



Modeling of Basin Scale Hydro-meteorological association by Hybrid Wavelet-ARX approach

Mayank Suman & Rajib Maity

Indian Institute of Technology Kharagpur

May 8, 2020



Introduction

- Global Circulation Model (GCM) prediction for climatic/meteorological variable is better than hydrological variable.
- Inter-relation (hydro-meteorological association) is challenging to model due to sptio-temporal variability; however if modeled, it will help in ensuring future water security with changing climate.
- The hydro-meteorological association should be continuously evolving with time.
- Meteorological forcing might not necessarily be affecting all frequency ranges of the hydrological variable.

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Literature	Conclusion \ Purpose					
Ahmad and Si- monovic, (2005)	Analysis of hydrograph					
Lau and Kim, (2012)	Teleconnection of hydrometeo- rological extremes (2010 Pak- istan flood and Russian heat wave)					
lshak <i>et al.</i> , (2013)	Decreasing error in wind speed calculation					
Shih <i>et al.</i> , (2014)	Development of early flood warning system					
Durocher <i>et al.,</i> (2016)	Predicting mean annual streamflow in Quebec river us- ing North Atlantic Oscillation Index					
Adnan <i>et al.</i> , (2019)	Streamflow prediction in mountainous basin					

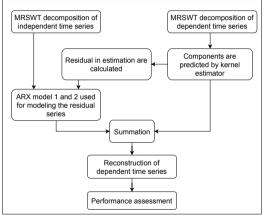
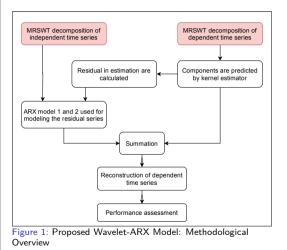
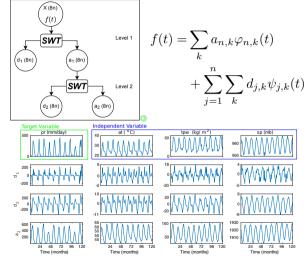


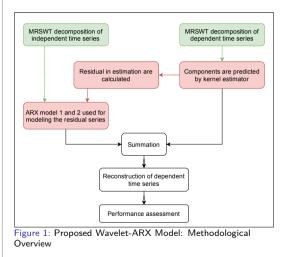
Figure 1: Proposed Wavelet-Auto-Regressive model with eXogenous inputs (ARX) Model: Methodological Overview

Hypothesis

- Due to continuous evolution, the hydro-meteorological association should be more pronounced at constituent wavelet level.
- Wavelet component may have two parts:
 - Memory: modeled using auto-regressive model.
 - Exogenous part: Driven by affecting/exogenous variables.
- After separating memory part, the exogenous part can be modeled as function of exogenous factors.





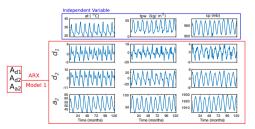


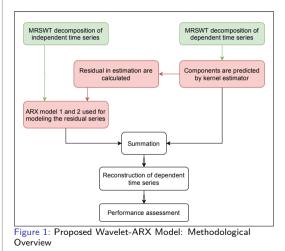
Kernel Estimator for dependent series components:0

$$\begin{split} \tilde{a}_n(m+1) &= \frac{\sum_{t=1}^m a_n(t) K\left(\frac{t-(m+1)}{h}\right)}{\sum_{t=1}^m K\left(\frac{t-(m+1)}{h}\right)} \\ \tilde{d}_j(m+1) &= \frac{\sum_{t=1}^m d_j(t) K\left(\frac{t-(m+1)}{h}\right)}{\sum_{t=1}^m K\left(\frac{t-(m+1)}{h}\right)} \end{split}$$

ARX models for residuals: $A_{c,y}(t) = \sum_{k=1}^{p} S_k A_{c,y}(t-p) + \sum_{l=1}^{c} \sum_{j=0}^{q} T_{l,j} C_{x,l}(t-j) + E$

The models differ in selection of exogenous components:



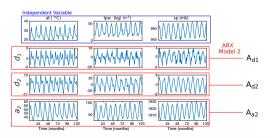


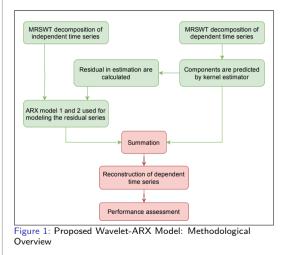
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The models differ in selection of exogenous components:





$$\hat{f}(t) = \sum_{k} \hat{a}_{n,k} \varphi_{N,k}(t) + \sum_{j=1}^{n} \sum_{k} \hat{d}_{j,k} \psi_{j,k}(t)$$

