

Model bias in climate sensitivity tied to biases in surface temperature patterns over the tropical Indo-Pacific.

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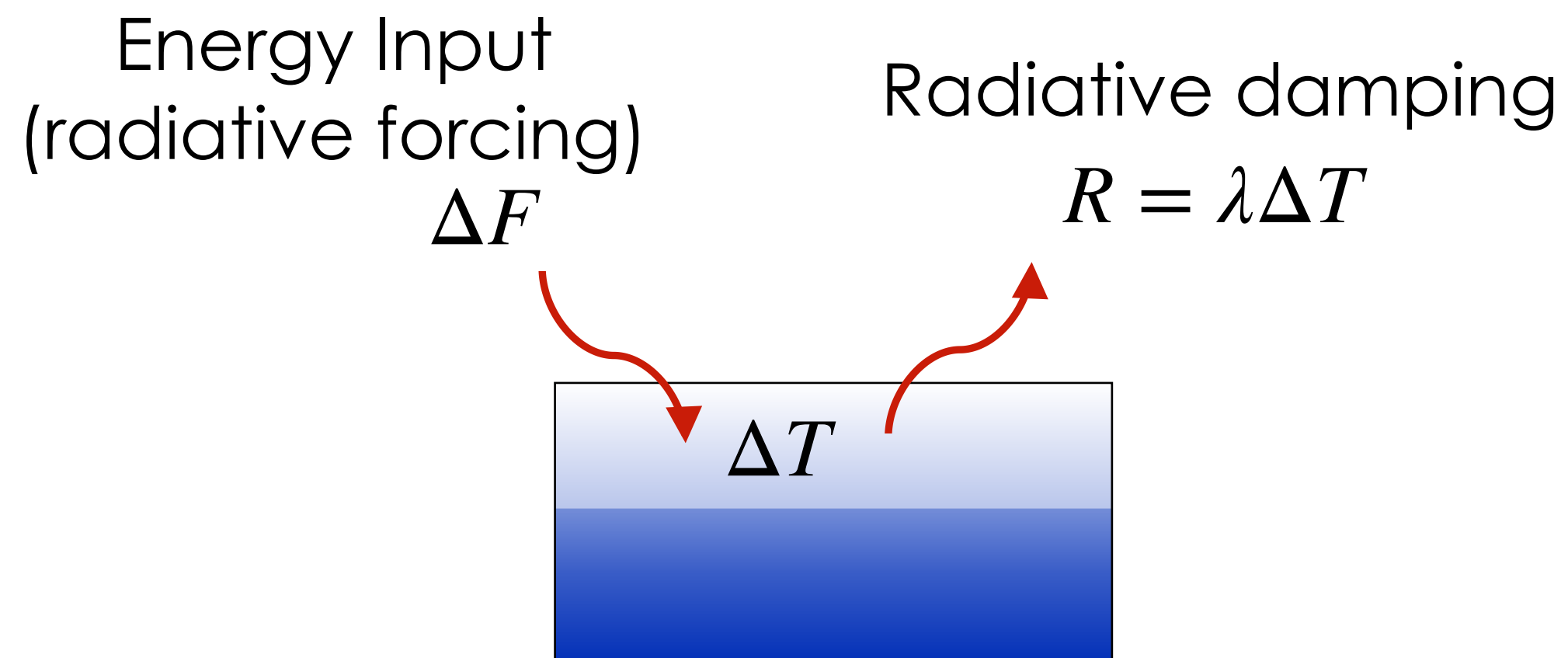


Earth's **Equilibrium** Energy Balance

$$\Delta F = \Delta R$$

Radiative feedback
(Efficiency of radiative damping)

$$\lambda = \frac{\Delta R}{\Delta T}$$



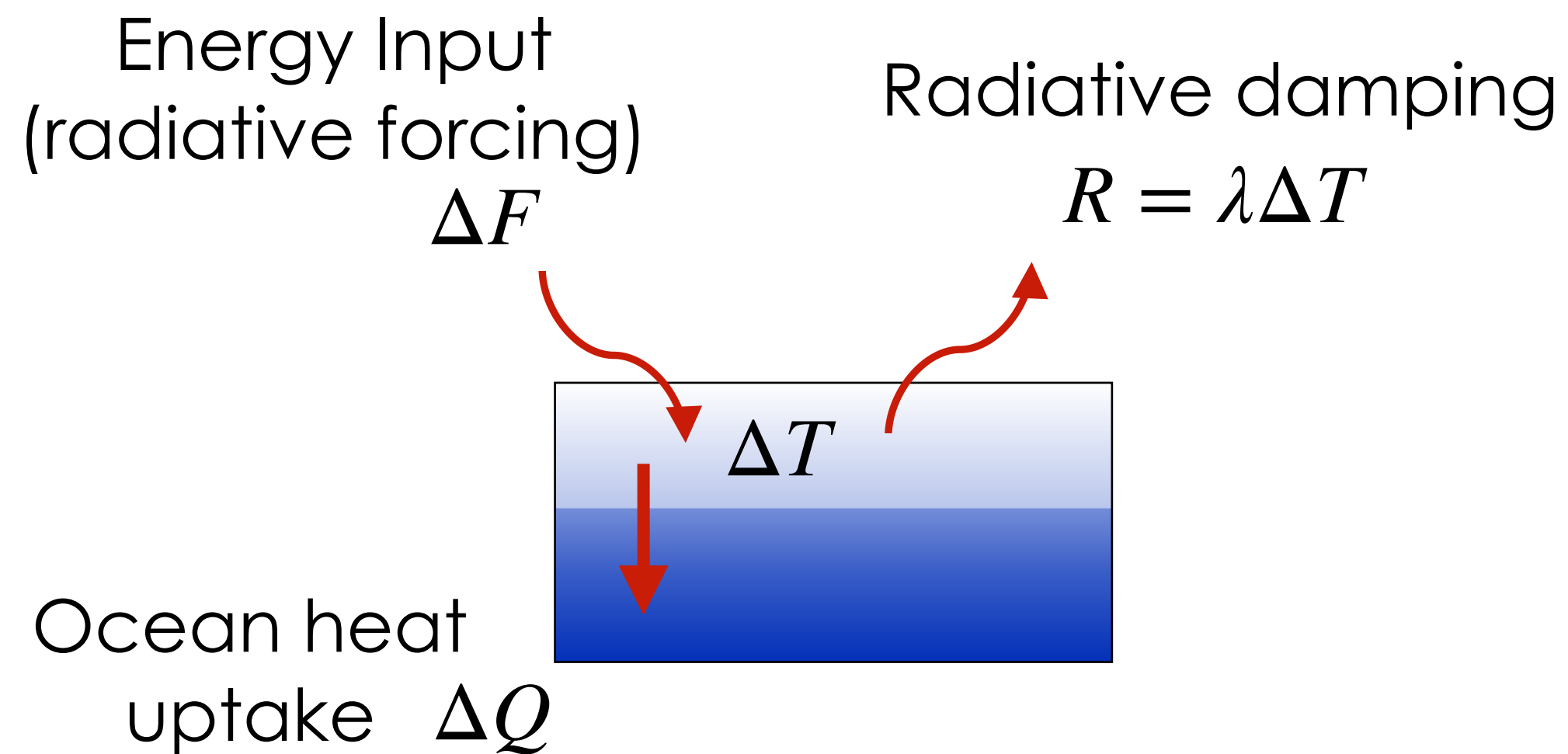
Equilibrium Climate Sensitivity:

$$\Delta T_{2\times} = \frac{\Delta F_{2\times}}{\lambda}$$

Determined by “How efficient is the Earth at restoring energy balance through warming?”

Earth's **Transient** Energy Balance

$$\Delta F(t) + \Delta Q(t) = R(t)$$



Instantaneous Radiative feedback
(Efficiency of radiative damping)

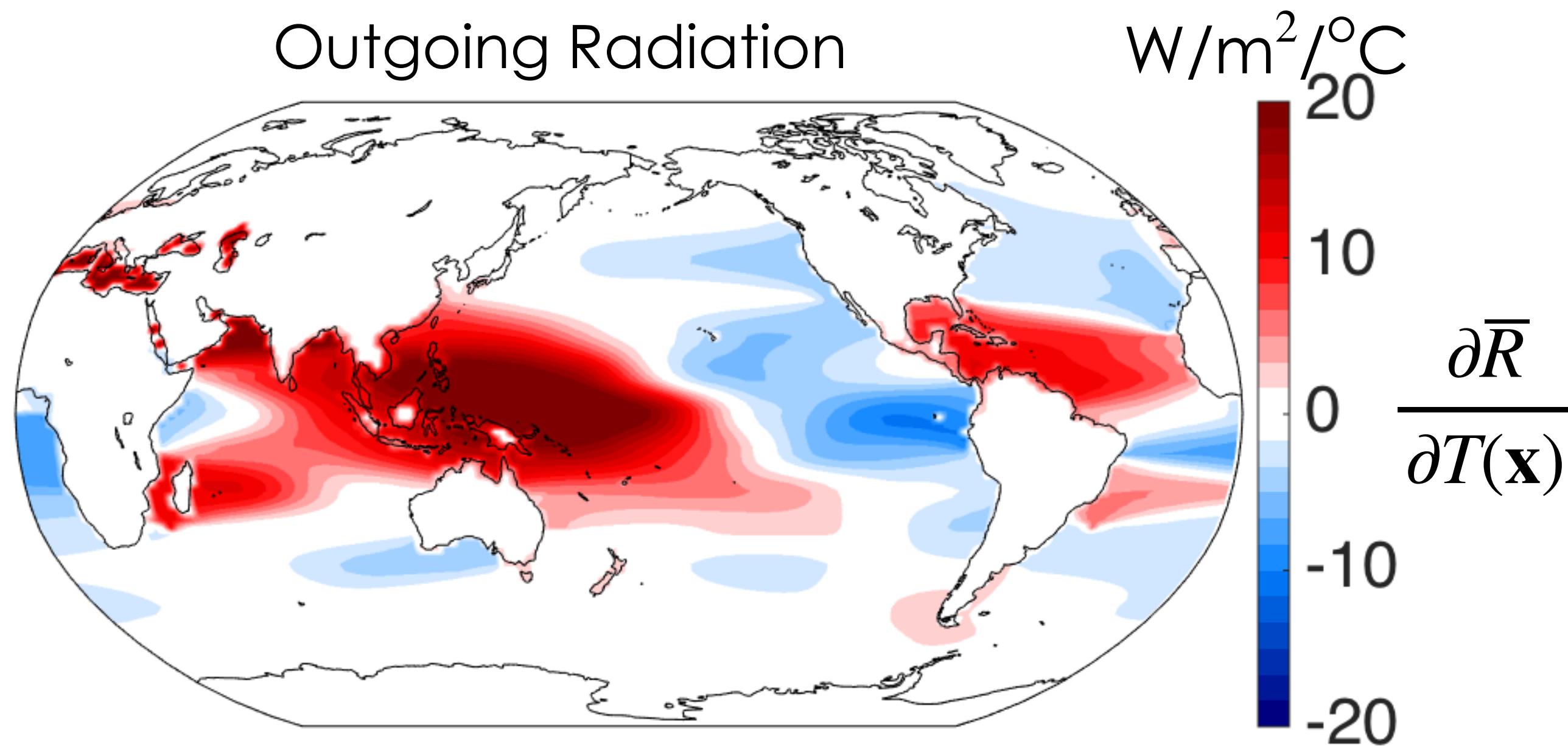
$$\lambda(t) = \frac{\Delta R(t)}{\Delta T(t)}$$

Inferred Climate Sensitivity:

$$\Delta T_{2\times} = \frac{\Delta F_{2\times}}{\lambda(t)}$$

Determined by “How efficient is the Earth at restoring energy balance through warming?”

Radiative feedback depends on **pattern** of warming



Global radiation response to

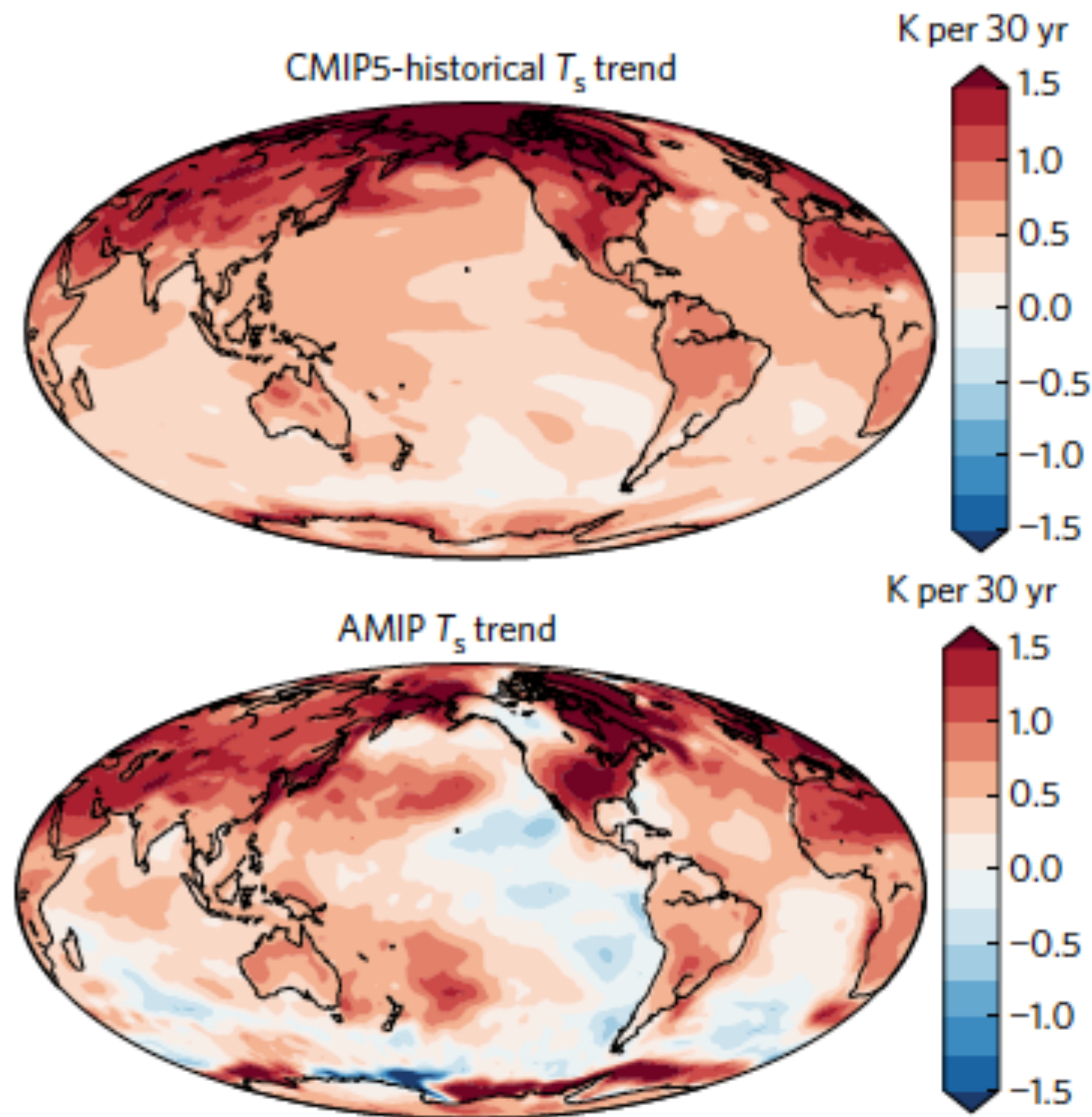
- 1°C of warming in West Pacific: +30 W/m²
- 1°C of warming in East Pacific: -10 W/m²

