

Thermodynamic consistent formulation for the multiphysics of a brittle ductile lithosphere: semi-brittle semi-ductile deformation and damage rheology

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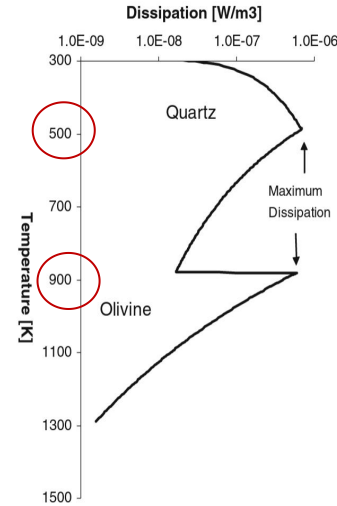
Revisit the concept of crustal rheology

Classical EVP concept

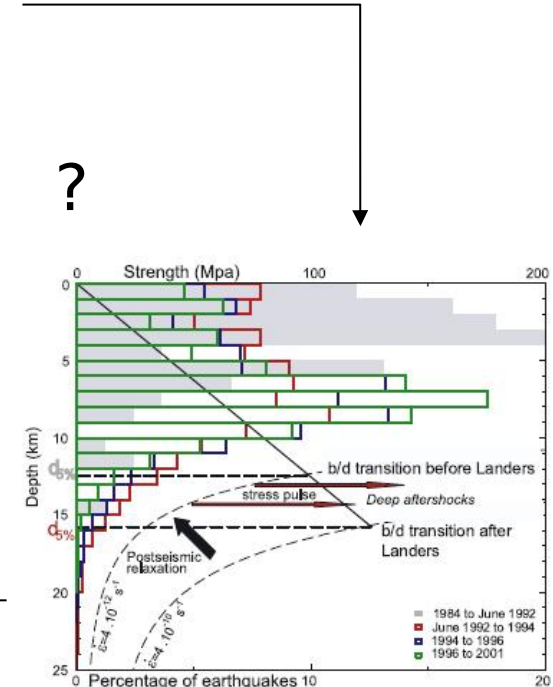
- Peak of stress @ BDT (maximum dissipation/entropy)
- No constraints on stored elastic strain (dissipative fluid behaviour)
- Sharp BDT (fast accommodation of dissipative deformation by thermal activated creep)

Experimental evidence

- Semi brittle behavior @ mid-lower crustal & upper mantle conditions (Mancktelow and Pennacchioni, 2005)
- Dilatant rock behavior @ mid-lower crust T-P conditions (Fischer and Paterson, 1989, Violay et al., 2019)
- Transitional, transient semi-brittle/ductile domain due to the system thermodynamic evolution



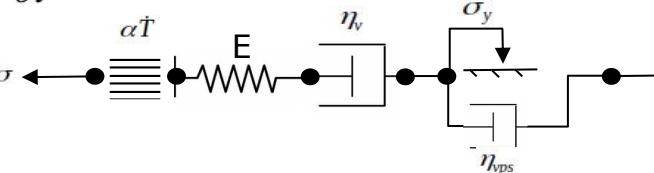
Regenauer-Lieb and Yuen, 2008



Rolandone et al., 2002

$$\nabla \cdot (\boldsymbol{\tau}_{ij} - p\delta_{ij}) + \rho \mathbf{g} = 0$$

$$+ \frac{\partial c_i}{\partial t} + \mathbf{v} \cdot \nabla c_i = 0.$$

$$\frac{\partial T}{\partial t} + \mathbf{v} \cdot \nabla T - \nabla \cdot (\kappa \nabla T) = \dot{H}_r + \dot{H}_s$$


$$\dot{\epsilon} = \dot{\epsilon}^e + \dot{\epsilon}^v + \dot{\epsilon}^{vp} + \dot{\epsilon}^{\theta}$$

