

It is key to **cope with water scarcity** for both sustaining agricultural production (Rockström et al., 2002) and **restore degraded landscapes** (Oweis, 2017).

The main effect of LRWH is to **retain rain water and runoff** in an **Agroecosystem**, in open storage reservoirs, in the soil or for aquifer recharge.

Critchley, W., et al. 1991. Water harvesting (AGL/MISC/17/91). FOOD AND AGRICULTURE ORGANIZATION OF THE UNITED NATIONS, Rome, Italy.

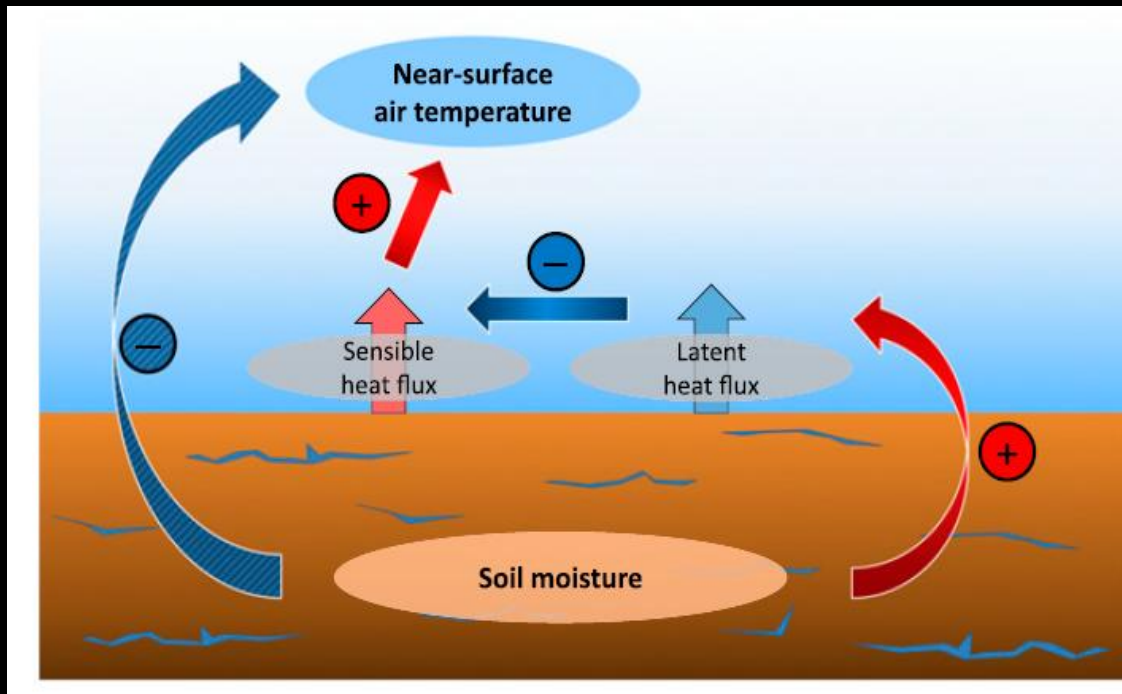
Rockström, J., et al., 2002. Rainwater management for increased productivity among small-holder farmers in drought prone environments. Phys. Chem. Earth 27, 949–959.

Oweis, T.Y., 2017. Rainwater harvesting for restoring degraded dry agro-pastoral ecosystems: A conceptual review of opportunities and constraints in a changing climate. Environ. Rev. 25, 135–149.



# Soil Moisture-Temperature Coupling (SMTC)

**Soil moisture ( $\theta$ )** can influence **near surface air-temperature ( $T$ )** (Schwingshackl et al., 2017, and cited literature)



Taken from Schwingshackl et al. (2017)

$$LH + SH + G = R_{net}^t$$

LH – Latent Heat flux

SH – Sensible Heat flux

G – Ground heat flux

$R_{net}$  – Net incoming Radiation

Feedback “dry”:  $\theta \downarrow$   $LH \downarrow$   $SH \uparrow$   $T \uparrow$

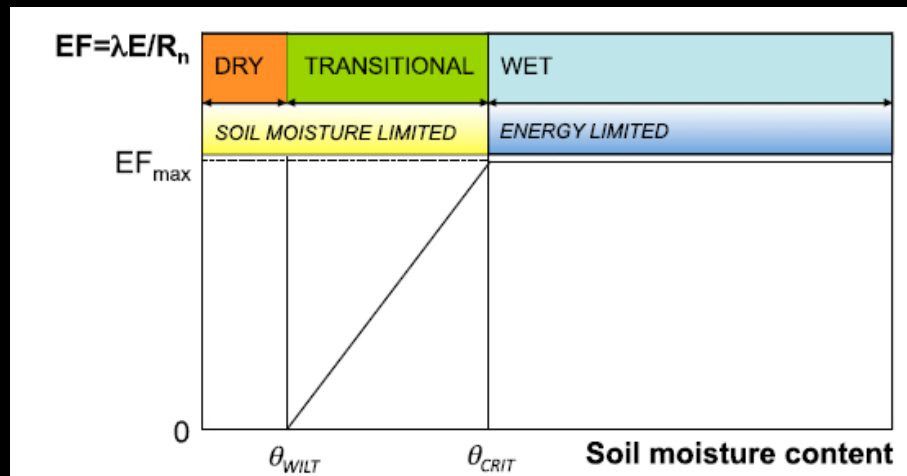
Feedback “wet”:  $\theta \uparrow$   $LH \uparrow$   $SH \downarrow$   $T \downarrow$

# Soil Moisture-Temperature Coupling (SMTC)

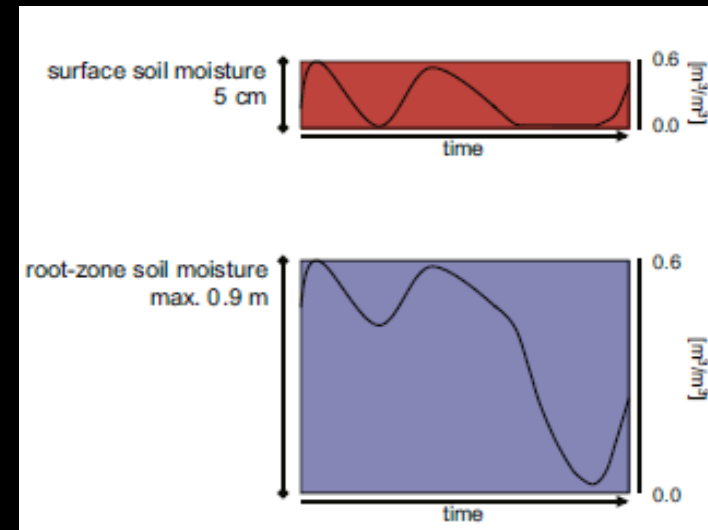
**Soil moisture deficit - heat waves feedback** has been largely discussed and documented:

**Hirschi et al. (2014)** SMTC dynamics are mostly evident when considering **root-zone soil moisture** (evaluated with SPI), rather than surface soil moisture (~5-10 cm, evaluated with remote sensing)

Mostly evident in locations with **Transitional Soil moisture and evapotranspiration regimes**  
Regions including Sahelian areas and Mediterranean climates.



Taken from Schwingshackl et al. (2017)



Taken from  
Hirschi et al. (2014)

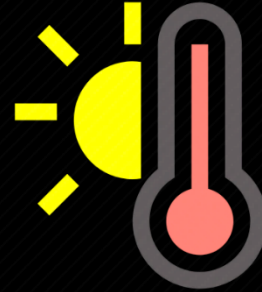
# Problem Statement

- Many studies have been carried out on the increase of soil moisture given by LRWH, but at local level and/or with modelling. **A data-based assessment** on possibility of **storing water** at **landscape (agroecosystem) level** with LRWH is still lacking.
- Studies on agricultural microclimate, and on meso-climate at urban and forest scale have been realised. Few studies on **meso-climatic modification at agroecosystem** scale are present.
- Studies demonstrate how SMTC affects continent and regional level heatwaves. Can it be used in a **proactive way (meso-climate modification)**?

