

Partitioning of halogens between hydrated silicates

Analysis and first principles modelling

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Part I: Main objectives

Studying halogen partitioning by analytics and modelling

Why studying the partitioning of halogens between minerals in metamorphic and mantle rocks?

Constraints on:

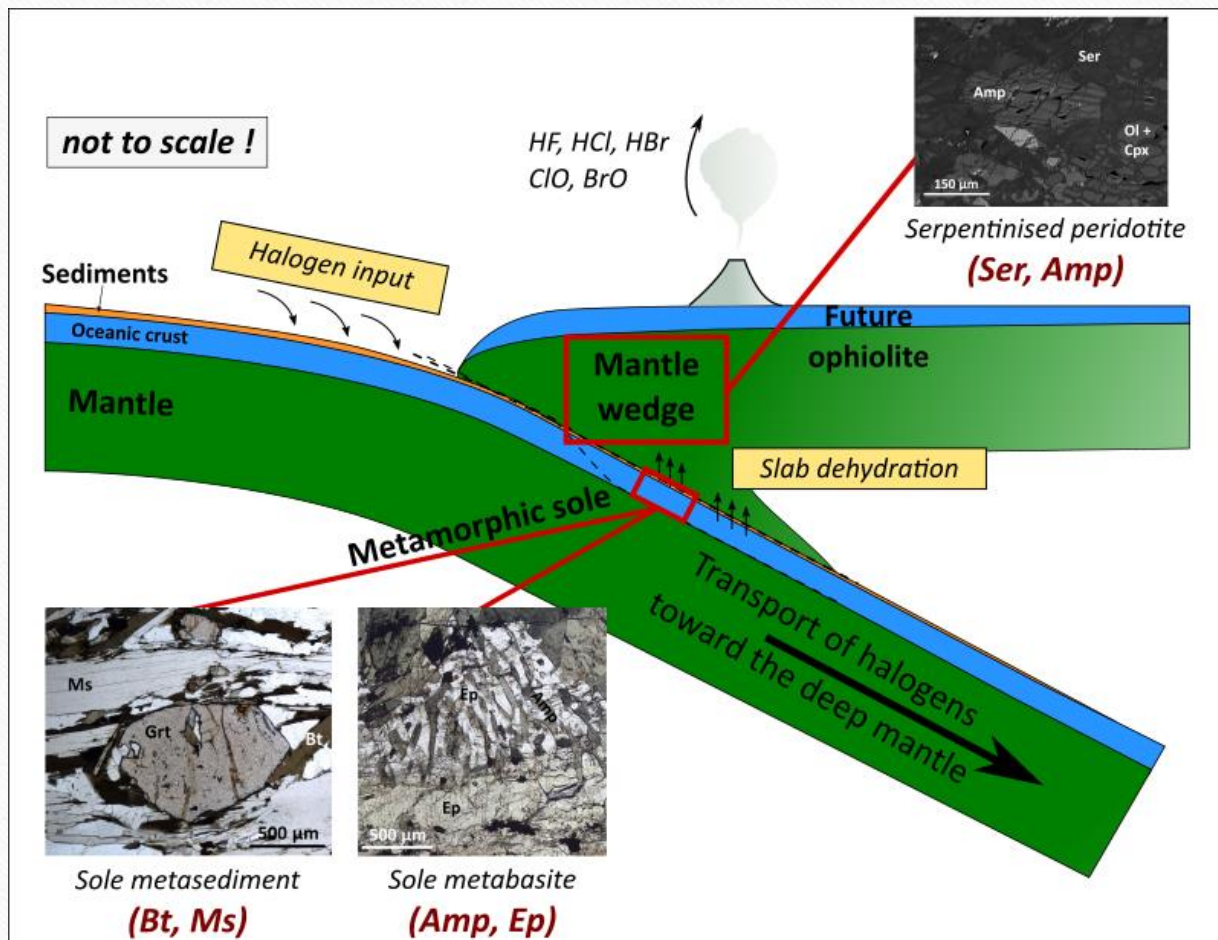
- ✓ Geochemical cycle
- ✓ Metasomatism & fluid/rock interactions

Minerals of interest: hydrated silicates

Biotite (Bt), Muscovite (Ms), Amphibole (Amp), Epidote (Ep), Serpentine (Ser)...

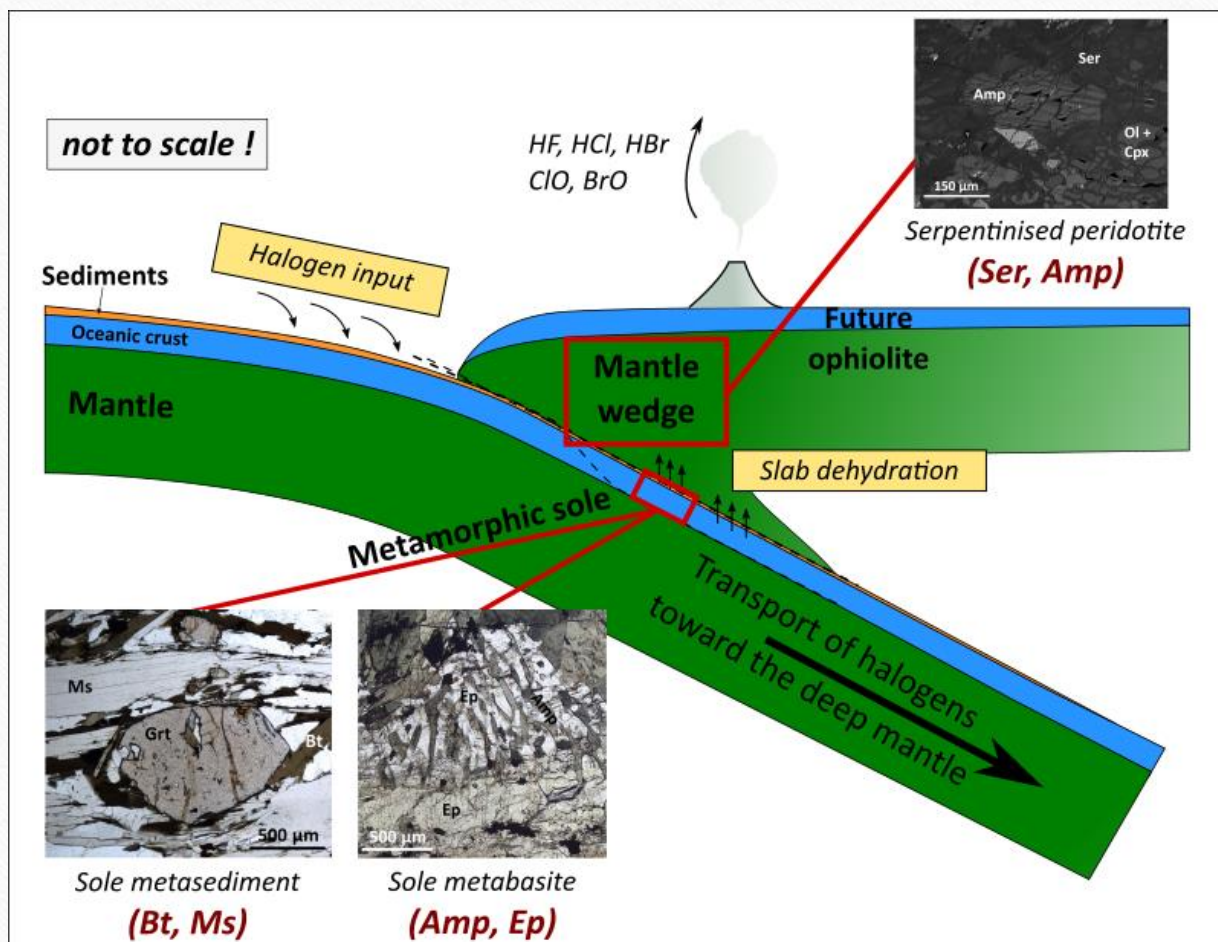
- ➔ Very abundant ! (main role in deep storage)
- ➔ But contain very low amount of halogens...

