

On the Nature of Fault Slip: From the Field to the Lab

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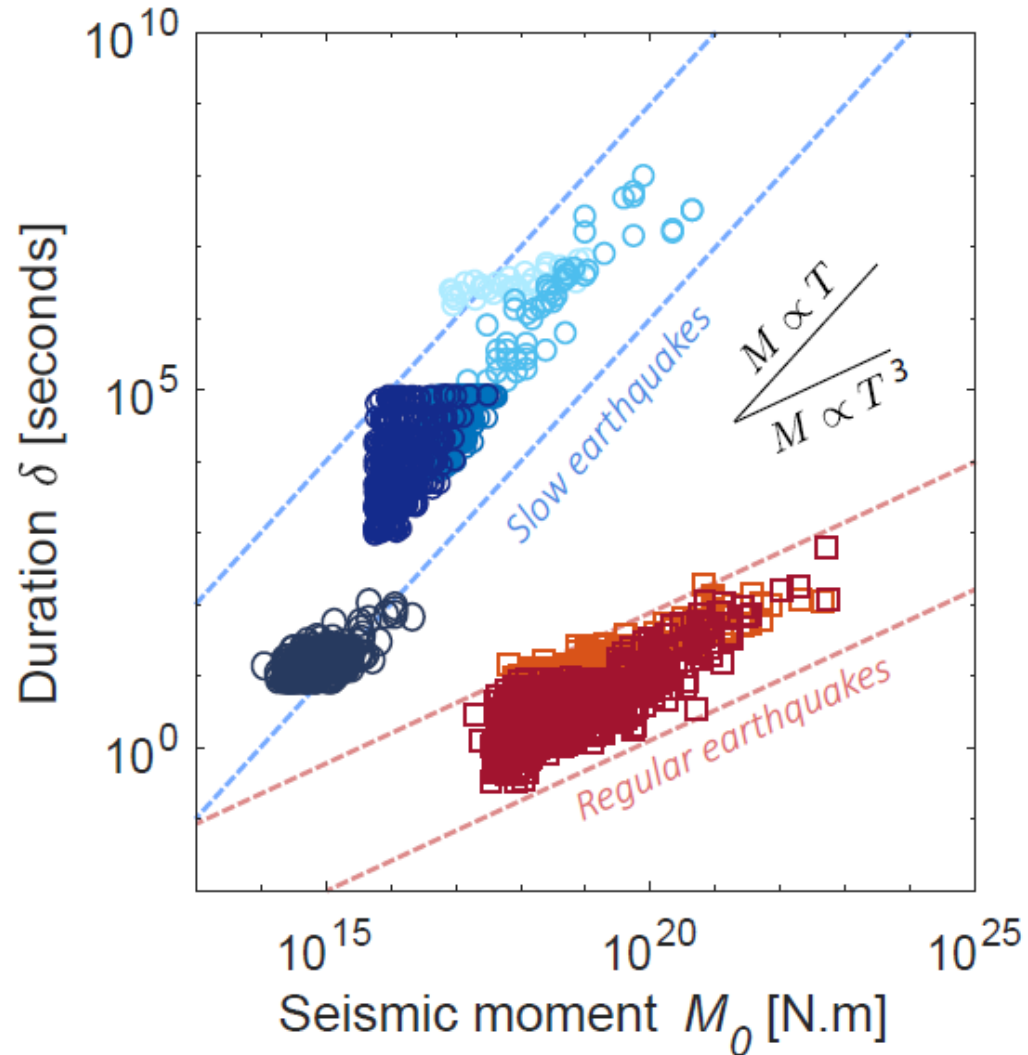
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Various types of seismic ruptures