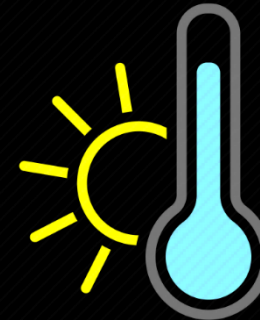
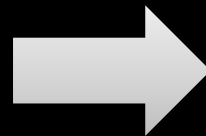


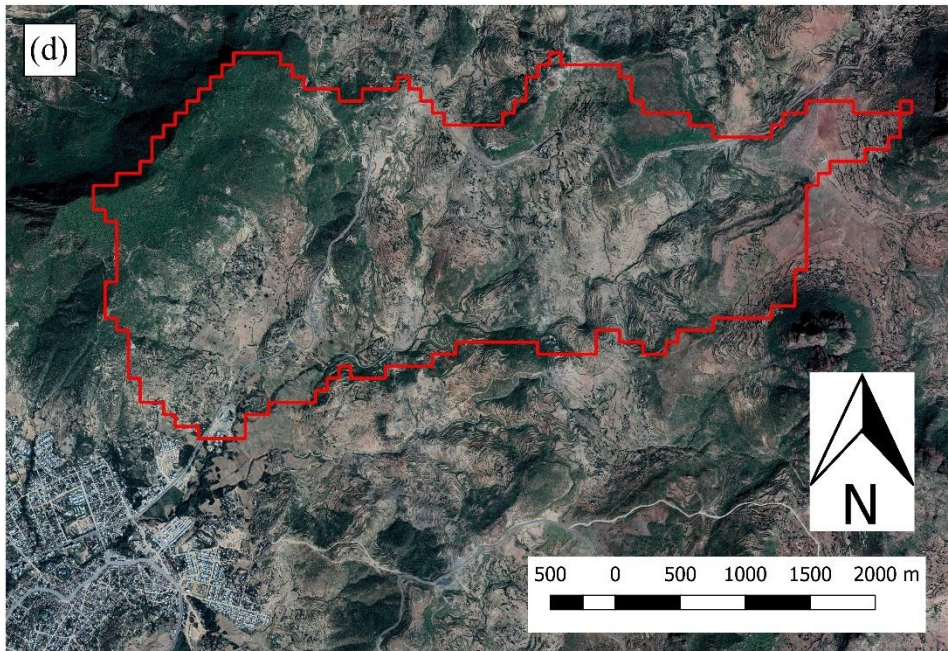
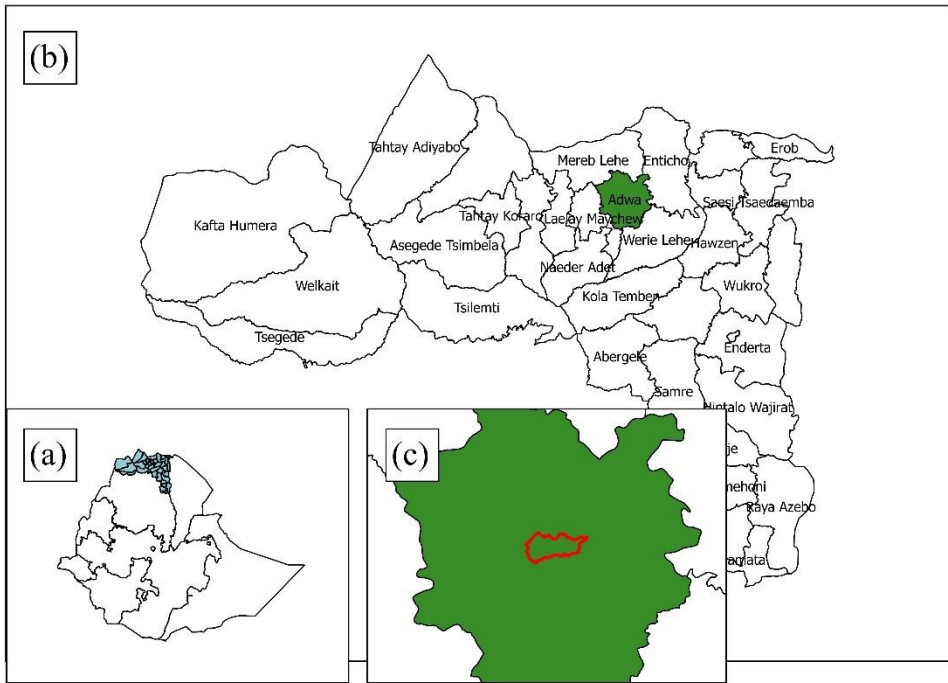


# Research Question



Mueller and Seneviratne (2012)  
Hirschi et al. (2014)  
Schwingshackl et al. (2017)





# Case Study

**Enabered Watershed**, Adwa district, Tigray Region, Ethiopia.

Between 38°53' to 38°57'E and 14°08' to 14°11'N

**Elevations:** from 1,850 to 2,540 m a.s.l.

Average annual precipitation (1998-2008): 742 mm

Average daily temperature (1998-2008): 19.8 °C

Rainy season: **June-September** (85% of rainfall)

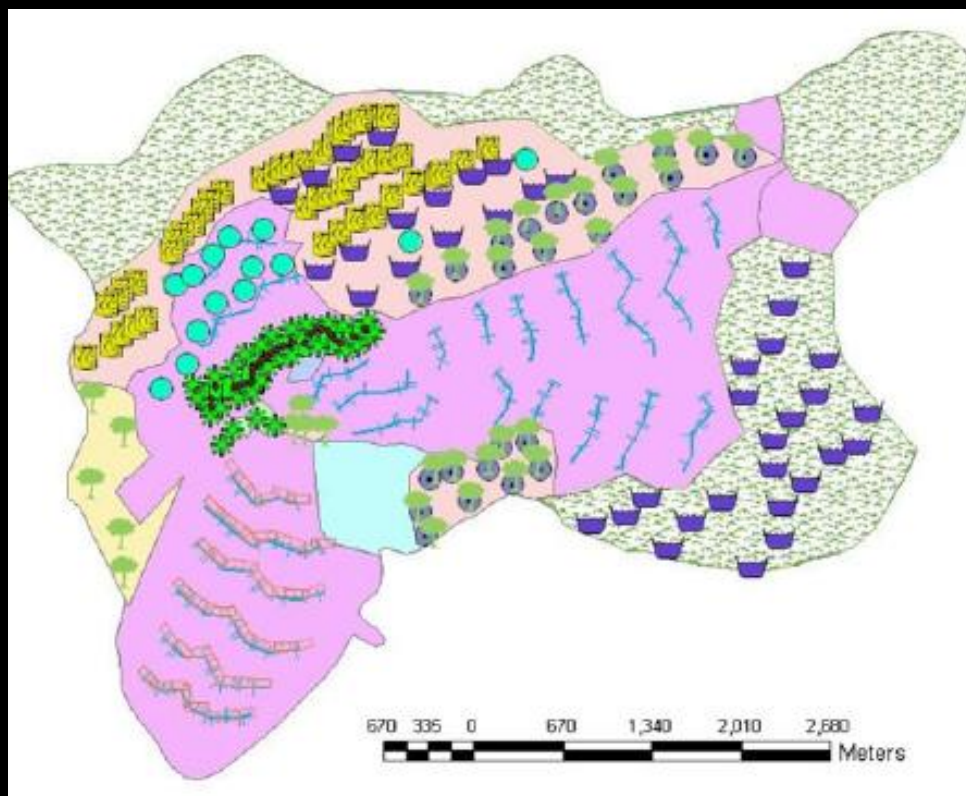
**Transitional soil moisture and evapotranspiration regime**

LRWH interventions implemented **between 2004 and 2008**

Haregeweyn et al. (2012) reported the **full list of the techniques implemented in the area**







Taken from Haregeweyn et al. (2012)

Type of LRWH	unit	Extent of LRWH			Total
		Hillside	Gully	Cultivated and grazing land	
Physical measures	ha	1,108	8	1,036	2,152
Stone-faced bunds with trench	km	135			135
Stone and soil bunds	km	472		205	677
Deep trenches	km	1,592			1,592
Trenches	km			555	555
Loose-stone check dams	m <sup>3</sup>	38,999	23,150		62,149
Gabion check dams	m <sup>3</sup>		20,231		20,231
Retention walls	km		0.5		0.5
Sediment storage dams	m <sup>3</sup>		498		498
Microbasins	no.	50,629			50,629
Gully reshaping	m <sup>3</sup>		90,788		90,788
Pond construction	no.			10	10
Bund stabilization	km			516	516
Biological measures	ha	1,201	28	635	1,931
Exclosures	ha	601			601
Grass/split planting	ha		8		8
Grass sowing	ha	545	5	308	850
Enrichment plantations	ha	55	8		63
Fruit trees	ha		2	7	9
Forage trees	ha		8	320	400

