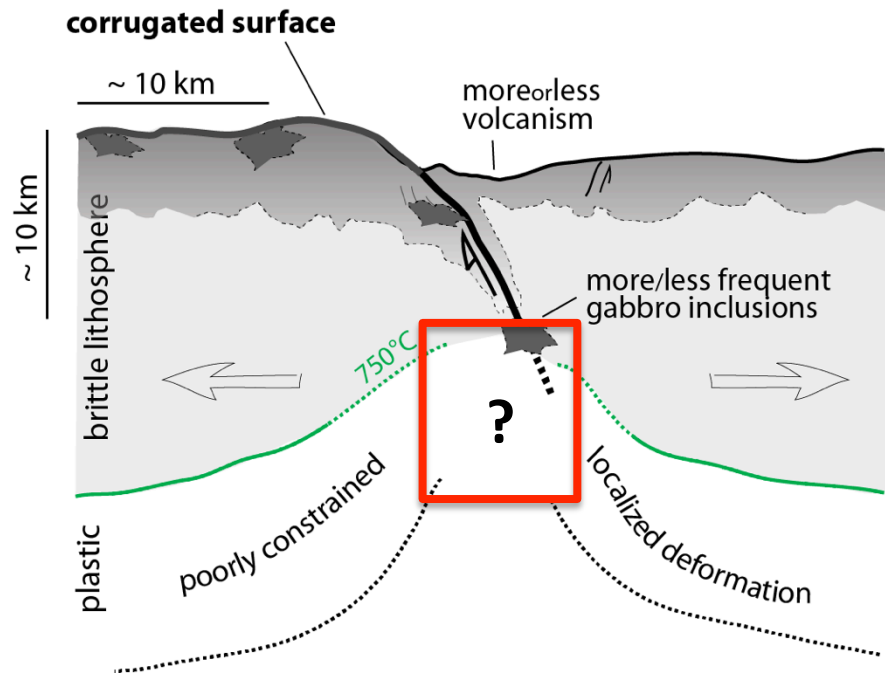


Strain localization processes at a magma-starved ridge: from micro-scale to macro-scale

Bickert¹, M., Cannat¹, M., Tommasi², A., Jammes³, S., Lavier⁴, L.

Detachment faults exhume mantle-derived rocks from the base on the brittle lithosphere to the seafloor.

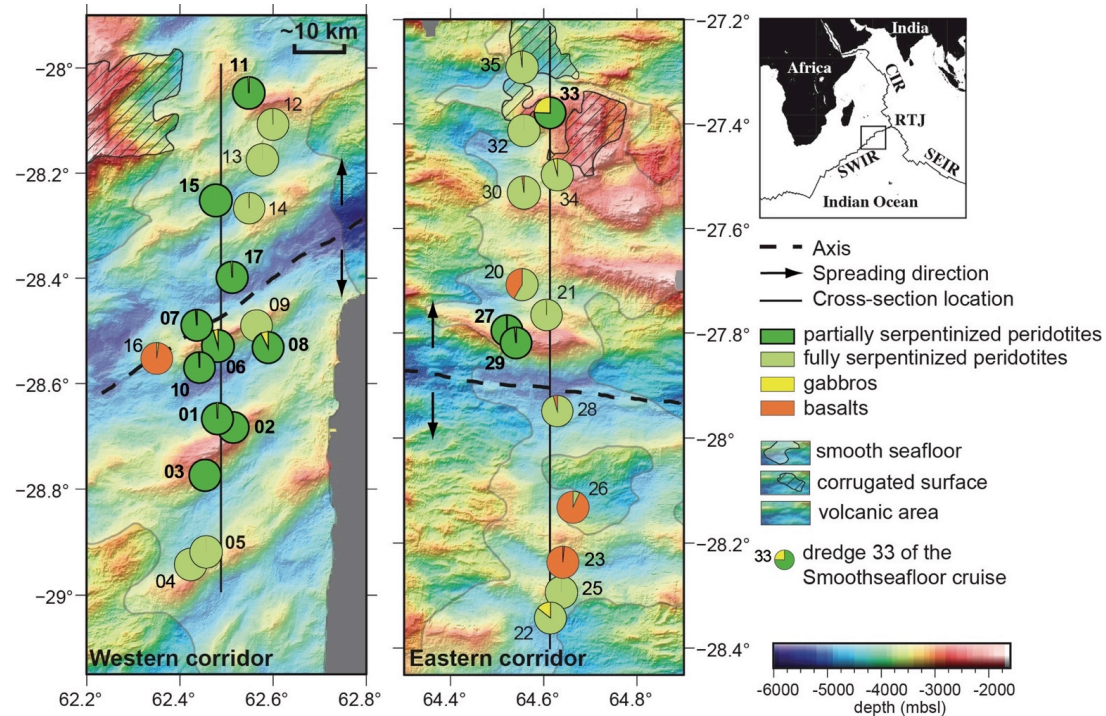
- What are the strain localization mechanisms in the deep axial lithosphere when there is **no magma** ?
- How to make detachments in a **thick** lithosphere (≥ 15 km) ?