Passelègue et al., submitted



Scaling relations

$$M_0 \propto \Delta\sigma L^3$$

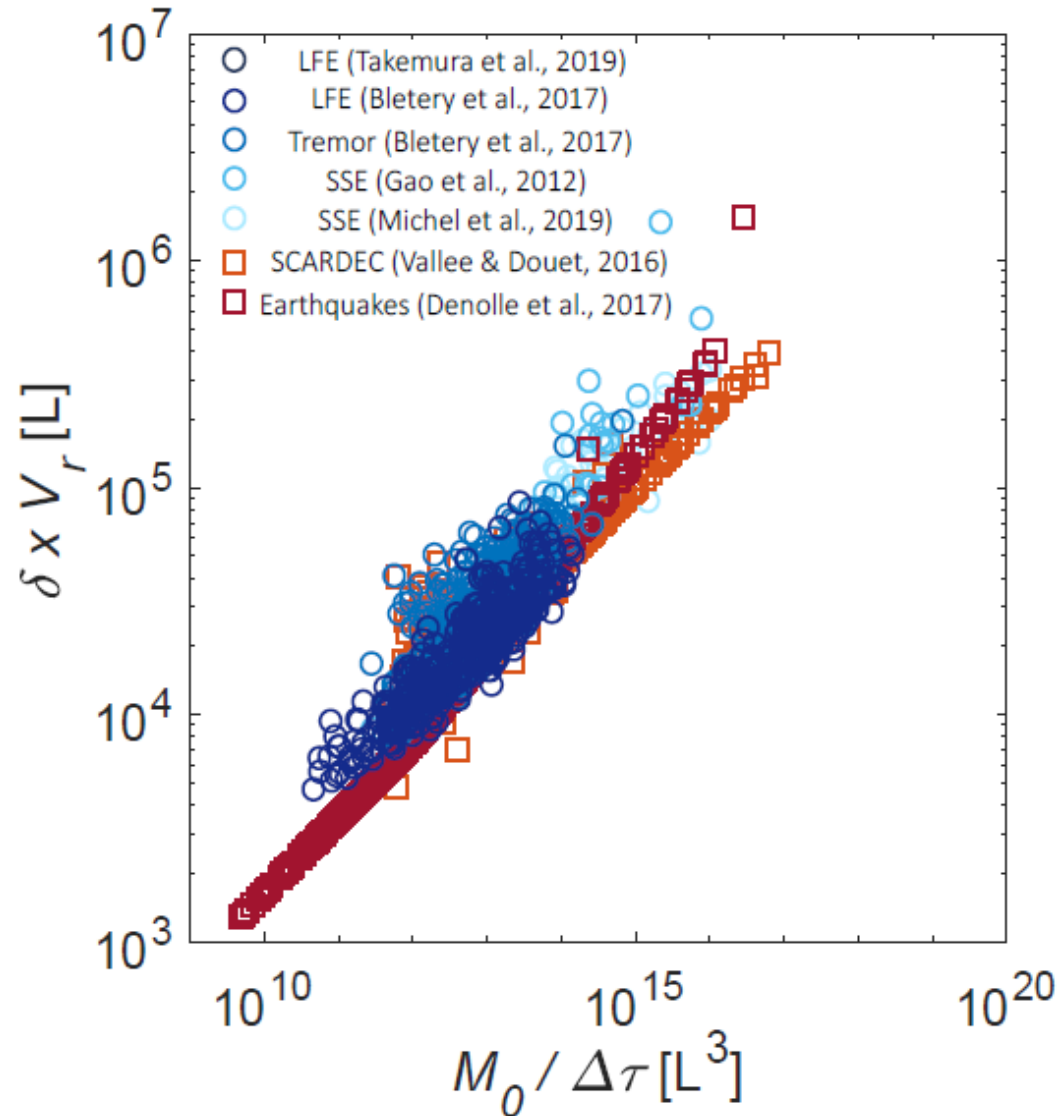
$$M_0 \propto \mu D L^2$$

$$\Delta\sigma \propto \mu \frac{D}{L}$$

$$\delta \propto L / V_r$$

Various types of seismic ruptures

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Scaling relations