Instantaneous Radiative feedback
(Efficiency of radiative damping)

$$\lambda(t) = \frac{\Delta R(t)}{\Delta T(t)}$$

$$\Delta R = \int_y \frac{\partial \bar{R}}{\partial T(x)} dT$$

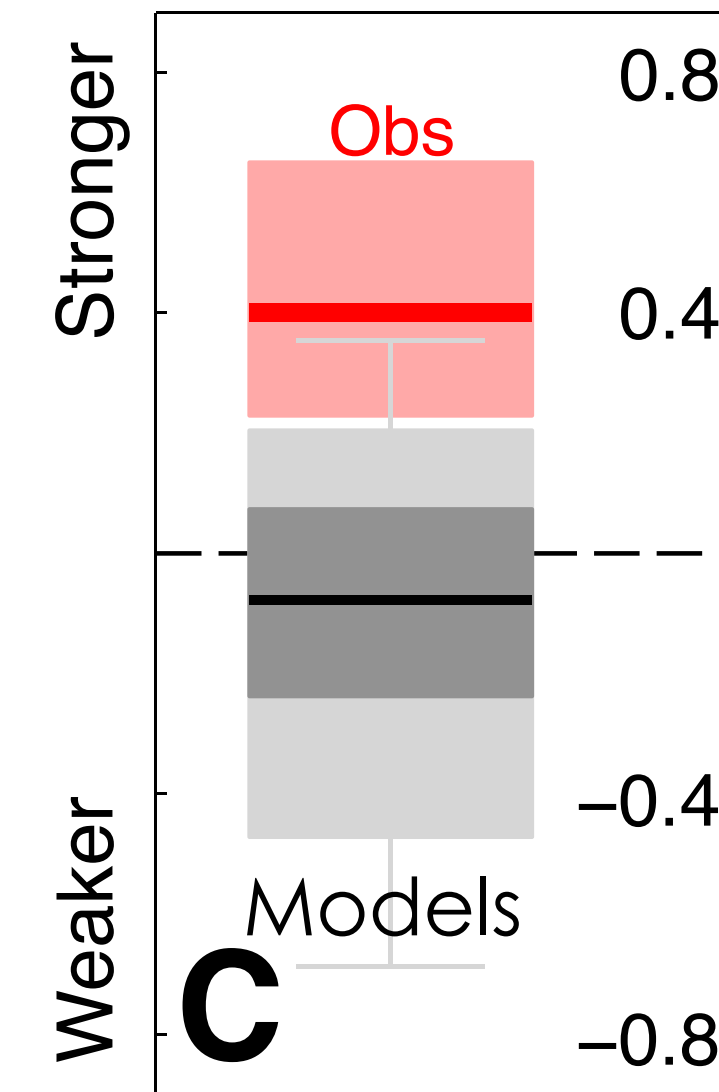
Models do not reproduce the observed pattern of warming



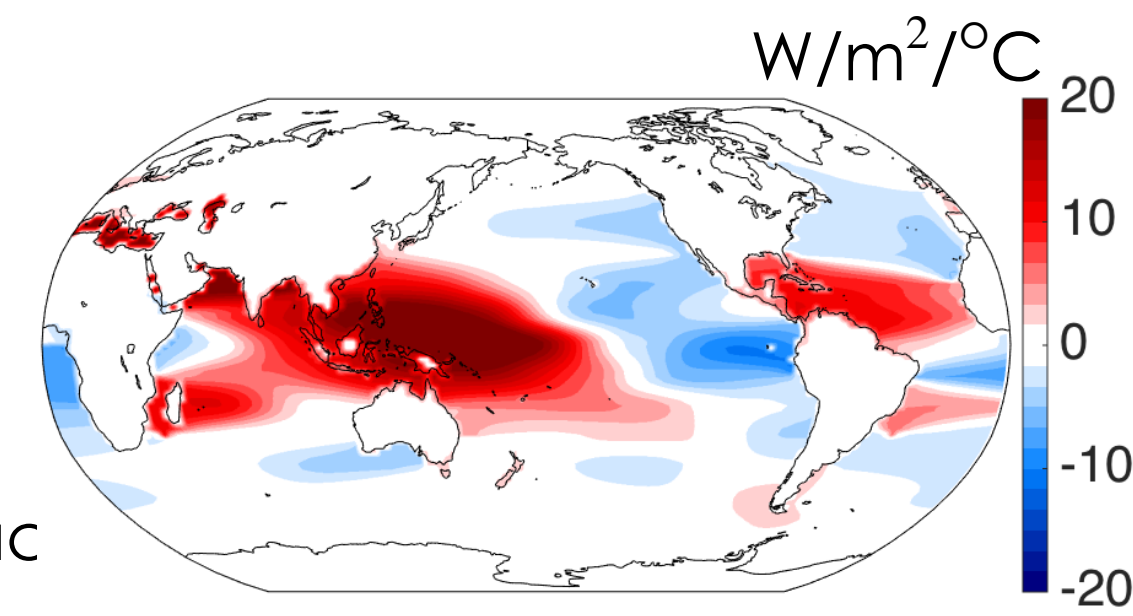
Zhou et al 2016

- *historical*: coupled models with dynamical oceans, forced with historical radiative forcing
- *AMIP*: atmosphere models forced with the *observed* history of SSTs

Changes in W-E Equatorial Pac
SST gradient (C/century)

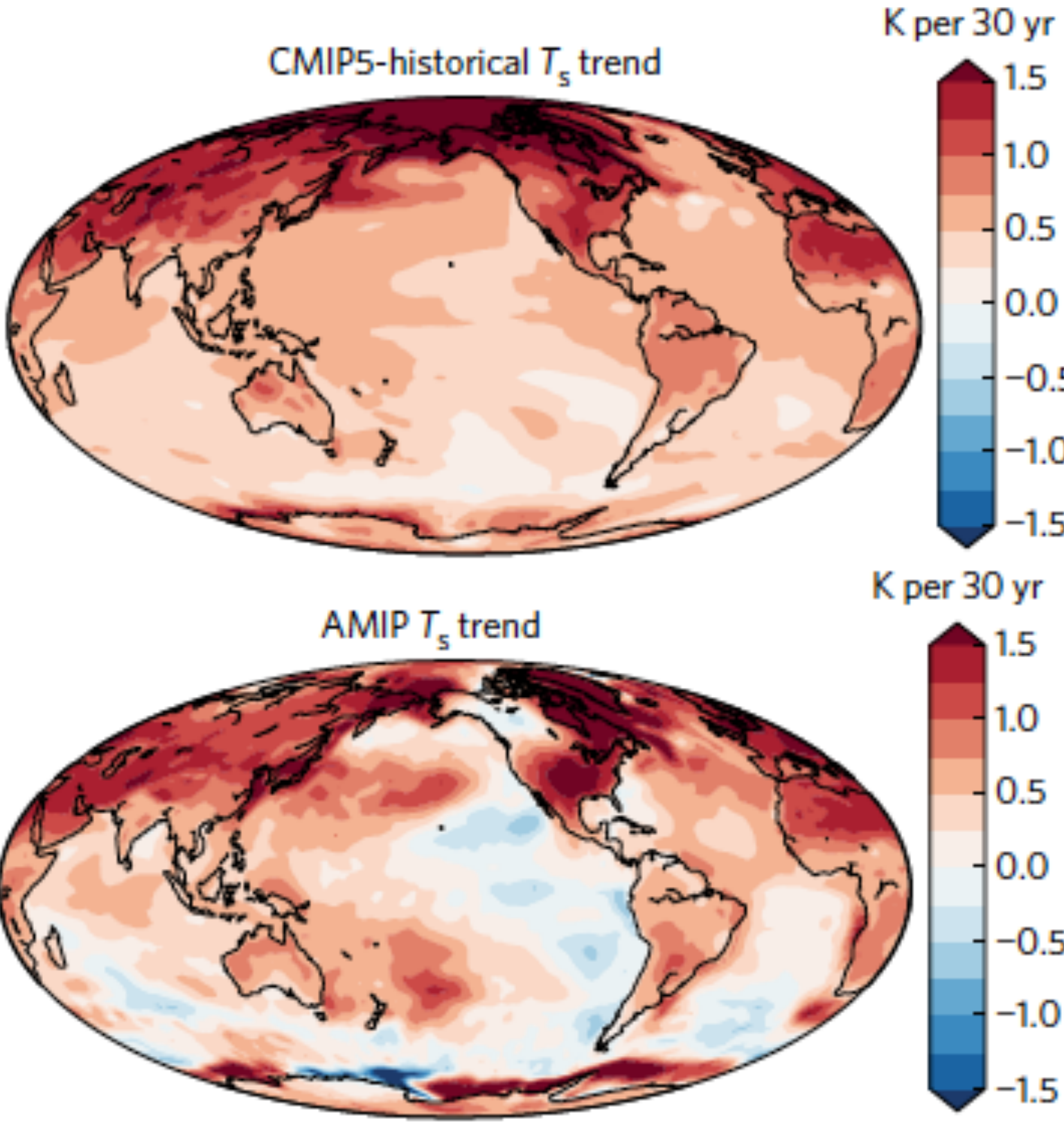
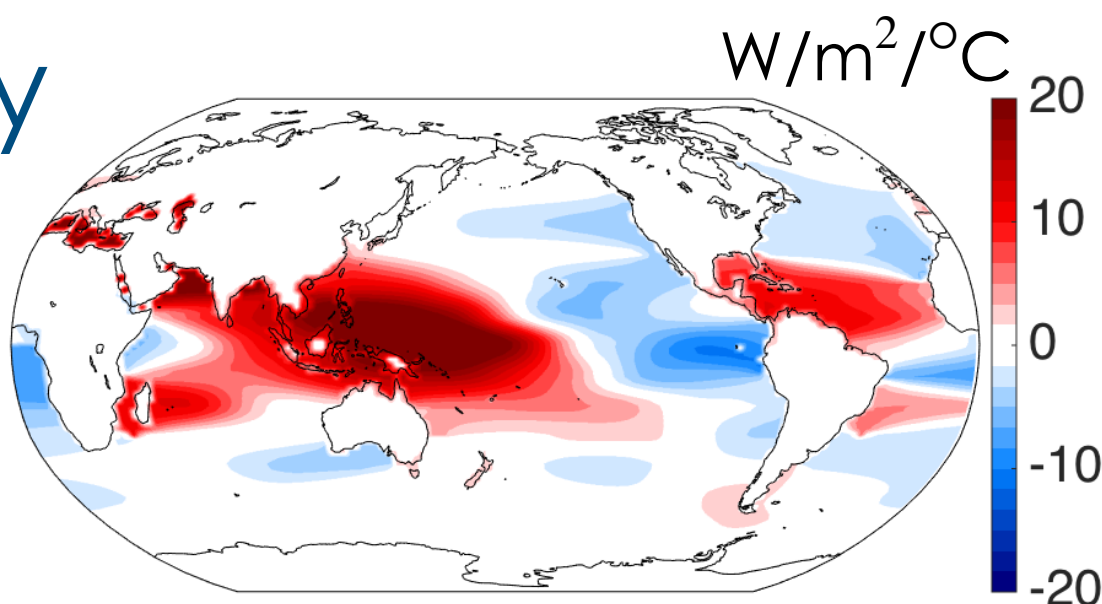