PETSc SNES cycle

$$\bar{p}^{e(n+1)} = p^{(n)} - K (\Delta \epsilon_{kk} - \alpha_{\theta} \Delta T) \quad \bar{\tau}_{ij}^{e(n+1)} = \bar{\tau}_{ij}^{(n)} + 2G \Delta \epsilon_{ij}$$

$$\mathcal{F}(p, \sigma_e) \leq 0$$

$$\bar{\tau}_{ij}^{ve(n+1)} = \bar{\tau}_{ij}^{e(n+1)} - 3G \Delta e^v \frac{\bar{\tau}_{ij}^{e(n+1)}}{\bar{\sigma}_e^{e(n+1)}}$$

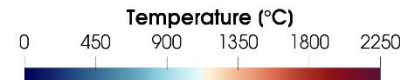
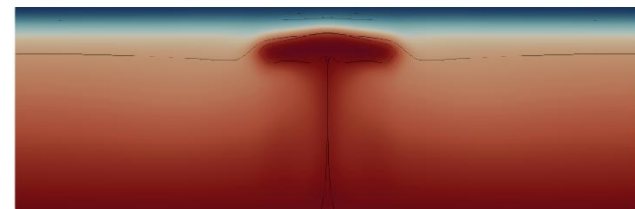
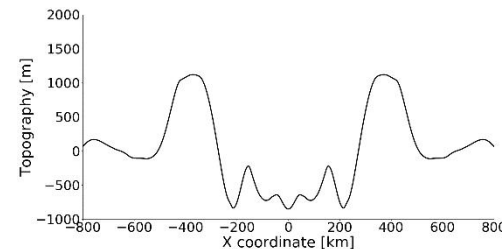
$$\Delta e^v = \frac{\frac{\Delta t}{t^*}}{1 + \frac{\Delta t}{t^*}} \frac{\bar{\sigma}_e^e}{3G}$$

$$p = \bar{p}^e + K \beta \Delta \gamma, \quad \sigma_e = \bar{\sigma}_e^{ve} - 3G \Delta \gamma$$

$$\Delta \gamma^p = \frac{\langle \mathcal{F}(\bar{p}^e, \bar{\sigma}_e^{ve}) \rangle}{3G + \alpha \beta K + H}$$

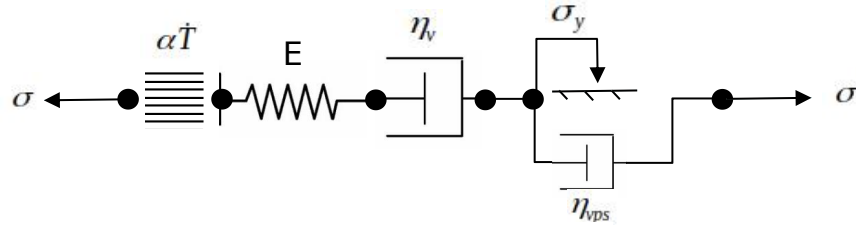
$$p^{(n+1)} = \bar{p}^{e(n+1)} + \beta K \Delta \gamma, \quad \tau_{ij}^{(n+1)} = \bar{\tau}_{ij}^{ve(n+1)} - 3G \Delta \gamma \frac{\bar{\tau}_{ij}^{ve(n+1)}}{\bar{\sigma}_e^{ve}}$$

Coupling of mantle plume with a EVP lithosphere



Time = 10 Myr

- Finite elastic strength leads to compressional and extensional stress accumulation
- Viscous weakening (relaxation) features in a decoupling in the middle crust and multi-wavelength topography (basins)



PETSc SNES cycle

$$\bar{p}^{e(n+1)} = p^{(n)} - K (\Delta \epsilon_{kk} - \alpha_\theta \Delta T) \quad \bar{\tau}_{ij}^{e(n+1)} = \bar{\tau}_{ij}^{(n)} + 2G \Delta \epsilon_{ij}$$

$$F(p, \sigma_e) \leq 0$$

$$\bar{\tau}_{ij}^{ve(n+1)} = \bar{\tau}_{ij}^{e(n+1)} - 3G \Delta e^v \frac{\bar{\tau}_{ij}^{e(n+1)}}{\bar{\sigma}_e^{e(n+1)}}$$

$$\Delta e^v = \frac{\frac{\Delta t}{t^*}}{1 + \frac{\Delta t}{t^*}} \frac{\bar{\sigma}_e^e}{3G}$$

