

3D convection, phase change, and solute transport in mushy sea ice

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Summary:

- Simulated brine drainage via 3-D convection in porous mushy sea ice
- Shallow region near ice-ocean interface is desalinated by many small brine channels
- Full-depth "mega-channels" allow drainage from saline layer near top of ice.



What is a mushy layer?

Dense brine drains convectively from porous mushy sea ice into the ocean.

- What is spatial structure of this flow in 3 dimensions?

Upper fig.: Sea ice is a porous mixture of solid ice crystals (white) and liquid brine (dark). *H. Eicken et al. Cold Regions Science and Technology* 31.3 (2000), pp. 207–225

Lower fig.: Trajectory (\rightarrow) of a solidifying salt water parcel through the phase diagram. As the temperature *T* decreases, the ice fraction increases and the residual brine salinity S_{I} increases making the fluid denser, which can drive convection.

Using a linear approximation for the liquidus curve, the freezing point is

$$T_f(S_I) = 0.1S_I$$





Problem setup

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Cold upper boundary T=-10^oC, no normal salt flux, g no vertical flow **Initial conditions** S=30g/kg $T=T_{freezing}$ (S=30g/kg) + 0.2^oC U = 0Plus small random O(0.01^oC) temperature perturbation 4m Х 4m **Open bottom boundary**

Inflow/outflow, with constant pressure Inflow: S=30g/kg, $T=T_{freezing}$ (S=30g/kg) + 0.2^oC Movie Contours of:

Ice permeability -function of ice porosity; (red lower, green higher ~ice-ocean interface)

Velocity (blue lower, purple higher).

https://drive.google.com/file/d/ 1JBltmurLZ1zHKXT-Qt-8pmV EHuJIYdKI/view?usp=sharing eil&ts=5eaef8d6





Results -- Permeability and Velocity



Fine channels coarsen, and mega-channel forms as time progresses.

Qualitative similarities with experiments



Fig 6d from Cottier & Wadhams (1999)

Photograph of dye entrainment in sea ic Fig 3c from Eide & Martin (1975)

Results -- Vertical Salinity Flux





Dark Blue - Strong Downward

Light Blue - Weak downward

Pink - weak upward.



Salt flux weakens in smaller channels as mega-channel develops

Liquid-region salinity and salt flux



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After strong initial desalination pulse, salt flux weakens over time

Discussion

- Initially many small brine channels form, then are consolidated into a single "mega-channel"
 - Single channel is robust over a range of domain sizes
- Shallow region near ice-ocean interface is desalinated by an array of many small brine channels
- Full-depth "mega-channels" allow drainage from saline layer near top of ice
- Comparison with observations:
 - Laboratory: Eide and Martin (1975),
 - "Icy fingers of death" -- BBC Earth











Fig. 3. A sequence of photographs illustrating the entrainment of dye and the formation of brine channels in our test cell. The horizontal wires are an apart. The dye was injected 170 min after the ice grouth began. (30) 420 min; (30) 364 min.



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Adaptive mesh refinement

Adaptive mesh refinement focuses computational effort where needed to resolve the problem while using lower resolution in less-dynamic regions.

Initial results are promising but more work needed to fine-tune mesh-refinement criteria



Shaded regions show refinement. Clear is base resolution, green is 2x finer, and purple is 2x even finer (4x base resolution). Over-aggressive refinement in early phases leads to refining the entire domain and slows computation.



Conclusions

- We have simulated brine drainage via 3-dimensional convection in porous mushy sea ice
- Shallow region near ice-ocean interface is desalinated by an array of many small brine channels
- Full-depth "mega-channels" allow drainage from saline layer near top of ice
- Adaptive mesh refinement capability is implemented and is being fine-tuned.

Reference: J.R.G. Parkinson, D.F. Martin, A.J. Wells, R.F. Katz, "Modelling binary alloy solidification with adaptive mesh refinement", *Journal of Computational Physics: X*, Volume 5, 2020, <u>https://www.sciencedirect.com/science/article/pii/S2590055219300599</u>



Appendix: Governing Equations

Continuous equations for conservation of momentum (1), mass (2), salt (3) and energy (4) are found by averaging over lengths greater than the pore scale of sea ice [4, 5].

1. $\vec{U} = -\frac{k_W(\chi)}{n} (\nabla p - \rho_l \vec{g})$ 2. $\nabla \cdot \vec{U} = 0$ 3. $\frac{\partial S}{\partial t} = \vec{U} \cdot \nabla S_l = \nabla \cdot \chi D_l \nabla S_l$ 4. $\frac{\partial H}{\partial t} + \rho_0 c_{p,l} \vec{U} \cdot \nabla T = \nabla \cdot [k_l \chi + (1 - \chi) k_s] \nabla T$ \hat{U} (Darcy Velocity), χ (porosity), p (pressure), T (temperature), S_l (liquid salinity), $S = \chi S_l$ (bulk salinity), η (viscosity), D_l (salt diffusivity), α , β (thermal, haline expansion), $c_{p,l}$, $c_{p,s}$ (liquid/solid specific heat), k_l, k_s (liquid, solid heat conductivity), K_0 (reference permeability) $H = \rho_0 \{ L\chi + [\chi c_{p,l} + (1 - \chi)c_{p,s}]T \} (enthalpy)$ $\rho_1 = \rho_0 [1 - \alpha T + \beta S_1 \text{ (liquid density)}]$ $K(\chi)^{-1} = (\frac{d^2}{12})^{-1} + [K_0\chi^3/(1-\chi)^2]^{-1}$ (permeability)



Appendix: Computational Approach

Solve (1)-(4) using Chombo finite volume toolkit:

- Momentum and mass: projection method [3].
- Energy and solute:
 - Advective terms: explicit, 2nd order unsplit Godunov method.
 - Nonlinear diffusive terms: semi implicit, geometric multigrid.
 - Timestepping: 2nd order Runge-Kutta method. *Twizell, Gumel, and Arigu (1996).*

Reference:

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