→ Huge analytical & computational challenge !



Main objectives:

- ✓ Demonstrating the main interest of **LA-ICP-MS/MS** combined to **EPMA** in halogen quantification
- ✓ Quantifying the role of **crystal-chemistry and geometry** in halogen uptake by *ab initio* modelling
- ✓ Quantifying and comparing **Gibbs free energies of halogen exchange** between modelled minerals



Part II: Analysis

Quantifying halogens in metamorphic and mantle rocks by EPMA & LA-ICP-MS/MS

State of the art: analytics

1) *Bulk* measurement

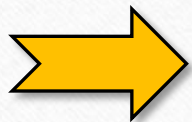
e.g.: ICP-MS solution

- ✓ Very low detection limit
- ✓ Pagé et al. (2018): 0.03 ppm Br and I in serpentines

Inadequate for mineral-mineral partitioning

Huge need to develop a method:

- ✓ for point analyses
- ✓ « en routine »
- ✓ *in situ*
- ✓ suited for trace to ultra-trace quantification



Candidate: ICP-MS/MS (laser ablation mode)

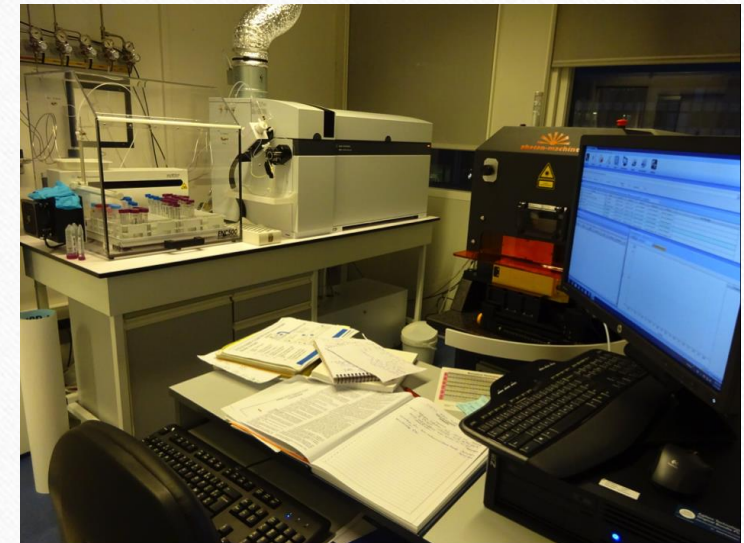
Photo: LA-ICP-MS/MS 8800 Agilent (ALIPP6)

2) Point measurement

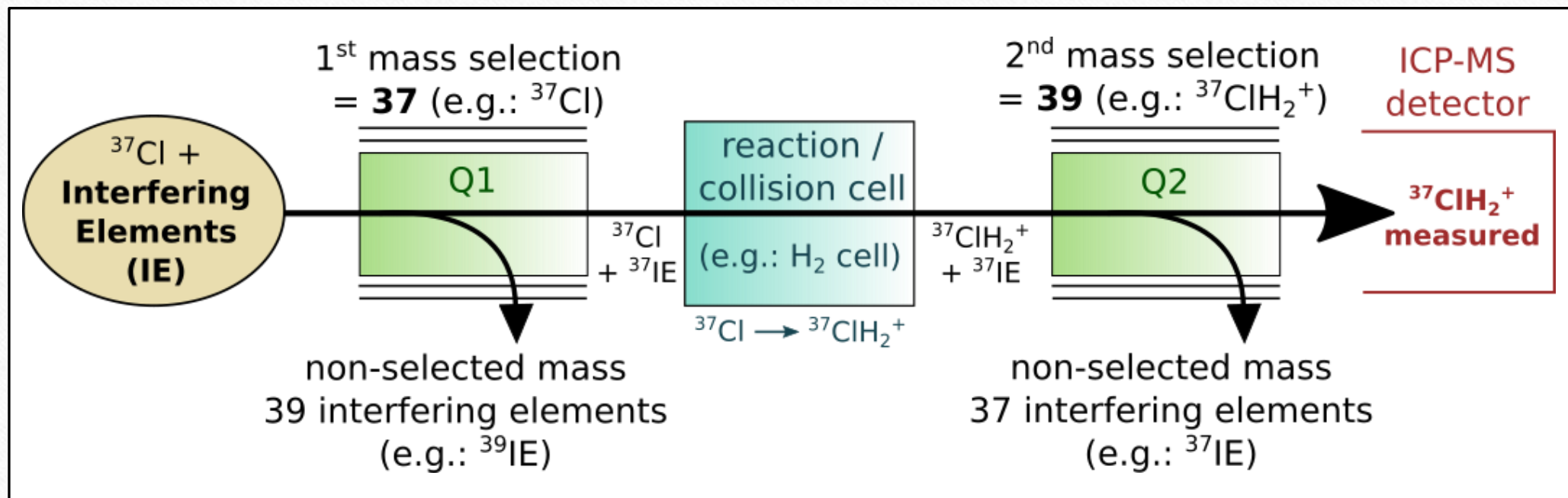
e.g.: SIMS

- ✓ In glasses, minerals, fluid inclusions
- ✓ Hughes et al. (2018): 350 ppb Br in olivine