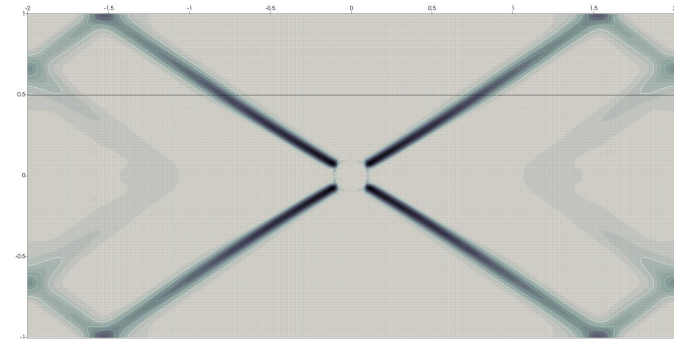
$$p = \bar{p}^e + K \beta \Delta \gamma, \quad \sigma_e = \bar{\sigma}_e^{ve} - 3G \Delta \gamma$$

$$\Delta \gamma^{vp} = \chi^{vp} \Delta \gamma^p$$

$$\chi^{vp} = \frac{\frac{(3G + \alpha \beta K + H) \Delta t}{\eta_p}}{1 + \frac{(3G + \alpha \beta K + H) \Delta t}{\eta_p}}$$

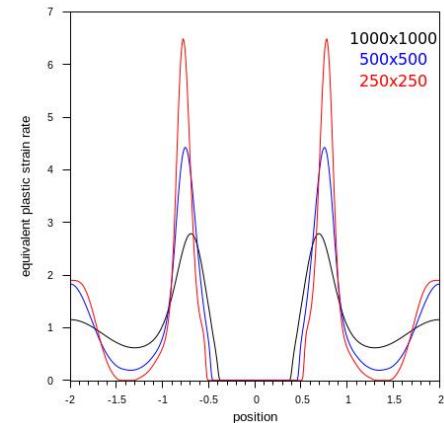
$$p^{(n+1)} = \bar{p}^{e(n+1)} + \beta K \Delta \gamma, \quad \tau_{ij}^{(n+1)} = \bar{\tau}_{ij}^{ve(n+1)} - 3G \Delta \gamma \frac{\bar{\tau}_{ij}^{ve(n+1)}}{\bar{\sigma}_e^{ve}}$$

Elastic-viscous-viscoplastic shear banding



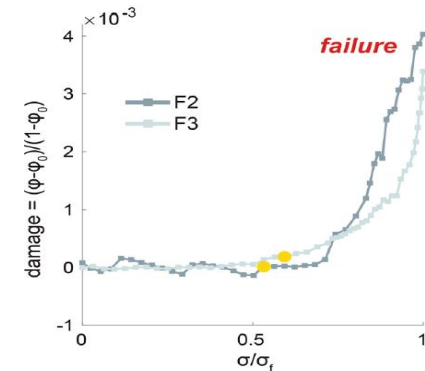
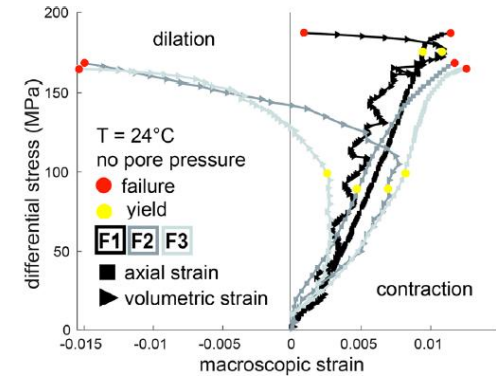
equivalent plastic strain rate (-)
0.0e+00 2 4 6 8 10 12 14 1.6e+01

plastic strain rate within the deformed bands



Nonlinear dynamics for the BDT - some aspects

- Transition from brittle faulting to ductile flow controls by porosity evolution upon (quasi) dynamic loading
- Brittle failure associated with dilatancy, cataclastic flow can be sustained under compactive mode
- Microstructure evolution during loading is essential feedback mechanism for the macroscopic evolution
- Brittle failure of rocks (coalescence of microcracks) sensitive to strain rate and effective confining pressure
- Experimental evidence of transient (un)drained response of rocks even at creep loading
- Transient inelastic modes influence subcritical crack growth, precursory phase to earthquake rupture and/or volcanic edifice collapse, recovery of engineered reservoirs



