### Using P- to S- wave conversions from controlled sources to determine the shear-wave velocity structure along Hikurangi Margin Forearc, New Zealand

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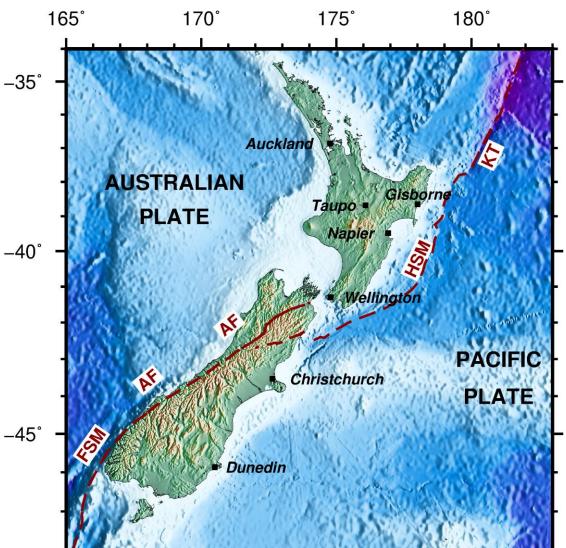
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### Introduction

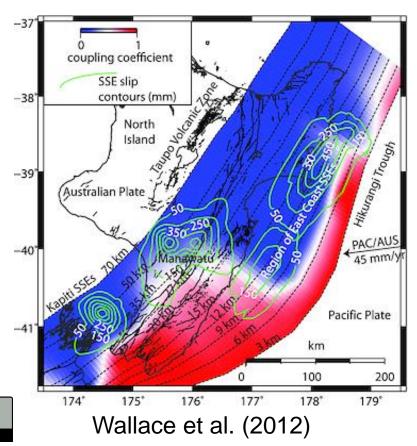
- Hikurangi margin marks the subduction of the Pacific Plate under the Australian Plate off the east coast of the North Island of New Zealand.
- Geodetic observations indicate along-strike variations in subductionthrust slip behavior along the Hikurangi margin.
  - Subduction-thrust of the southern segment of the margin is locked on the 30-100-year scale
  - In the northern segment it displays periodic slow-slip on the 1-2-year scale.



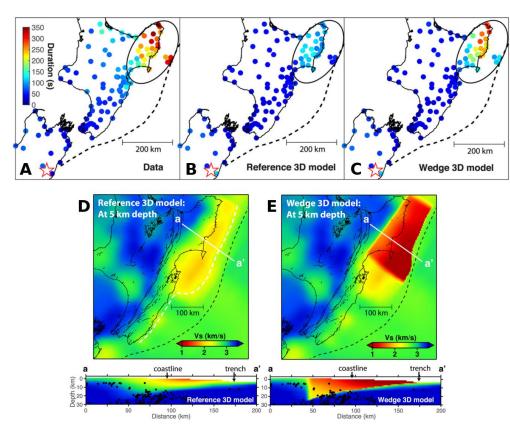
Tectonic setting around New Zealand. AF-Alpine Fault FSM-Fiordland subduction margin HSM-Hikurangi subduction margin KT-Kermadec Trench

### **Context and Rationale**

- Along-strike variation in subduction thrust-slip behavior
  - Hypothesised to be due to spatial variations in porosity, potentially linked with elevated pore-pressure



- Ultra-long duration seismic ground motion in the northern Hikurangi margin
  - Attributed to be due to a sediment wedge with low shear-wave speeds



Kaneko et al. (2019)

### **Seismic Wave Velocity**

• V<sub>P</sub>

- Compressional (P) wave velocity
- Function of bulk modulus, shear modulus
  Function of shear modulus and density
- Ambiguous indicator of a rock's lithology

### • V<sub>P</sub>/V<sub>S</sub>

- Directly related to the Poisson's ratio
- Diagnostic property of a rock's degree of consolidation and porosity
  - Consolidated sediments and crystalline rocks = 1.6 1.9
    - Unconsolidated sediments = 2.0 4.0