$$M_0 \propto \Delta \sigma L^3$$

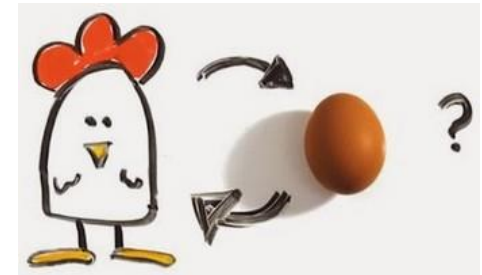
$$M_0 \propto \mu D L^2$$

$$\Delta \sigma \propto \mu \frac{D}{L}$$

$$\delta \propto L / V_r$$

Same physic

$$V_r \propto \Delta \sigma$$

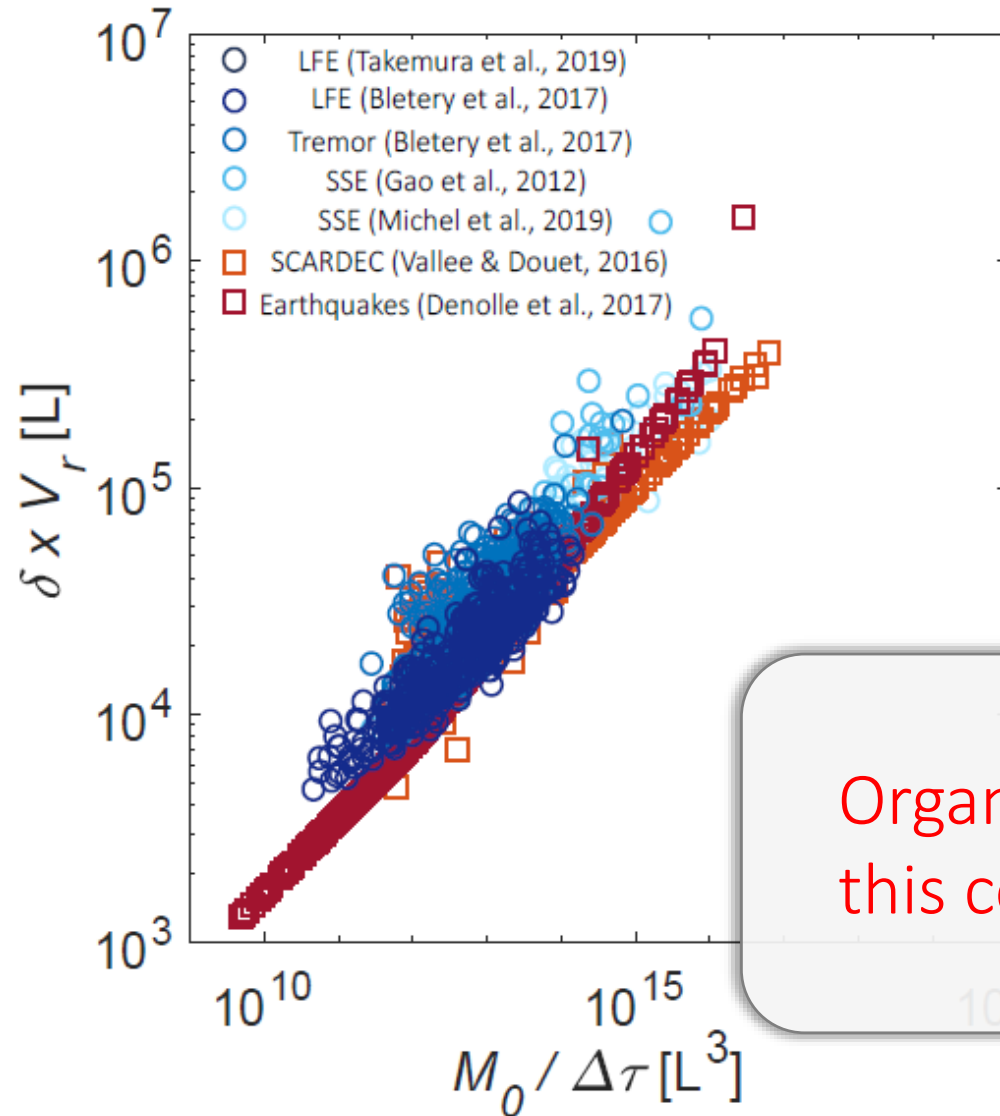


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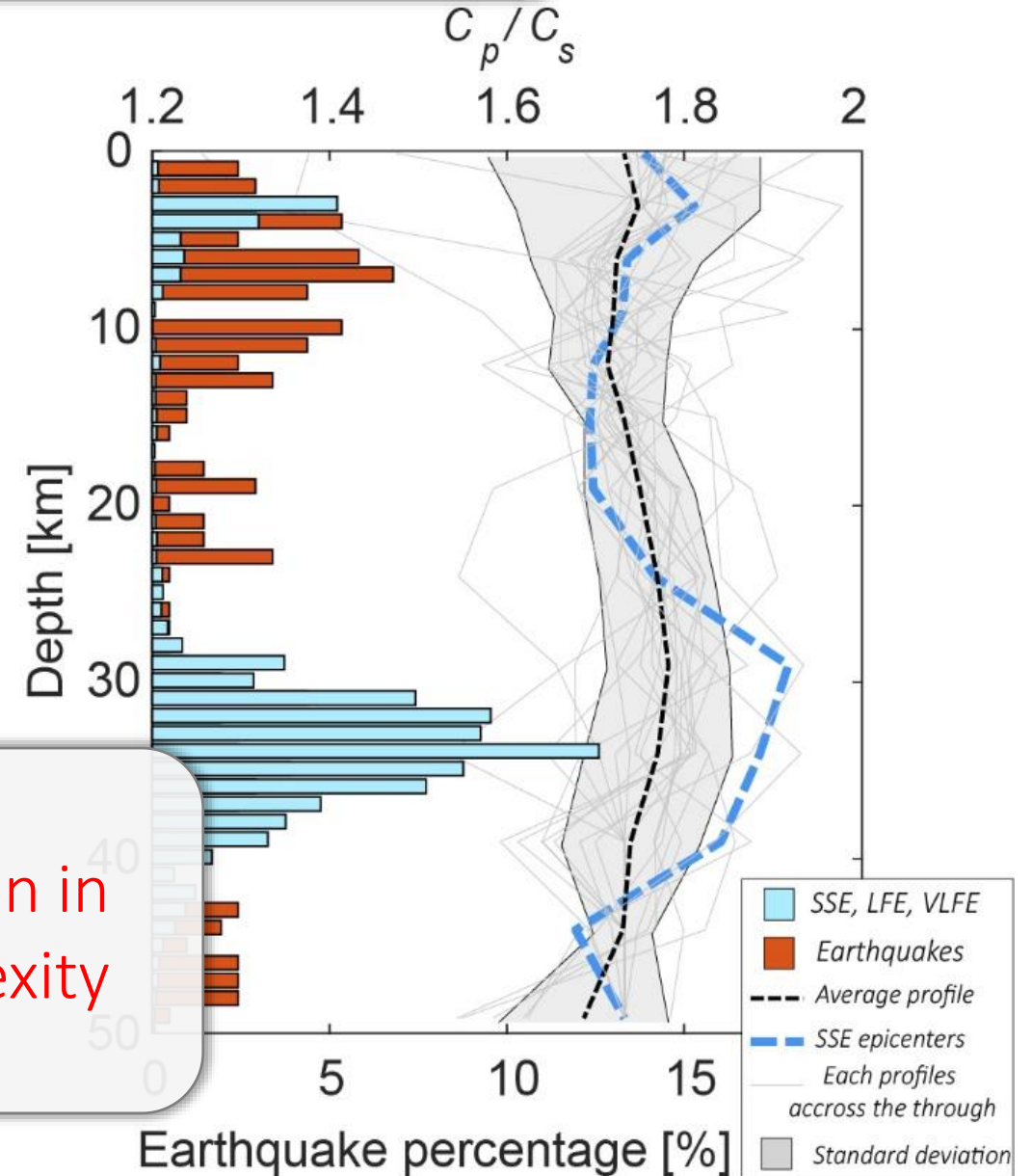