Renard et al. (2018)

porous damage rheology - theoretical framework

Main governing equations

$$\left[\begin{array}{l} \nabla \cdot (\boldsymbol{\tau}_{ij} - (p' + \alpha_B p_f) \delta_{ij}) + \rho \mathbf{g} = 0 \\ \frac{1}{M_B} \frac{\partial p_f}{\partial t} + \alpha_B \frac{\partial \epsilon_{kk}^e}{\partial t} + \frac{\partial \epsilon_{kk}^{in}}{\partial t} + \nabla \cdot \mathbf{q}_D = 0 \\ \frac{\partial \phi}{\partial t} - (\alpha_B - \phi) \left(\frac{(1 - \alpha_B)}{K} \frac{\partial p_f}{\partial t} + \frac{\partial \epsilon_{kk}^e}{\partial t} \right) - (1 - \phi) \frac{\epsilon_{kk}^{in}}{\partial t} = 0 \\ \frac{\partial T}{\partial t} + \mathbf{v} \cdot \nabla T - \nabla \cdot (\kappa \nabla T) = \dot{H}_r + \dot{H}_s \\ \frac{\partial c_i}{\partial t} + \mathbf{v} \cdot \nabla c_i = 0. \end{array} \right.$$

Thermodynamics of damage rheology

Helmholtz Free Energy (Lyakhovsky et al., 1993)

$$\Psi \left(\epsilon_{ij}^e, \alpha, \zeta, \epsilon_v^{in} \right) = \Psi_{dr} \left(\epsilon_{ij}^e, \alpha \right) + \Psi_{\phi} \left(\epsilon_{ij}^e, \zeta, \epsilon_v^{in} \right)$$

$$\Psi_{dr} \left(\epsilon_{ij}^e, \alpha \right) = \frac{1}{2} K(\alpha) \epsilon_v^{e^2} + \frac{3}{2} G(\alpha) \epsilon_d^{e^2} - \Gamma(\alpha) \epsilon_v^e \|\epsilon_{ij}^e\|$$

$$\Psi_{\phi} \left(\epsilon_{ij}^e, \zeta, \epsilon_v^{in} \right) = \frac{1}{2} M_B [\alpha_B \epsilon_v^e - (\zeta - \epsilon_v^{in})]^2$$

$$\sigma_{ij} = \frac{\partial \Psi}{\partial \epsilon_{ij}^e} = \frac{\partial \Psi_{dr}}{\partial \epsilon_{ij}^e} + \frac{\partial \Psi_{\phi}}{\partial \epsilon_{ij}^e} = \sigma'_{ij} - \alpha_B p_f \quad p_f = \frac{\partial \Psi}{\partial \zeta} = -M_B [\alpha_B \epsilon_v^e - (\zeta - \epsilon_v^{in})]$$

$$\sigma'_{ij} = \sigma_{ij} + \alpha_B p_f = \bar{\sigma}'_{ij} - \alpha \sigma_{ij}^{\alpha}$$

$$\bar{\sigma}'_{ij} = K_0 \epsilon_v^e \delta_{ij} + 2G_0 \mathbf{e}_{ij}^e \quad \sigma_{ij}^{\alpha} = \Gamma \left[\|\epsilon_{ij}^e\| \delta_{ij} + (\xi - 2\xi_0) \epsilon_{ij}^e \right]$$

$$\xi = \frac{\epsilon_v^e}{\|\epsilon^e\|} \quad \text{Elastic strain ratio } [-\sqrt{3}; \sqrt{3}]$$

ξ_0 critical strain ratio (internal friction)

porous damage rheology - flow laws and damage correction

Additive decomposition

$$\dot{\varepsilon} = \dot{\varepsilon}^e + \dot{\varepsilon}^v + \dot{\varepsilon}^{vp} + \dot{\varepsilon}^\theta$$

visco-plasticity + damage

$$\mathcal{F} = \sqrt{(\bar{p}' - \bar{p}'^0)^2 + (\bar{\sigma}_e - \bar{\sigma}_e^0)^2}$$

