

Introduction.

To define a background in the electromagnetic emissions above seismic regions, it is necessary to define the statistical distribution of the wave energy in absence of seismic activity. Nemec et al. [2008], built a map of electromagnetic emissions containing the statistical description of the intensity of electromagnetic waves obtained from the entire DEMETER satellite data set. Then, they estimated the probability of occurrence of a signal during a seismic event whose intensity results to be different to the background level defined by the map.

CSES EFD Data

The orbits of the CSES satellite were sun-synchronous, had an inclination of 97.4° and a local time of 14.00 on day-side and 02.00 on night-side. The instruments were operational at geomagnetic latitudes between -65° to +65° [Shen et al., 2018]. The data from the electric field experiment EFD [Huang et al., 2018] were used in order to detect any electromagnetic waves. We selected the ELF band (from 13 Hz to 5 kHz, with a sample rate of 2.5kHz). We analysed 1 month data of the satellite observations (August 2018), focusing on L'Aquila (Italy) and Bayan (Indonesia) location, evaluating both the electric field instrumental and environmental background to be used for the detection of possible electric field signal of seismic origin. For this purpose, we applied a completely the method of Bertello et al. [2018] to determine the environmental background based on ALIF technique [Piersanti et al., 2018].

Multiscale statistics via ALIF and Standardized Mean (SM) test

The algorithm for the evaluation of the energy background is based on Adaptive Local Iterative Filtering (ALIF) technique [Piersanti et al. (2018)]. ALIF first decomposes a signal (s(t)) into several Intrinsic Mode Functions (IMFs), which are functions oscillating around zero, but not necessarily with constant frequencies.

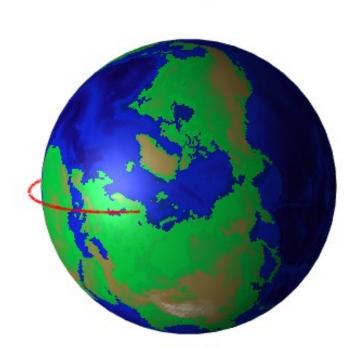
$$s(t) = \sum_{\ell=1}^{m} IMF_{\ell}(t) + r(t)$$

and the scale of variability of s(t) have been analysed by using the method proposed by Flandrin et al. (1998):

$$s(t) = s_0(t) + \delta s(t)$$

where $s_0(t)$ is the baseline and $\delta s(t)$ are the variations around $s_0(t)$. The basic idea is that $\delta s(t)$ has a standardized mean (SM: defined as the mean divided by the standard deviation) which is close to zero and represents the fluctuating part of s(t). That is:

The k value represents the last mode index for which the reconstruction given by Eq. (1) has an SM close to zero.



CSES Orbit Example: 2018/08/04

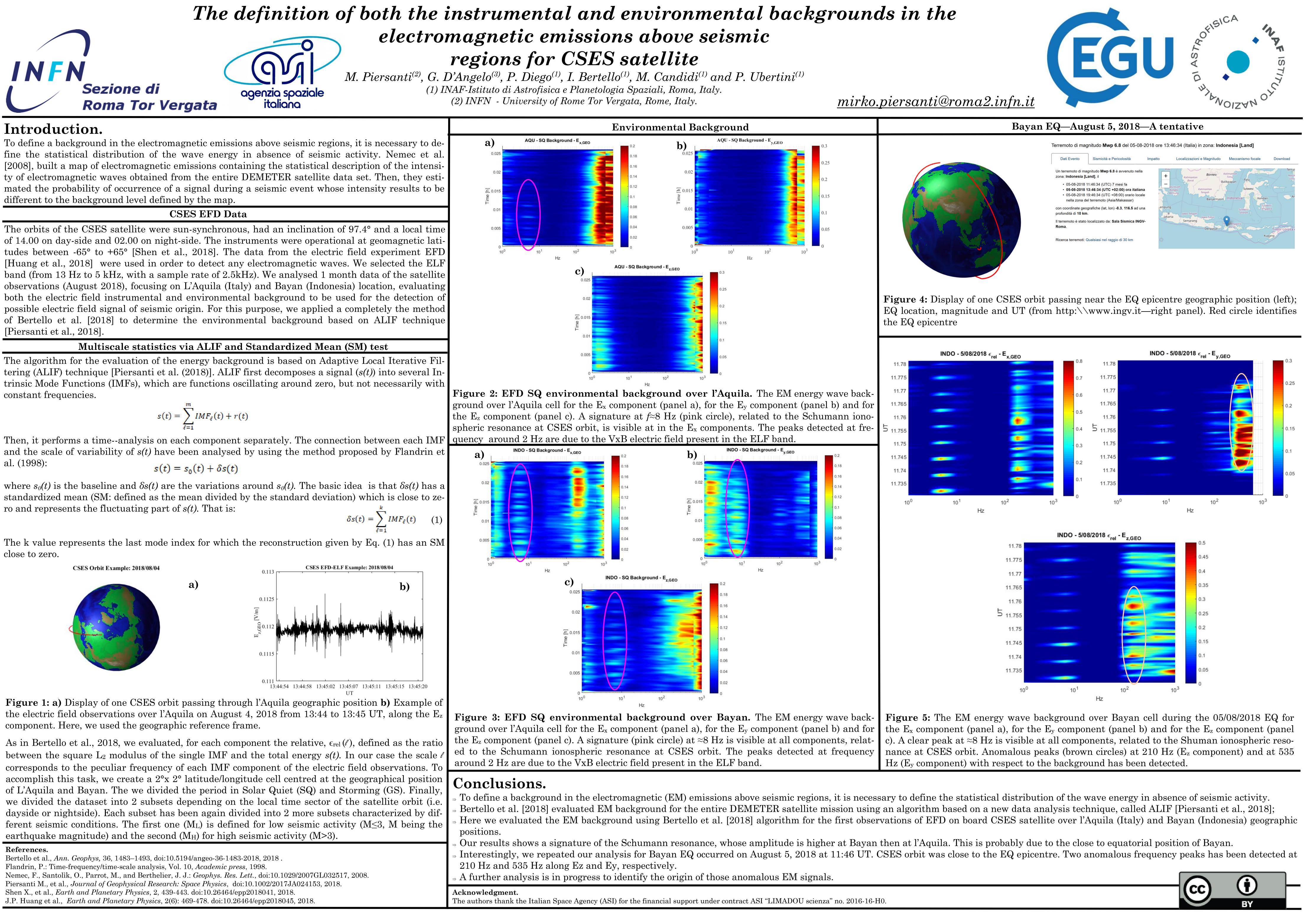


Figure 1: a) Display of one CSES orbit passing through l'Aquila geographic position b) Example of the electric field observations over l'Aquila on August 4, 2018 from 13:44 to 13:45 UT, along the E_z component. Here, we used the geographic reference frame.

As in Bertello et al., 2018, we evaluated, for each component the relative, $\epsilon_{rel}(\ell)$, defined as the ratio between the square L₂ modulus of the single IMF and the total energy s(t). In our case the scale corresponds to the peculiar frequency of each IMF component of the electric field observations. To accomplish this task, we create a 2°x 2° latitude/longitude cell centred at the geographical position of L'Aquila and Bayan. The we divided the period in Solar Quiet (SQ) and Storming (GS). Finally, we divided the dataset into 2 subsets depending on the local time sector of the satellite orbit (i.e. dayside or nightside). Each subset has been again divided into 2 more subsets characterized by dif ferent seismic conditions. The first one (M_L) is defined for low seismic activity (M \leq 3, M being the earthquake magnitude) and the second (M_H) for high seismic activity (M>3).

References.

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