Coates and Karnauskas 2018

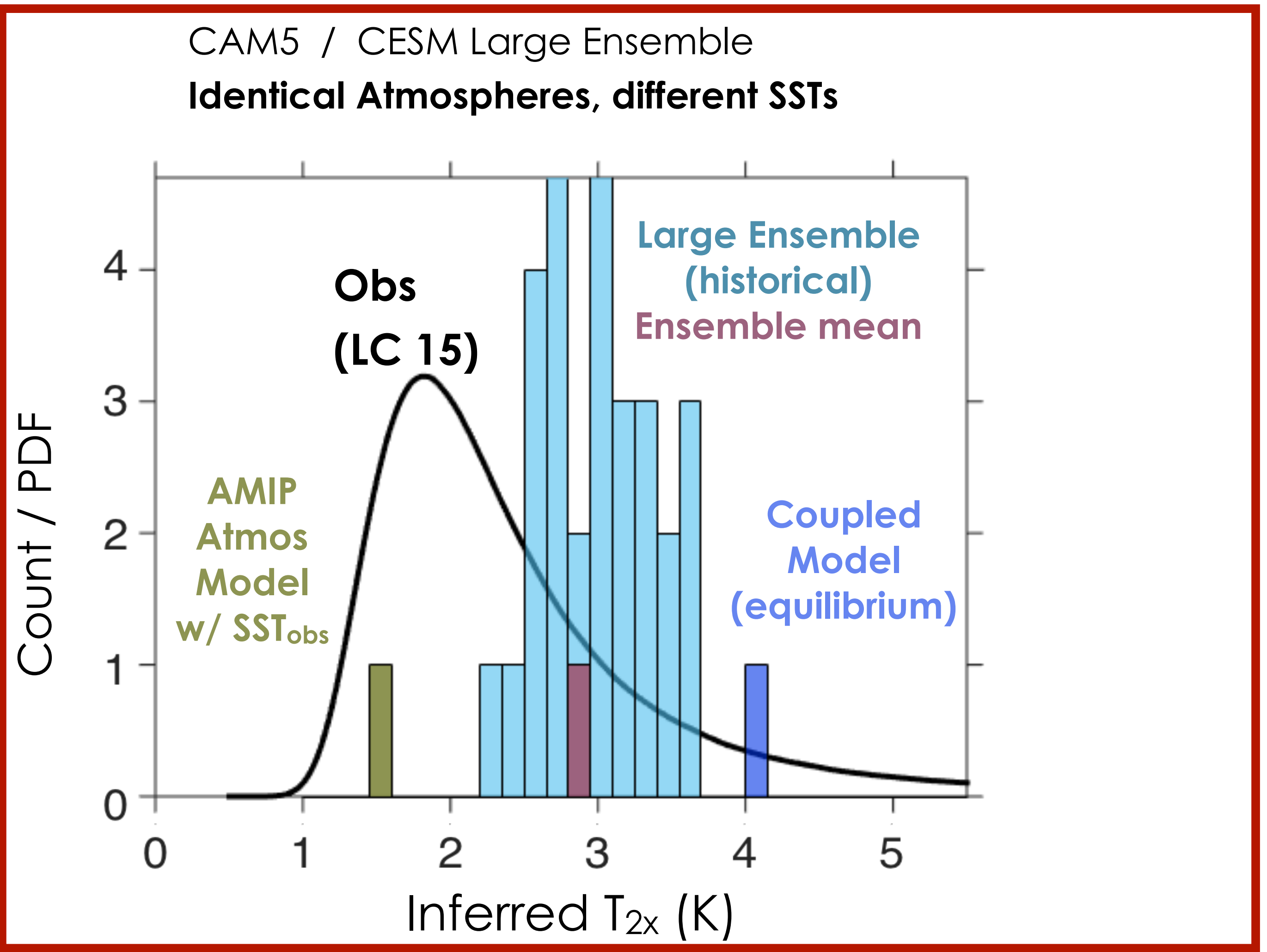


Coupled climate models are generally not able to reproduce the observed evolution of Sea Surface Temperature (SST) patterns

SST mismatches lead to large differences in inferred climate sensitivity



Zhou et al 2016



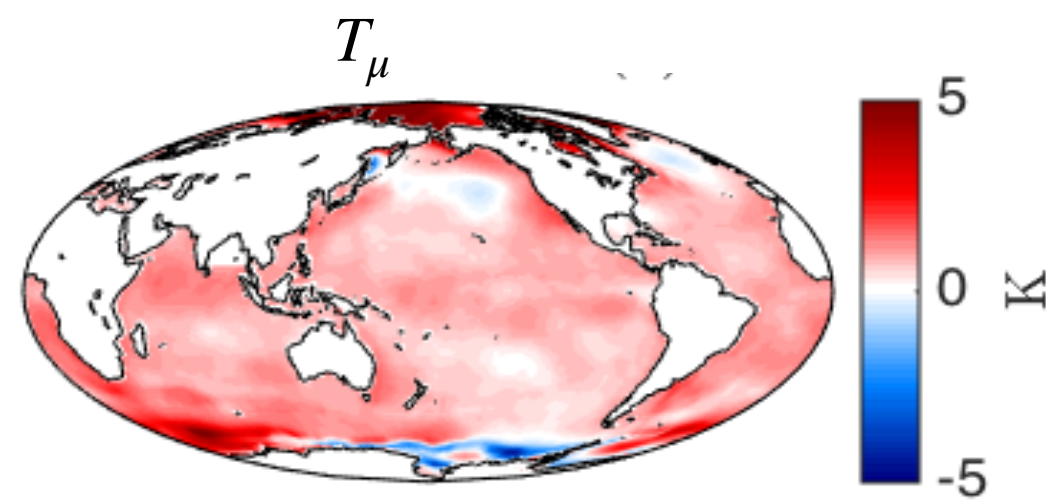
Mean pattern
(constant feedback)

Patterned SST field

$$\lambda(t) = \lambda_\mu + \frac{R_{\text{patterned}}(t)}{\bar{T}(t)}$$

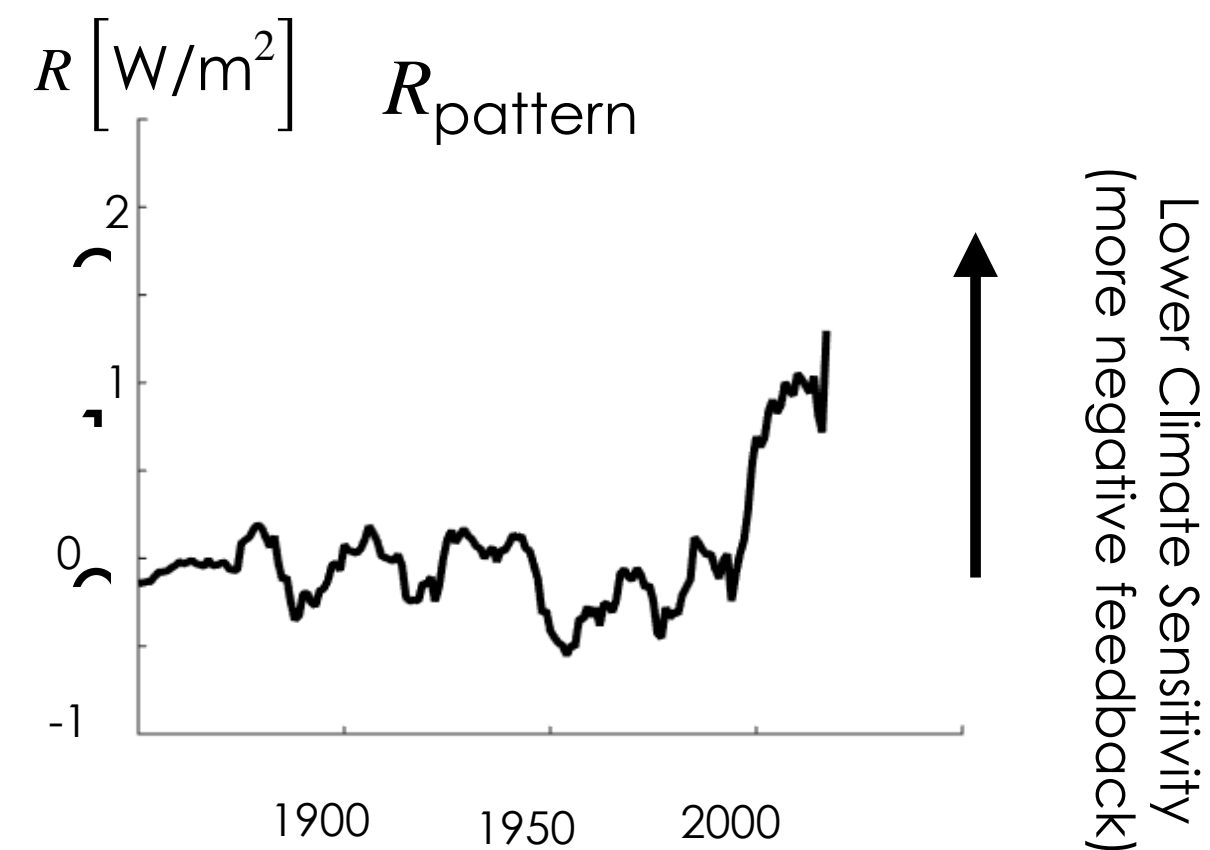
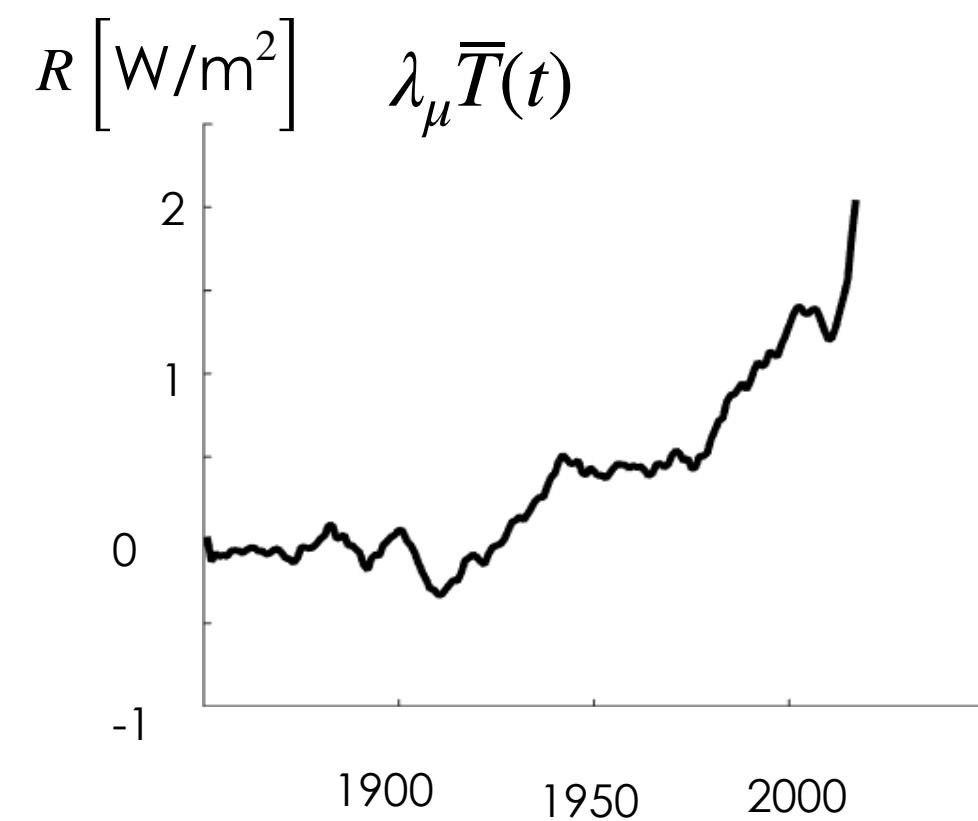
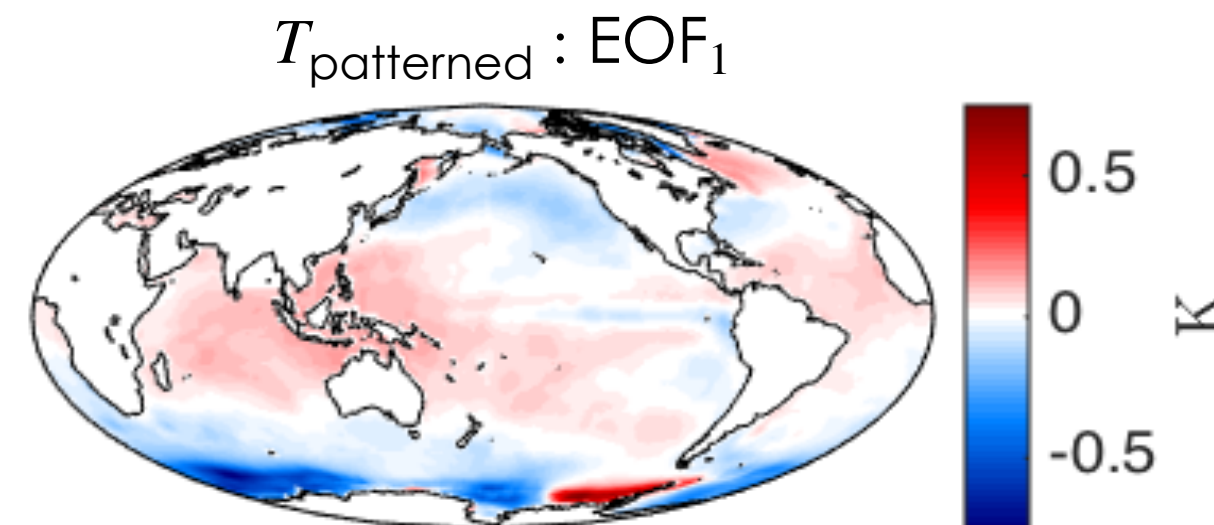
Constant pattern
Constant feedback

$$\lambda_\mu = \int \frac{\partial \bar{R}}{\partial T(\mathbf{x})} T_\mu(\mathbf{x})$$



Changes in SST pattern
(Zero global mean)