[modified from Cannat et al., 2008]

The Eastern SouthWest Indian Ridge (SWIR)

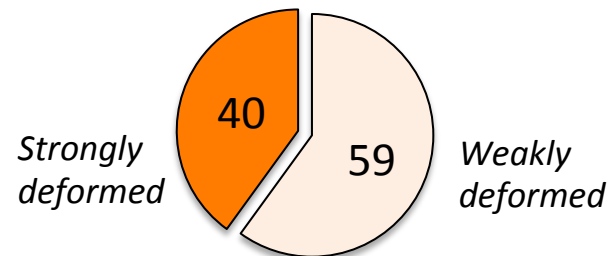
Dredges realized on and off-axis recovered **variably serpentinized peridotites**, with minor amounts of gabbros (< 4%) and basalt (16%).



385 samples of ultramafic rocks

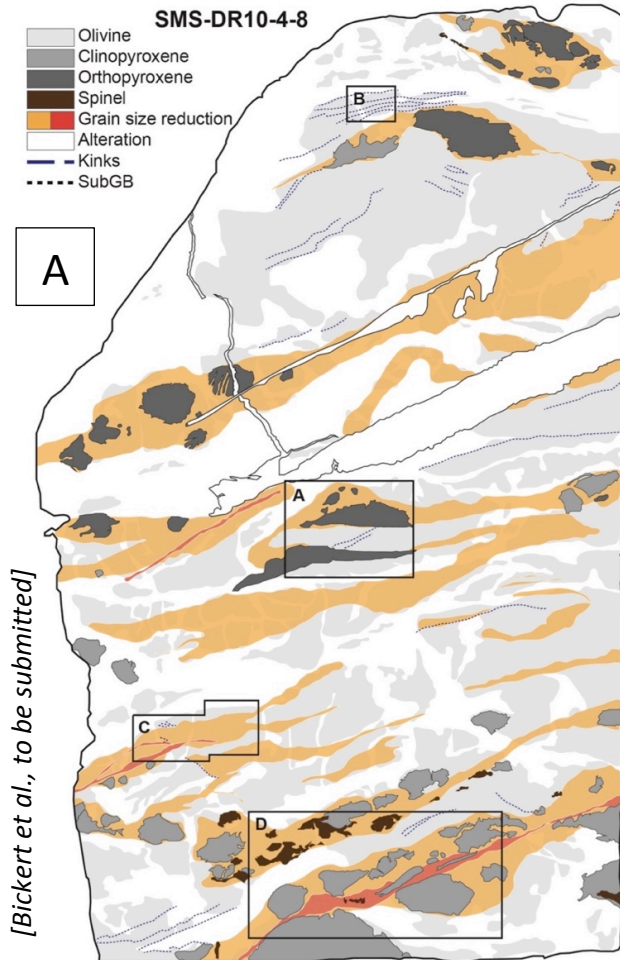
99 partially serpentinized samples

Partially serpentinized samples (99)



40% of the samples are **strongly deformed**, with **planar fine-grained zones**.