Coefficient of Determination (R^2) , Refined Index of Agreement (D_r) , Mean Absolute Error (MAE) and unbiased Root Mean Square Error (uRMSE) are used for assessing model performance.

$$\begin{split} R^2 &= (r)^2; \qquad r = \frac{Cov(Y,\hat{Y})}{\sigma_y \sigma_{\hat{y}}} \\ Dr &= \begin{cases} 1 - D_{r_frac} & \text{for } D_{r_frac} \leq 1\\ \frac{1}{D_{r_frac}} - 1 & \text{for } D_{r_frac} > 1 \end{cases} \\ D_{r_frac} &= \frac{\sum_{i=1}^n \left|Y_i - \hat{Y}_i\right|}{2\sum_{i=1}^n \left|Y_i - \overline{Y}\right|} \\ MAE &= \frac{\sum \left|Y_i - \hat{Y}_i\right|}{n} \\ uRMSE &= \sqrt{\frac{\sum \left((Y - \overline{Y}) - (\hat{Y} - \overline{\hat{Y}})\right)^2}{n}} \end{split}$$

Relative Importance Analysis

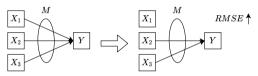
Dominance Analysis

Measure the average association contributed by individual input.

Birnbaum Importance Measure

Measures the probability of a input being critical for proper functioning of the model.

Assumption:



Hence,

$$DARIM_{i} = \frac{\sum_{j=1}^{2^{N-1}-1} \left(R^{2}_{(M(\{PO_{r},I_{i}\}))} - R^{2}_{(M(PO_{r}))} \right)}{2^{N-1}-1}$$

$$\begin{array}{c|c} RMSE(M(X_1)) & RMSE(M(X_1, X_2)) \\ RMSE(M(X_2)) & RMSE(M(X_2, X_3)) \\ \hline \\ RMSE(M(X_3)) & RMSE(M(X_1, X_3)) \\ \hline \\ u_1 = 1 - \frac{3}{4} = 0.25 & u_2 = 0.25 & u_3 = 0.5 \\ \hline \\ BIM_2 = u_1 u_3 = 0.125 & \\ BIM_i = u_1 u_2 u_3 \cdots u_{i-1} u_{i+1} \cdots u_N \end{array}$$

 $\mathbf{DMCE}(\mathbf{M}(\mathbf{Y})) = \mathbf{DMCE}(\mathbf{M}(\mathbf{Y} | \mathbf{Y})) = \mathbf{DMCE}(\mathbf{M}(\mathbf{Y} | \mathbf{Y} | \mathbf{Y}))$

Prediction of total monthly precipitation in UMB

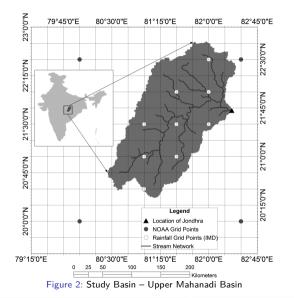
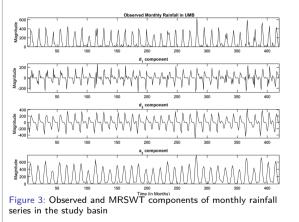


Table 1: Details of data utilized for predicting monthly rainfall over UMB

Independent Time Series	Source	Temporal Res.	R^2 with rainfall		
Avg. Surface Air Temp.	NOAA*	Monthly	0.0015		
Precipitable Water Content	NOAA	Monthly	0.7682		
Surface Pressure	NOAA	Monthly	0.4576		
Avg. Air Temp. (925 mb)	NOAA	Monthly	0.0102		
Avg. Air Temp. (700 mb)	NOAA	Monthly	0.2739		
Avg. Air Temp. (500 mb)	NOAA	Monthly	0.5996		
Avg. Air Temp. (200 mb)	NOAA	Monthly	0.5799		
Avg. Sp. Humidity (925 mb)	NOAA	Monthly	0.7265		
Avg. Sp. Humidity (850 mb)	NOAA	Monthly	0.7444		
Avg. Geop. Height (925 mb)	NOAA	Monthly	0.5039		
Avg. Geop. Height (500 mb)	NOAA	Monthly	0.0063		
Avg. Geop. Height (200 mb)	NOAA	Monthly	0.5120		
Avg. U Wind (925 mb)	NOAA	Monthly	0.4277		
Avg. U Wind (200 mb)	NOAA	Monthly	0.6403		
Avg. V Wind (925 mb)	NOAA	Monthly	0.0003		
Avg. V Wind (200 mb)	NOAA	Monthly	0.2426		