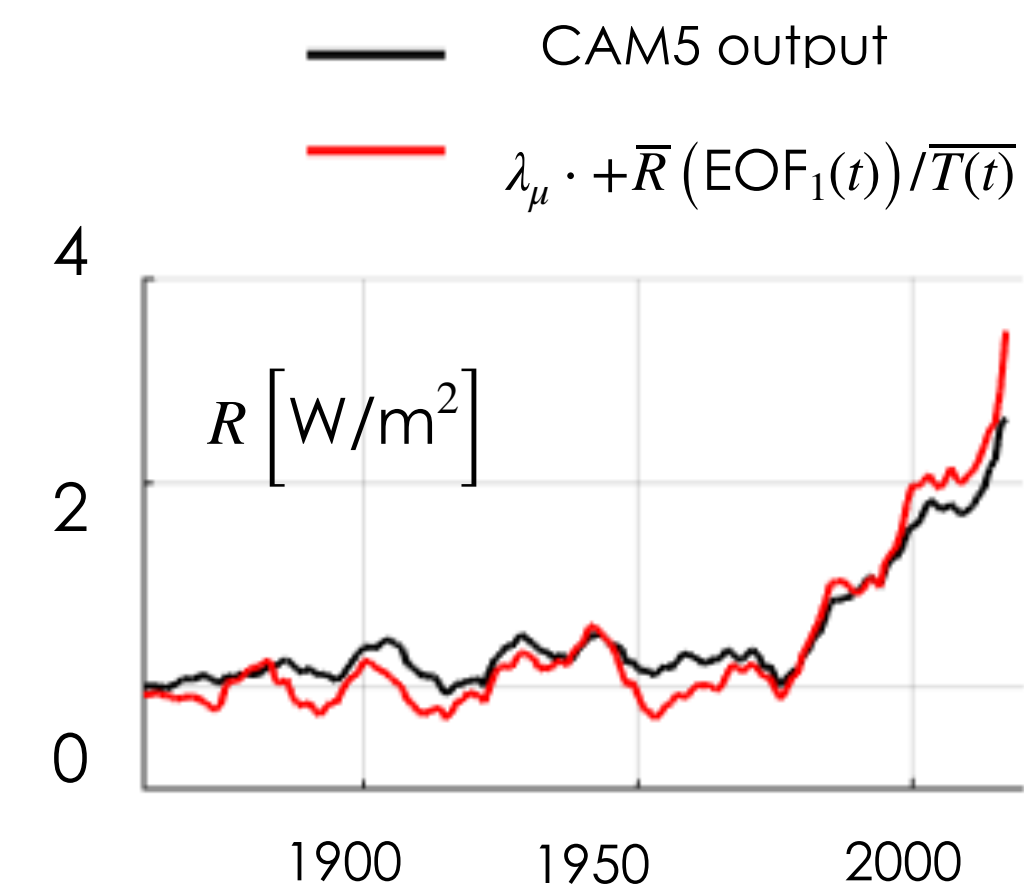
$$R_{\text{pattern}}(t) = \int \frac{\partial \bar{R}}{\partial T(\mathbf{x})} T_{\text{pattern}}(\mathbf{x}, t)$$



Pattern change in **historical** SSTs

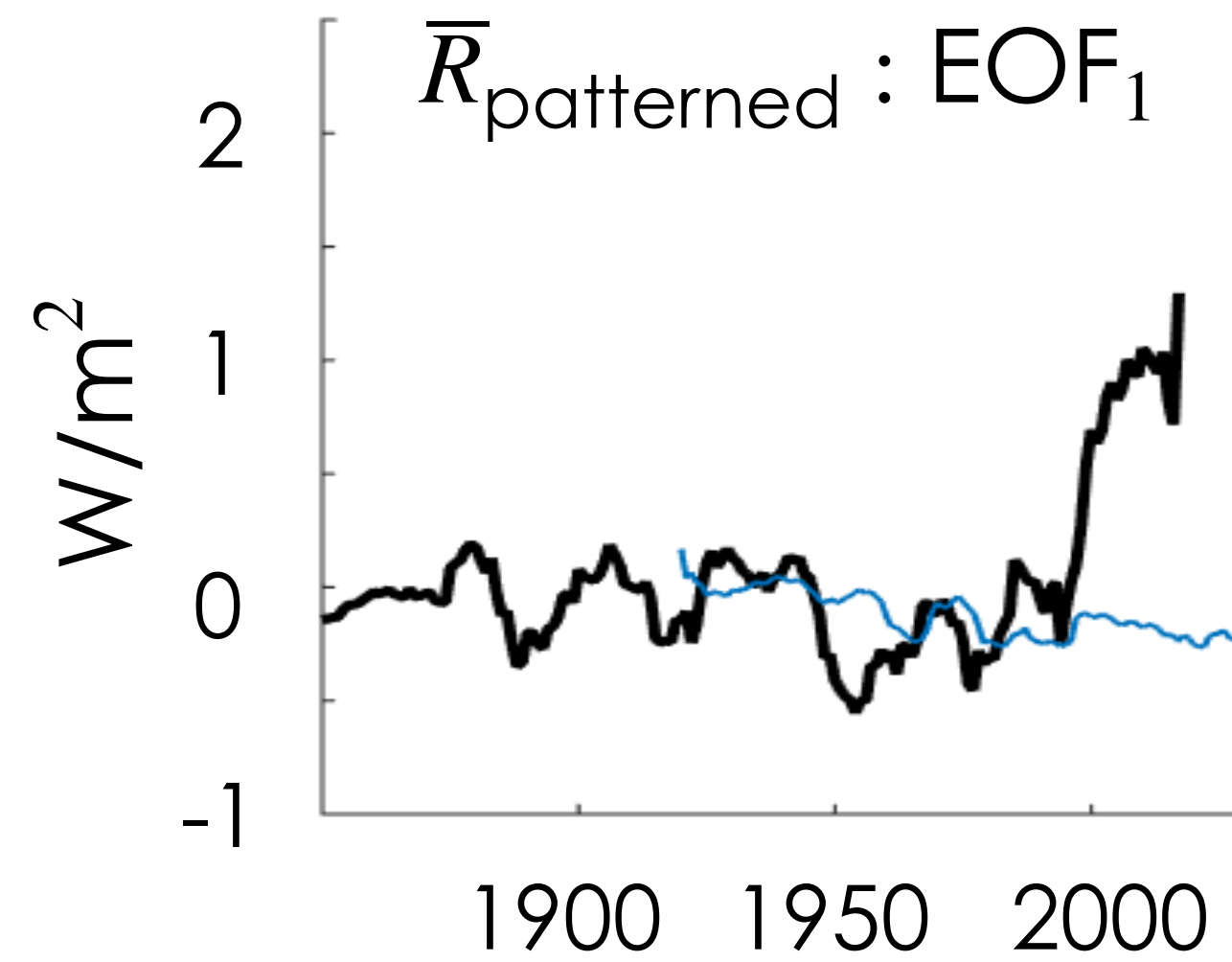
We decompose the feedback into a constant feedback plus a contribution from changing SST patterns (the pattern effect).

Mean Pattern + 1st EOF of *SSTs weighted by green's function* are sufficient to explain changes in outgoing radiation in CAM5-AMIP

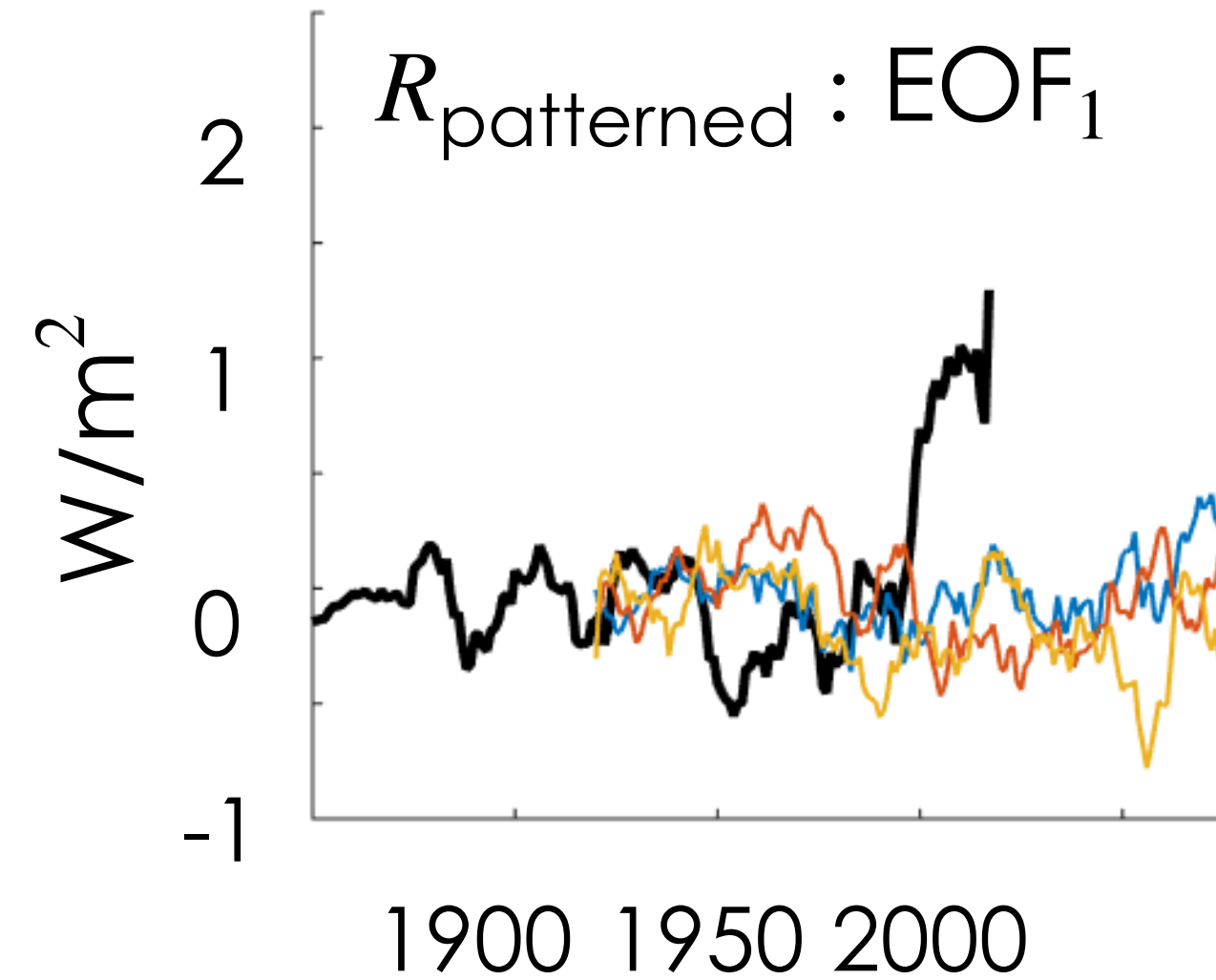


Coupled model SSTs lack significant pattern effect

Large Ensemble Average



Large Ensemble deviations



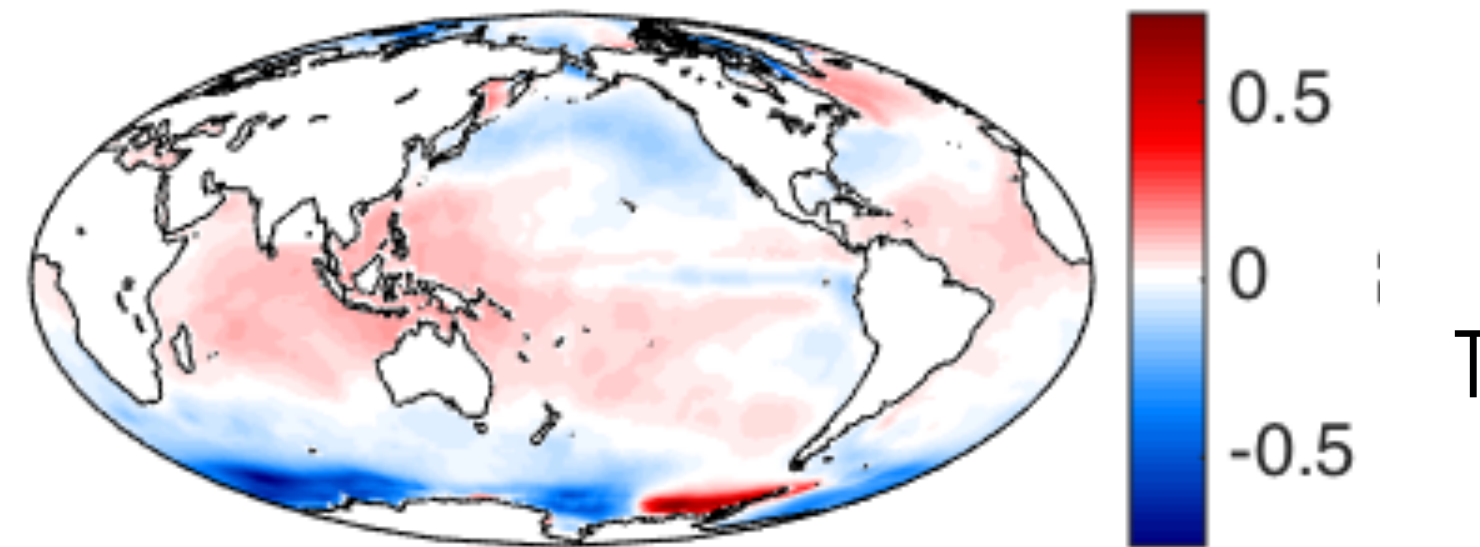
Lower Climate Sensitivity
(more negative feedback)

- AMIP run (CAM5 + historical SST)
- LE Ensemble average (forced)
- — — Anomalies from LE Ensemble average