# Materials and Methods: Water Conservation Index (WCI)

$$WC_i(y) = 100 \frac{NDI_i(y)}{R_{rs}(y)}$$

- **WCI<sub>i</sub>(y)** - WCI for the i-th month of the year y
- **R<sub>rs</sub>(y)** - rainfall in the rainy season (June-August) in the year y (mm), from **CHIRPS dataset** (Funk et al., 2015);
- **NDII<sub>i</sub>(y)** - Normalised Difference Infrared Index for the i-th month of the year y

$$NDI = \frac{\rho_{B4} - \rho_{B5}}{\rho_{B4} + \rho_{B5}}$$

- $\rho_{B4}$  - reflectance in **Landsat 7 ETM+** sensor Band 4 (0.77-0.90  $\mu\text{m}$ )
- $\rho_{B5}$  - reflectance in **Landsat 7 ETM+** sensor Band 5 (1.55-1.75  $\mu\text{m}$ )
- NDII - '**Landsat 7 Collection 1 Tier 1 8-Day NDWI Composite**' on **Google Earth Engine** (Gorelick et al., 2017). *De facto* NDII



# Materials and Methods: Water Conservation Index (WCI)

$$WC_i(y) = 100 \frac{NDI_i(y)}{R_{rs}(y)}$$

- WCI time series calculated for respectively for the months of **September (WCI<sub>9</sub>)**, **October (WCI<sub>10</sub>)** and **November (WCI<sub>11</sub>)**, ranging from 2000 to 2017.
- Data "Before full LRWH implementation": 2000-2008
- Data "After full LRWH implementation": 2009-2017
- Good accordance for values of **NDII and root-zone soil moisture during the dry season** (Sriwongsitanon et al., 2016) [ $R^2 = 0.87$ ]

Funk, C., et al., 2015. The climate hazards infrared precipitation with stations—a new environmental record for monitoring extremes. Sci. Data 2, 150066.

Gorelick, N., et al., 2017. Remote Sensing of Environment Google Earth Engine : Planetary-scale geospatial analysis for everyone. Remote Sens. Environ. 202, 18–27.

Sriwongsitanon, N., et al., 2016. Comparing the Normalized Difference Infrared Index (NDII) with root zone storage in a lumped conceptual model. Hydrol. Earth Syst. Sci. 20, 3361–3377.



# Materials and Methods: Normalised temperature index (t)

$$t_i(y) = \frac{T_{LS_i(y)}}{T_{850,i}(y)}$$

- $T_{LS_i(y)}$ , average Land Surface Temperature (°C) for the i-th month of the year y  
MODIS MYD11A2.006 Aqua Land Surface Temperature and Emissivity 8-Day Global at 1 km from Google Earth Engine (NASA LP DAAC, 2018).
- $T_{850,i}(y)$  - average the temperature at 850 hPa at 12:00 a.m. (°C) obtained from **ERA-INTERIM climatic reanalysis dataset** (Balsamo et al., 2015)
- Data "Before full LRWH implementation": 2002-2008
- Data "After full LRWH implementation": 2009-2017



# Materials and Methods: SMTC

Based on the framework of Schwingshackl et al. (2017):

$$\frac{\partial T}{\partial \theta} = \frac{\partial T_F}{\partial E} \frac{\partial E}{\partial \theta} F$$

With a proxy approach:

$$\partial t / \partial WCI$$

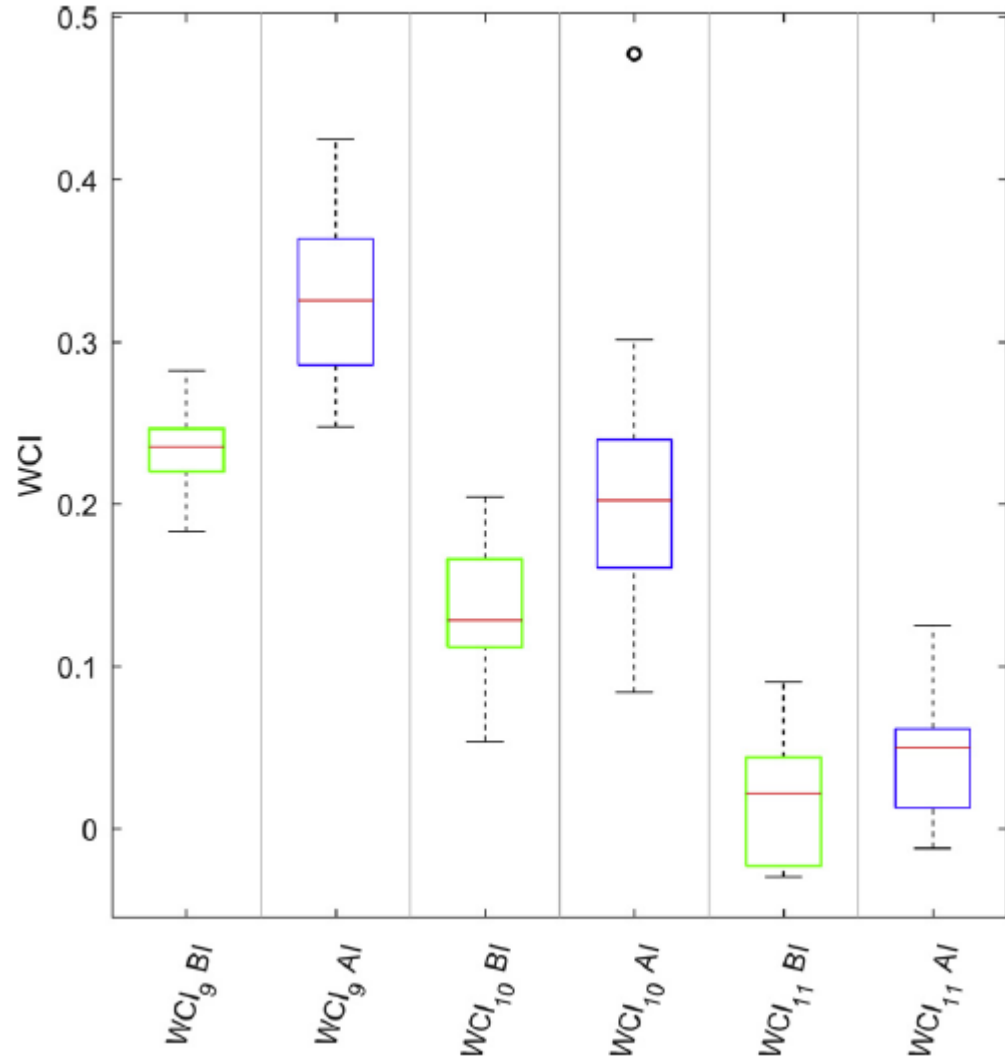
To detect possible lag effects, two version of a linear model have been investigated:

- (i)  $t_i = f(WCI_{i-1})$  (with lag of one month);
- (ii)  $t_i = f(WCI_i)$  (without lag).



