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Investigation of Aerosol Optical Properties in the European Arctic using Lidar remote sensing technique



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Motivation-Research Questions

- **Persistent lofted layers** over Ny-Ålesund and Fram Strait
- **Optical properties** in different transport events?
- **Radiative impact** of different transport events?
- **Contribution to Arctic Amplification (AA)?** [1]

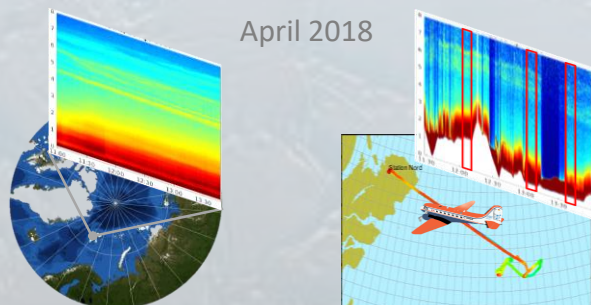


Fig 1: Lidar range-corrected signal at 532 nm (arbitrary units) measured over Ny-Ålesund (left) and Fram Strait (right).

Instrumentation

Ground-based system KARL [2]

- Raman aerosol Lidar
- “3β+2α+2δ+2wv” system
- Nd: YAG laser (1064, 532, 355 nm)

Air-borne system AMALI [3]

- Elastic aerosol Lidar
- “2β+1δ” system
- Nd: YAG laser (532, 355 nm)

Baseline Surface Radiation Network [4]

- Pyranometer → SW diffuse, global, reflected (0.2-3.6 μm)
- Pyrhelimeter → SW direct (0.2-4 μm)
- Pyrgeometer → LW fluxes (3.5-50 μm)

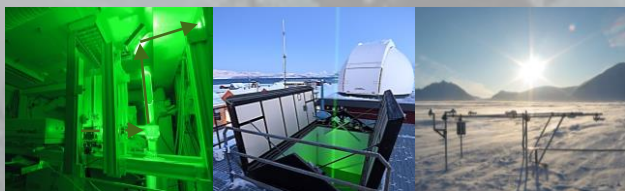


Fig 2: Emitting unit, comprising laser head, optical components and beam widening telescope (left), and hatch of KARL system (middle). The Baseline Surface Radiation Network (BSRN) station, which is located in Ny-Ålesund (right).

Elevated Layers over Ny-Ålesund in the past decade

Ground-based Lidar Observations

April 2009

March 2003

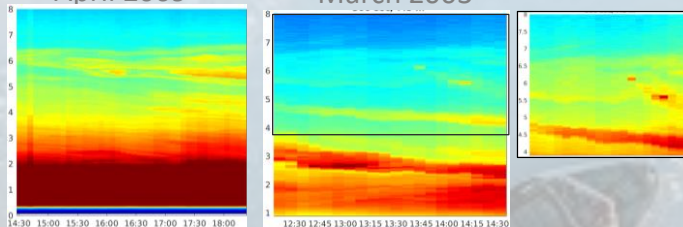


Fig 3: Lidar range-corrected signal at 532 nm (arbitrary units) over Ny-Ålesund on 1st April 2009 (left) and 14th March 2003.

Inversion of Aerosol Optical Properties

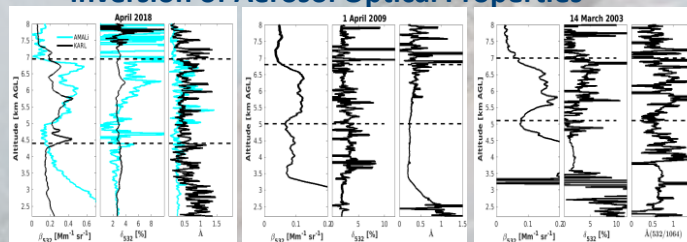


Fig 4: Aerosol optical properties on 5th April 2018 (left), 1st April 2009 (middle) and 14th March 2003 (right).

$$\delta^{aer}(\lambda) = \frac{\beta^{aer}(\lambda)}{\beta_{||}^{aer}(\lambda)}$$

$$\hat{A}(\lambda_1, \lambda_2) = -\frac{\log \frac{\lambda_1}{\lambda_2}}{\log \frac{\beta_{\lambda_1}}{\beta_{\lambda_2}}}, \lambda_1 < \lambda_2$$