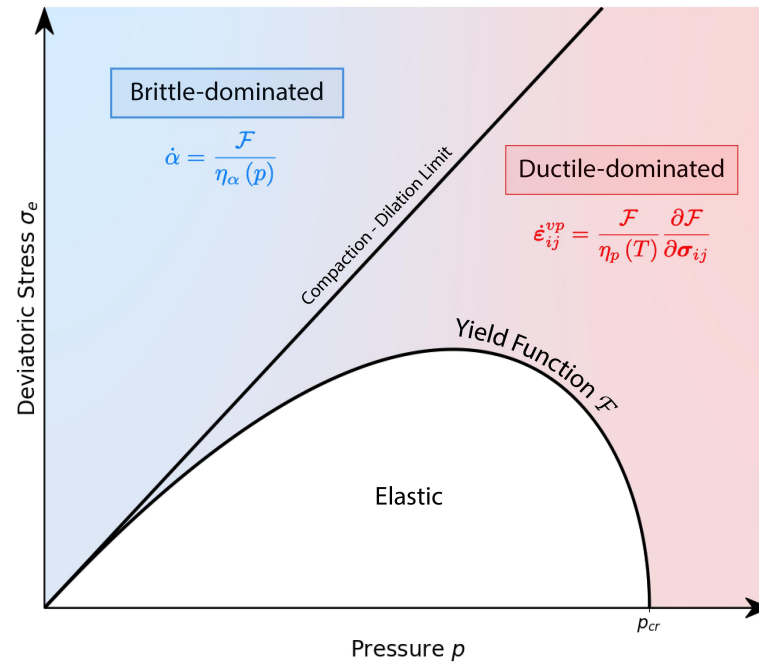
$$\dot{\varepsilon}_{ij}^{vp} = \dot{\gamma}^{vp} \frac{\partial \mathcal{F}}{\partial \bar{\sigma}_{ij}}, \quad \dot{\gamma}^{vp} = \frac{\langle \mathcal{F} \rangle}{\eta_p}$$

$$\dot{\alpha} = \frac{\langle \mathcal{F} \rangle}{\eta_\alpha} = \frac{\eta_p}{\eta_\alpha} \sqrt{(\dot{\varepsilon}_v^{vp})^2 + (\dot{\varepsilon}_d^{vp})^2}$$

$$\eta_p = \eta_p^0 \exp\left(\frac{Q}{RT}\right)$$

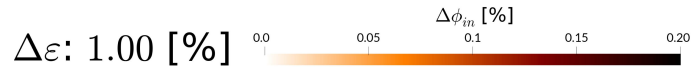
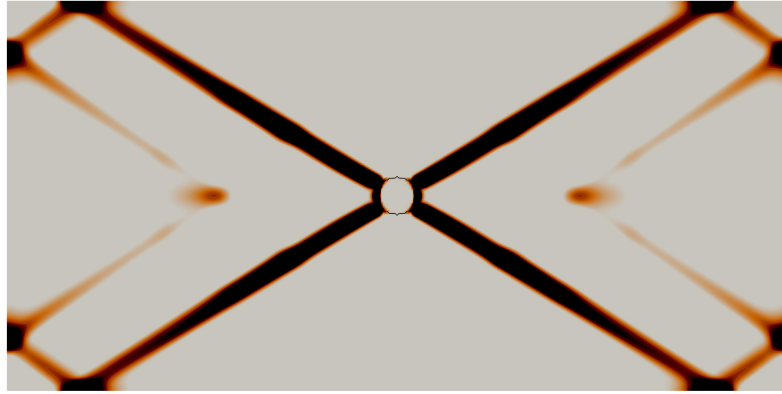
$$\eta_\alpha = \eta_\alpha^0 \exp\left(\frac{p_c}{p^\star}\right)$$

Yield function (undamaged space)