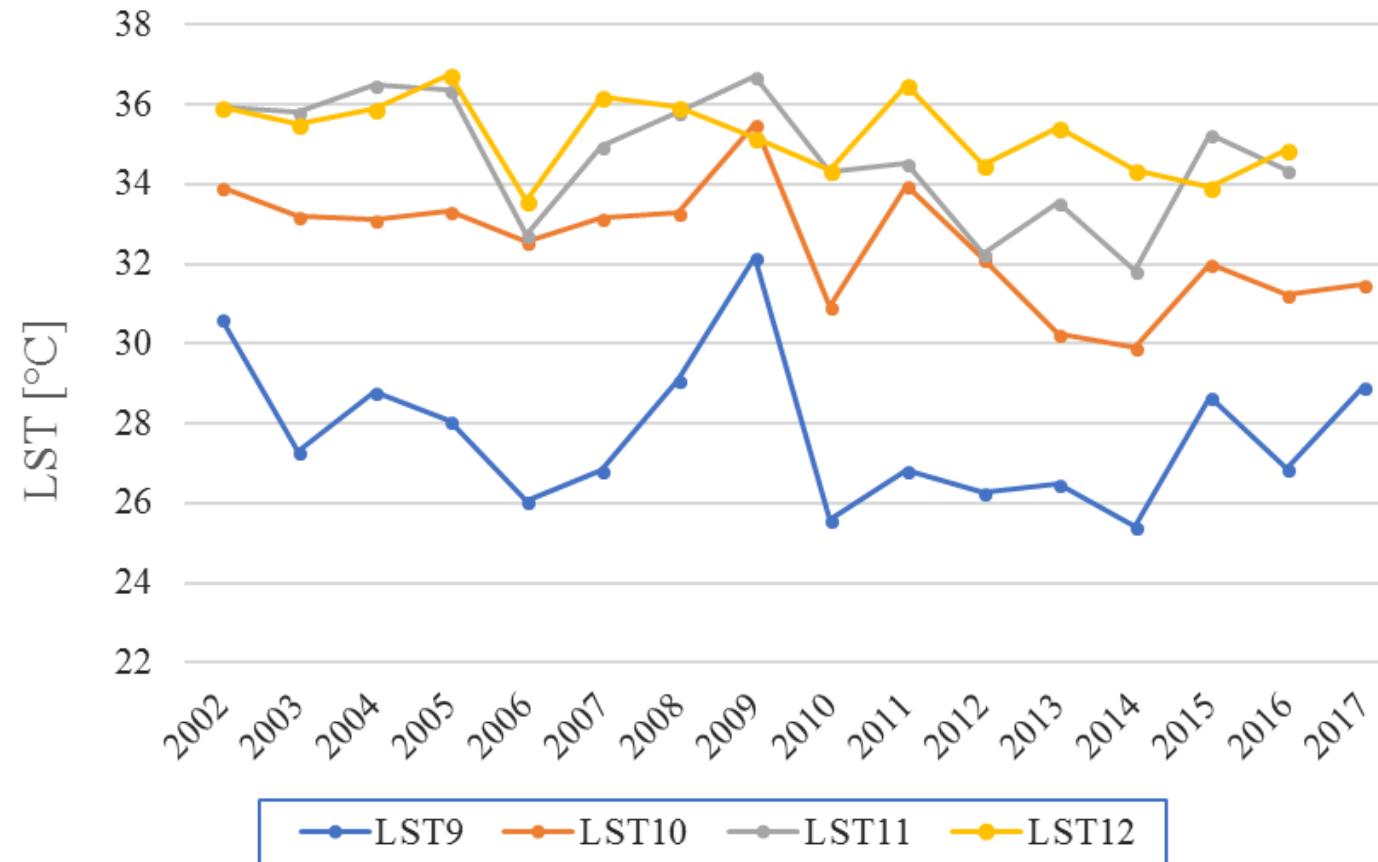


# Results: Water Conservation Index (WCI)

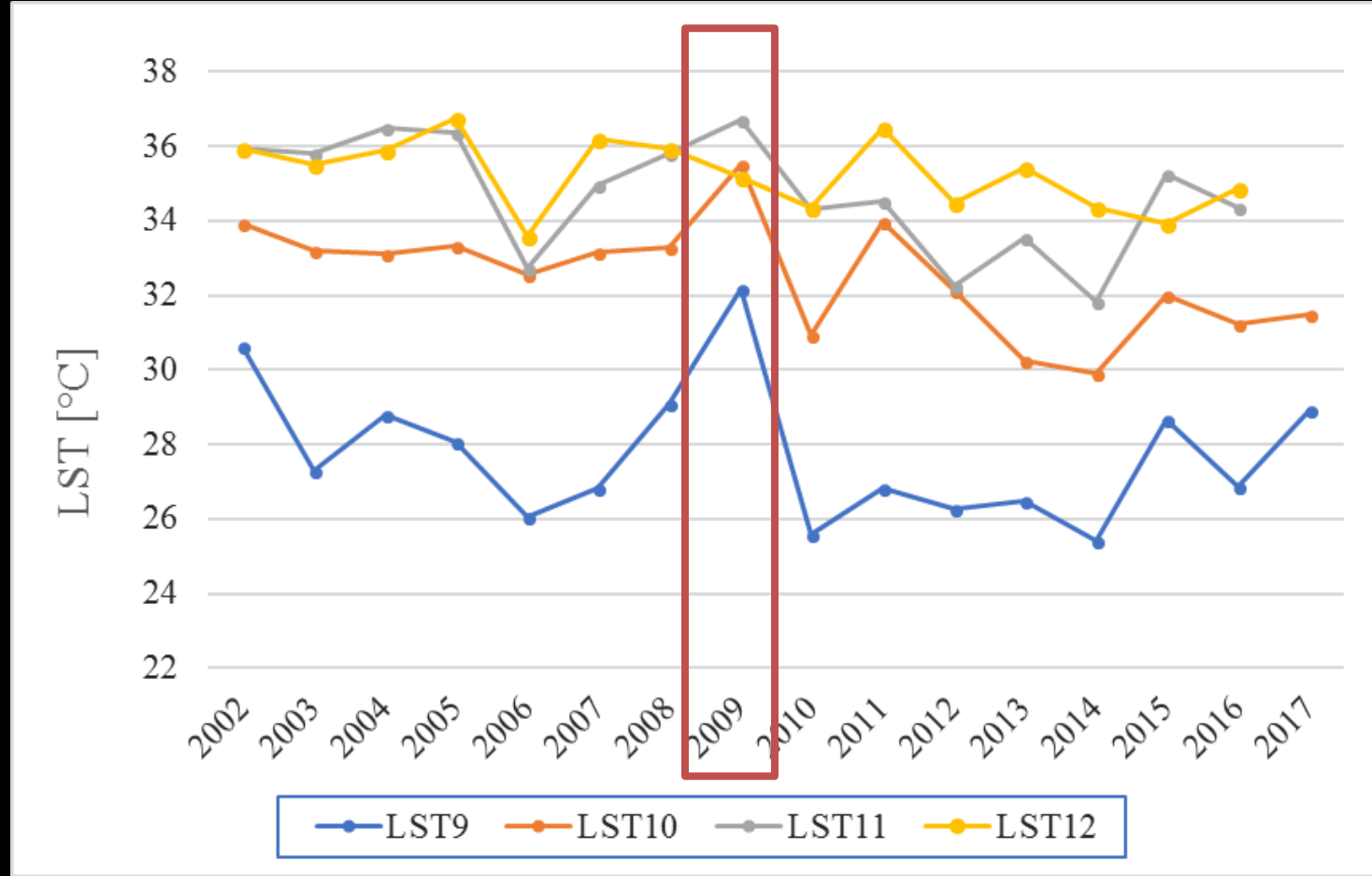


	September	October	November
WCI Average 2000-2008	0.235 (0.028)	0.134 (0.049)	0.016 (0.044)
WCI Average 2009-2017	0.325 (0.038)	0.221 (0.095)	0.045 (0.034)
WCI Difference before and after full implementation	0.090	0.087	0.029
WCI Difference before and after full implementation (%)	38%	65%	181%
P- value, test on differences	0.00047	0.08330	0.21833
Statistical signifiante	> 99%	91%	78%