Work in progress

Various types of seismic ruptures

Passelègue et al., submitted



Organisation in
this complexity



Which parameter(s) control(s) the nature of seismicity?

Something that:

- **changes spacially along faults**

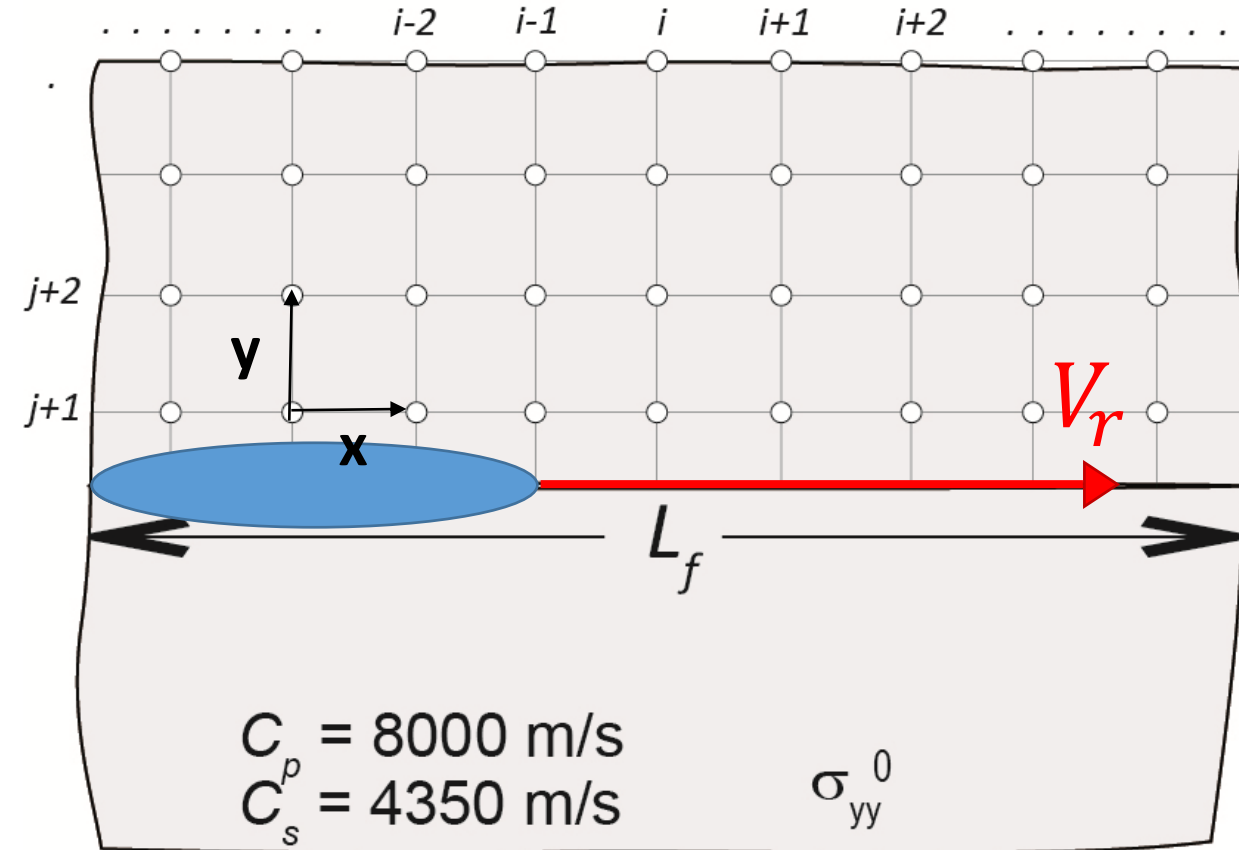
Fault jogs can present both slow and fast ruptures