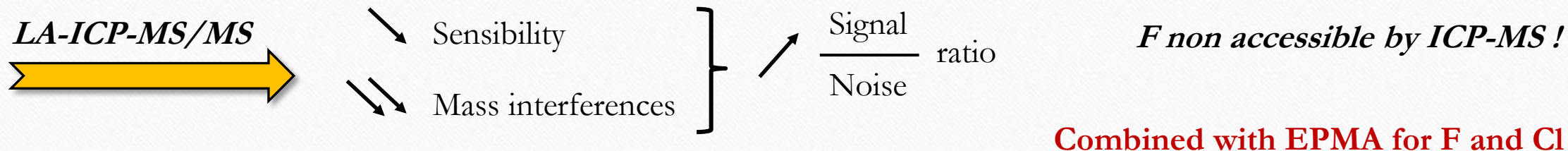
Complex sample preparation and signal processing + low availability



The LA-ICP-MS/MS (triple quad)



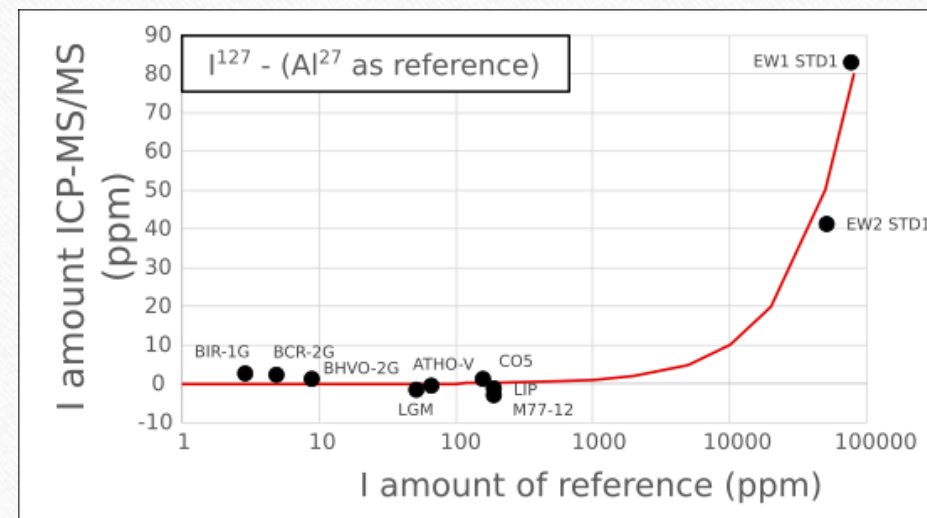
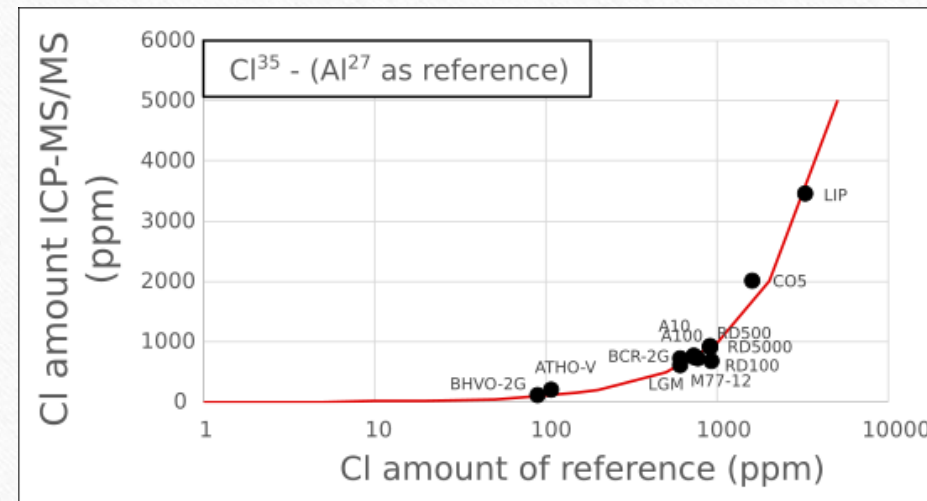
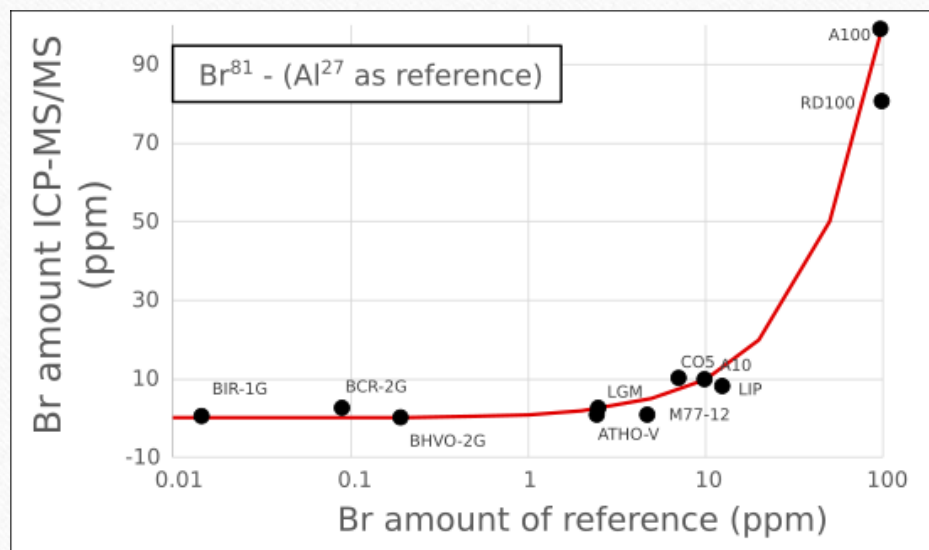
two quadrupoles + reaction/collision cell to optimise the mass/charge separation of isotopes

