# Effect of latitudinal displaced gravity wave forcing in the stratosphere on the circulation of the middle atmosphere

N. Samtleben<sup>1</sup>, C. Jacobi<sup>1</sup>, P. Šácha<sup>2</sup>, P. Pišoft<sup>2</sup> and A. Kuchar<sup>2</sup>

Leipzig Institute for Meteorology, Universität Leipzig, Leipzig, Germany<sup>1</sup> Department of Atmospheric Physics, Charles University, Prague, Czech Republic<sup>2</sup> Email: Nadja.Samtleben@uni-leipzig.de



H7 - Ref

H3 - Ref

### Introduction

- Gravity waves (GWs) are the major contributor to vertical coupling of atmospheric layers by distributing energy and momentum throughout the whole atmosphere.
- GWs have horizontal wavelengths of tens to hundreds of kilometers so that GWs are mostly parameterized in global circulation models as zonal means.
- Breaking GW hotspots have been observed in the stratosphere [Šácha et al., 2015].
- Objective: Effect of non-zonal GW distribution has to be analyzed.
- Hypothesis: Preconditioning of the polar vortex by impact of GWs at midlatitudes [Albers and Birner, 2014].

### MUAM – Middle and Upper Atmosphere Model

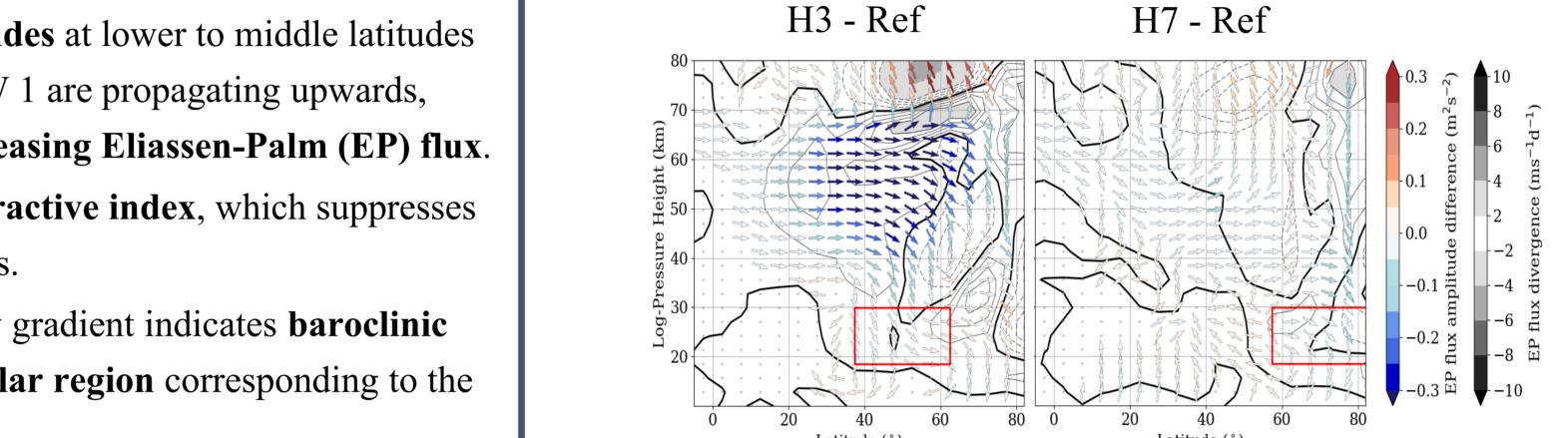
- Primitive equation 3D grid point model [Lilienthal et al., 2017; Samtleben et al., 2019]
- horizontal resolution: 5°x5.625°
- Upper boundary: 160 km (log-p) with  $\Delta z=2.842$  km
- Nudging of ERA-Interim zonal mean temperature (decadal mean: 2000-2010) below 10 km
- GW parameterization: linear scheme with multiple breaking levels
- GW initialized at 10 km (1 cm s<sup>-1</sup> vertical velocity perturbation)

Fig. 1: Horizontal resolution of MUAM.