Radiative Impact of Aerosol

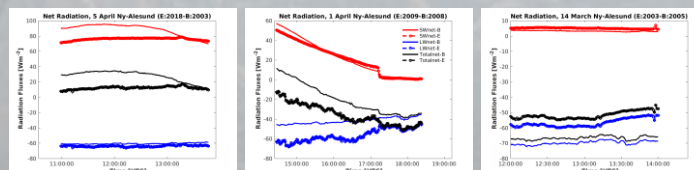


Fig 5: Radiative fluxes on ground level at Ny-Ålesund in April 2018/2003 (left), in April 2009/2008 (middle) and in March 2003/2005 (right).

$$F_{net}^{TOTAL} = (F_{in}^{SW} - F_{out}^{SW}) + (F_{in}^{LW} - F_{out}^{LW})$$

$$\Delta F_{net}^{TOTAL} = F_{netE}^{TOTAL} - F_{netB}^{TOTAL} = (F_{netE}^{SW} - F_{netB}^{SW}) + (F_{netE}^{LW} - F_{netB}^{LW})$$

Findings

- Elevated aerosol layers (Fig. 1; Fig. 3) observed in the Arctic over the past decade (2003-2018) on an infrequent basis
- Inversion of optical properties [5], [6] in 2018 yielded higher aerosol backscatter coefficient over Fram Strait ($\beta_{532}^{AMALI} = 0.42 \pm 0.15 \text{ Mm}^{-1} \text{sr}^{-1}$; $\beta_{532}^{KARL} = 0.34 \pm 0.08 \text{ Mm}^{-1} \text{sr}^{-1}$)
- In 2009 and 2003, lower β ($\beta_{532}^{KARL} = 0.3 \pm 0.09 \text{ Mm}^{-1} \text{sr}^{-1}$; $\beta_{532}^{KARL} = 0.2 \pm 0.09 \text{ Mm}^{-1} \text{sr}^{-1}$)
- Nearly spherical particles (low depolarization)
- Aerosol size parameter indicates low spectral dependency (Ångstrom exponent ≈ 1)
- Lidar ratio ($LR_{355nm}^{2018} = 40 \text{ sr}$; $LR_{355nm}^{2009} = 38 \text{ sr}$) suggesting mixture of smoke and polluted continental particles [7], [8], [9], similar to iAREA campaign in spring 2014 [10]
- In 2018, moderate total net warming of 12 Wm^{-2} at ground level (Fig. 5) along with negative total aerosol forcing ($AF = -15.3 \text{ Wm}^{-2}$)
- In 2009, total net ground cooling of -36 Wm^{-2} (Fig. 5) and negative total AF (-15.4 Wm^{-2})
- In 2003, significant total net cooling of -48 Wm^{-2} at the surface, mainly due to LW (Fig. 5), while positive total AF (15 Wm^{-2})

Conclusions

- ✓ Rare occurrence of elevated layers in the Arctic but impact on radiation budget
- ✓ Similar optical properties indicating smoke-polluted continental aerosol type
- ✓ Similar AF in layers occurring in April, although total net radiation balance different
- ✓ Layers with similar optical properties potentially have different impact on radiation balance. The role of insolation is potentially critical

Future Work

- Contribution to Arctic Amplification needs further investigation
- Aerosol radiative impact at different atmospheric levels by means of SCIATRAN RTM simulations [8]
- Inversion of aerosol microphysical properties and closure with in-situ measurements [9], [10]

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[1] M.C. Serreze et al. Glob. Planet. Change 77, 85-96 (2011)
 [2] A. Hoffmann, PhD Thesis Uni.Potsdam (2011)
 [3] I. Stachlewska et al. Atmos. Chem. Phys. 10, 2947-2963 (2010)
 [4] M. Maturilli, et al. Theor. Appl. Climatol. 120, 333-339 (2015)
 [5] J. D. Klett, Appl. Opt. 20, 211-220 (1981)

[6] A. Ansmann et al. Appl. Opt. 31, 7113-7131 (1992)
 [7] D. Müller et al. JGR 112 D16202 (2007)
 [8] S. Groß et al. Atmos. Chem. Phys. 13, 2487-2505 (2013)
 [9] A. Illingworth et al. Bull. Amer. Meteorol. Soc. doi:10.1175/BAMS-D-12-00227.1

[10] C. Ritter, et al. Atmos. Env. 141, 1-19 (2016)
 [11] V.V. Rosanov et al. JQSRT 197, 65-85 (2017)
 [12] C. Böckmann et al., Appl. Opt. 40, 1329-1342 (2001)
 [13] S. Samaras et al. J. Comput. Phys. 299, 156-174 (2015)