* (Kalnay et al., 1996)

Target Series: Basin precipitation recorded by India Meteorological Department (IMD) (0.5° latitude \times 0.5° longitude)(Rajeevan *et al.*, 2008)



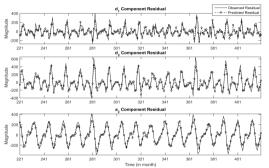


Figure 4: Predicted components residuals from ARX model 2 (p = 3 and q = 2) as compared to the originally calculated residuals during testing period.

Model Performance Statistics	No. of Auto- Regressive input (p)	Lag in exogenous time series input (q)											
		ARX (for Comparison)			Model 1				Model 2				
		0	1	2	3	0	1	2	3	0	1	2	3
R^2	1	0.873	0.864	-	-	0.818	0.783	-	-	0.752	0.821	-	-
	2	0.877	0.864	0.862	-	0.901	0.877	0.862	-	0.883	0.897	0.895	-
	3	0.874	0.860	0.859	0.851	0.911	0.885	0.873	0.763	0.887	0.905	0.907	0.895
D_r	1	0.858	0.856	-	-	0.835	0.806	-	-	0.810	0.833	-	-
	2	0.858	0.853	0.853	-	0.871	0.845	0.831	-	0.857	0.864	0.861	-
	3	0.856	0.853	0.851	0.842	0.873	0.847	0.839	0.782	0.855	0.864	0.864	0.855
MAE	1	57.85	59.43	-	-	61.00	67.58	-	-	70.16	61.26	-	-
	2	56.70	59.20	58.32	-	46.15	50.99	54.10	-	50.17	47.05	47.52	-
	3	56.50	59.15	58.84	59.11	44.37	49.46	51.90	71.15	50.01	45.64	45.28	47.89
uRMSE	1	56.39	57.60	-	-	60.95	68.82	-	-	71.13	60.63	-	-
	2	55.40	57.55	56.67	-	46.13	50.34	53.23	-	51.04	46.70	47.20	-
	3	55.41	57.67	57.17	58.19	44.35	48.87	51.18	69.90	51.00	45.19	44.83	47.40

Table 2: Statistics showing model performance for prediction of monthly rainfall over the study basin during the testing period

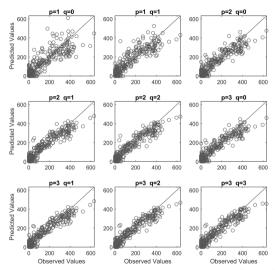


Figure 5: Scatter plot between observed and predicted monthly rainfall over the study basin (in mm/month) for the different combinations of p and q using hybrid Wavelet-ARX model 2.

BIM DARIM Table 3: Relative Importance Measure for independent Variables variables using Wavelet-ARX model 2 (p = 3 and Surface Surface q = 2) for monthly rainfall prediction. Pressure Pressure (Table 2) Air Tem-Geon $BIM (\times 10^{-6})$ $DARIM (\times 10^{-3})$ Independent Time Series Height perature . (925mb) (925mb) Avg. Surface Air Temp. 6.48 4.10 Precipitable Water Content 4.83 3.94 Basin Basin Surface U-wind Surface Pressure 8.13 6.39 Air Tem-Precip-Precip-(925mb) Avg. Air Temp. (925 mb) 7.14 4.76 perature itation itation Avg. Air Temp. (700 mb) 4.77 2.00 Surface Air Tem-Avg. Air Temp. (500 mb) 4.20 -1.38 Air Temperature Avg. Air Temp. (200 mb) 5.03 1.04 . (925 mb) perature Geop. Geop. Avg. Sp. Humidity (925 mb) 5.64 3.88 Height Height Avg. Sp. Humidity (850 mb) 4.12 1.06 (925mb) (500mb) Avg. Geop. Height (925 mb) 6.10 5.47 Avg. Geop. Height (500 mb) 5.64 4.53 \mathbb{R}^2 0.910 0.907 0.907 Avg. Geop. Height (200 mb) 5.47 1.83 Avg. U Wind (925 mb) 6.06 3.29 D_r 0.867 0.869 0.864 Avg. U Wind (200 mb) 5.48 2.47 Avg. V Wind (925 mb) 5.22 0.82 MAE44.25 45.28 44.72 Avg. V Wind (200 mb) 3.62 -1.83 uRMSE44.53 44.01 44.83