### Results: Background and stationary Planetary Waves

- H3 (37.5-62.5°N) GW hotspot, which is half located in the easterlies (not acting against the zonal mean zonal wind) as well as in the westerlies leads to: - a weakening of the Aleutian high pressure system,
  - a displacement of the polar vortex towards Canada (see Fig. 5).
- Changes in Background conditions:
  - decreasing zonal mean zonal wind at middle to high latitudes,
  - heating of the lower stratosphere,
  - increasing geopotential height.
- In case of the H7 (57.5-82.5°N) GW hotspot the polar vortex is less disturbed so that the differences are much smaller.
- Decreasing zonal wind SPW 1 amplitudes at lower to middle latitudes between 20 and 60km, because less SPW 1 are propagating upwards, which can be seen by means of the decreasing Eliassen-Palm (EP) flux.
- GW hotspot is leading to a **negative refractive index**, which suppresses the propagation of SPW 1 at midlatitudes.
- A reversed meridional potential vorticity gradient indicates baroclinic instability generating SPW 1 in the polar region corresponding to the strong positive EP divergence.

## - Preconditioning of the polar vortex -



(middle and right panel)) box refers to the H3 (H7) simulation

37.5-62.5N - H3

Fig. 7: Zonal mean EP flux (arrows) and divergence (isolines and shaded areas, dashed lines show negative values) of SPW 1. Shown is the difference between all H3 (left)/H7 (right) simulations and the Ref simulation.

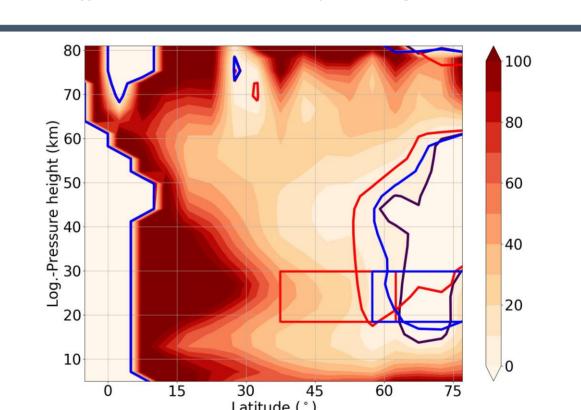
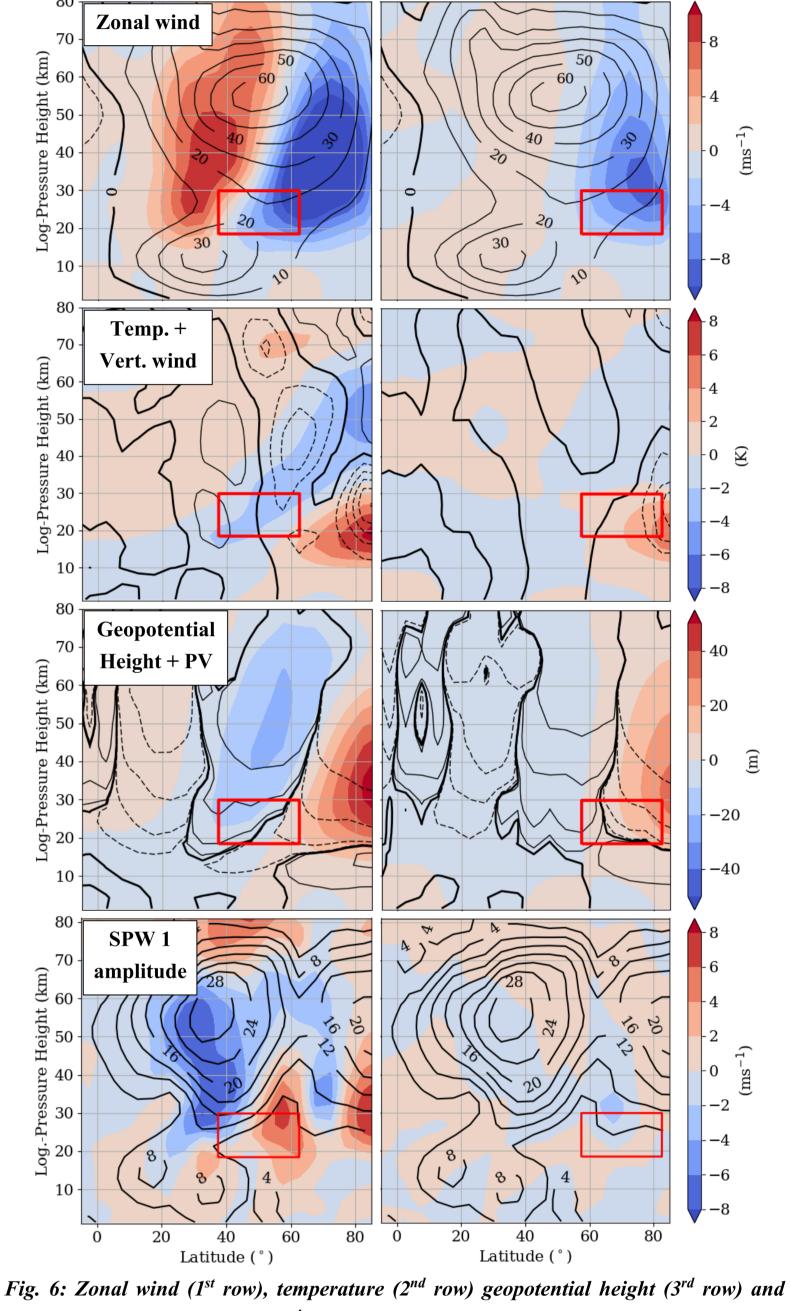