• V<sub>S</sub>

• Shear (S) - wave velocity

# Mode-converted waves from controlled sources in OBS data

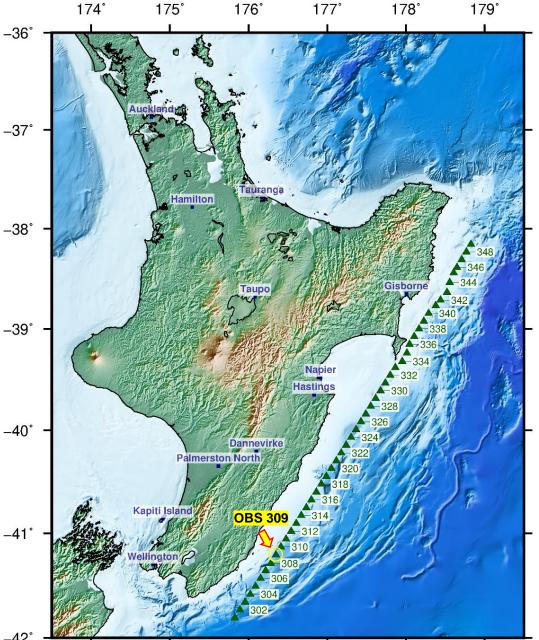
- Most controlled-source ocean bottom seismic studies focus on determining the compressional (P-) wave velocity structure.
- By identifying mode-converted waves in the ocean bottom seismometers, the shear (S-) wave velocity structure can be estimated.
- V<sub>P</sub>/V<sub>S</sub> ratio can be determined more accurately than passive-source seismic tomographic methods.
  - Time and location of controlled-sources known accurately from GPS clocks.
- Uncertainties can be quantified.
- Closely spaced controlled-sources provide higher resolution and better ray coverage in offshore regions.



### Controlled-source Seismic Data

### Seismogenesis at Hikurangi Integrated Research Experiment (SHIRE)

- Controlled-source seismic data acquired in 2017 by *R/V Marcus G Langseth* and *R/V Tangaroa*
- 49 ocean bottom seismometers (OBS) of SHIRE03 transect along Hikurangi forearc
  - Multicomponent
    - Triaxial seismometer
    - Hydrophone
  - ~10 km spacing along Hikurangi forearc
- Multichannel seismic (MCS) acquisition
  - 12.7 km long streamer
  - Airgun source spaced 50 100 m



Ocean bottom seismometer stations along SHIRE03 transect

### **Methodology**

#### Processing OBS gathers

- Determining orientation of horizontal components of each OBS
- Rotation of the horizontal components into radial and transverse components

### Improving signal to noise ratio

- Bandpass filtering (1-20 Hz)
- Automatic gain control
- Predictive deconvolution
- Identification of mode-converted waves
- Estimation of  $V_P/V_S$



# **Types of mode-converted waves**

#### • PSS

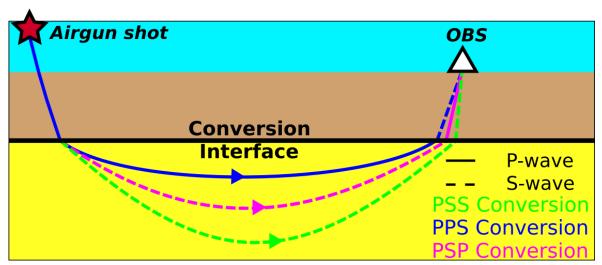
- Down-going P-wave converts to S-wave at an interface
- Slow apparent velocity
- Not recorded by hydrophone

#### • PPS

- Up-going P-wave converts to S-wave at an interface
- Lags behind the P-phase
- Same apparent velocity as P-phase
- Not recorded by hydrophone

#### • PSP

- Down-going P-wave converts to an S-wave at an interface and again to P-wave on its way up to the OBS
- Slower apparent velocity
- Recorded by hydrophone



Types of mode-converted waves observable in an ocean bottom seismometer from airgun sources



### **Identification of mode-converted waves**

- Mode converted waves are identified in the radial and transverse components
- Following approaches are used to identify different mode-converted waves
  - Polarisation angle of arrivals
    - A measure of the polarisation angle of the particle motion from the three seismograph components (Flinn, 1965)
    - Distinguish incoming S-waves to OBS
      - Incoming S-waves have higher polarization angles (e.g.  $\ensuremath{\mathsf{PP}_nS_b}\xspace)$

#### Rectilinearity of arrivals

- A measure of the the linearity of the particle motion from the three seismograph components.
- Can be define as 1 ellipticity (Flinn, 1965)
- Distinguish water column multiples
  - Water column multiples are linear (e.g. P<sub>o</sub> mul)