Take Home

- The association between meteorological and hydrological variables are more prominent on wavelet component level, resulting in Wavelet-ARX model outperforming ARX model in most of the cases.
- WT-ARX model 2 outperforms model 1 when exogenous input components are considered, hence, the WT components of input/target variables at same level (same frequency bins) are more associated.
- Additionally, the relative importance analysis of meteorological variables helps in identifying the forcings having stronger hydro-climatic association with hydrological variables.
- For instance, the meteorological variable having highest association with total monthly precipitation in UMB are surface pressure, average geo-potential height at 925mb, average air temperature at 925mb, average geo-potential height at 500mb and average surface air temperature.

Further Reading

Suman, M. and Maity, R., 2019. Hybrid Wavelet-ARX approach for modeling association between rainfall and meteorological forcings at river basin scale. Journal of Hydrology, 577, p.123918.



Research papers

Hybrid Wavelet-ARX approach for modeling association between rainfall and meteorological forcings at river basin scale

Church for speciation

Mayank Suman^a, Rajib Maity^{b,+}

^a School of Water Resources, Indian Institute of Technology Kharagpur, Kharagpur 721302, West Bengal, India ^b Department of Civil Engineering, Indian Institute of Technology Kharagpur, Kharagpur 721302, West Bengal, India

ARTICLE INFO

ABSTRACT

This manuscript was handled by A. Bardossy, Editor-in-Chief

Keywards: Hydroclimatic association Rainfall prediction/simulation Climate change Hybrid Wavelet-ARX model Relative importance analysis

Interaction between meteorological and hydrologic processes is challenging to model owing to their high spatiotemporal variability. The understanding of their associations can help to ensure future fresh water security with changing climate. In this study, due to continuously evolving nature of these interactions, the hydrological and meteorological variables are studied on wavelet component level, Multi-Resolution Stationary Wavelet Transformation (MRSWT) is used to transform the independent (climatic variable) and dependent (hydrological variable) time series into their components. The components of the dependent time series are modeled using a kernel-based auto-regressive (AR) model for separating their memory part. The residuals are hypothesized to be the effect of interaction of predictor variables and thus are modeled using the MRSWT components of meteorological variables in an auto-regressive model with exogenous inputs (ARX). Finally, the predicted residuals (effect of climatic variables) are added to the component estimated by kernel-based AB estimator (memory of dependent series components) to obtain the predicted components of the dependent hydrologic variable, which are then inverse-transformed to obtain the predicted dependent hydrologic variable. The developed hybrid Wavelet ARX is found to capture the information about relationship between synthetically generated data better than a simple ARX model. The model is then applied to predict total monthly rainfall over Upper Mahanadi Basin and is found to effectively extract the information from the poorly associated hydro meteorological variables. While the potential of Wavelet-ARX is found to be impressive for hydro meteorological applications, additionally, discarding some climatic inputs on the basis of their relative importance may lead to better prediction by the developed model. The developed model is suitable for extracting climatic forcings and is desirable in a changing climate.

1. Introduction

association as far as rainfall is concerned in the basin. It is hypothesized that the inter-relationship between meteor



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Thank You

Additional Slides

Wavelet Transform

- Need of Transform The usual representation (Amplitude vs Time) is not always the best for analysis.
- Help in separating slow moving and fast moving components of time series.
- Wavelet Transform (WT) : transforms signals into coefficients of scaled and shifted version of wavelet function.
- Due to finite domain, scaling and shifting of wavelet function enable it to catch most of intermittent disturbances of different durations.
- WT represents the signal in terms of frequency and time domain.