- **Is variable with time**

A same asperity can break seismically and aseismically

- **Dominates at shallow depth or close to BDT,
affected by fluid pressure**

Reproduce earthquakes in the laboratory: **controlled conditions!**



Measurement of stress along the fault

Regular triaxial experiments

$$\sigma_{xy}(i, j) \text{ \& \; } \sigma_{yy}(i, j) = \textit{constant}$$

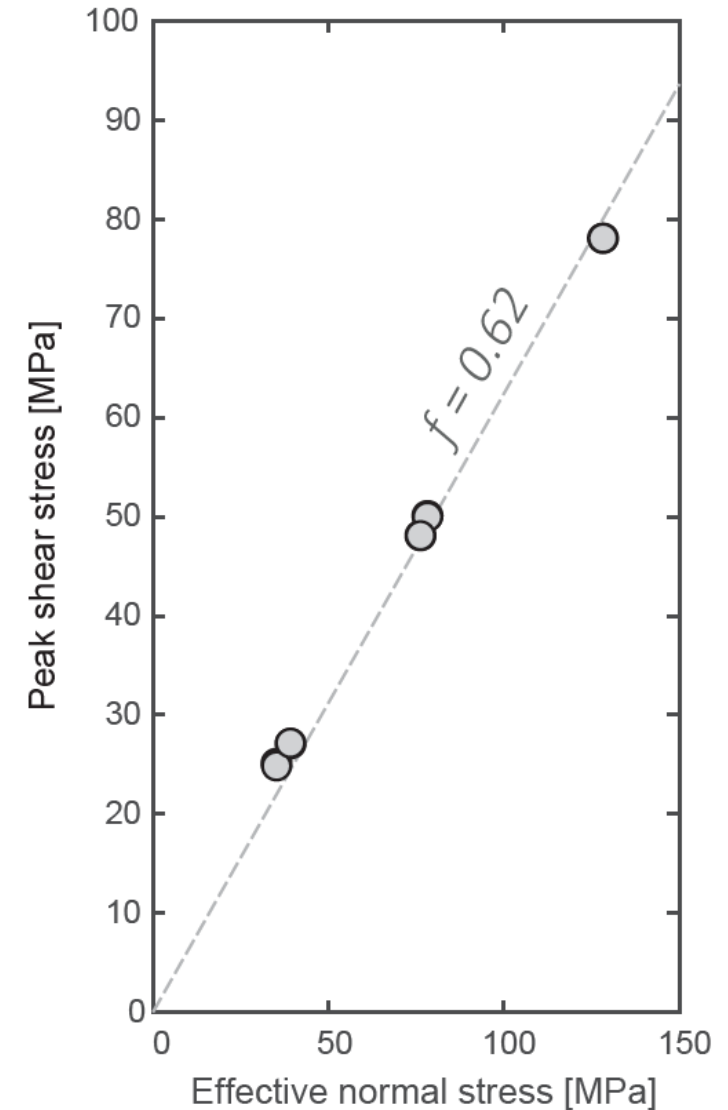
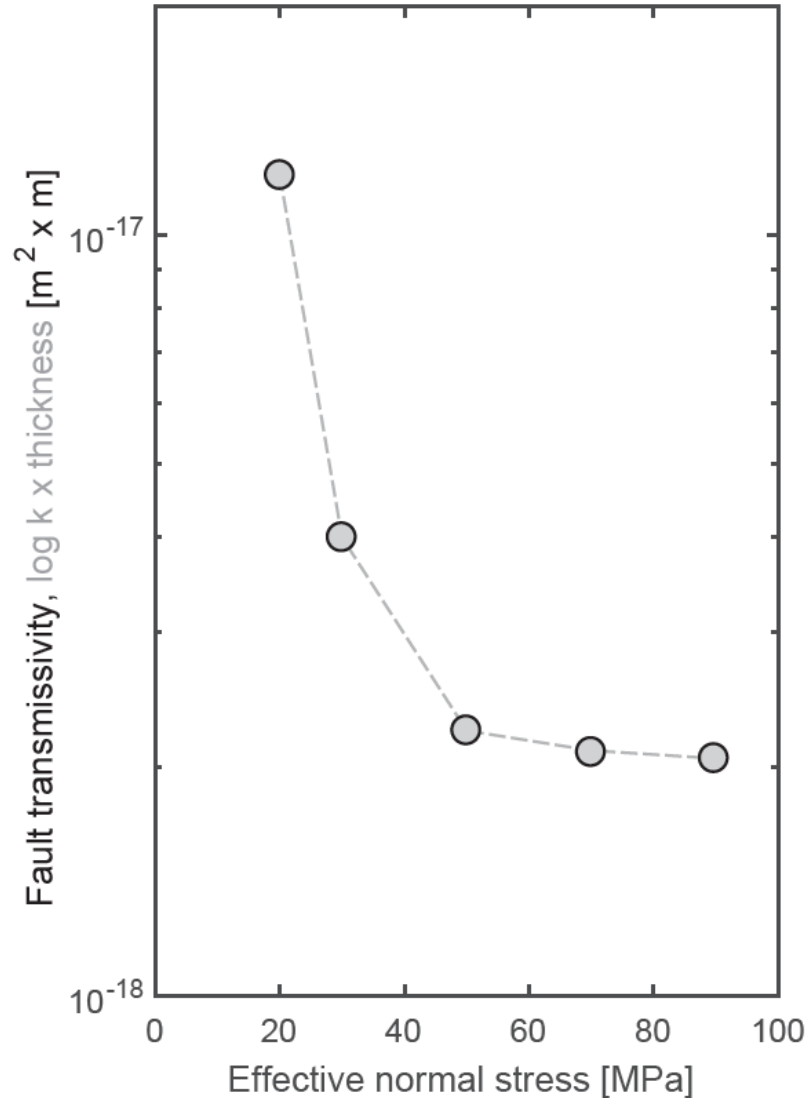
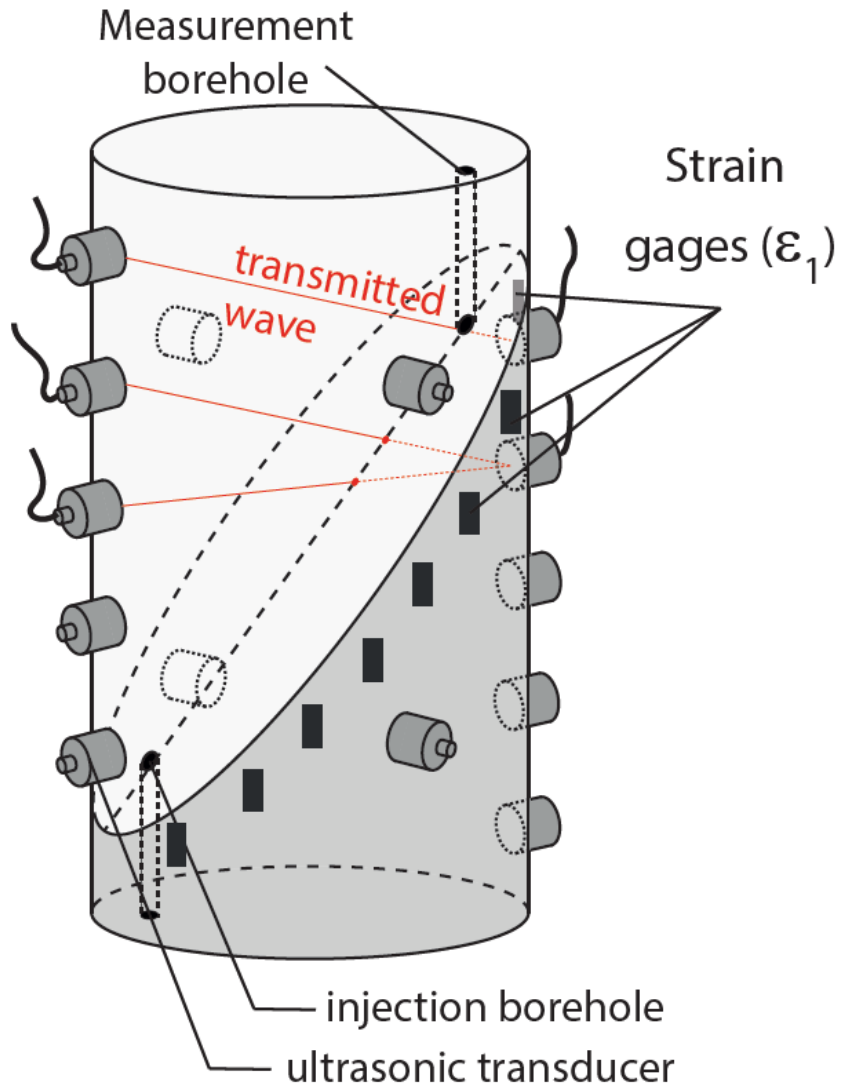
Here