#### Hydrophone component

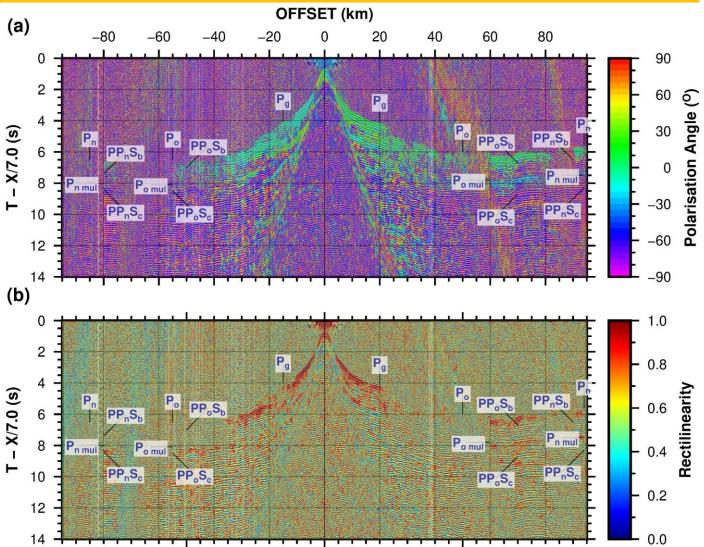
Does not record incoming S-waves

#### Apparent velocity

(†)

(cc

- Horizontal slowness
- Identify S-wave refractions



Polarisation angles (top) and rectilinearity (bottom) of the arrivals windowed every 50 msec, overlain on top of the radial component at OBS309

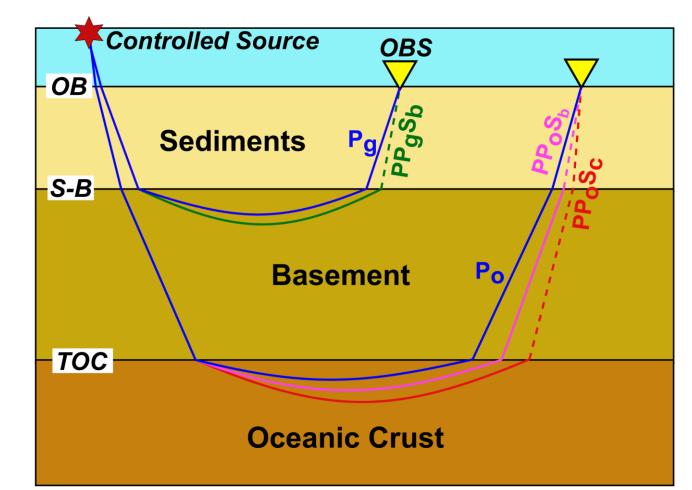
### **Identified mode-converted phases**

### PPS Converted Waves

- Can be used to estimate  $V_{\rm P}/V_{\rm S}$  above the converting interface
  - PP<sub>g</sub>S<sub>b</sub>
  - PP<sub>o</sub>S<sub>b</sub>
  - PP<sub>o</sub>S<sub>c</sub>