# Results: LST



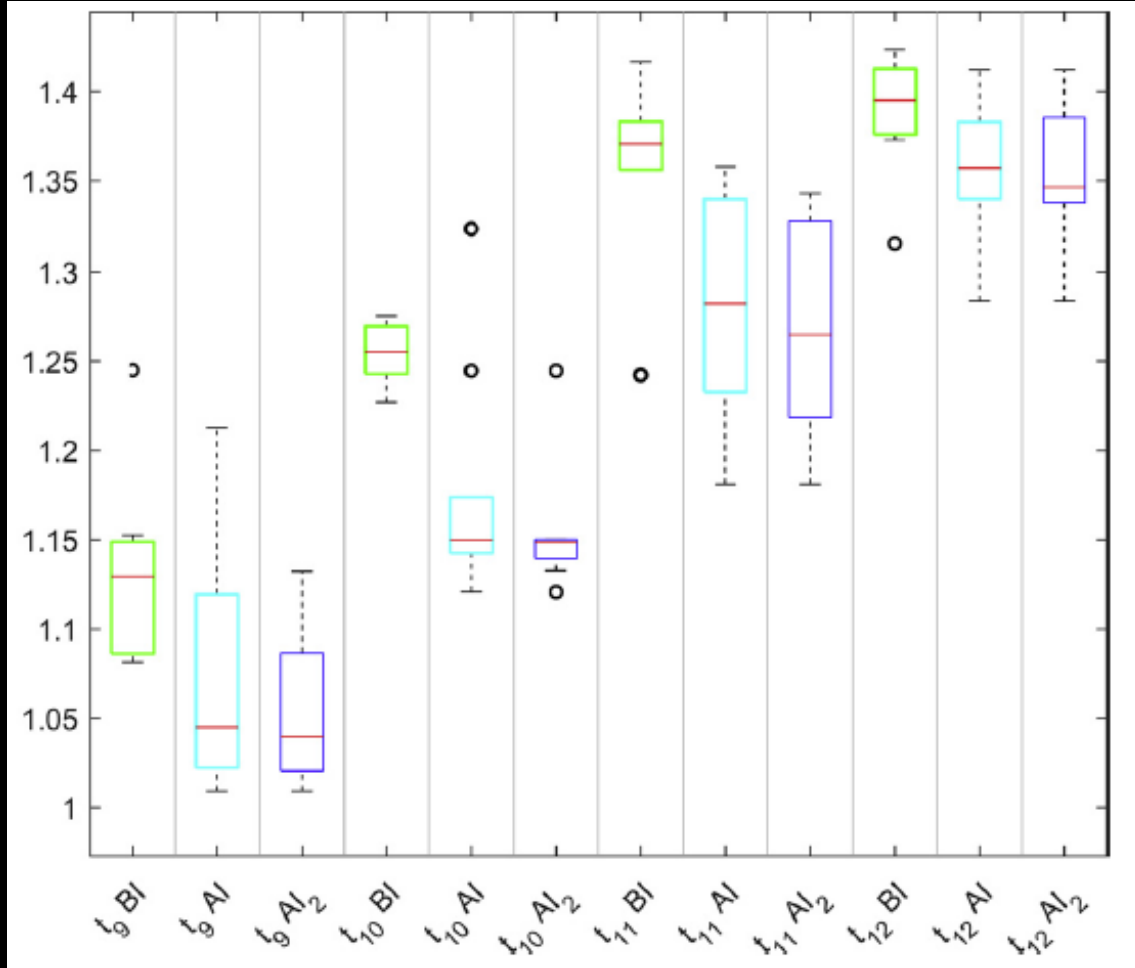
# Results: LST



# Results: LST

Month	September	October	November	December
Average LST (2002-2008)	28.13	33.24	35.45	35.68
Average LST (2009-2017)	27.48	31.94	34.10	34.89
Average LST (2010-2017)	26.89	31.49	33.73	34.85
Difference LST (2002-2008) – LST (2009-2017)	0.65	1.30	1.35	0.80
Difference LST (2002-2008) – LST (2010-2017)	1.24	1.74	1.72	0.84

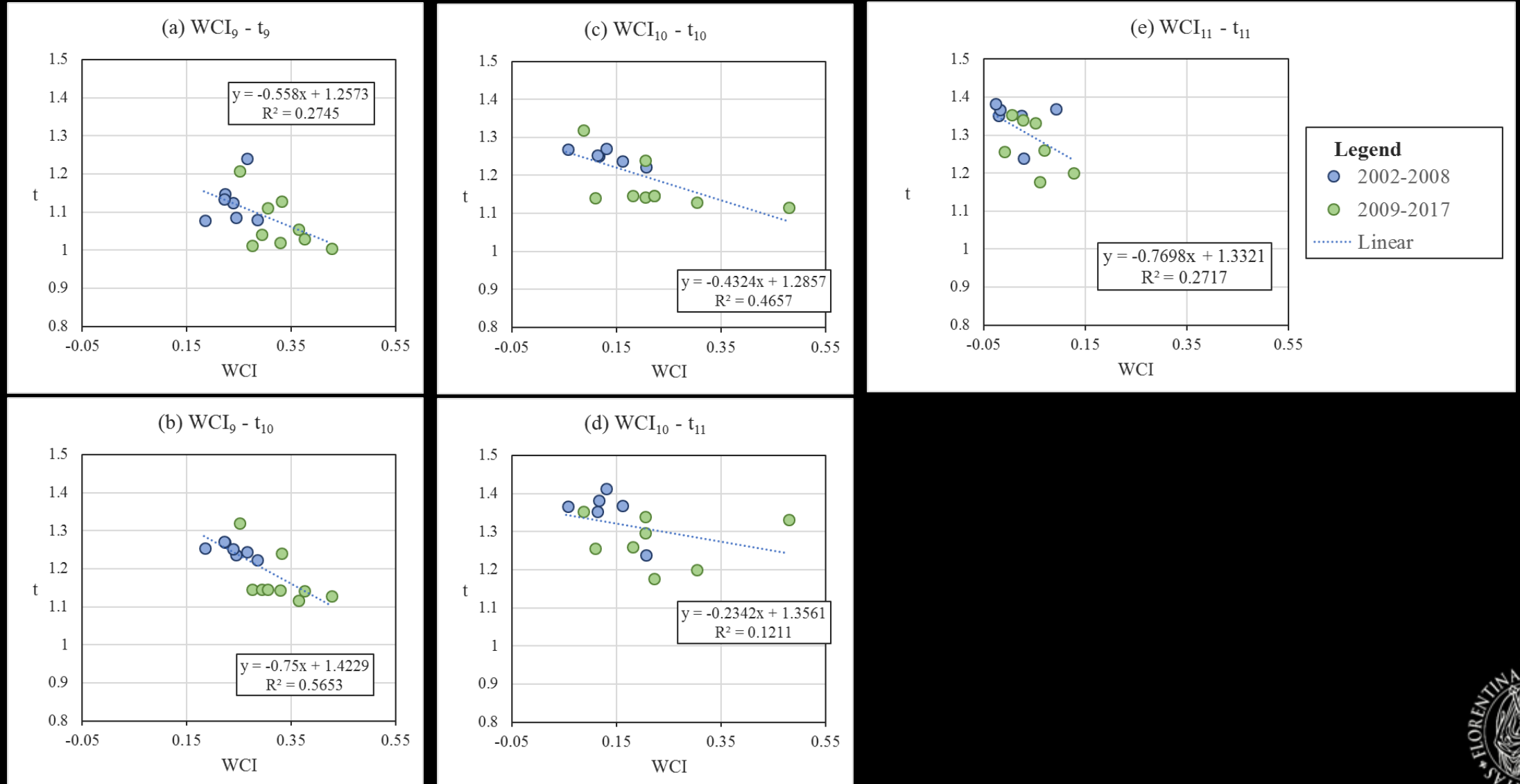
# Results: Normalised temperature index (t)