$$\sigma_{xy}^0(i, j) = \textit{constant} \quad \sigma_{yy}^{eff}(i, j) \text{ vary spacially}$$

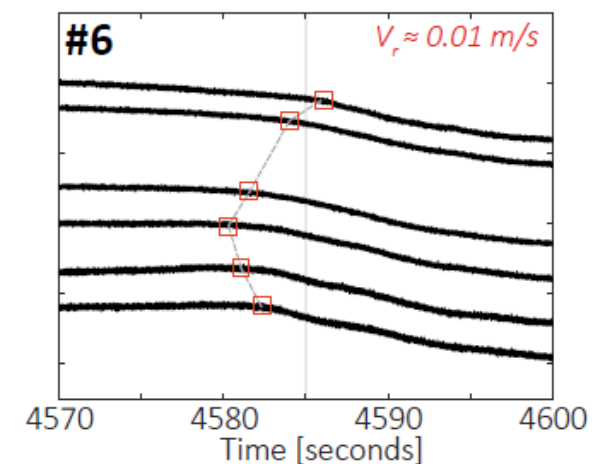
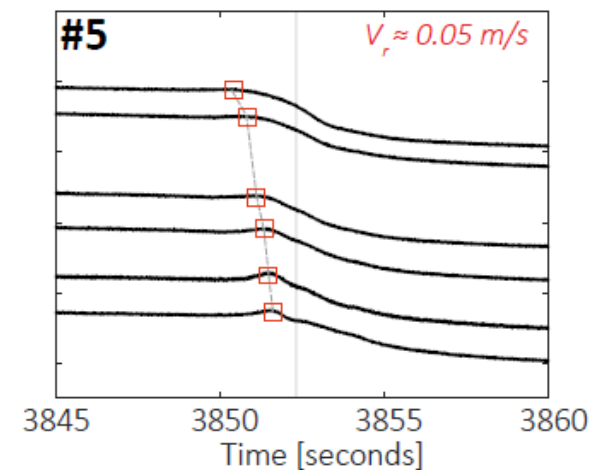
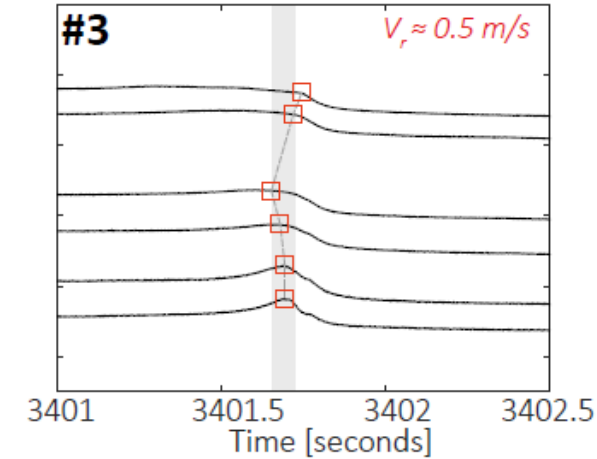
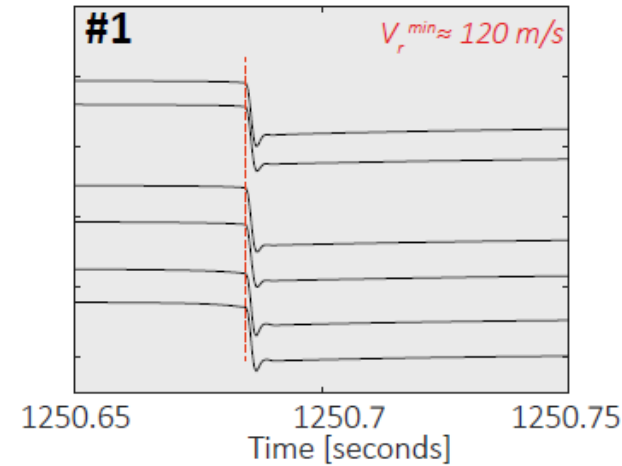
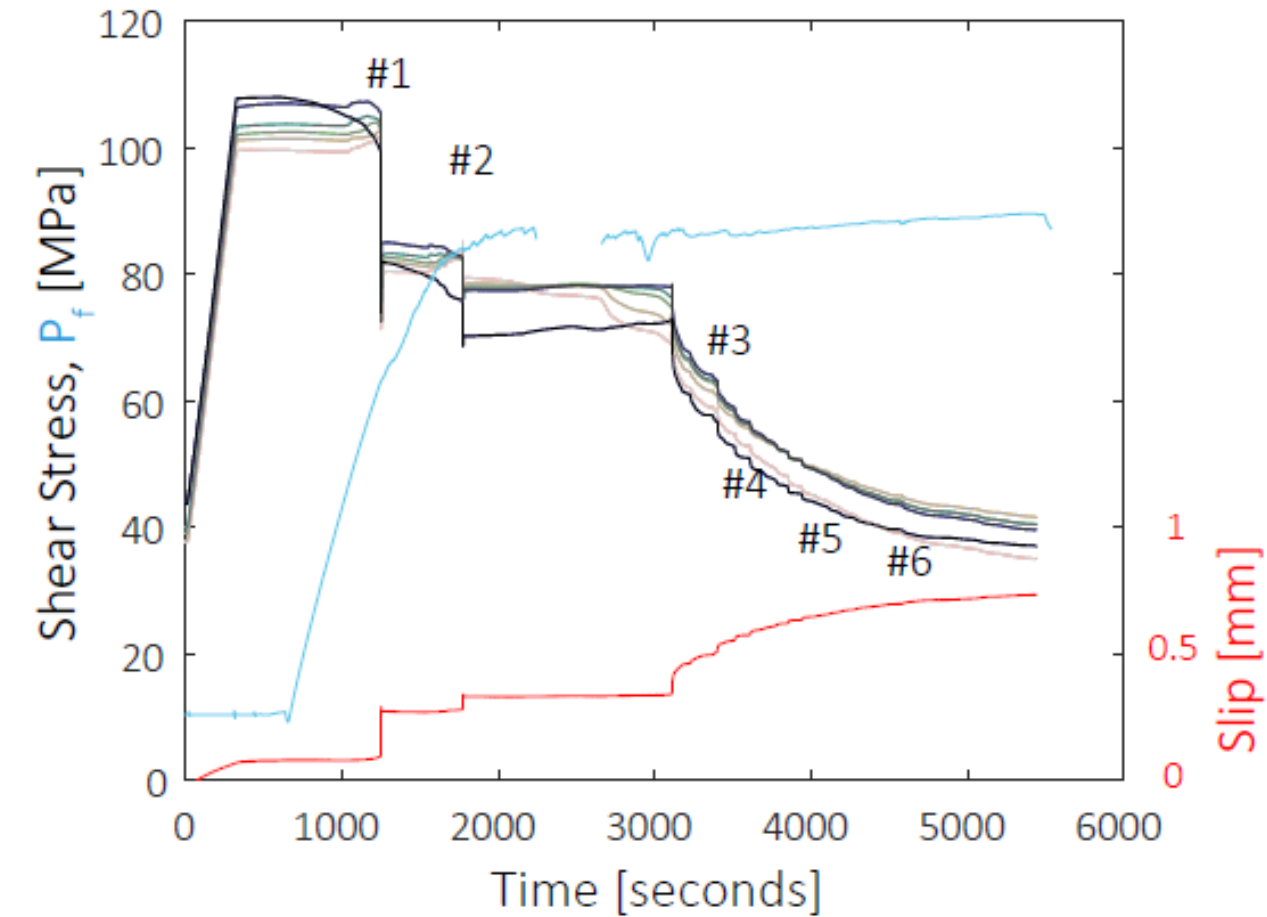
Fluid injection at the edge of the fault
Low fault permeability, no bulk diffusion

Inversion of the fluid pressure
(Almakari et al, in prep)