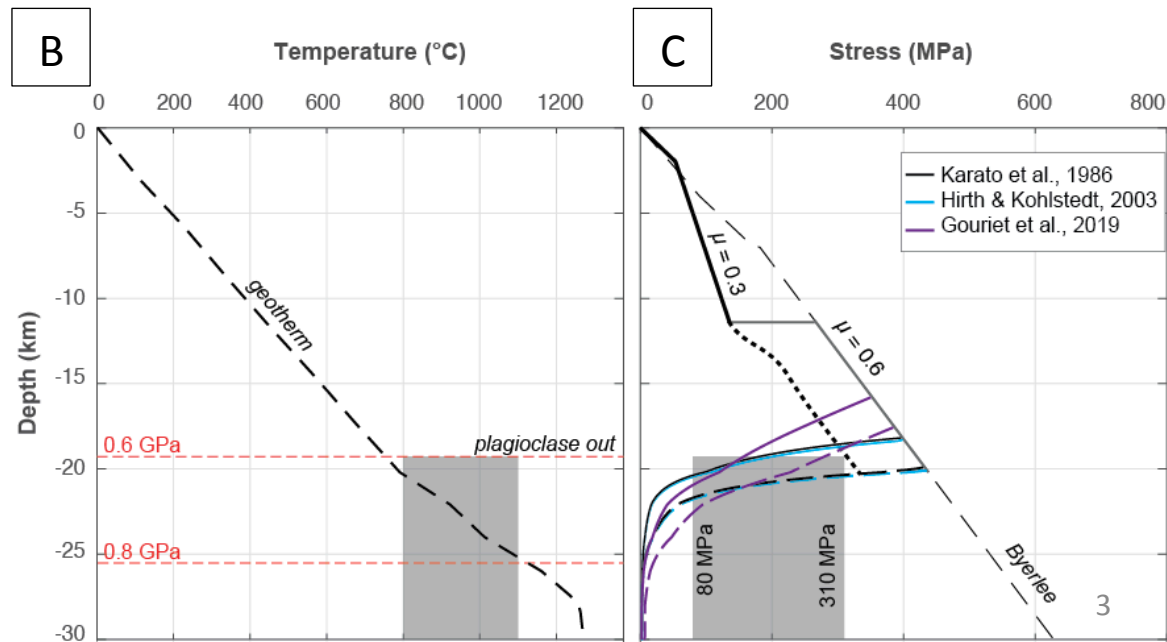
Extracting P-T conditions from grain size reduction zones



Heterogeneous deformation combining brittle and ductile mechanisms (Fig. A):

- Pressures in the **spinel field** i.e. at depths > 20 km (Fig. B)
- **High temperature** (800-1000°C, Fig. B) from orthopyroxene thermometry
- **High deviatoric stresses** (80-270 MPa, Fig. C) from olivine grain size

We propose that this deformation characterizes **the root of SWIR axial detachments**.