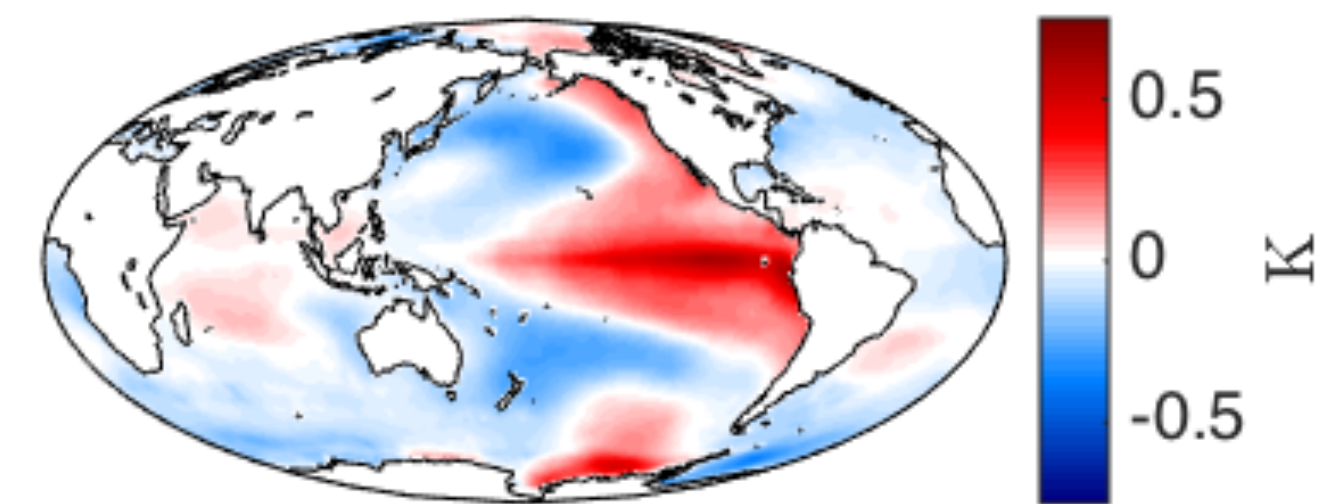
Coupled model SSTs lack sufficient variability in Indo-Pacific Warm pool

$$\text{EOFs of : } T(x, t) \cdot \frac{\partial \bar{R}}{\partial T(x)}$$

EOF 1

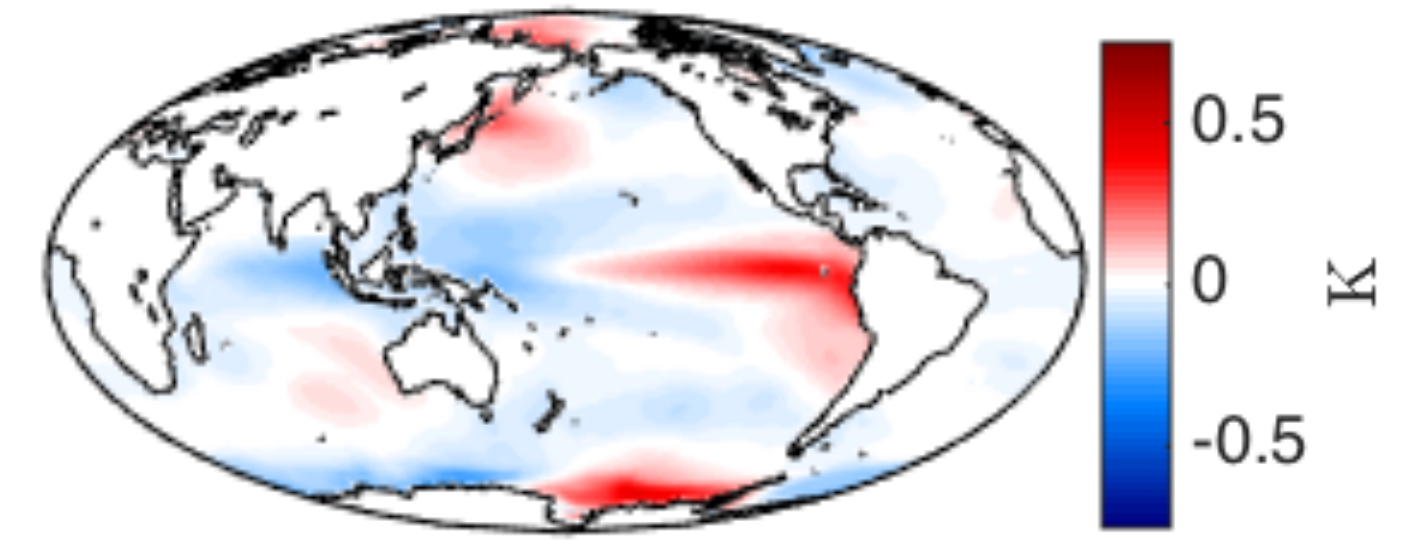
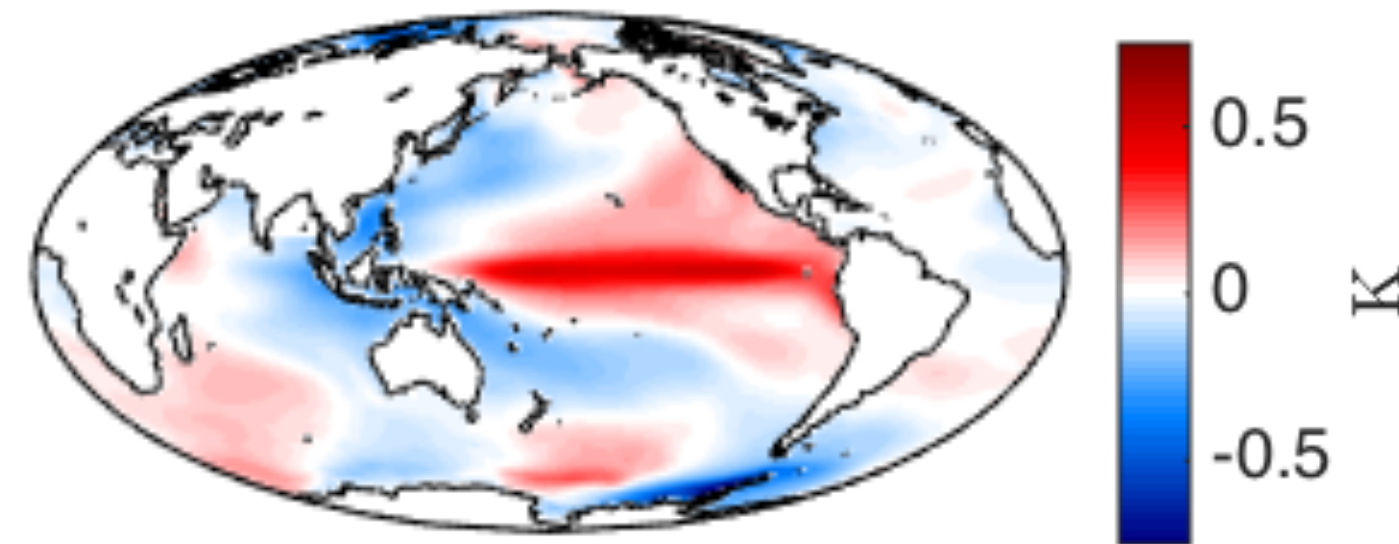


EOF 2



Observed SSTs

Large Ensemble



Thank You!

Questions?

Collaborators:

Yue Dong

Malte Stuecker

Kyle Armour

David Battisti

Rob Wills

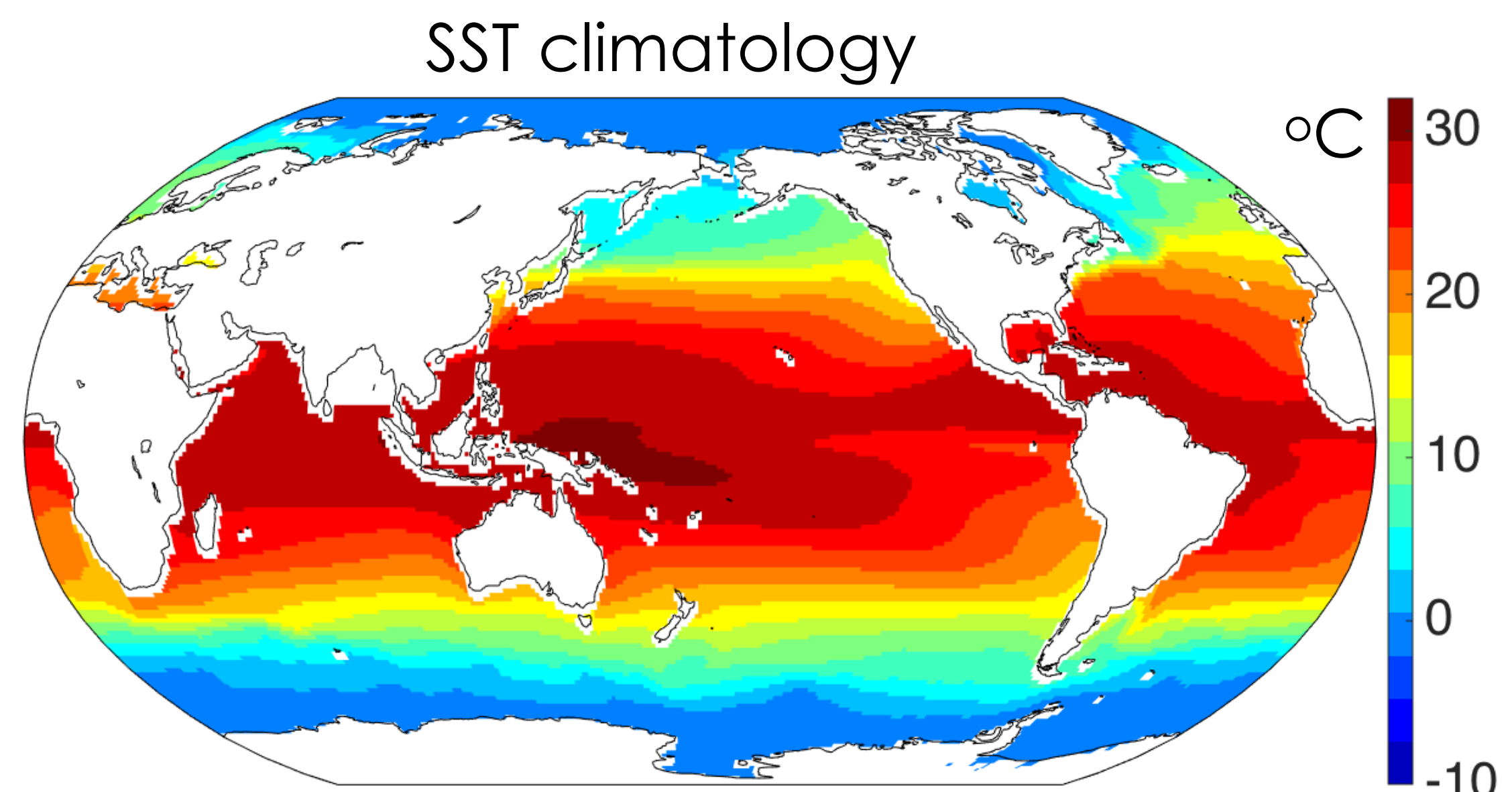
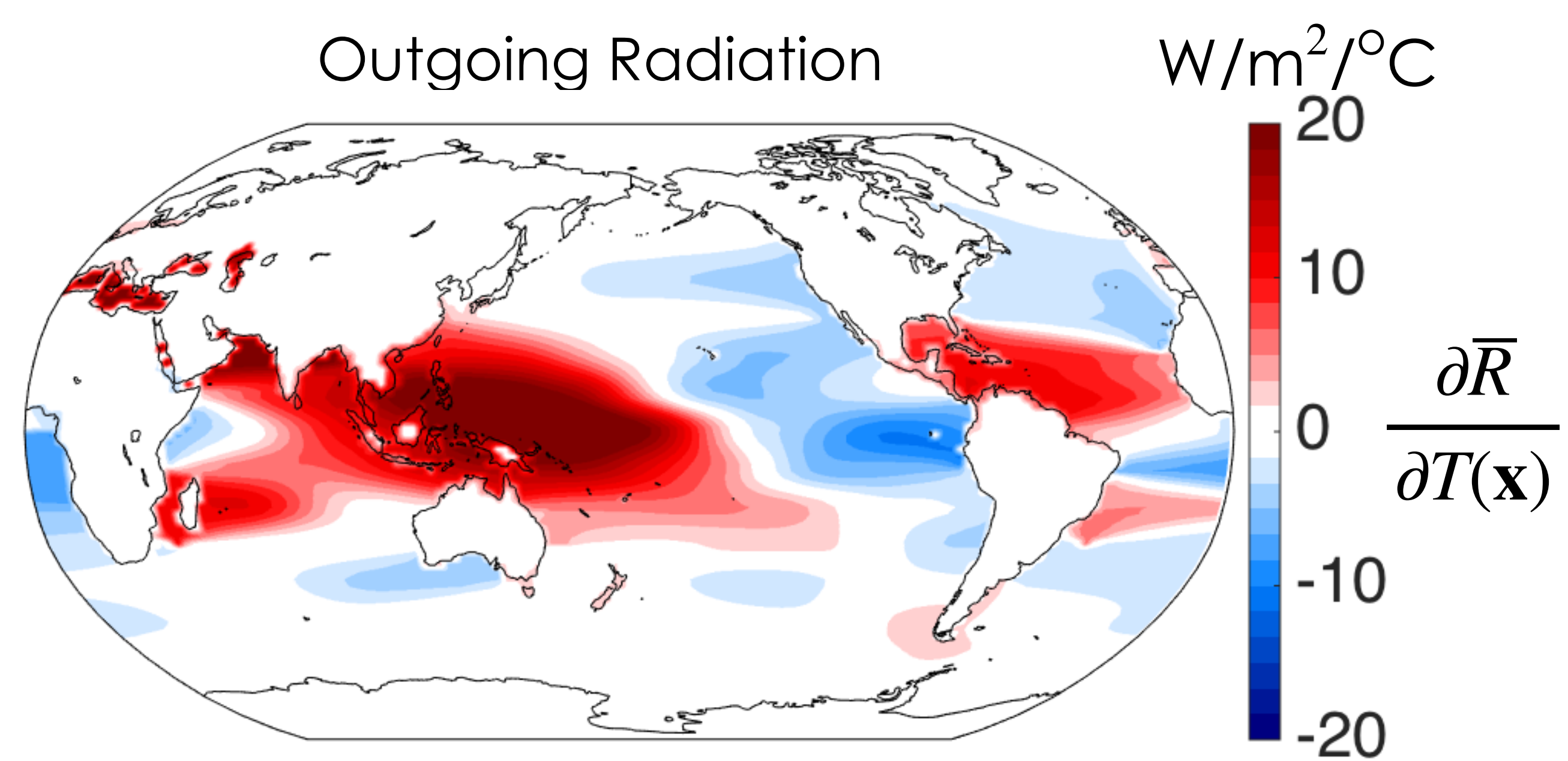
Luke Parsons