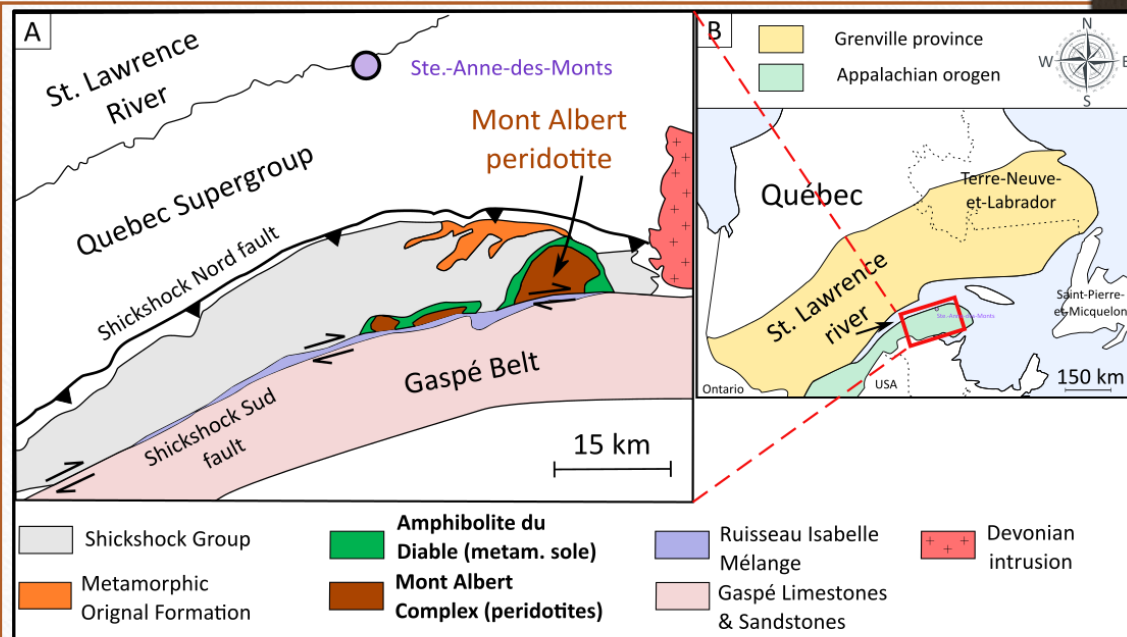


ICP-MS calibrations of Cl, Br and I

With international and in-house glass standards

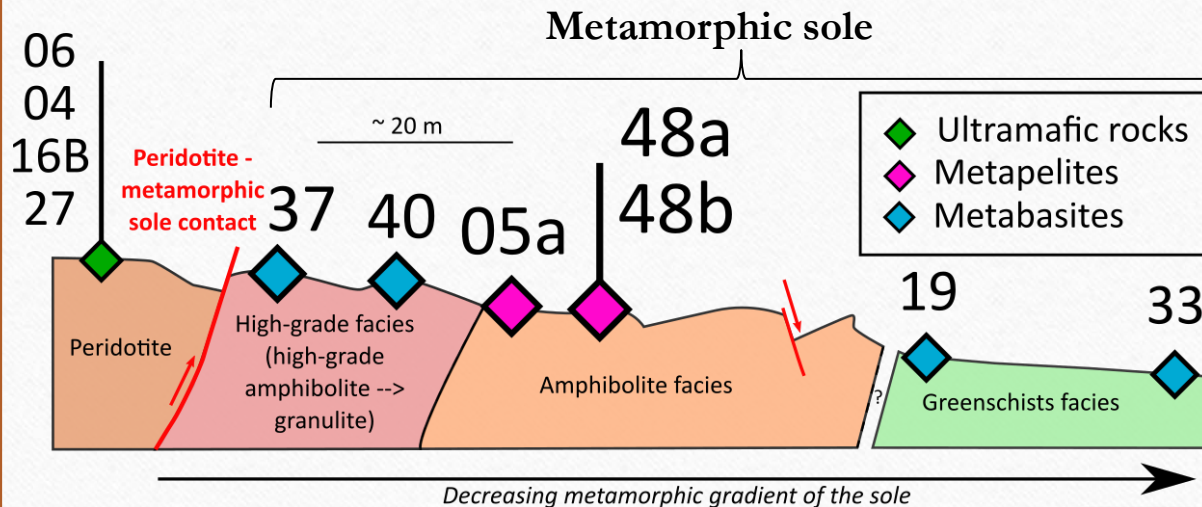
- ✓ Cl : about 80 ppm
- ✓ Br : about 0.1 ppm
- ✓ I : < 1 ppm (work in progress)





Case study: Mont Albert (Gaspésie, Québec, Canada)

Metamorphic sole (metasediments + metabasites) at the base of the peridotite body



After Dubacq et al., 2019

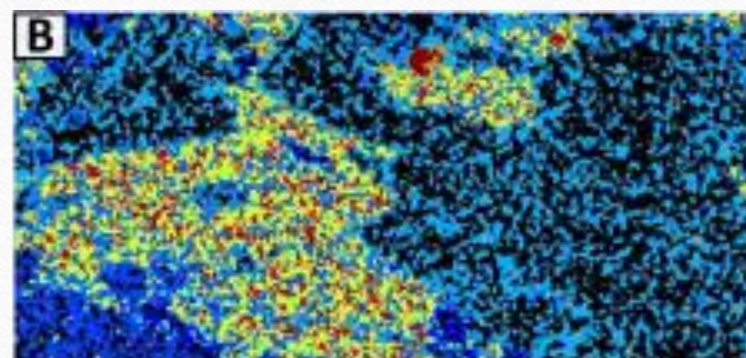
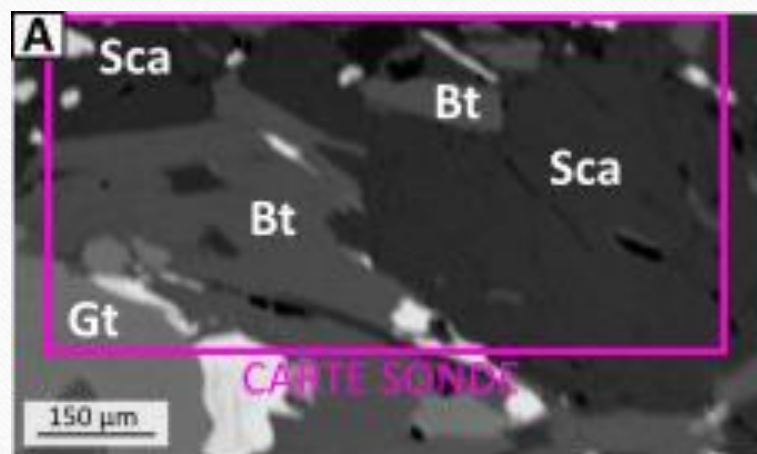
Metamorphic sole (hydrated)

Fluid exchanges ??

Mantle wedge (should be anhydrous!)