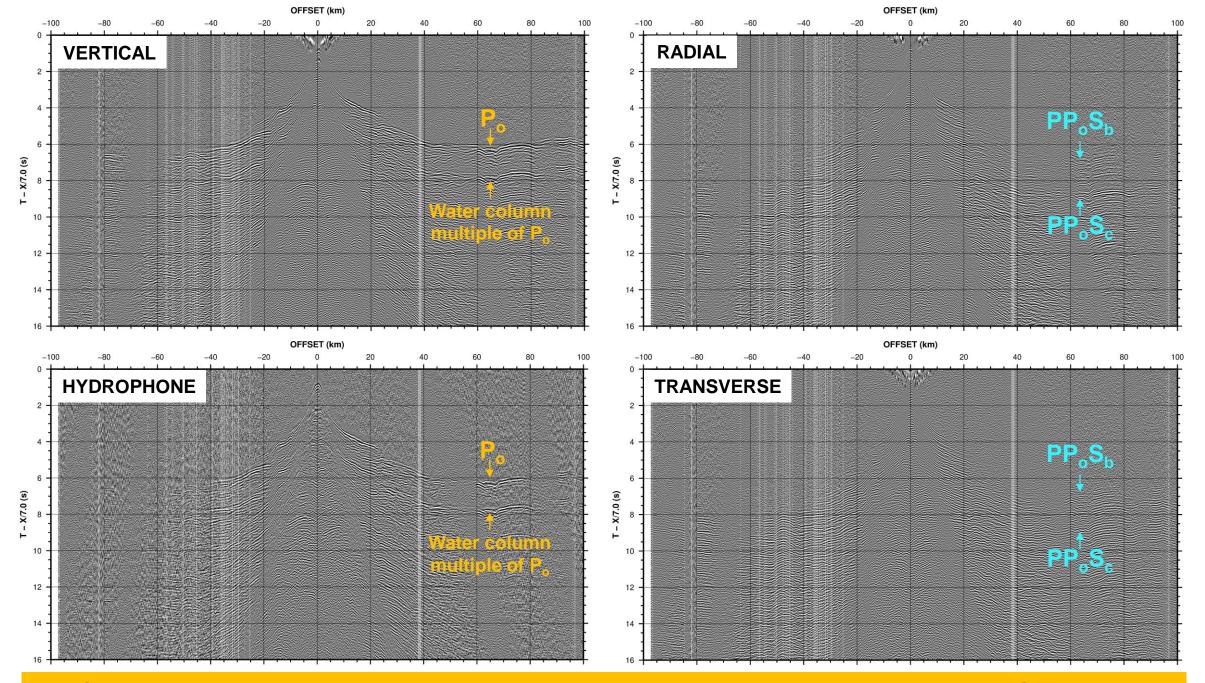
### PSS Converted Waves

• Some hardly distinguishable candidates



**OB** = Ocean Bottom, **S-B** = Sediment-Basement, **TOC** = Top of Oceanic Crust





OBS gather records of vertical, hydrophone, radial and transverse components at OBS station 309

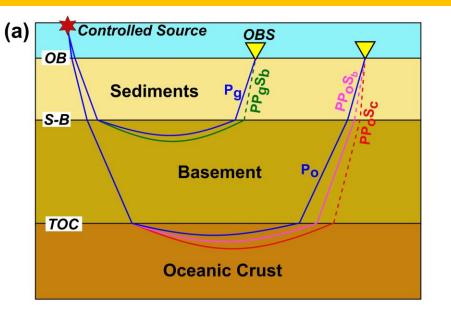
CCC D

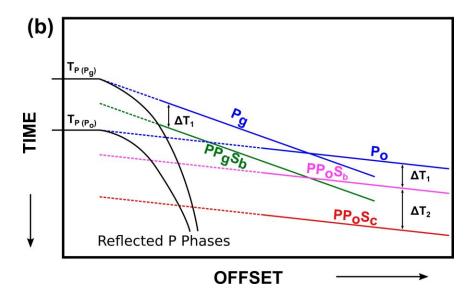
# Average V<sub>P</sub>/V<sub>S</sub> from PPS phases

$$\bullet \frac{V_P}{V_S} = \frac{2 \Delta T + (t_p - t_{psf})}{(t_p - t_{psf})} \quad \text{(Tsuji et al., 2011)}$$

• 
$$\Delta T$$
 = Time lag between P phase and PPS phase

- $t_p$  = Zero offset travel time of P phase
- $t_{psf}$  = Zero offset sea-floor reflection time







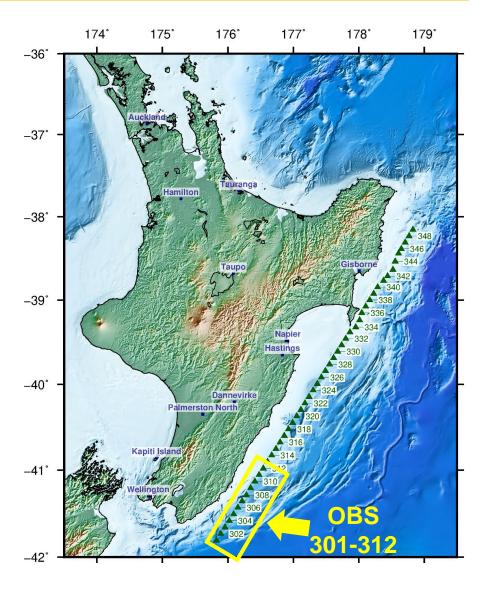
## **Results and Discussion**

- Average V<sub>P</sub>/V<sub>S</sub> of the forearc in the southern Hikurangi margin
  - ≈1.70 (between OBS301-312)
  - Indicates the presence of consolidated sediments with low pore-pressure
  - Determined from the time lag of the observed PPS converted waves
  - Impedance contrasts at the top of oceanic crust and the sediment-basement interface for an up-going wave are sufficient to generate S-waves

#### PSS converted waves

- Were not observed
- Impedance contrasts at the interfaces for a down-going wave are not sufficient to generate S-waves
- Next steps ...
  - Extending the study to the north to estimate of V<sub>P</sub>/V<sub>S</sub> in the northern Hikurangi forearc





### References

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