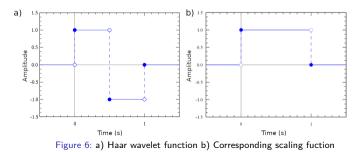
- Unlike Fourier transform, WT provide the temporal information along with frequency information.
- Well suited for the study of multi-scale, nonstationary signals (Daubechies, 1992; Burrus *et al.*, 1998).
- Multi-Resolution Wavelet Transform (MRWT): WT can again be applied on approximate component to produce components at even lower level or frequency ranges (Burrus *et al.*, 1998).
- WT are of many types based different shifting and scaling schemes
 - Continuous Wavelet Transform (CWT)
 - Discrete Wavelet Transform (DWT)
 - Stationary Wavelet Transform (SWT)

Wavelet Function

• Wavelet functions are finite domain square-integrable disturbance of unit energy and zero mean amplitude. Ex.- Haar, Daubechies, Morlet etc.

$$\int \psi(t)dt = 0 \qquad \qquad \iint |\psi(t)|^2 dt = 1$$

- Wavelet functions can separate only one half (higher) of frequency range, hence, they have another function called scaling function to separate the other half frequency range from signal.
- Haar wavelet:



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Continuous Wavelet Transform

If $\psi(t)$ is wavelet function then its shifted and scaled functions $(\psi_{a,b}(t))$ are obtained as:

$$\psi_{a,b}(t) = \frac{1}{\sqrt{a}}\psi\left(\frac{t-b}{a}\right) \tag{1}$$

where $a \in \mathbb{R}^+$, and $b, t \in \mathbb{R}$. a and b are scaling and shifting parameters respectively. Wavelet transform is given by:

$$W_f(a,b) = \frac{1}{\sqrt{C_{\psi}}} \int f(t)\psi_{a,b}^*(t)dt$$
⁽²⁾

where the * denotes complex conjugate, $C_{\psi} = 2\pi \int \left| \hat{\psi}(\omega) \right|^2 / \omega d\omega$. The $\hat{\psi}(\omega) = \int e^{-i\omega t} \psi(t) dt / \sqrt{2\pi}$.

If the basis wavelet or mother wavelet $\psi(t)$ is orthogonal (Daubechies, 1992), then the inverse of wavelet transformation is given by:

$$f(t) = \frac{1}{\sqrt{C_{\psi}}} \iint \frac{W_f(a,b)\psi_{a,b}(t)}{a^2} da \ db \tag{3}$$

Dyadic Discrete Wavelet Transform

DWT is formed when shifting and scaling parameters are taken discrete. The discrete wavelet if sampled over dyadic space, time grid, then they are called dyadic discrete wavelets (Cao *et al.*, 1995).

$$\psi_{a,b}(t) = \frac{1}{\sqrt{2^j}} \psi\left(\frac{t}{2^j} - k\right) \tag{4}$$

where $j, k \in \mathbb{Z}$. The DWT is similar to CWT but it is applied in discrete terms.

As per Nyquist- Shannon sampling theorem, subband coding (downsampling by avoiding every second sample) is carried out to avoid redundancy in component series as suggested by following relation:

$$N_o B_c = N_c B_o \tag{5}$$

where N and B shown length and bandwidth of signal. Subscript o and c mean parent and component series. Hence, $N_c={}^{N_o}\!/_2$

Stationary Wavelet Transform

- SWT is specially designed to avoid the transition-invariance of DWT.
- SWT achieves transition-variance by avoiding down-sampling of components as per Nyquist–Shannon sampling theorem and up-sampling the filter coefficient.
- Despite redundancies in components, SWT reduces the complexity of signal analysis as both input signal and its components have equal length.
- In this study, the SWT is used as preferred WT.

Filter theory basis of DWT and SWT

- DWT or SWT can be applied as pair of high pass and low pass filters.
- High pass filter is obtained from wavelet function and low pass filter is obtained from corresponding scaling function.
- Component from high pass filter is called detailed component and other is called approximate component.
- The approximate and detailed components show trend and local disturbance respectively in the parent signal.

Wavelet components at time t is projection of f(t) over daughter wavelet functions.

$$W_f(a,b) = \frac{1}{\sqrt{C_{\psi}}} \int f(t)\psi_{a,b}^*(t)dt$$
 (6)

Given constant j, the equation is convolution of f(t) with dilated, reflected and normalized mother wavelet $h(t) = \frac{1}{2^j}\psi\left(\frac{-t}{2^j}\right)$. For reference convolution between p(t) and q(t) is given by

$$(p*q)(t) = \int_{-\infty}^{\infty} p(\tau)g(t-\tau)d\tau$$
 (7)

Hence, for Haar wavelet, the dilated reflected and normalized filters are $h = \frac{1}{\sqrt{2}}[1, -1]$ and $g[n] = \frac{1}{\sqrt{2}}[1, 1].$