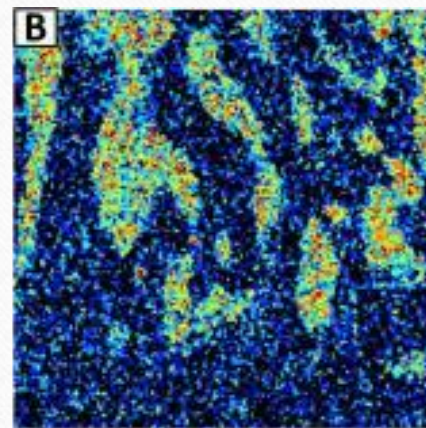
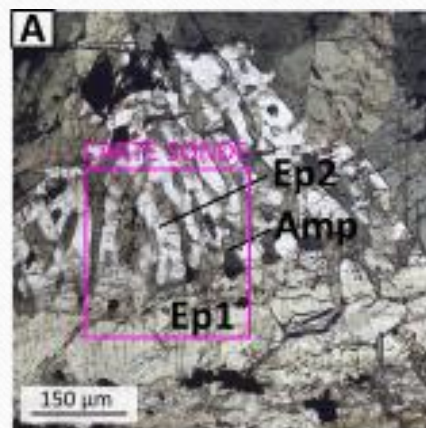
EPMA: F and Cl amount in the metamorphic sole

Metasediment (spl 05a)

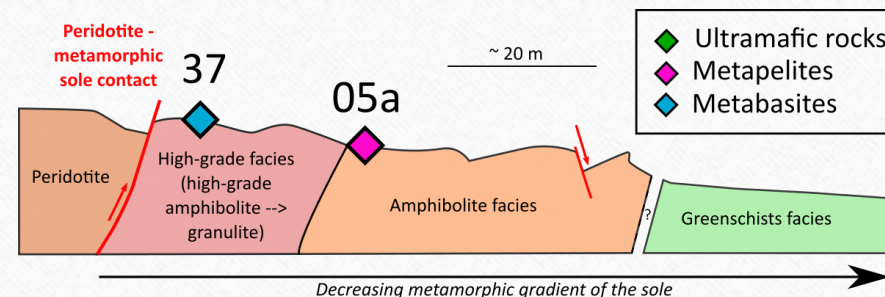


0.05 F (wt. %) 0.25

Metabasite (spl 37)



0 Cl (wt. %) 0.05



Metamorphic sole = enriched in F and Cl

EPMA detection limits :

✓ F : 200 ppm ✓ Cl : 25 ppm

Main hosts in metasediments

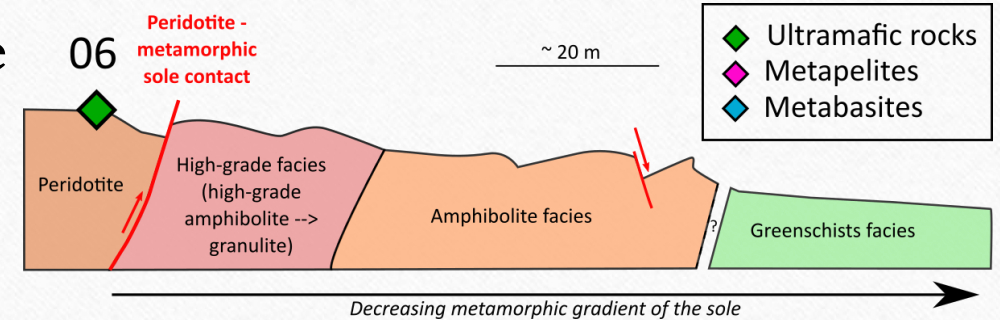
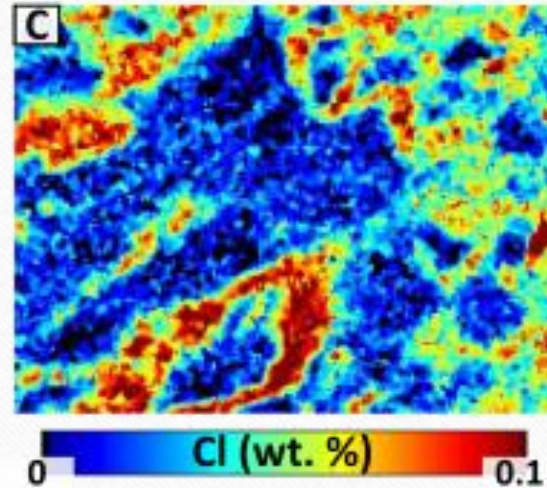
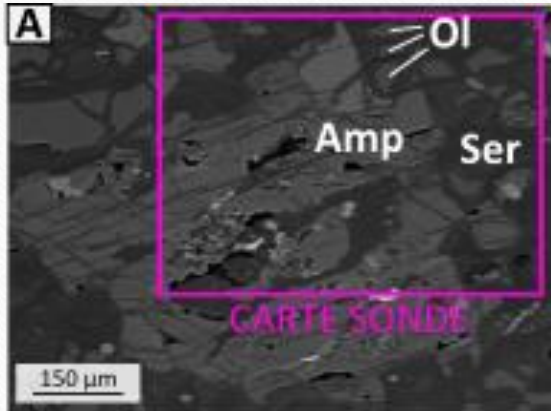
- ✓ Biotite : 1500 ppm F, 130 ppm Cl
- ✓ Muscovite : 330 ppm F, < 30 ppm Cl
- ✓ Scapolite : < 200 ppm F, 200 ppm Cl

Main host in metabasites

- ✓ Amphibole : 250 ppm F, 100 ppm Cl

EPMA: F and Cl amount in the mantle wedge

Harzburgite (spl 06)



Main hosts in mantle rocks

- ✓ Serpentine: 500 ppm F, 600-700 ppm Cl
- ✓ Amphibole: <200 ppm F, **40-100s ppm Cl**

- ✓ Serpentine are more enriched in Cl than F
- ✓ Amphiboles : enriched in Cl in serpentinised harzburgites.