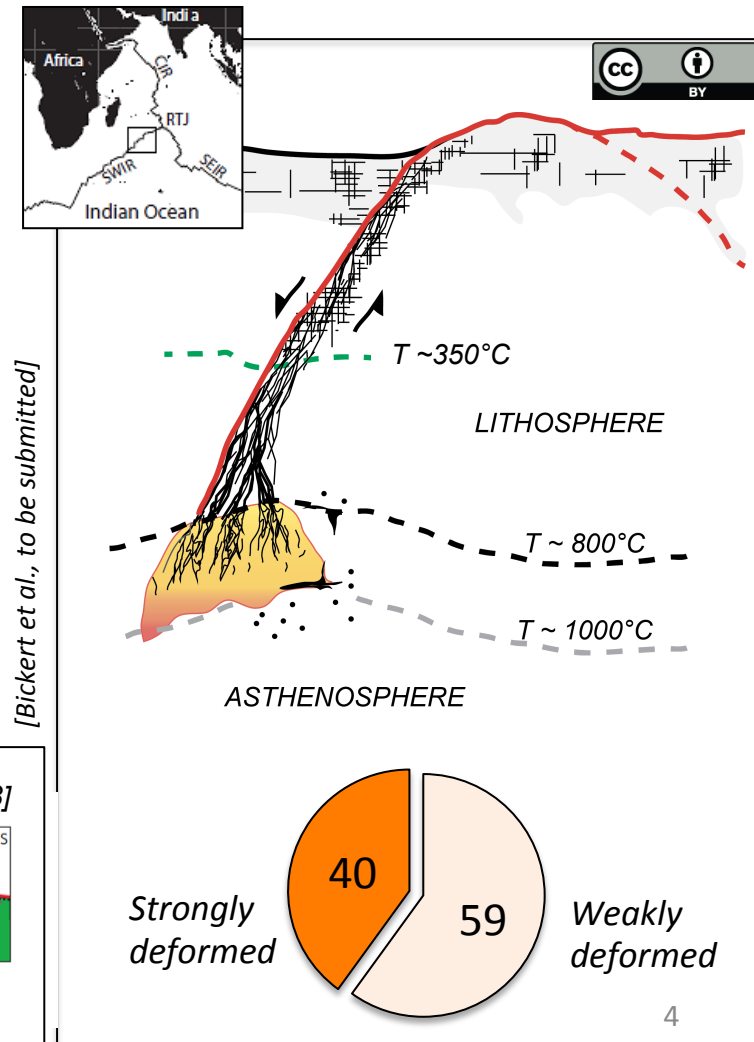
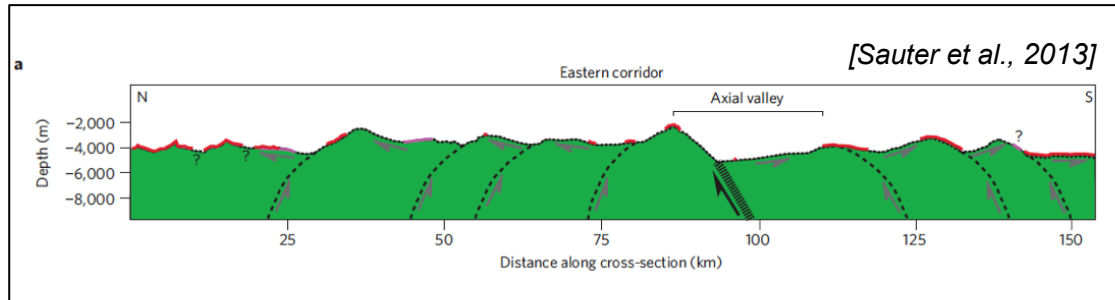


From the microscale to the macroscale:

Plate scale deformation models

Is it possible to obtain flipflop detachment faults with similar life-spans as those observed in nature by considering:

- A thick brittle lithosphere (~ 20 km)
- Two strain softening mechanisms observed in the samples:
 - ✓ Serpentinization ($T \leq 350^\circ\text{C}$)
 - ✓ Grain size reduction ($800\text{--}1000^\circ\text{C}$)



Serpentinization alone ... does not produce flipflop detachments.