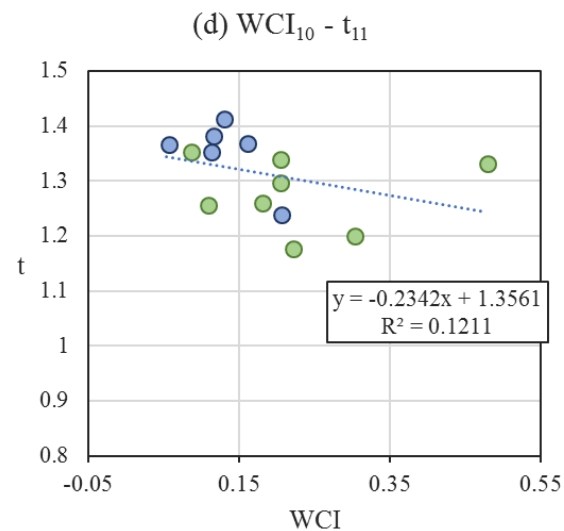
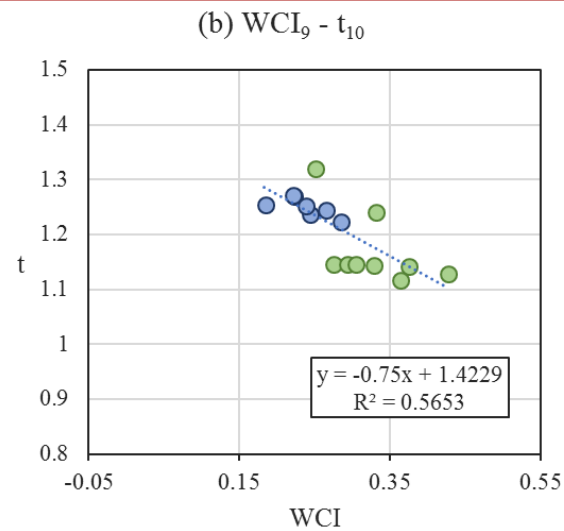
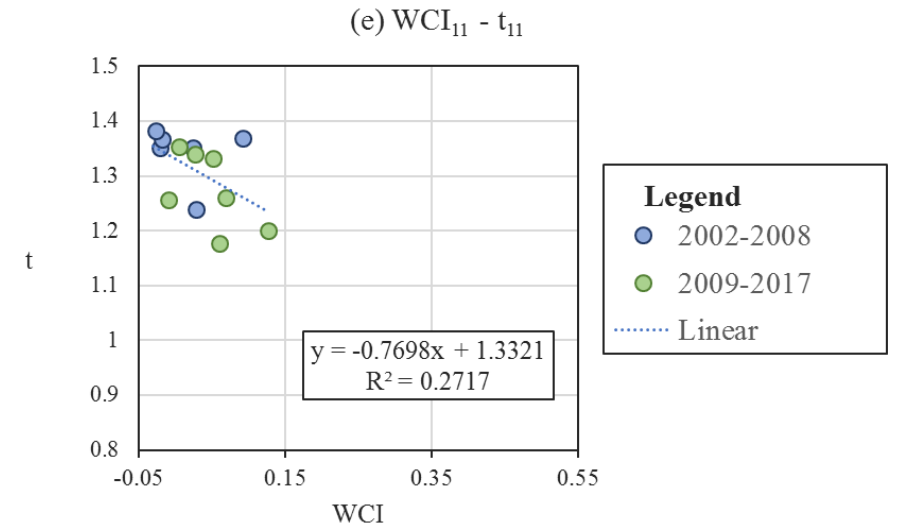
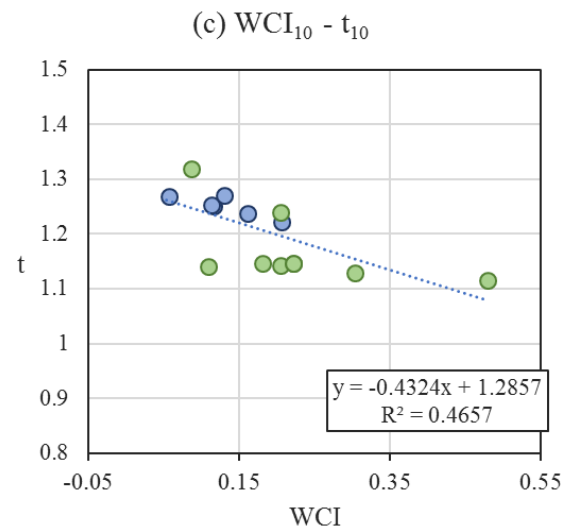
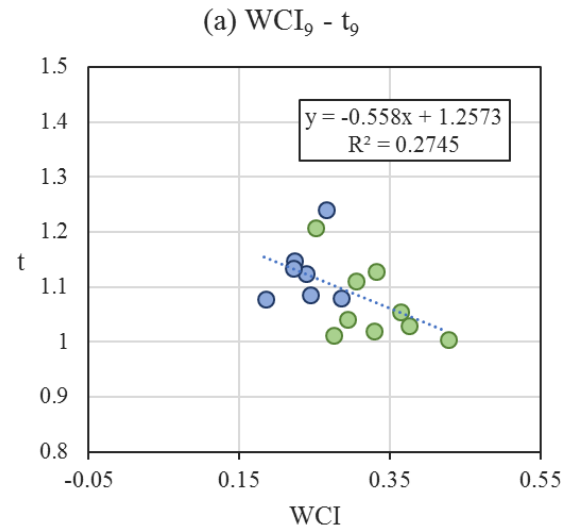
Month	September	October	November	December
Average t (2002-2008)	1.132 (0.057)	1.254 (0.017)	1.357 (0.055)	1.387 (0.036)
Average t (2009-2017)	1.072 (0.068)	1.174 (0.066)	1.281 (0.065)	1.357 (0.039)
p-value	0.083	0.008	0.030	0.266
Statistical significance	<b>90 %</b>	<b>&gt; 99 %</b>	<b>&gt; 95 %</b>	<b>73 %</b>
Difference t (2002-2008) – t (2009-2017)	0.06	0.08	0.076	0.03
Relative difference	5%	6%	6%	2%