Reproduce earthquakes in the laboratory: **controlled conditions!**



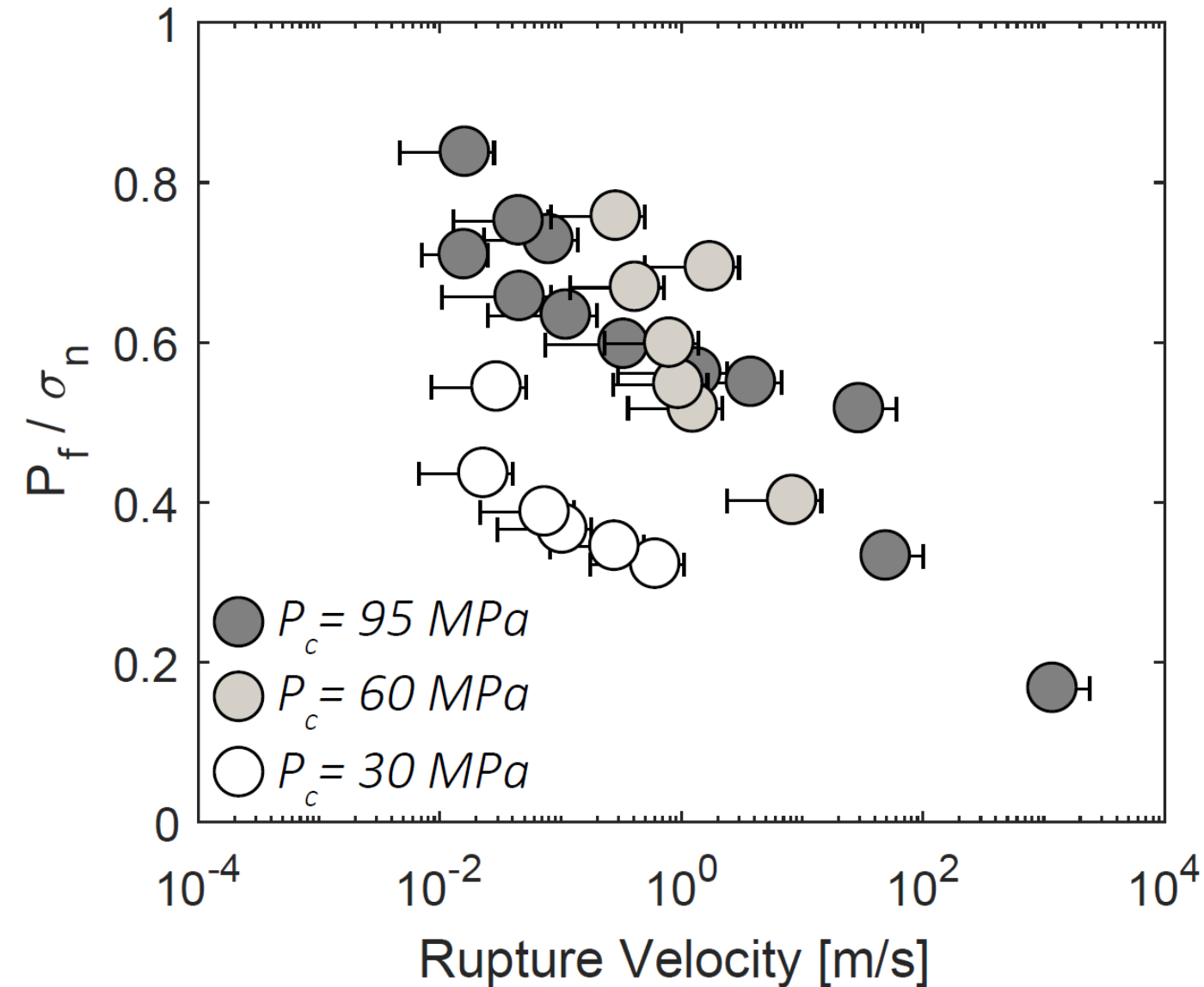
Experimental results: From fast to transient slow slip events



Slip rate and stress drop decrease with decreasing τ_0

Rupture speed also: transition from fast to slow rupture

What about the control of the rupture speed?



$$V_r \searrow Pf \propto L_c?$$

$$L_c < L_f$$

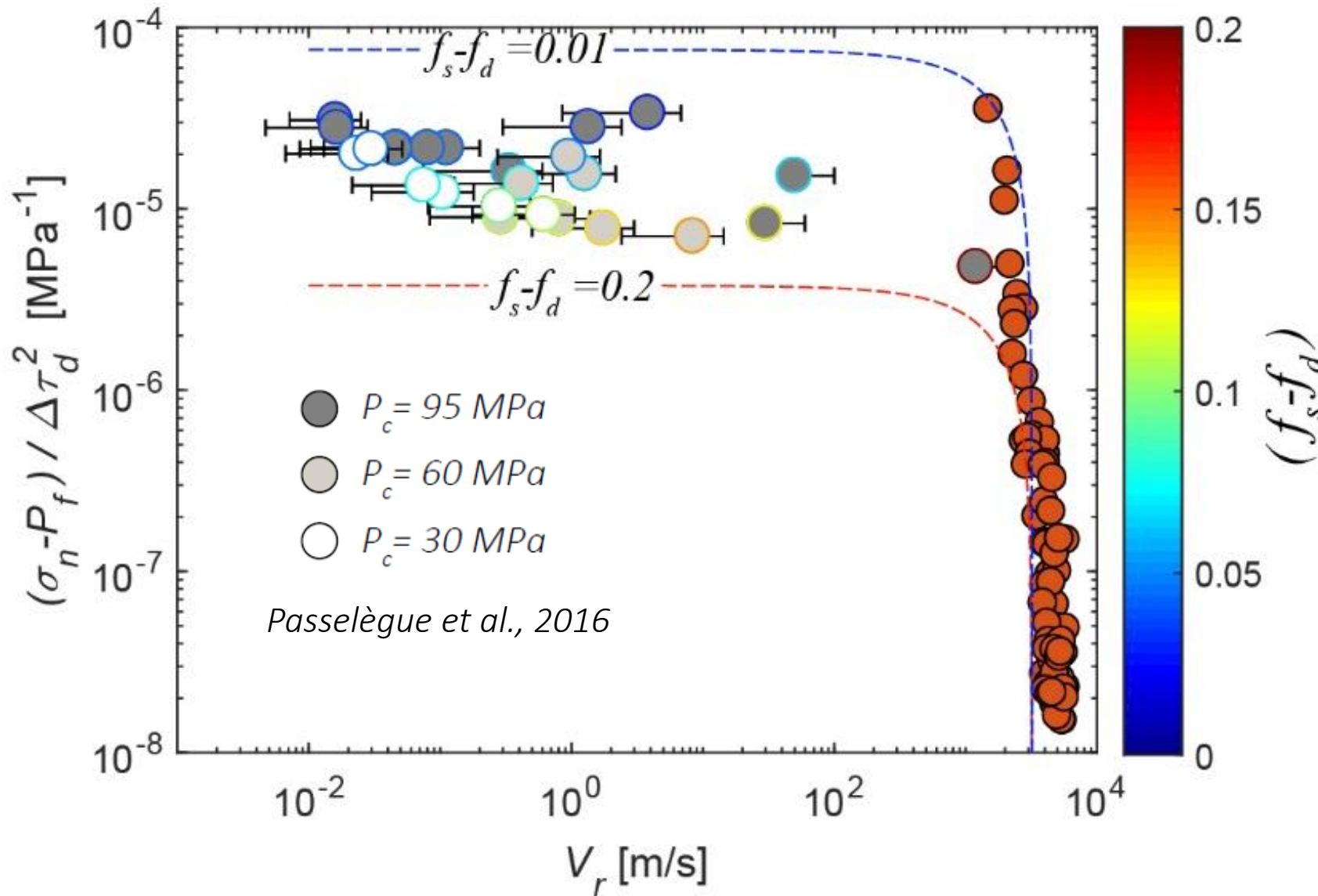
From LEFM (*Freund, 1990*)

$$V_r = C_R \left(1 - \frac{G_c}{\Gamma} \right)$$



$$V_r = C_R \left(1 - \frac{(\overline{\sigma_n} - \overline{P_f})}{(\tau_0 - \tau_r)^2} \frac{\Omega E^*}{\pi l / 2} \right)$$

What about the control of the rupture speed?



$$V_r \searrow Pf$$

$$V_r \nearrow \Delta \tau^d$$

Passelègue et al., 2016

Initial stress τ_0
& Energies are keys!