1) When conditions are appropriate:

$$T \leq 350^{\circ}\text{C}$$

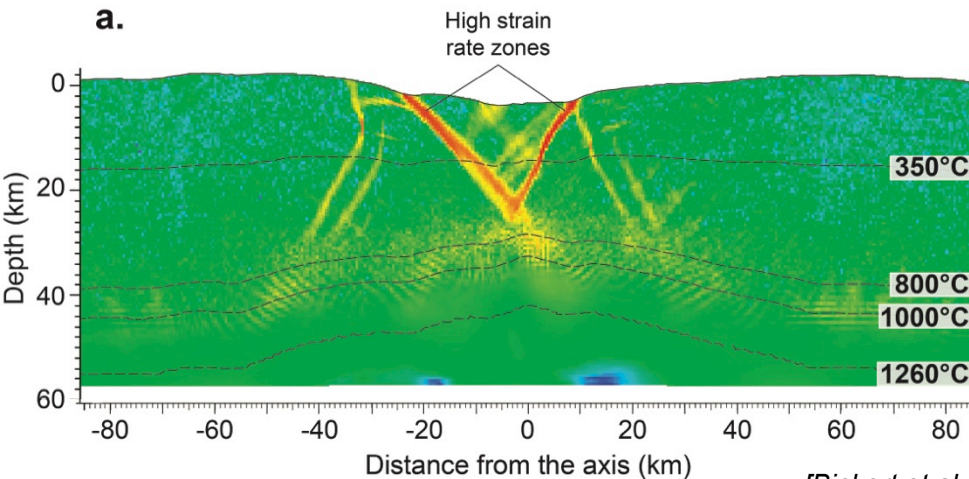
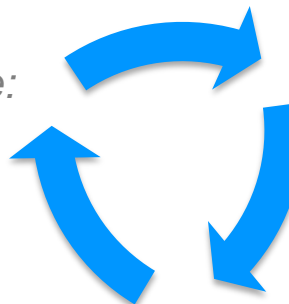
$$\dot{\epsilon}^{\text{II}} \cdot \sigma^{\text{II}} = 10^8 \text{ J}$$

2) Decrease of cohesion and friction:

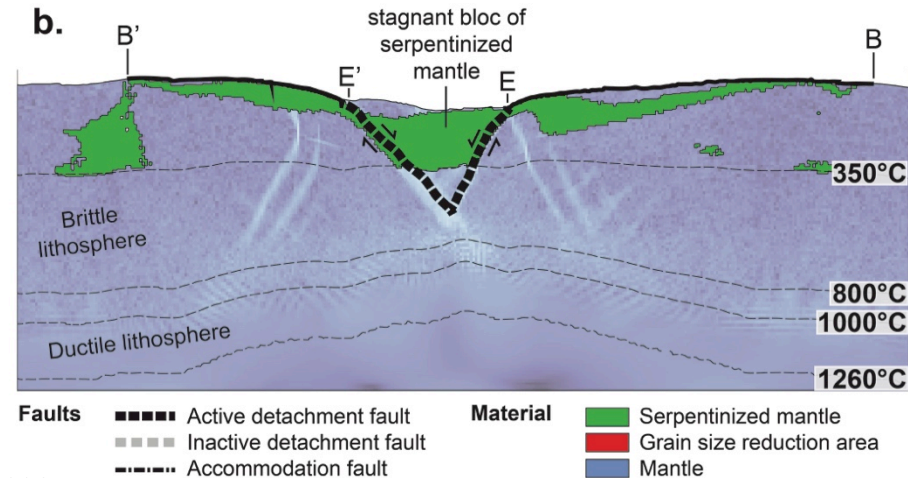
$$\tau = \mu \cdot \sigma_n + C$$

3) Decrease of activation energy:

$$\dot{\epsilon} = A \cdot \sigma^n \cdot \exp\left(\frac{-Q}{RT}\right)$$



[Bickert et al., 2020]



Grain size reduction alone ... does produce flipflop detachments.