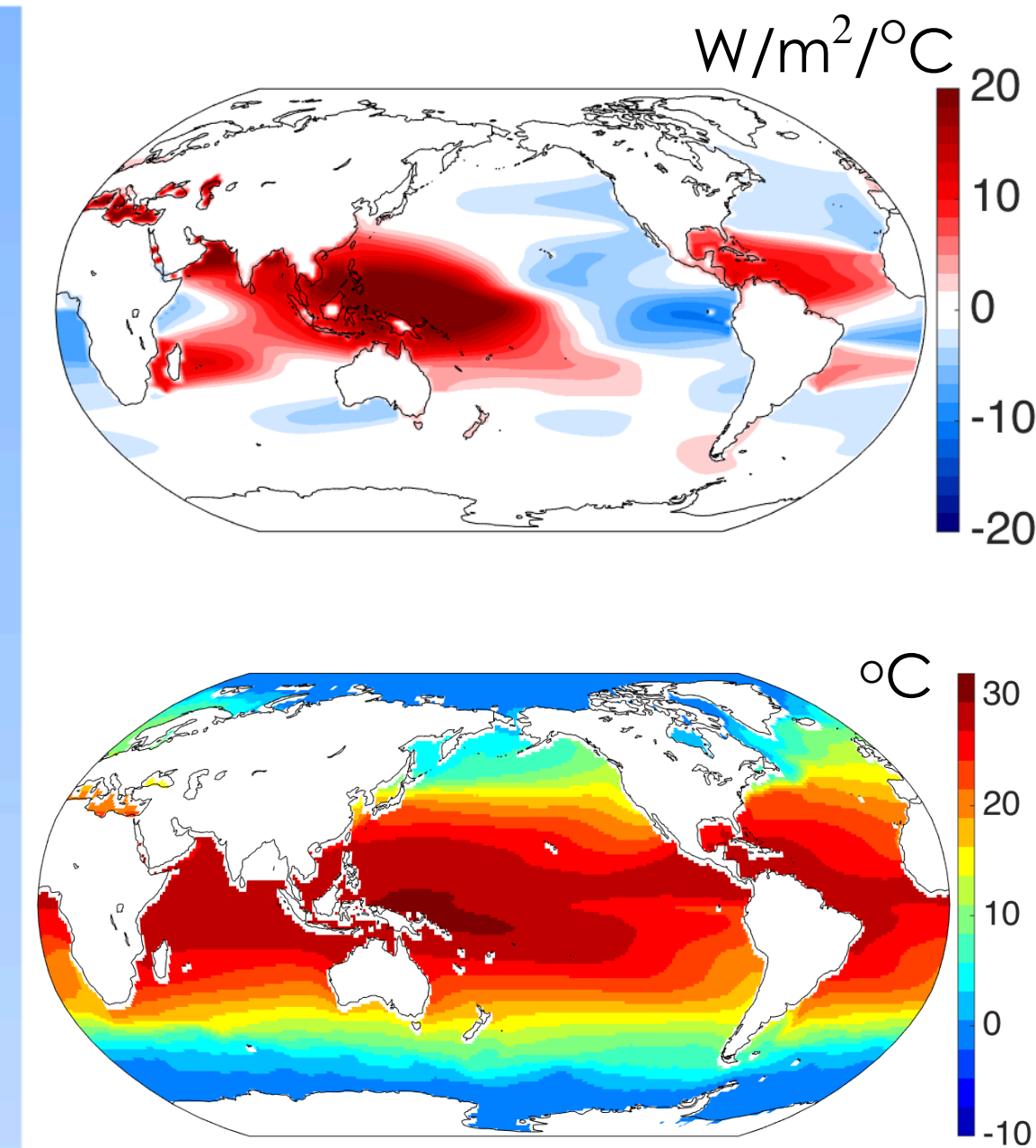
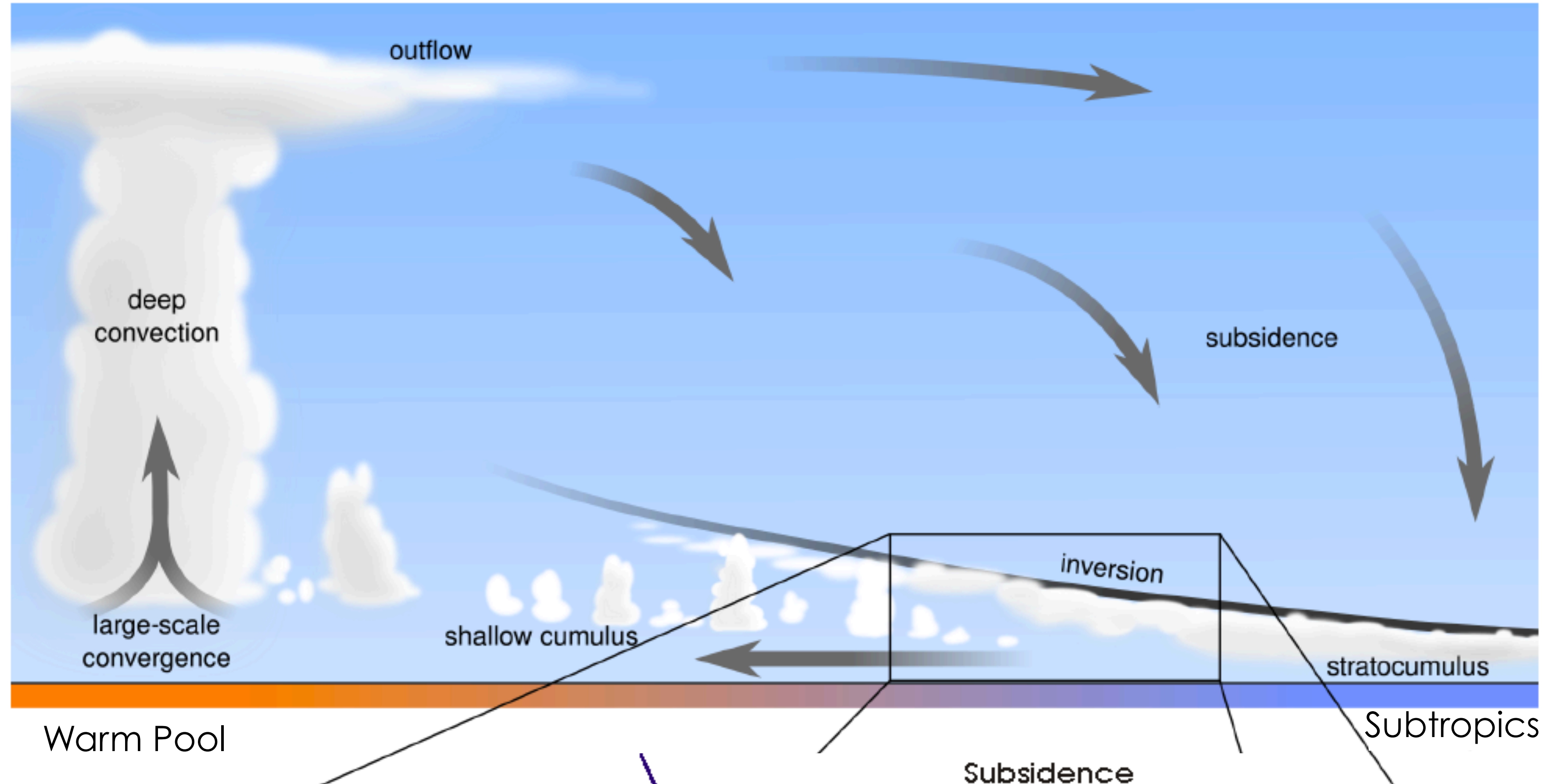
- University of Washington, Atmospheric Sciences,
- University of Hawaii, Oceanography
- University of Washington, School of Oceanography
- University of Washington, Atmospheric Sciences
- University of Washington, Atmospheric Sciences,
- University of Washington, Atmospheric Sciences

Supplementary slides

How does radiation depend on the pattern of warming?

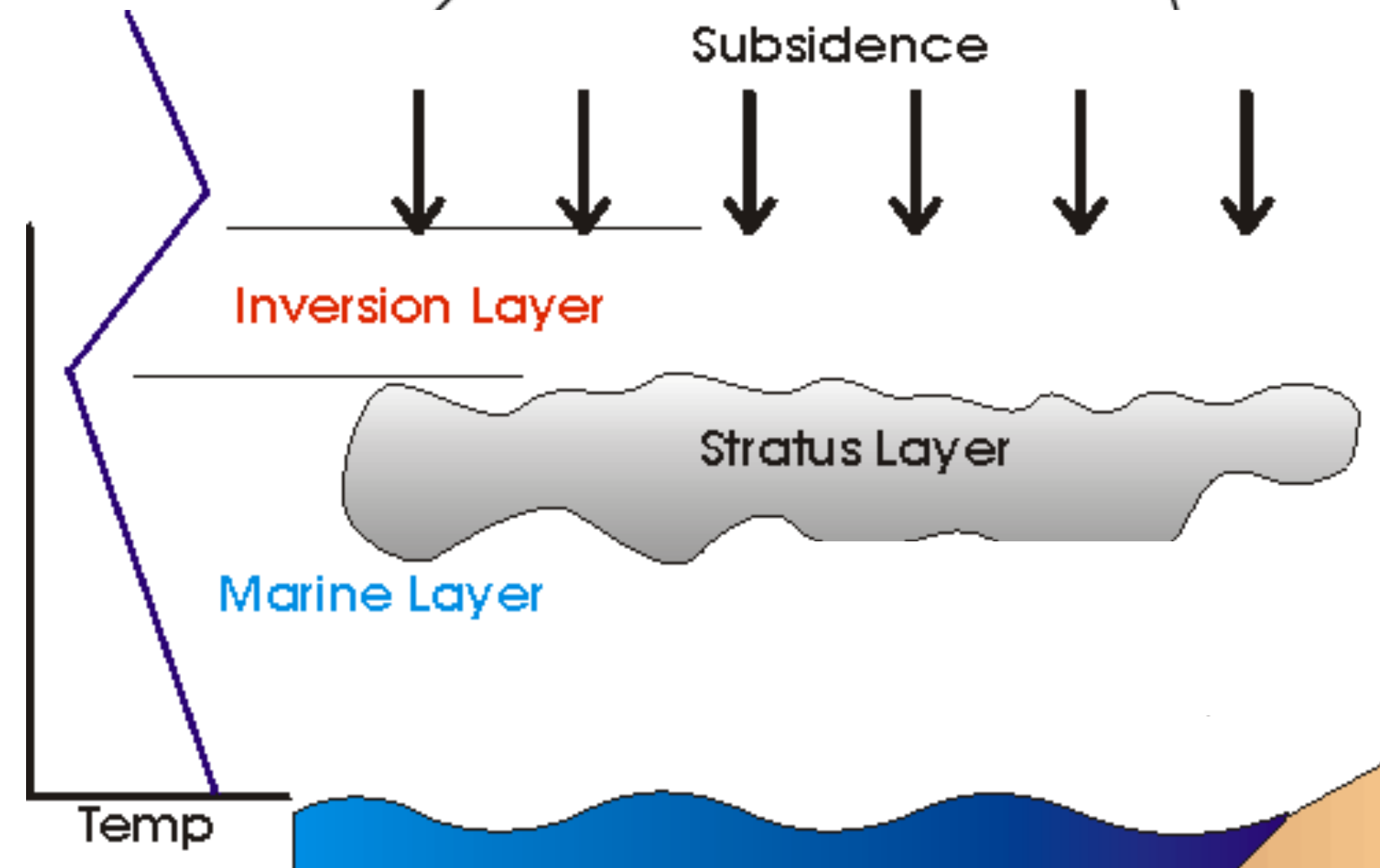
Radiative response to local warming **tied to SST climatology**