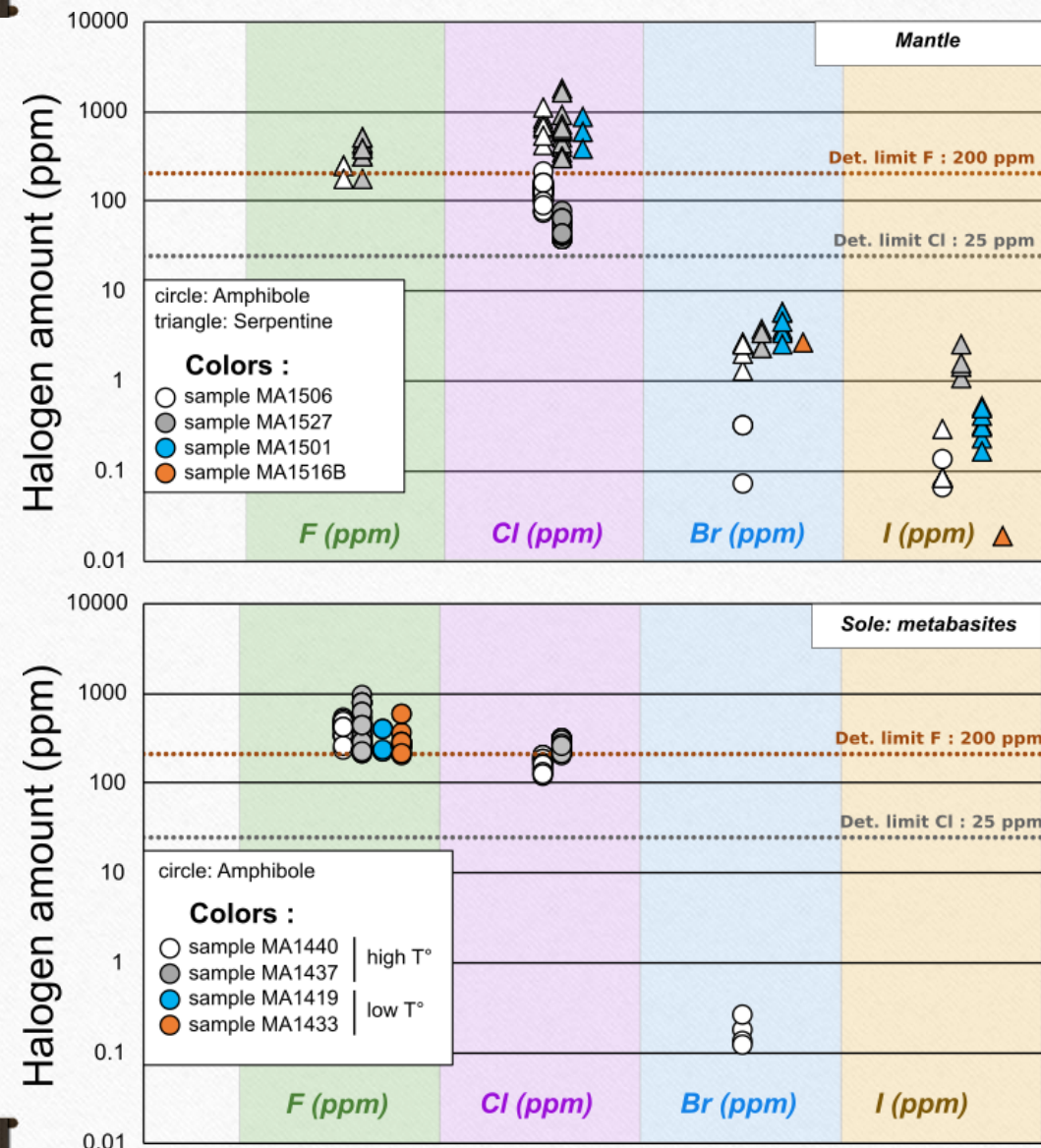
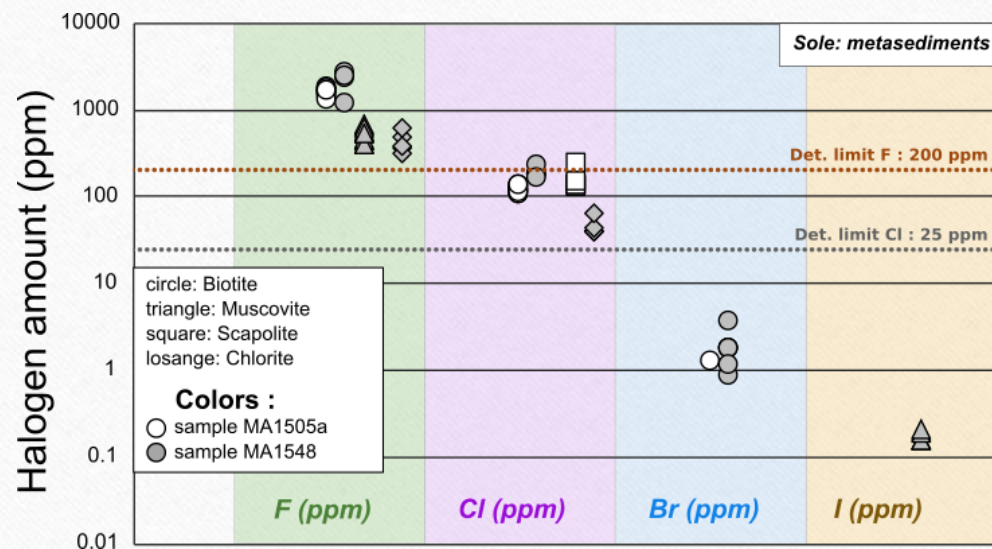
EPMA and ICP-MS/MS results

- ✓ **Metasediment:** biotite is the main host (about 10^3 ppm of F, 10^2 ppm of Cl, 1 ppm of Br)
- ✓ **Mantle:** amphibole are enriched in Cl, Br and I (up to 200 ppm of Cl)

➡ Dehydration of the metamorphic sole

➡ Fluid-rock interaction (from sole to mantle)

➡ Halogens incorporated in the mantle wedge



Detection limits indicated for EPMA (F and Cl)