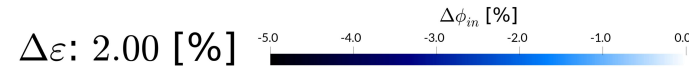
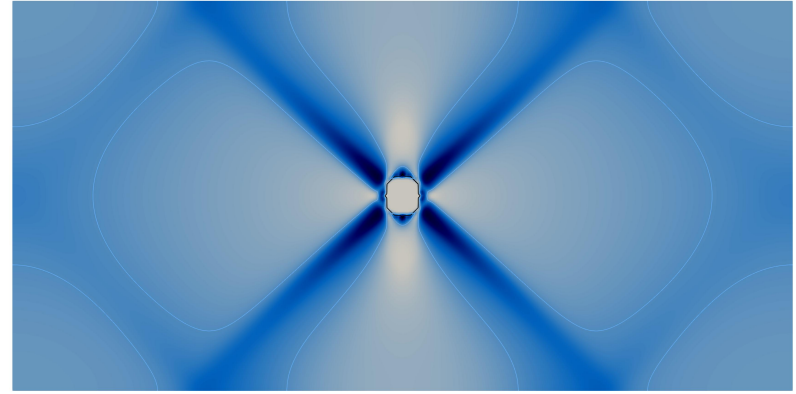
Shear bands evolution under dilation (a) and compactant (b) mode

(a)



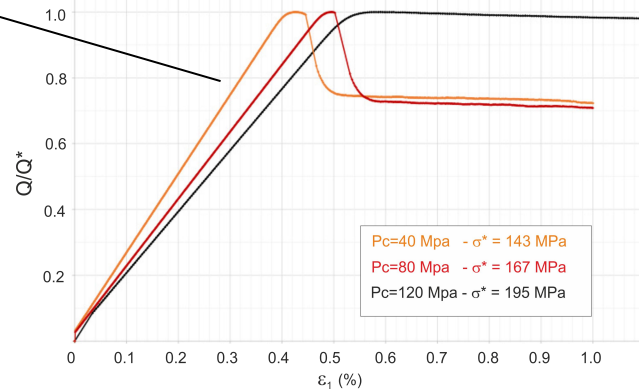
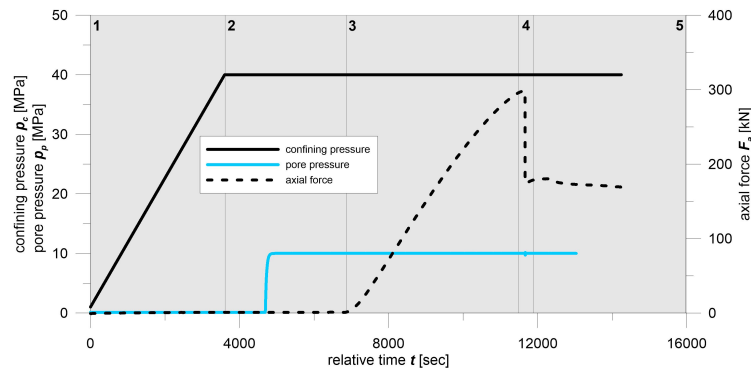
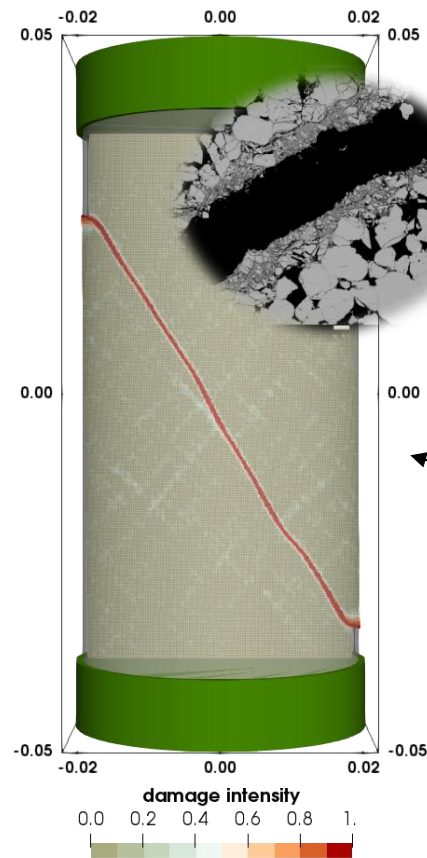
- increase in inelastic porosity
- localized deformation
- bands orientation at approx. 33°

(b)