On going work and perspectives

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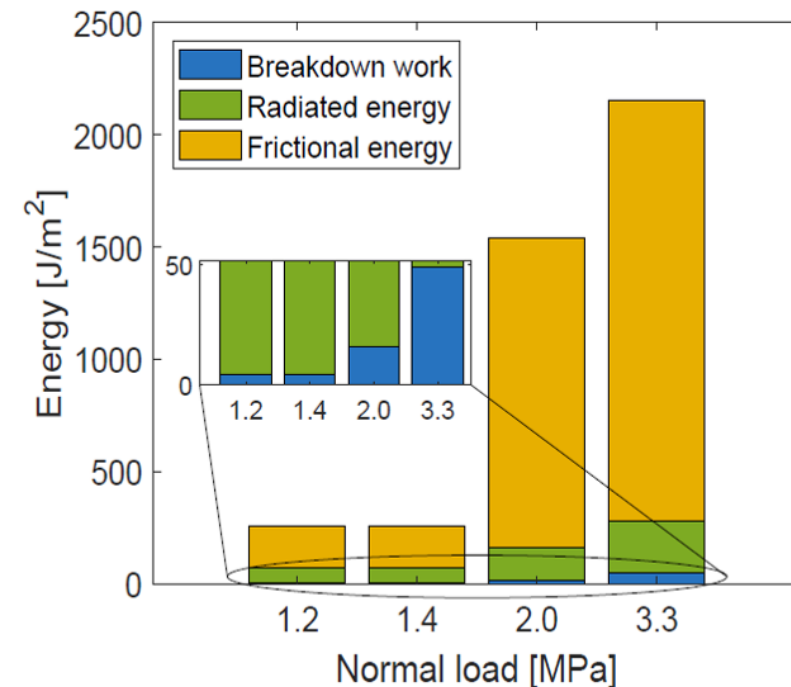
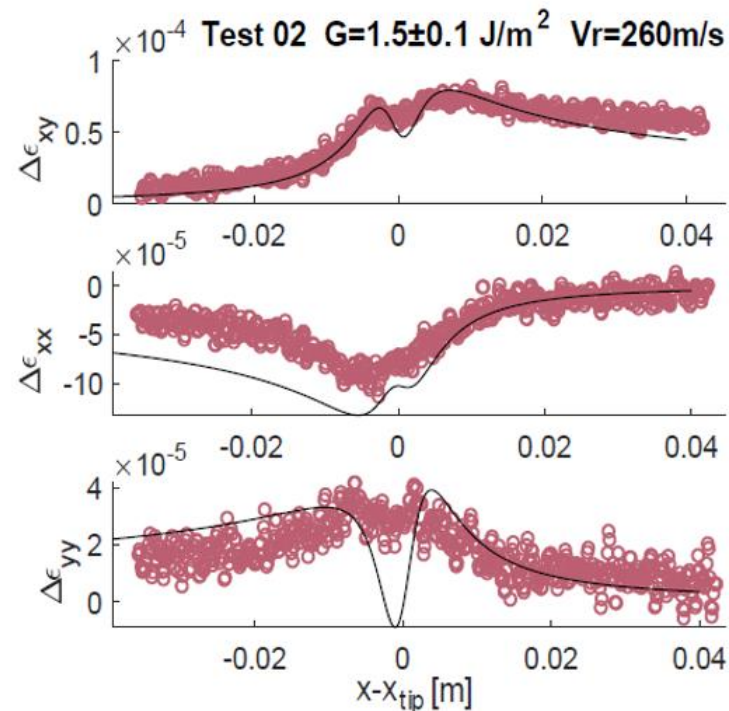
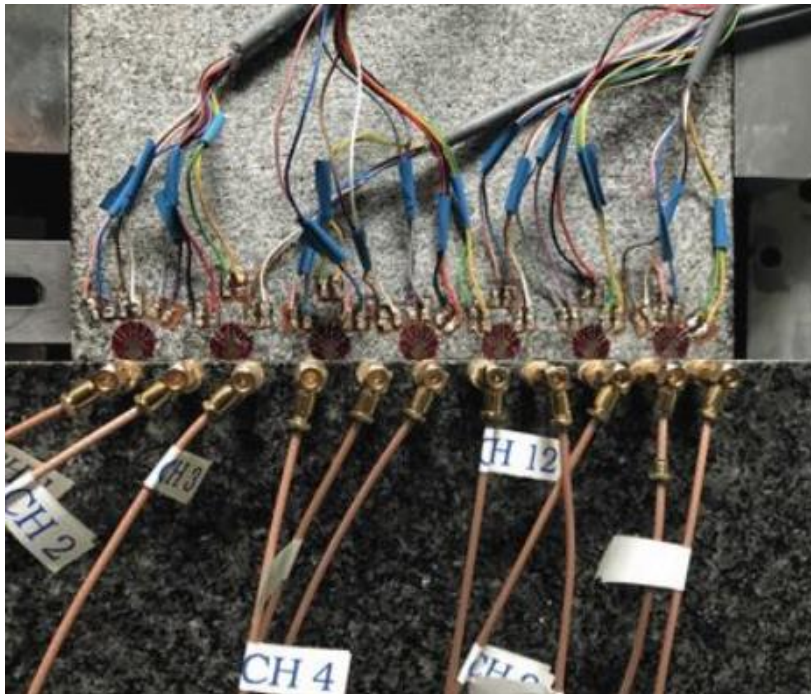
She is closing the energy budget in the lab: E_g , E_h , E_r

She compare the estimates of E_G (LEFM, Sismo, Stress-slip)

Study deceleration and rupture arrest (control on τ_0)

Wednesday

TS2.5/EMRP1.12



Conclusions and Take home message

Rupture speed, i.e. nature of the seismicity,
depends of
the **stress acting along the fault** ahead of the rupture tip, **the energy available**

- Explain slow rupture phenomena at shallow depth and in fluid over-pressurized area
- What about rupture length? Also predictable from LEFM!
- Low stress level inhibits radiation damping terms $\rightarrow V_r$ remains slow (*Barras et al., 2019*)

Small rupture of high stressed patches trigger slow rupture front?