Fig. 8: Refractive index for SPW 1 for the Ref simulation (right panel) with a thicker zero line colored in violet. The position of the H3 (H7) GW hotspot and the respective zero line is represented by the red (blue) line.



 $ISPW\ 1$  amplitude (4 $^{th}$  row) differences of the H3 (left) and H7 (right) GWalso including zonal mean zonal wind of the Ref simulation (contour lines in plot), vertical wind differences (contour lines in temperature plot starting with  $\pm -0.0005~{\rm ms}^{-1}$  up to  $\pm -0.0025~{\rm ms}^{-1}$  with increments of  $0.0005~{\rm ms}^{-1}$ ), potential vorticity differences (contour lines in geopotential height plot) and zonal wind SPW I amplitude of the Ref simulation (contour lines in SPW 1 amplitude plot).

## Model setup: Local non-zonal GW forcing

### 1. Gravity wave distribution in MUAM:

- Zonal mean GW distribution based on the observed potential energy (E<sub>pot</sub>) data derived from GPS radio occultations (FORMOTSAT3/COSMIC)
- E<sub>pot</sub> divided by its global mean E<sub>pot</sub> at each grid point

### 2. Implementation of the non-zonal GW forcing:

- Artificial enhanced GW drag in the GW parameterization output (period of simulation: 120 days)
- Position of observed breaking GW hotspot:

Longitude: 120-170°E **Latitude:** 37.5-62.5°N

18-30km **Height:** 

- Latitudinal displacement of the GW hotspot in 5° steps (H1-H8) with a southernmost (northernmost) GW hotspot between 27.5-52.5°N (62.5-87.5°N)
- Enhanced GW parameters:

zonal GW drag:

**Heating:** 

-10 ms<sup>-1</sup>day<sup>-1</sup>

meridional GW drag: -0.1 ms<sup>-1</sup>day<sup>-1</sup>

0.05 Kday<sup>-1</sup>

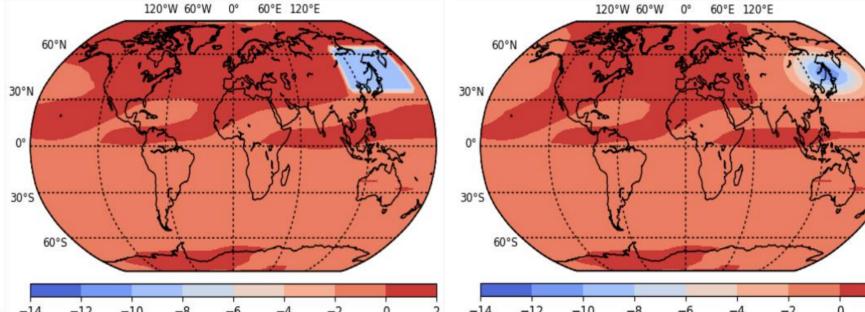


Fig. 4: Zonal GW drag (ms<sup>-1</sup>day<sup>-1</sup>) at 26.9 km for the reference (left) and the H3 hotspot simulation as box (middle) and as Gaussian distribution (right) for the last 30 days of analysis. Note the different scaling on the left panel.

Fig. 9: Meridional potential vorticity gradient difference between the H3 (left panel), H7 (right

panel) and the Ref simulation representing the last 30 days of the simulations. The positions of the

0.0 0.5 1.0 1.5 2.0 2.5 3.0 3.5 4.0 4.5 Fig. 2: Global GW weights derived from  $E_{not}$  data and the resulting zonal mean GW weights for MUAM (right, blue contour line).