2) Set up of new grain size:

$$d = 0.015 \cdot \sigma^{-1.33}$$

[Van der Wal et al., 1993]

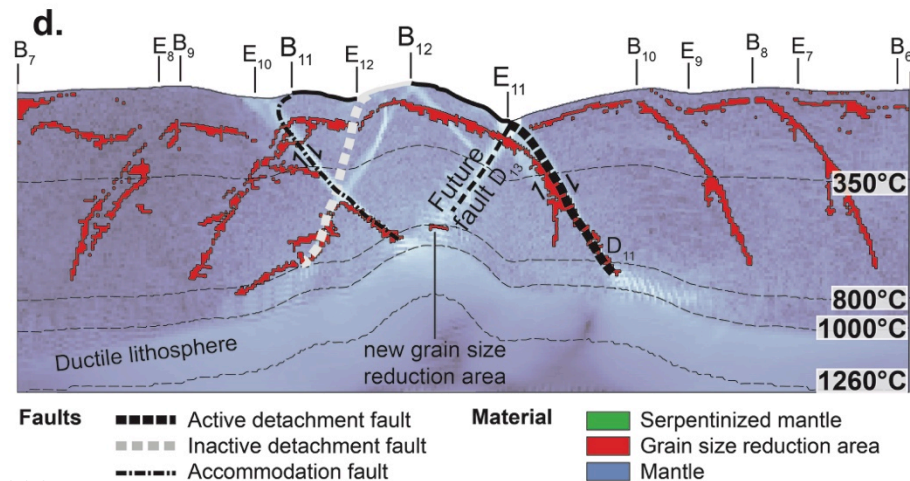
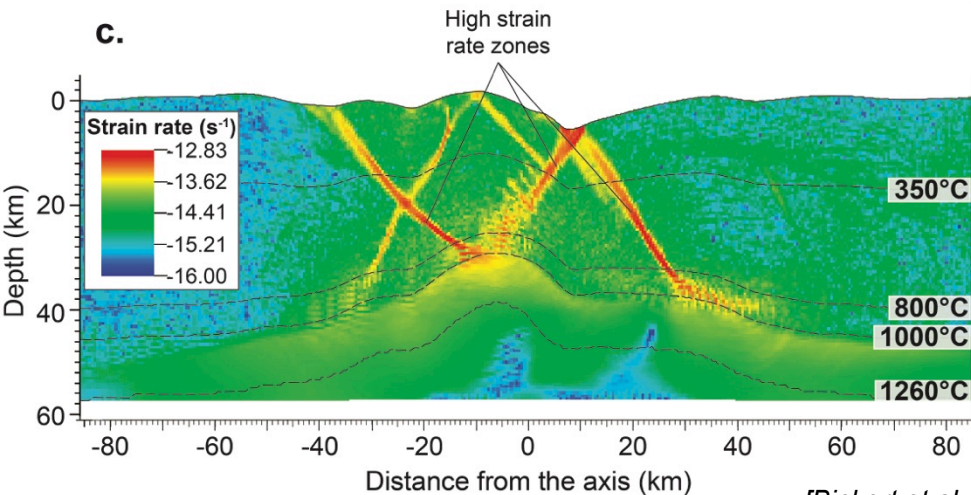
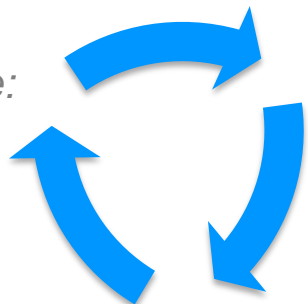
1) When conditions are appropriate:

$$T = 800 - 1000^{\circ}\text{C}$$

$$\dot{\epsilon} \cdot \sigma = 200 \cdot 10^{-8} \text{ Pa} \cdot \text{s}^{-1}$$

3) New viscosity using
diffusion creep law:

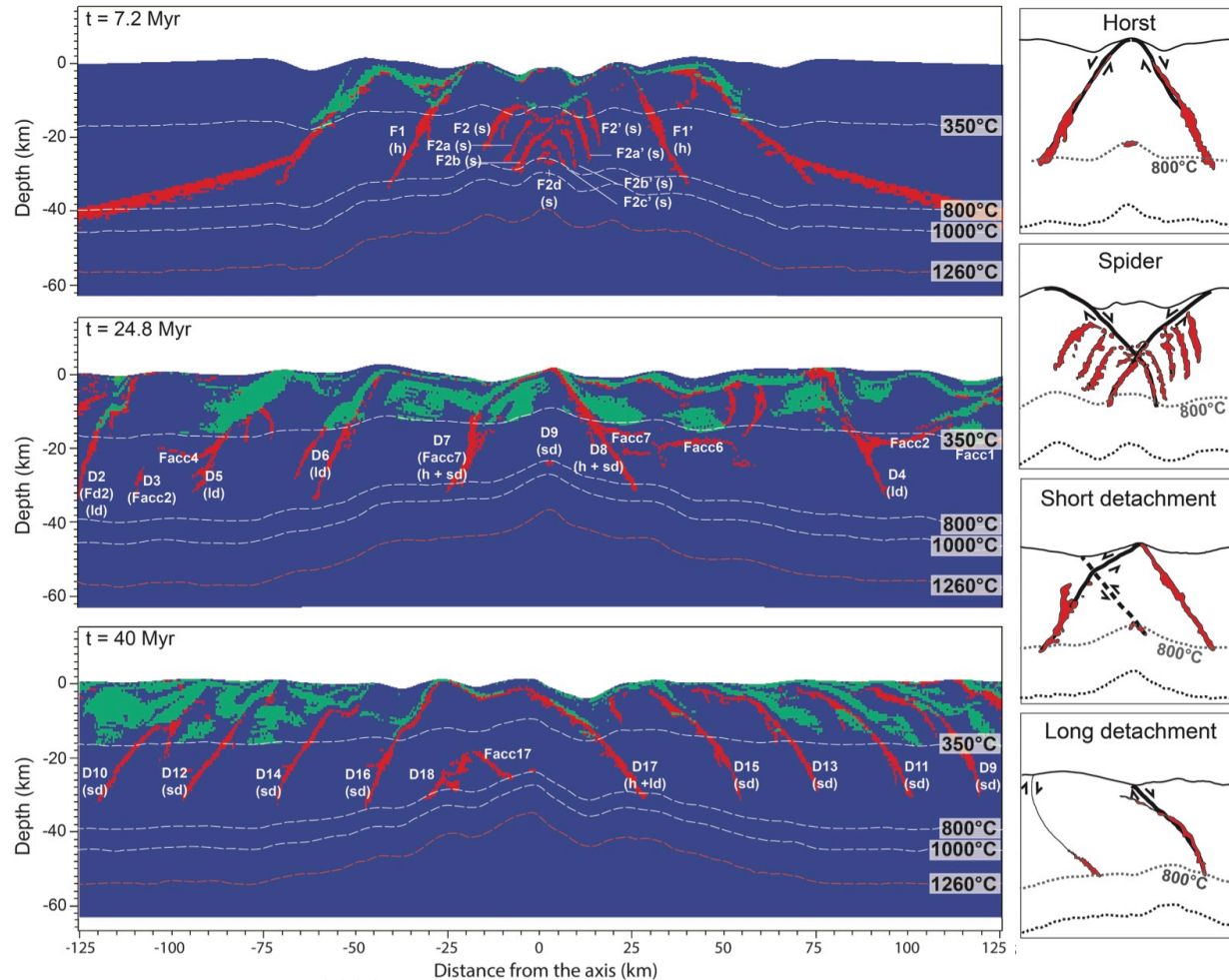
$$\dot{\epsilon} = A \cdot \sigma \cdot d^{-p} \cdot \exp\left(\frac{-Q}{RT}\right)$$



[Bickert et al., 2020]

Models with serpentinization + grain size reduction ...

- ... **develop four faulting modes**, including 2 detachment modes.
- Serpentinization alone does not produce flipflop detachments.
- Grain size reduction makes flipflop detachments possible in a thick lithosphere.
- **By combining the two**, the resulting fault patterns approach the flipflop detachment faulting pattern at the Eastern SWIR.



[Bickert et al., 2020]

Questions ? Comments ?
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