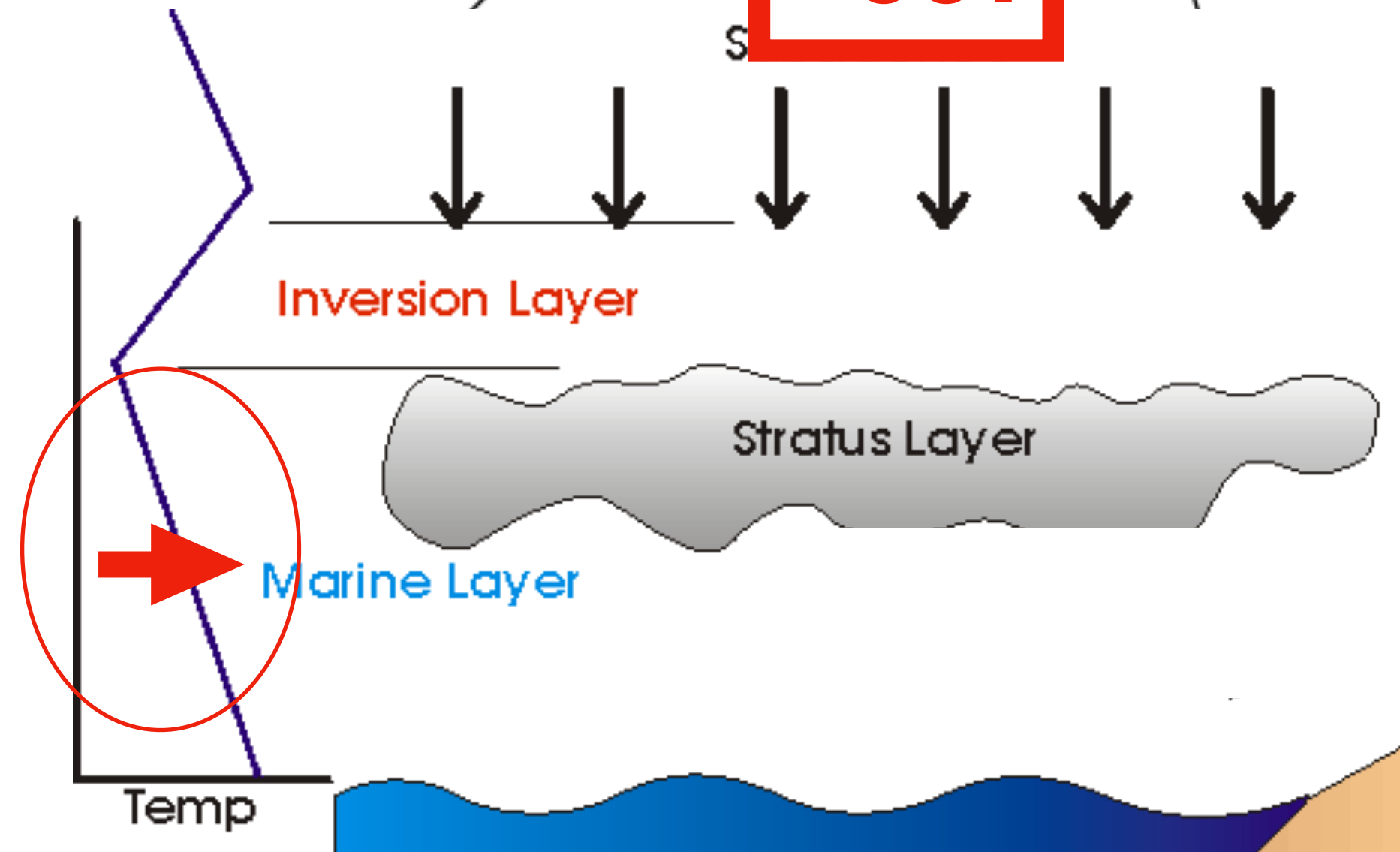
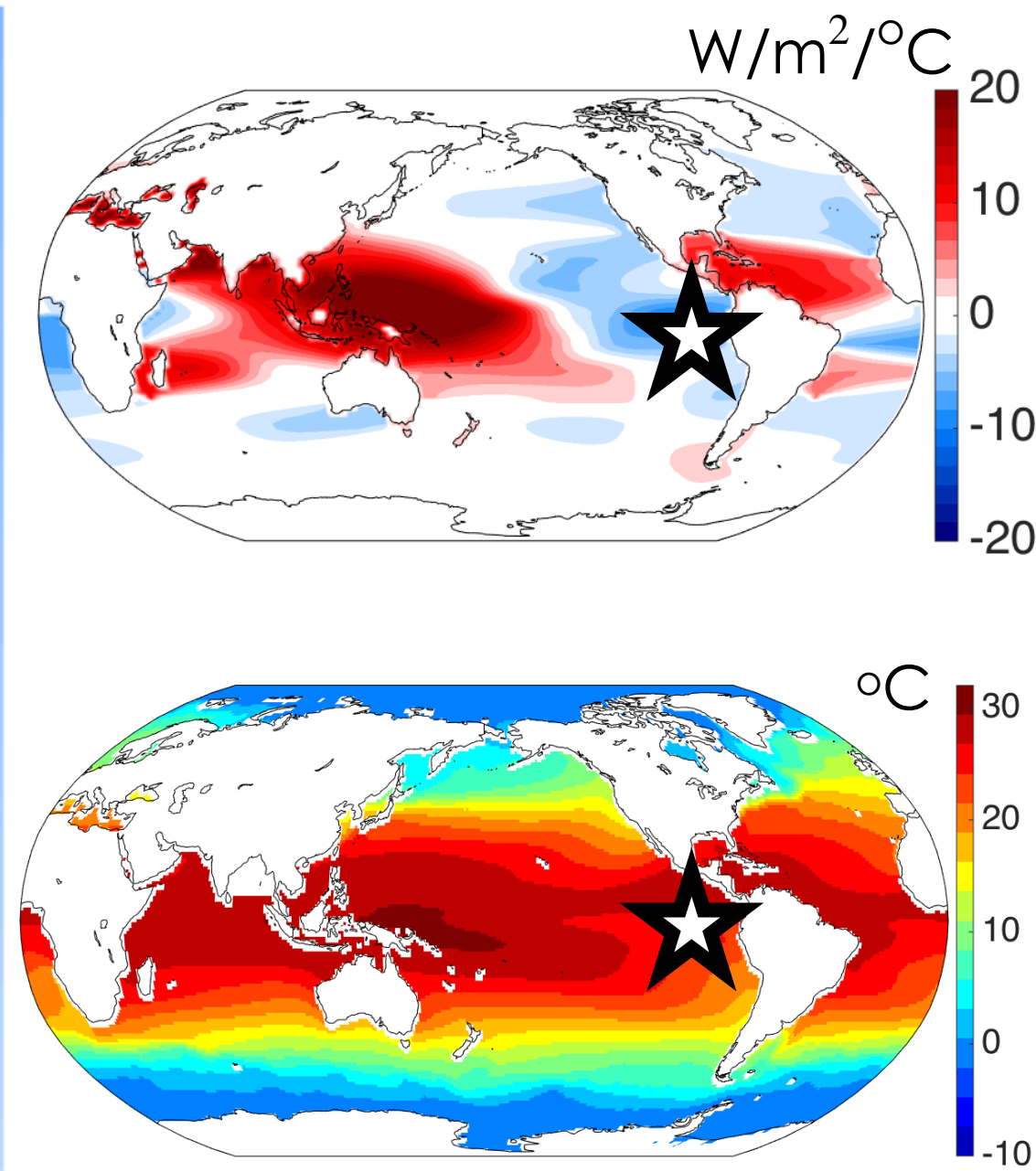
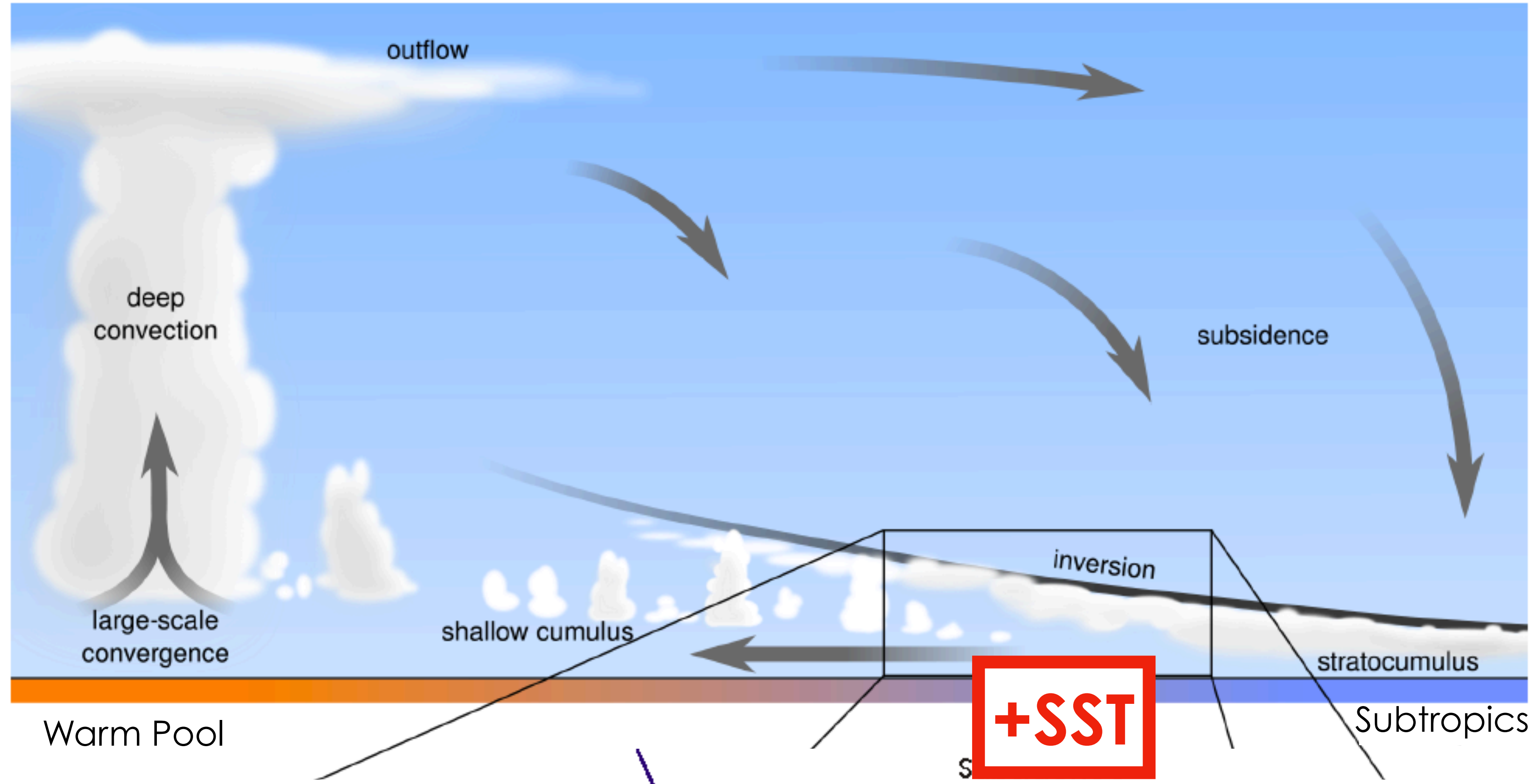




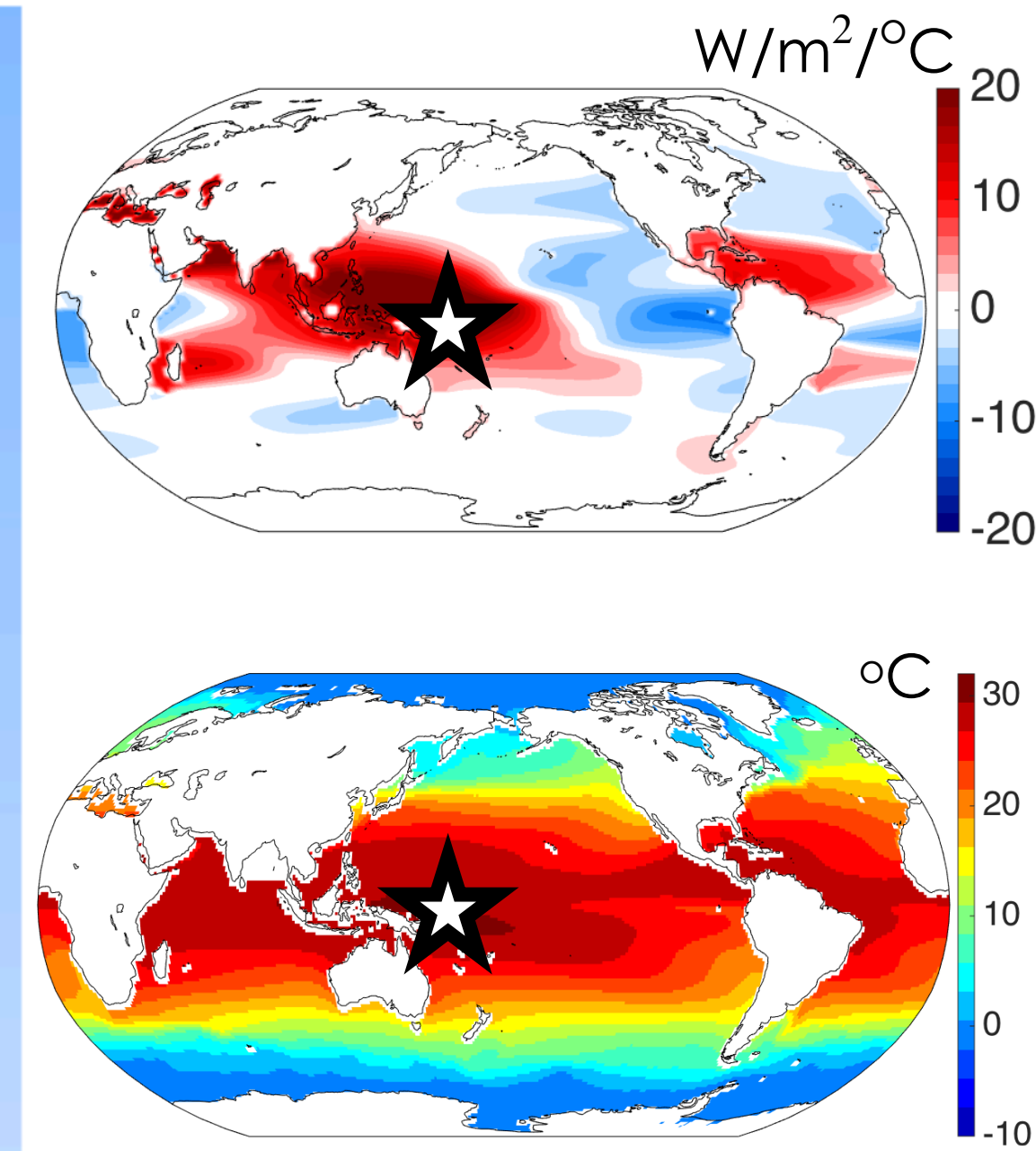
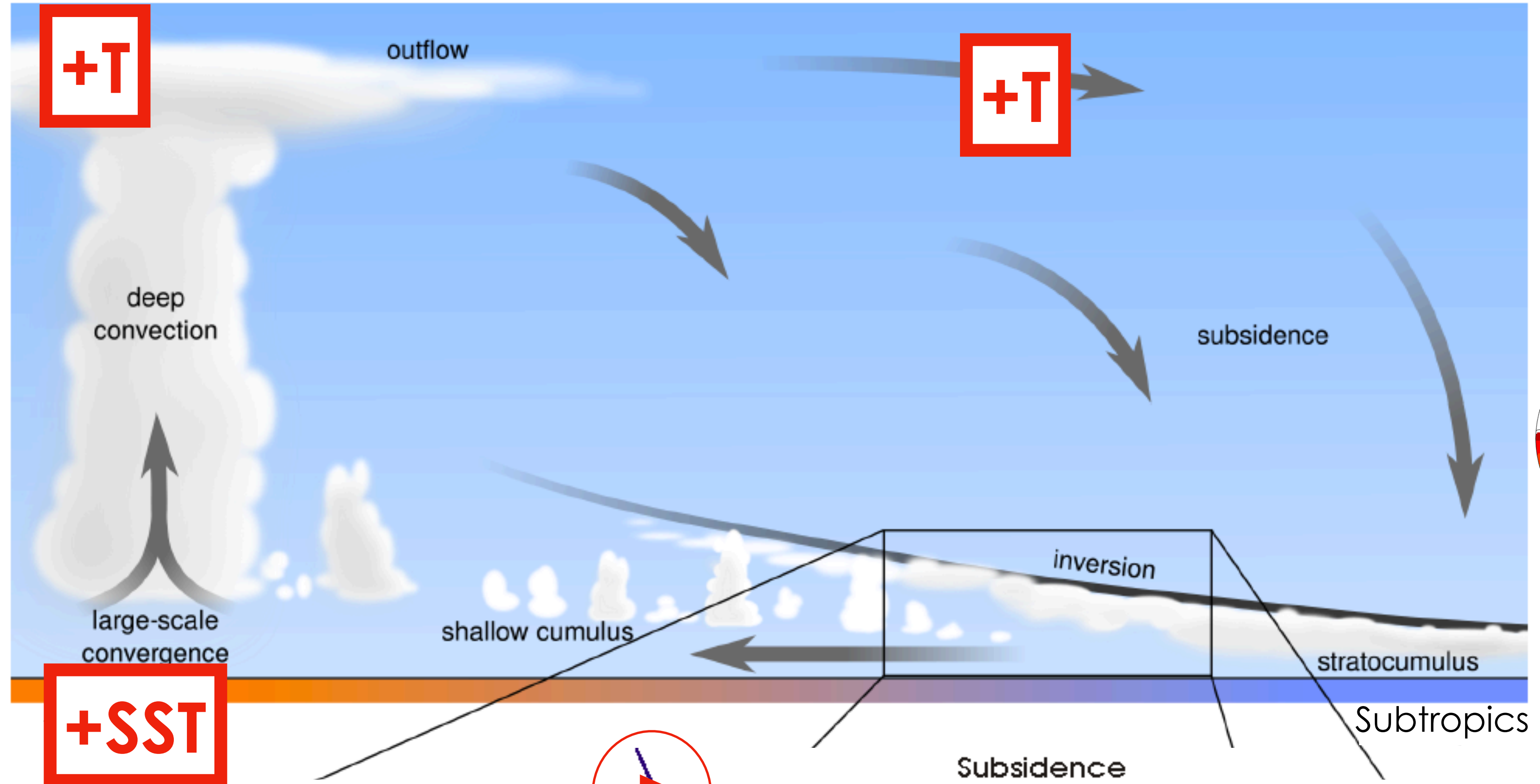
Low cloud amount depends on strength of inversion and local temperature