H3 and H7 GW hotspot is represented by the red boxes.

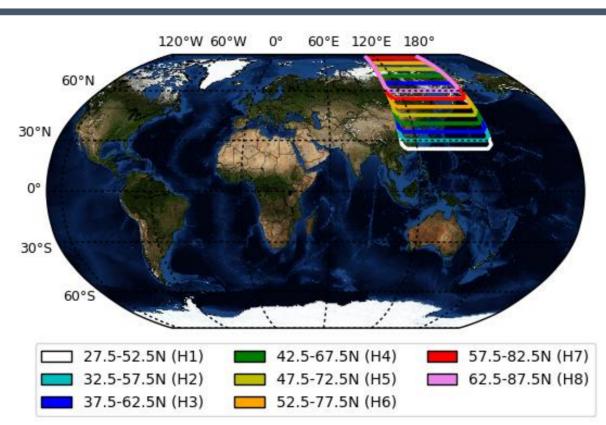


Fig. 3: Positions of the artificial GW hotspots, which have been analyzed in this sensitivity study.

### **Results: Shape of the forcing**

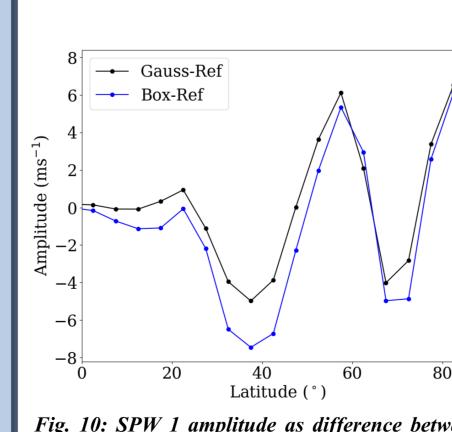


Fig. 10: SPW 1 amplitude as difference between the H3 and the Ref simulation (blue line) and th Gauss distribution and the Ref simulation (black line) at 35 km for the last 30 days extracted from the zonal wind.

- Suppression of the SPW might be caused by the shape of the GW hotspot.
- Integral of the three dimensional box and the Gauss distribution is the same (increased the maximum of GW parameters, which slightly increases SPW l amplitude).
- smoothly decreasing GW drag distribution towards lower and higher latitudes only slightly reduces the suppression of SPW 1 around 40°N and 70°N.

### **Conclusion and Outlook**

- Sensitivity study regarding the position of the enhanced GW forcing shows a preconditioning of the polar vortex caused by GW hotspots, which are mainly located southward of 50°N.
- Based on observations the locations of the GW hotspots apart from the H3 GW hotspot are not quite realistic (only jet sources may generate those).
- In another sensitivity study we displaced the H3 GW hotspot longitudinally in 45° steps. These GW hotspots are partly located above the Rocky Mountains and the Himalaya (orography) and along the polar front jet (jet sources).

Albers, J. R. and Birner, T.: Vortex Preconditioning due to Planetary and Gravity Waves prior to Sudden Stratospheric Warmings, J. Atmos. Sci., 71, 4028–4054, https://doi.org/10.1175/JAS-D-14-0026.1, 2014. Lilienthal, F., Jacobi, Ch., Schmidt, T., de la Torre, A. and Alexander, P., 2017: On the influence of zonal gravity wave distributions on the Southern Hemisphere winter circulation, Ann. Geophys., 35, 785-798, https://doi.org/0.5194/angeo-35-785-2017 Šácha, P., Kuchar, A., Jacobi, C., and Pišoft, P.: Enhanced 5 internal gravity wave activity and breaking over the Northeastern Pacific / Eastern Asian region, Atmos. Chem. Phys., 15, 13 097–13 112, https://doi.org/10.5194/acp-15-13097-2015, 2015. Samtleben, N., Jacobi, C., Pišoft, P., Šácha, P., and Kuchar, A.: Effect of latitudinally displaced gravity wave forcing in the lower stratosphere on the polar vortex stability, Ann. Geophys. Discuss., https://doi.org/10.5194/angeo-2019-15, in review, 2019. Acknowledgements This study has been supported by Deutsche Forschungsgemeinschaft (DFG) under the grant JA836/32-1. ECMWF reanalysis data are provided by apps.ecmwf.int/datasets/data/.