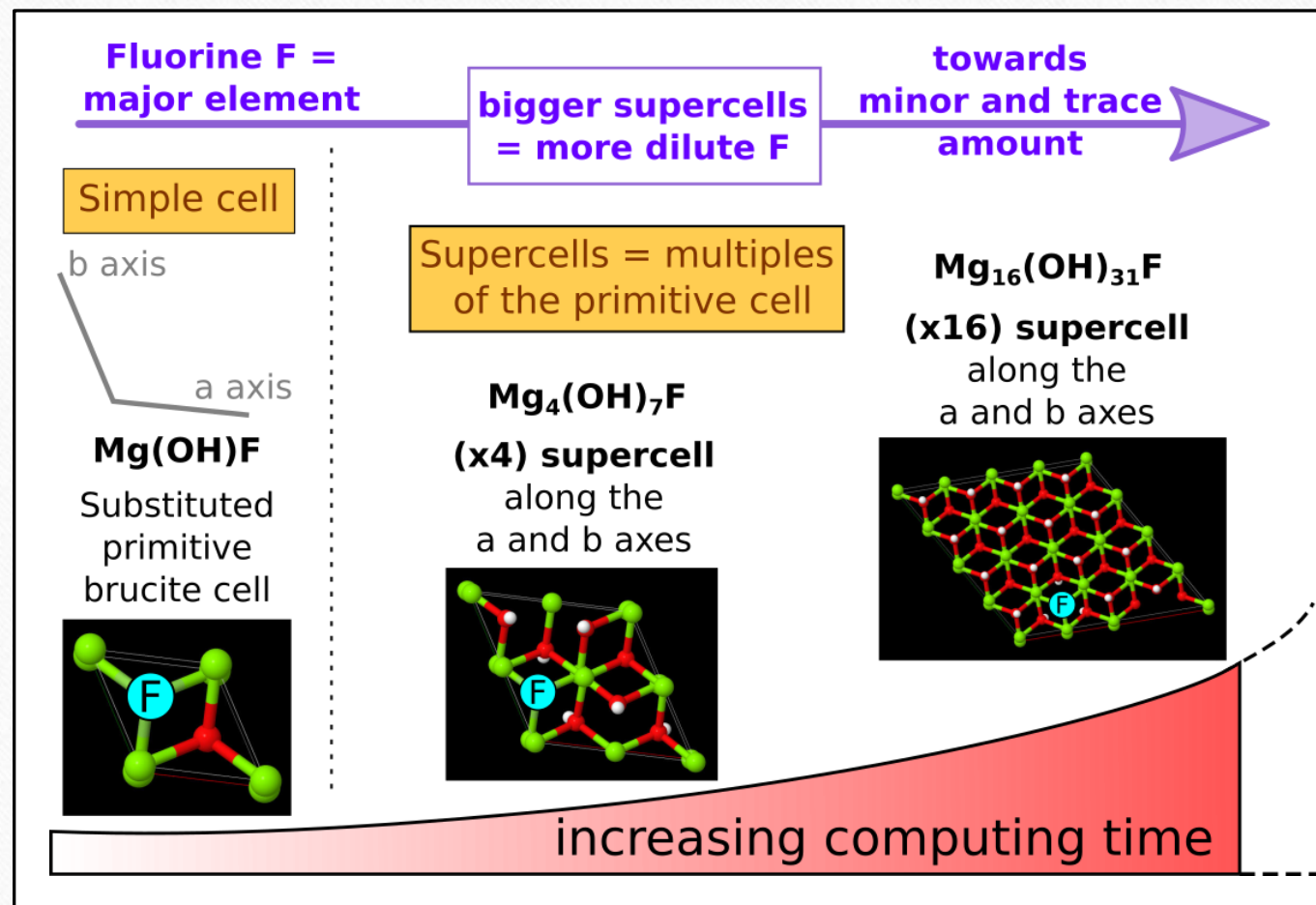
Part III: Modelling

Ab initio modelling of F = OH and Cl = OH
exchanges in phyllosilicates

See also the poster presented by Dubacq et al. :

The *ab initio* modelling of the $\text{OH}^- = \text{X}^-$ exchange

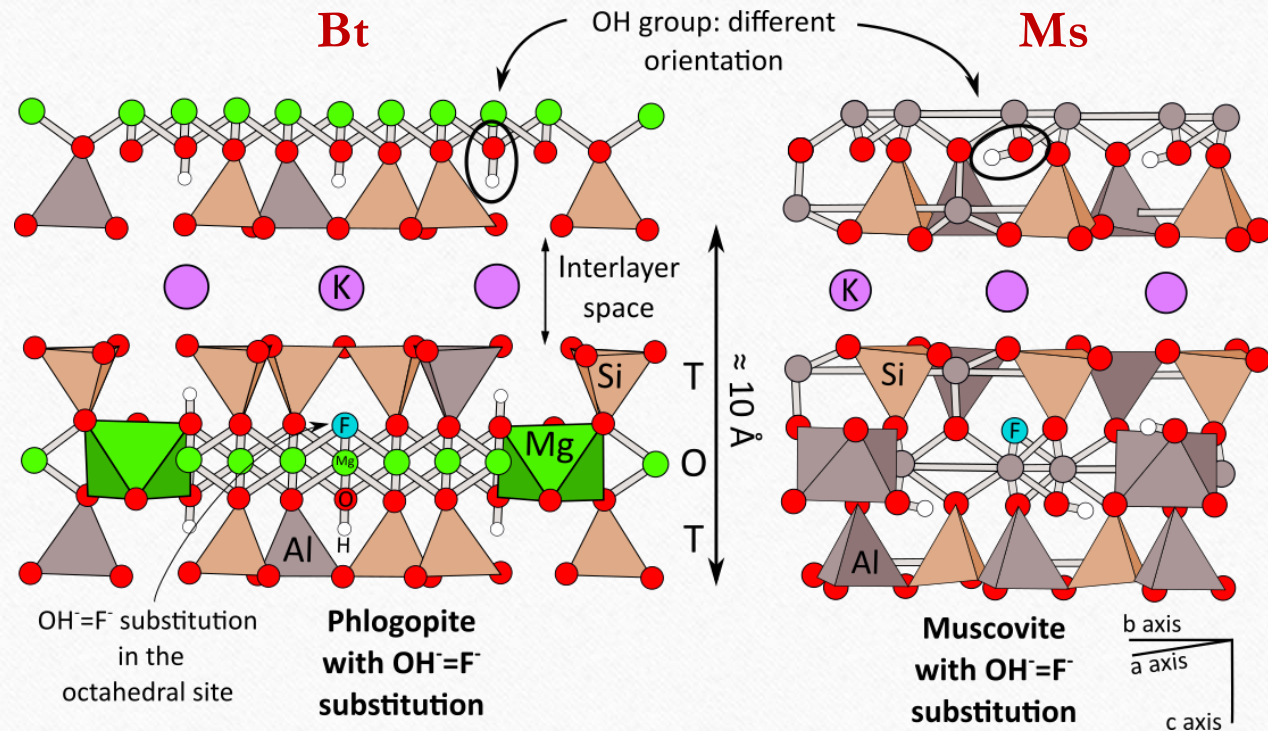
($\text{X}^- = \text{F}^-, \text{Cl}^-, \text{Br}^-, \text{I}^-$)



CRYSTAL17 *ab initio* program
(see Dovesi et al., 2004)

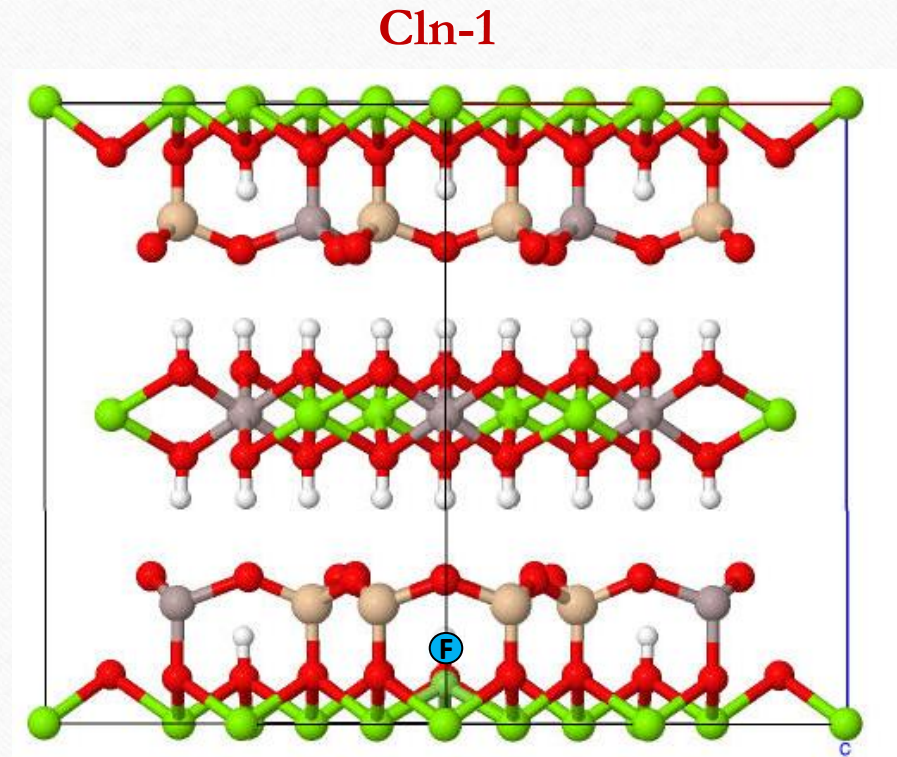
- ✓ Reaching lower concentrations: supercells
- ✓ Need to work on simplified systems: **end-members**