Multi-Resolution Wavelet Transform

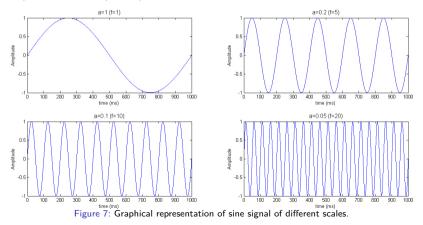
- Multi-Resolution Wavelet Transform (MRWT) provides the detailed and approximate components at even lower levels by using low pass filter component from higher level as input to wavelet transform at each subsequent level.
- MRWT enhances the accuracy of prediction.
- The MRWT is named on the basis of the wavelet transform algorithm being used repeatedly like Multi-Resolution Discrete Wavelet Transform (MRDWT) or Multi-Resolution Stationary Wavelet Transform (MRSWT).
- Irrespective of WT used, after application of MRWT a function can be represented as

$$f(t) = \sum_{k} a_{n,k} \varphi_{n,k}(t) + \sum_{j=1}^{n} \sum_{k} d_{j,k} \psi_{j,k}(t)$$
(8)

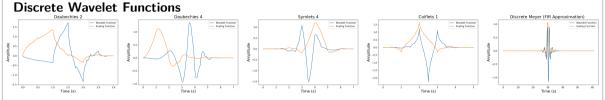
where $a_{n,k}$ is called the coarse or approximate coefficient and $d_{j,k}$ is called the detailed component or wavelet coefficient at level j. The n is the maximum level of decomposition of MRWT, k denotes the shift parameter of wavelet functions.

Scaling parameter

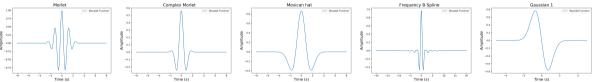
- Scale parameter (a) is inversely related to the wave frequency.
- High scale or low frequency corresponds to non-detailed global view (of the signal), and low scale or high frequency corresponds to detailed view.
- Dilate (a > 1) or compress (a < 1) mother wavelet function.



Common wavelet functions



Continuous Wavelet Functions



Kernel Estimator

- Non-parametric estimators for estimating probability distribution of a given data set.
- The components of hydrological time series are estimated as

$$\tilde{a}_{n}(m+1) = \frac{\sum_{t=1}^{m} a_{n}(t) K\left(\frac{t-(m+1)}{h}\right)}{\sum_{t=1}^{m} K\left(\frac{t-(m+1)}{h}\right)}$$
(9a)
$$\tilde{d}_{j}(m+1) = \frac{\sum_{t=1}^{m} d_{j}(t) K\left(\frac{t-(m+1)}{h}\right)}{\sum_{t=1}^{m} K\left(\frac{t-(m+1)}{h}\right)}$$
(9b)

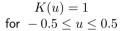
where K represents kernel function, and parameter h is the window of the kernel function. Naïve kernel function (Maity and Nagesh Kumar, 2008; Bosq, 2012) is used in this study.

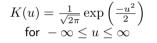
Different Kernel Functions

Naïve

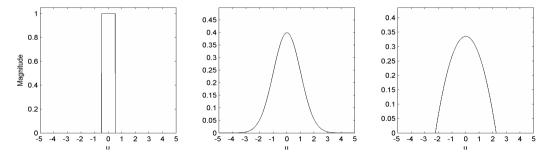
Normal

Epanechnikov





$$\begin{split} K(u) &= \frac{3}{4\sqrt{5}} \left(1 - \frac{u^2}{5} \right) \\ \text{for } &- \sqrt{5} \leq u \leq \sqrt{5} \end{split}$$



ARX Model

$$A_{c,y}(t) = \sum_{k=1}^{p} S_k A_{c,y}(t-p) + \sum_{l=1}^{c} \sum_{j=0}^{q} T_{l,j} C_{x,l}(t-j) + E$$
(10)

where the number of auto-regressive terms and exogenous inputs are represented by p and q respectively.

 $A_{c,y}(t)$ is t^{th} time step residual in any one component of the dependent variable Y, C_x represents the set of selected individual components of the independent variable set X, and c represent the number of members in set C_x or cardinal number of C_x . $C_{x,l}$ represents l^{th} member of C_x .

S and T represent the set of regression parameters estimated during model calibration period.