# Results: SMTC at catchment scale

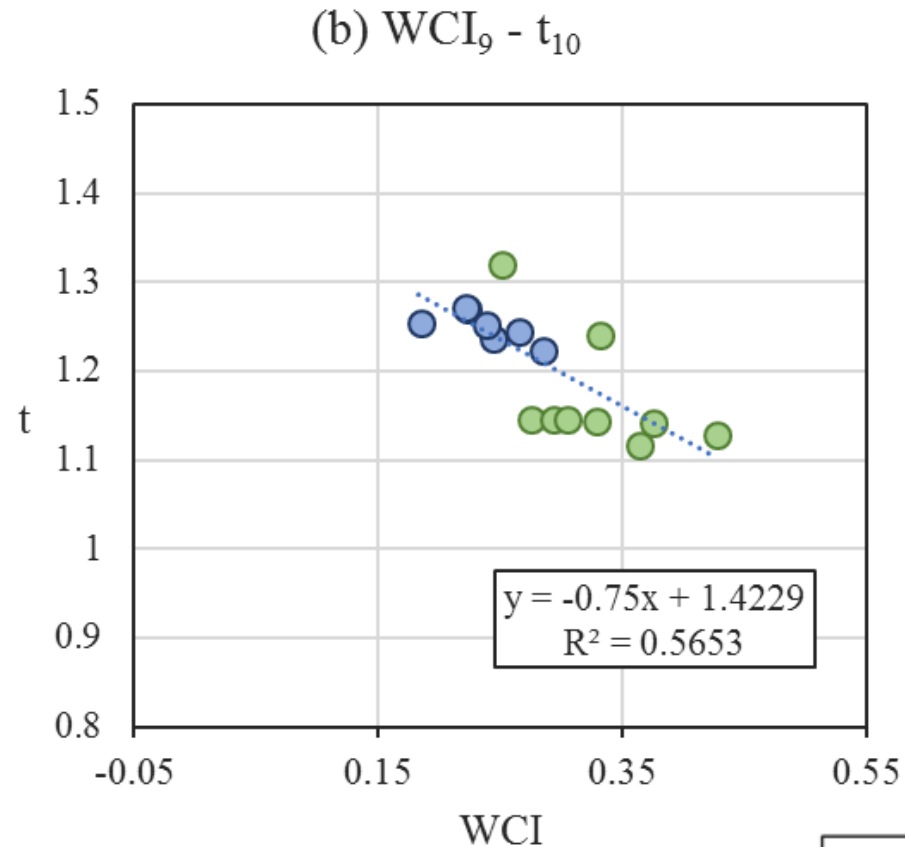


# Results: SMTC at catchment scale



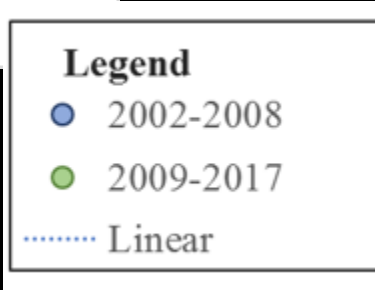
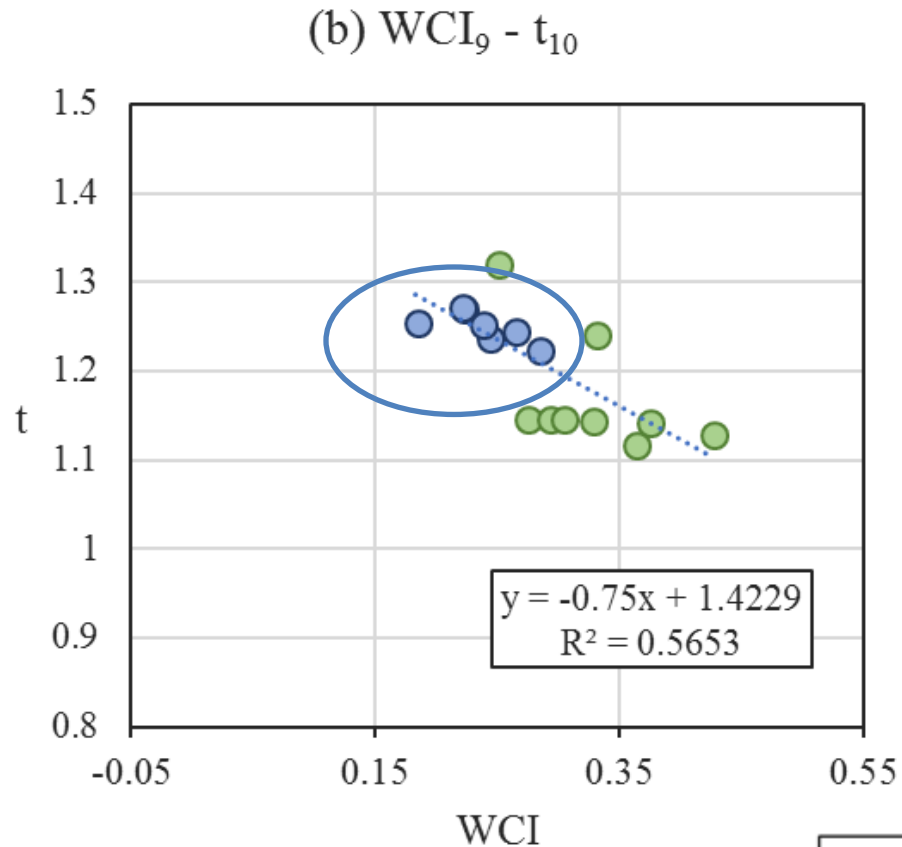
# Results: SMTC at catchment scale

- Highest SMTC is the one characterised by the relation  $t_{10} = f(WCI_9)$ .
- **Separation of populations, except 2009**



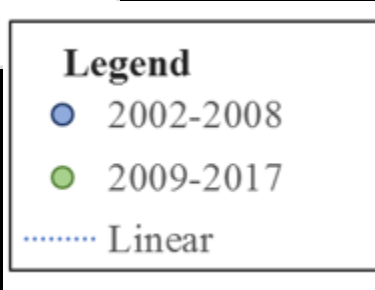
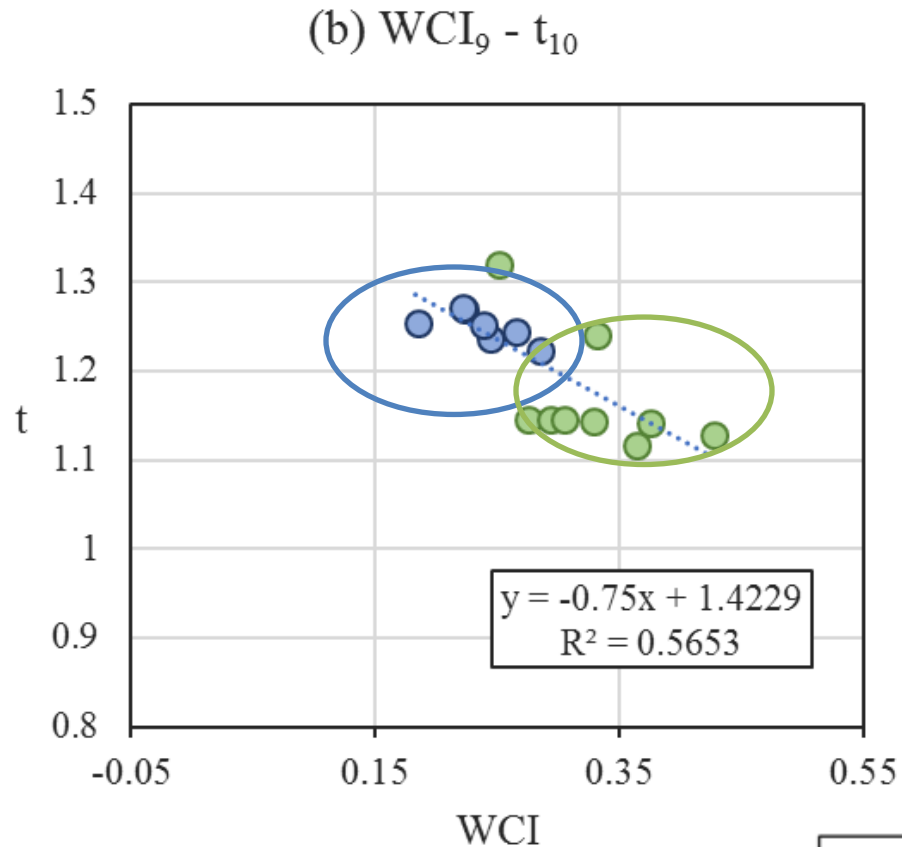
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# Results: SMTC at catchment scale

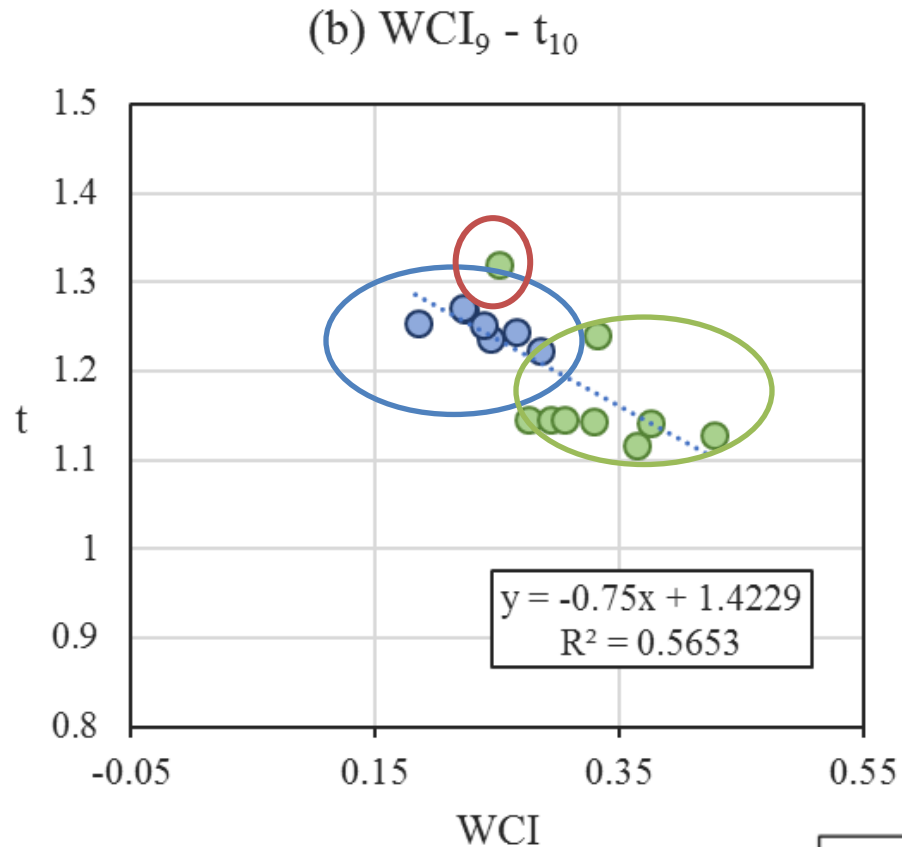
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# Results: SMTC at catchment scale