Simulation of F = OH and Cl = OH substitution in three magnesian end-members: phlogopite, muscovite and clinochlore (space group : C2/m)



F-Phlogopite
 $\text{KMg}_3(\text{AlSi}_3)\text{O}_{10}(\text{OH},\text{F})$
 O-H follows the c-axis

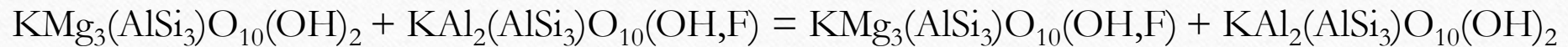
F-Muscovite
 $\text{KA l}_2(\text{AlSi}_3)\text{O}_{10}(\text{OH},\text{F})$
 O-H: almost perpendicular to c



- ✓ **Cln-1**: F in the TOT sheet
- ✓ **Cln-2**: F in the brucite sheet

Reaction for this example: *

Exchange from Muscovite to F-Biotite:



Biotite

F-Muscovite

F-Biotite

Muscovite

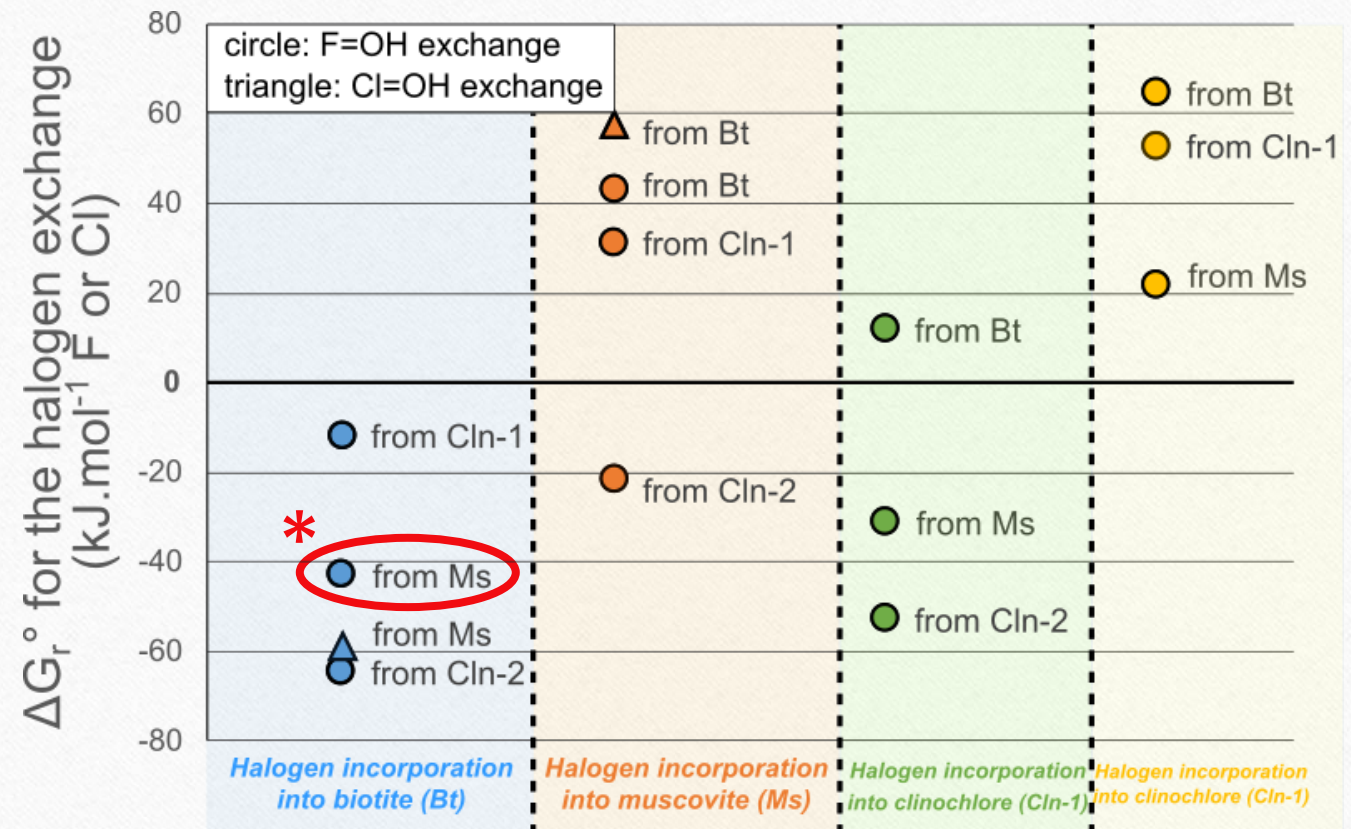
$$\Delta G_r^\circ = -11 \text{ kJ.mol}^{-1} \text{ p.s.f.}$$

Energy stabilization: phlogopite >
clinochlore 1 > muscovite > clinochlore 2

Mg-biotite is systematically favoured

Hypothesis (see also Munoz, 1984):

- ✓ O-H follows c-axis in biotite
- ✓ Repulsive effect between H^+ and K^+ ...
- ✓ ...substituted by an attraction between F^- and K^+ !



To summarize...

- ✓ **EPMA and LA-ICP-MS/MS** allow easy *in situ* measurement of **F, Cl and Br** in most hydrated silicate phases (Villemant et al., submitted to American Mineralogist). Detection limits are **very satisfying**. Calibration of I may be improved with more reference standards.
- ✓ ***Ab initio* modelling** clearly shows that **Mg-biotite** is a favoured **F- and Cl-host** compared with muscovite and Mg-chlorite, in agreement with natural analyses.
- ✓ ***In progress***: halogen incorporation differences between dioctahedral and trioctahedral micas may be linked to electrostatic effects between interfoliar cation, hydrogen and halogen atoms: **but what about vacancy-rich chlorite...? and amphibole...?**

Many thanks for your attention !