Klein and Hartman 1993,
Wood & Bretherton 2006 ,
Bretherton & Blossey 2014

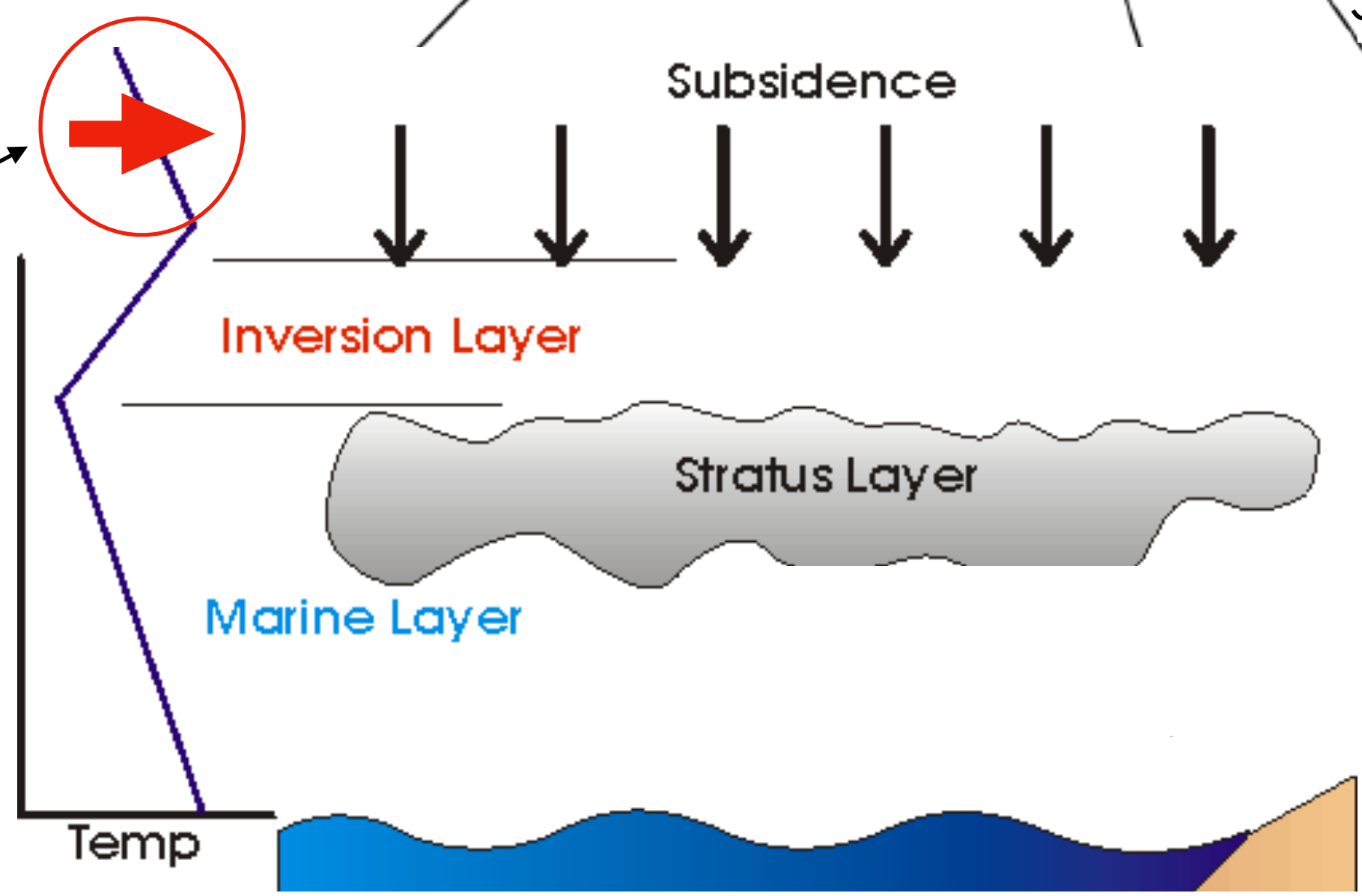




Controlled by East Pacific SST
 +EP SST - LCC

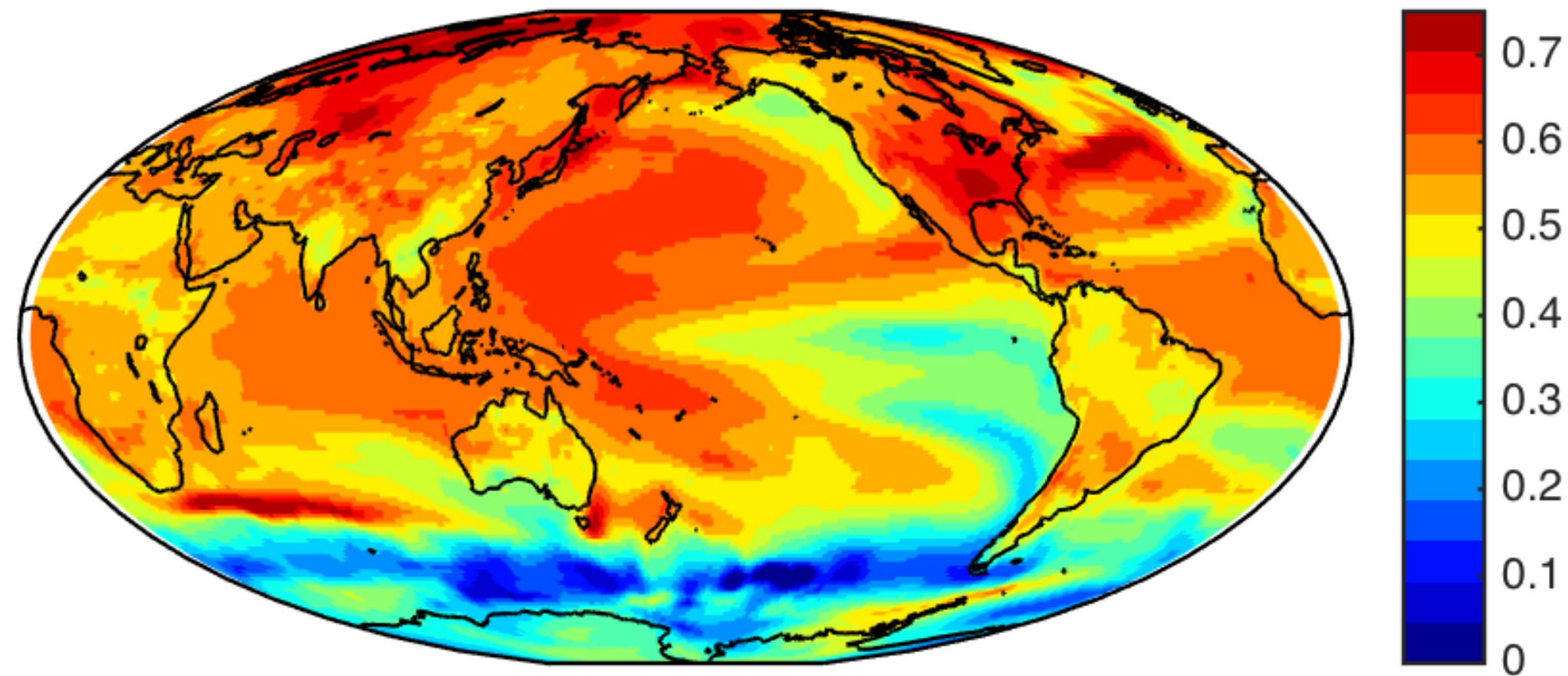


Controlled by West Pacific SST
 +WP SST + LCC

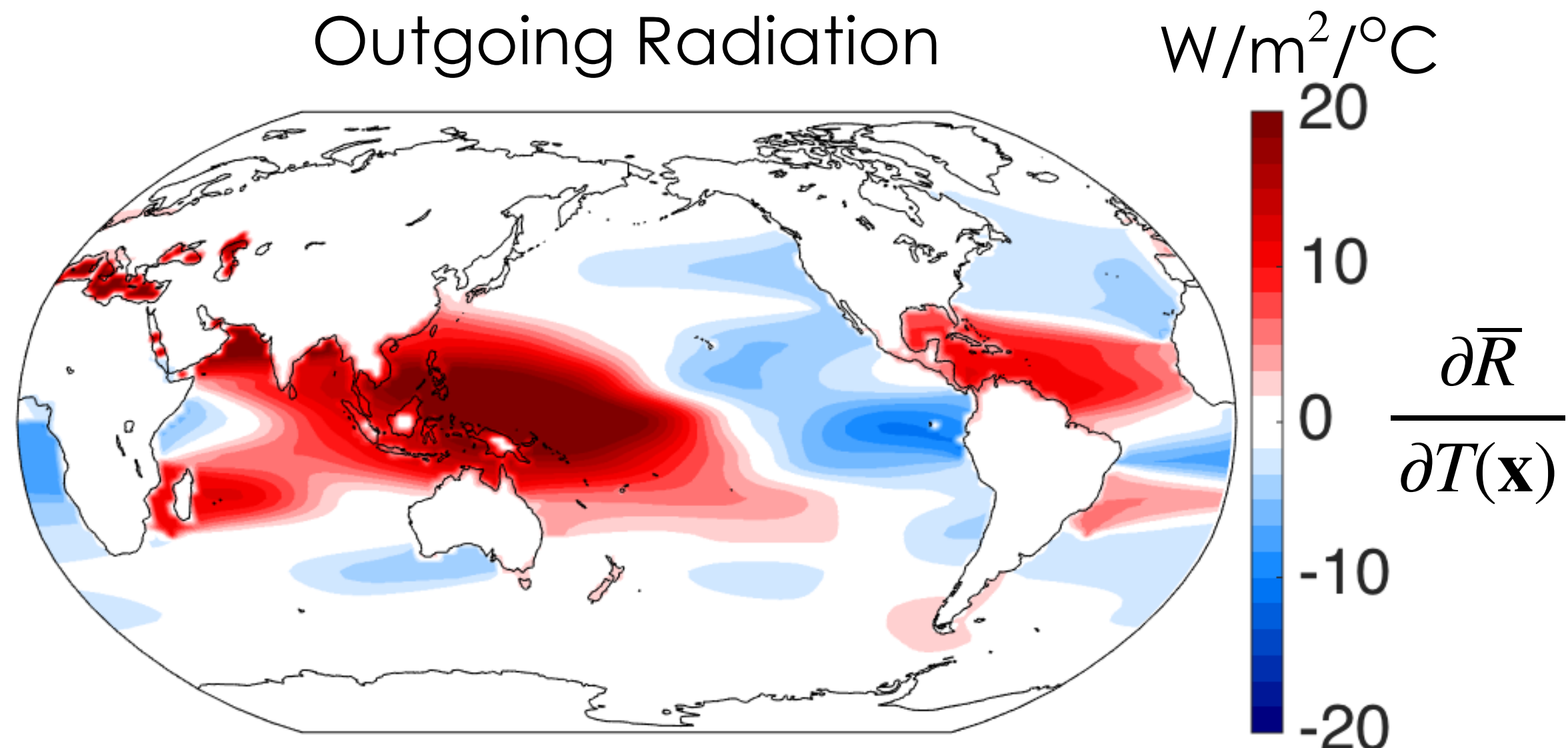


The “Pattern Effect” summary

Fast Warming / Total Warming



Outgoing Radiation



Pattern effect:

- **Fast modes of warming:**

- Land + Mixed Layer
- More West Pacific warming
- More efficient at radiating to space

- **Slow warming:**

- Regions of deep ocean heat uptake + East Pacific
- Less efficient at radiating to space

Proistosescu & Huybers 2017

Dong, Proistosescu, Armour, Battisti, 2019