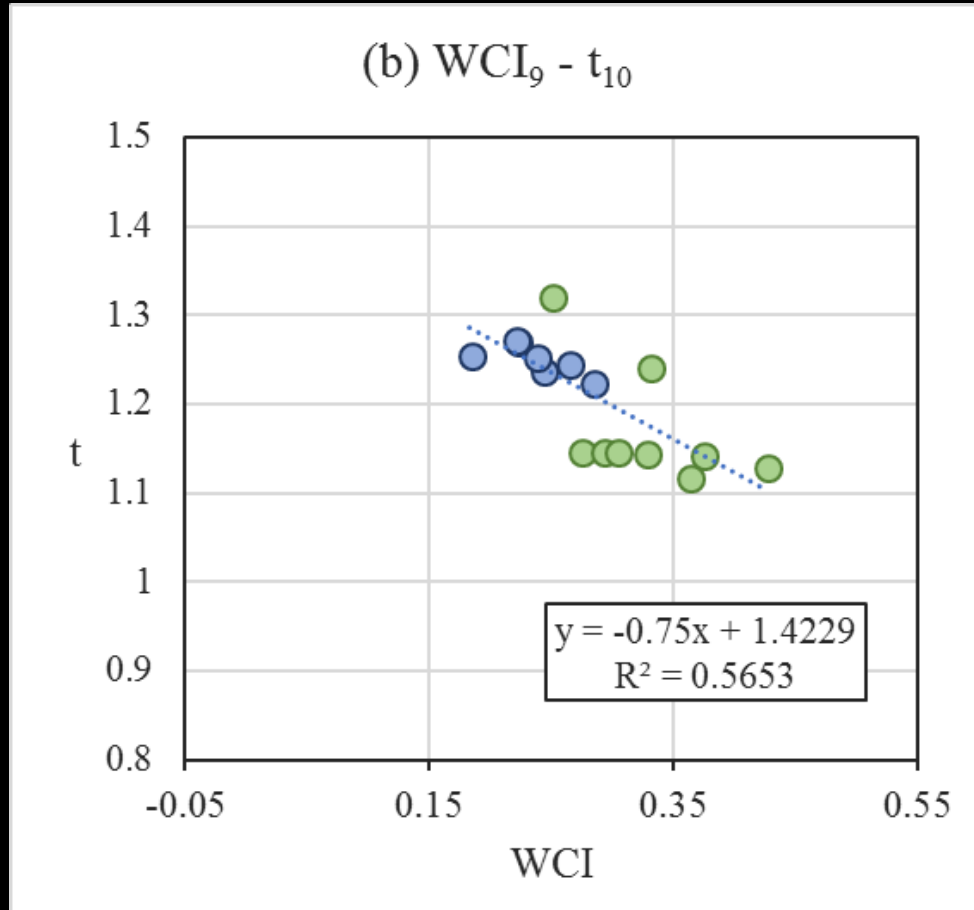
- Highest SMTC is the one characterised by the relation  $t_{10} = f(WCI_9)$ .
- **Separation of populations, except 2009**



**Legend**

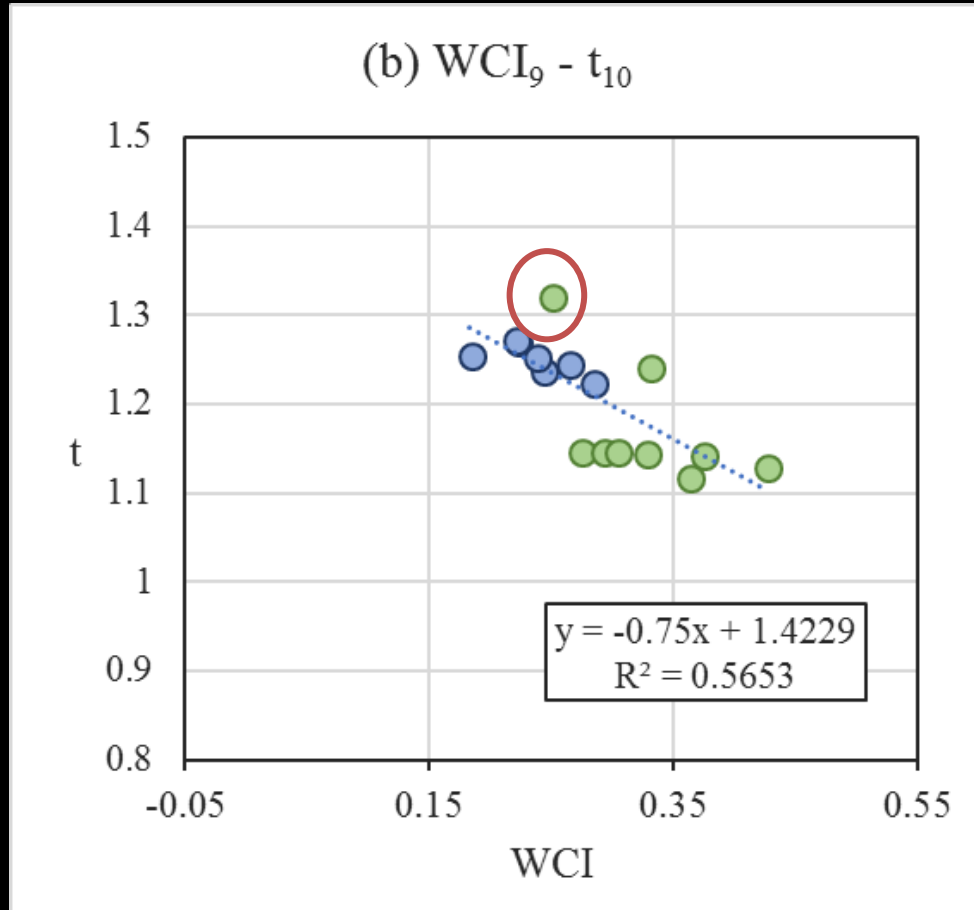
- 2002-2008
- 2009-2017
- ..... Linear

# Results: SMTC at catchment scale



- Highest SMTC is the one characterised by the relation  $t_{10} = f(WCI_9)$ .
- **Separation of populations, except 2009**
- Coupling of **the root zone soil moisture conserved at catchment scale in September** and the **catchment average temperature in October**.
- Soil moisture available in September is depleted as evapotranspiration from September to October.
- **1-month lag** can be expected (up to three months for heatwaves in central Europe)

# Results: SMTC at catchment scale



- Considering 2009, the coupling strength is the maximum analysed,  $\partial t / \partial WCI = -0.75$ , correspondent to an average decrease in LST of  $1.30^\circ\text{C}$ , with an  $R^2$  of **0.5653**.
- Without 2009, decrease in LST is of  **$1.74^\circ\text{C}$**
- **Extreme dry year occurred in 2009** as reported by Winkler et al., 2017. The work explains also the other peak of LST occurring in October 2011.

# Discussion

- **High LST and  $t$  in 2009**: despite the coupling dynamics, the soil moisture available at catchment scale in September 2009 was not sufficient to provide enough LH.
- LRWH interventions contributed to lower the average temperatures at the watershed scale, **their influence can be limited in the case of extreme events.**
- **Similar** to the role of **water harvesting** as a mean to **deal with water scarcity**: more effective in bridging short **dry spells of 5 to 15 days**, that represent the first source of crop failure, rather than allowing to **buffer prolonged droughts** (Rockström et al., 2002).



# Conclusions and further developments

- LRWH enhance the water retention capacity at catchment scale for **September ( $P < 0.01$ )** and **October ( $P < 0.1$ )**. Effects in November are not evident for this scale of analysis.
- After LRWH full implementation, **temperature decreased in September ( $P < 0.1$ ), October ( $P < 0.01$ ) and November ( $P < 0.05$ )**.
- The analysis has also taken into account the exceptional year of 2009, with extremely high temperatures.
- By removing 2009 from the analysis, the study shows an average decrease in **LST of  $1.74\text{ }^{\circ}\text{C}$** . The variation, in absolute terms, **is similar to the ones that can be induced in urban areas by the conversion of large areas of paved surfaces and built environment into green infrastructures and vegetated areas** (Di Leo et al., 2016; Zareie et al., 2016).
- **SMTC is evident at catchment scale.**
- WCI values of **September evidence a negative linear correlation to t values of October ( $R^2 = 0.59$ )**. The 1-month lag can be well justified by considering the framework for the modelling of SMTC presented by Schwingshackl et al. (2017).

**The implementation of LRWH measures provided a climate regulation effect in the watershed.**





# Conclusions and further developments

## Further Developments

Analysis of the evidence of similar dynamics in **other regions of the world**.

Use of **more advanced remote sensing datasets** such as the recent Sentinel-2 imagery, but available only from 2015.

Downscaling of **global** (Schwingshackl et al., 2017) or **regional** (Mohamed et al., 2005) size modelling tools.

Investments in **long-term experiments** for the analysis of SMTC **at catchment scale** may be considered if further studies will confirm this initial one.





# Thank You



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# Limitations of the analysis

LRWH in the catchment was implemented **between** 2004 and 2008. Some other interventions may be implemented afterwards

- Some interventions could have an actual effect in **some years after the implementation**
- In 2008 ALL infrastructures were in place and functioning
- **After 2008: minor interventions**

Proxy approach with LST and NDII

- Most of the catchments where LRWH is implemented recently are **often ungauged**
- NDII has good accordance with **root-zone soil moisture**
- LST is often used in landscape studies (e.g. urban environments)