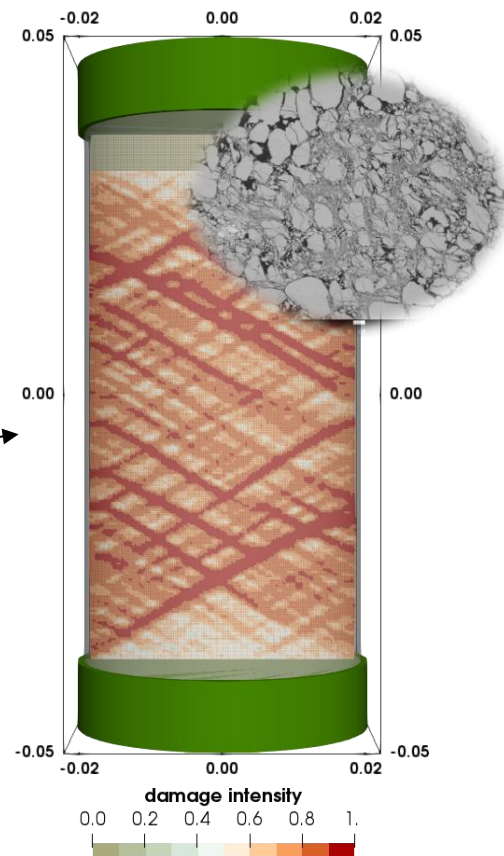


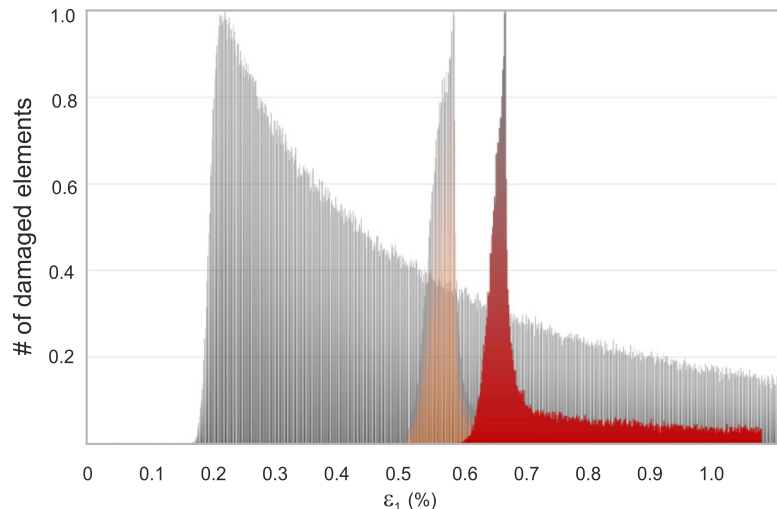
- decrease in inelastic porosity
- diffuse deformation
- bands orientation @ approx. 45°

Triax model - 40 MPa confinement



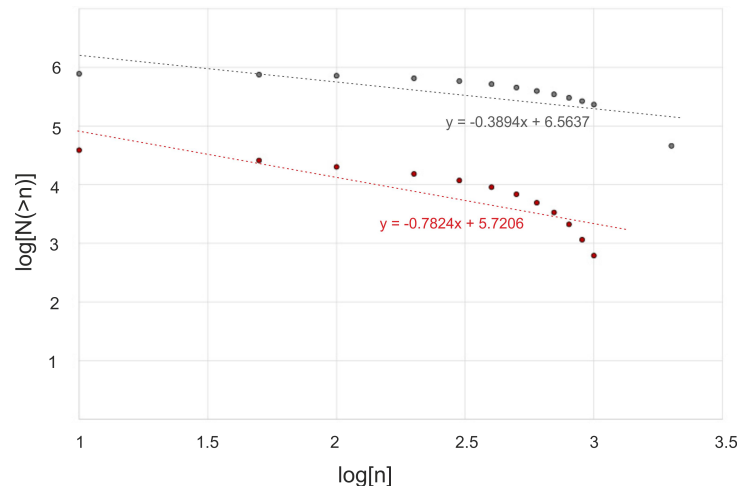
Triax model - 120 MPa confinement





Proxy for AEs

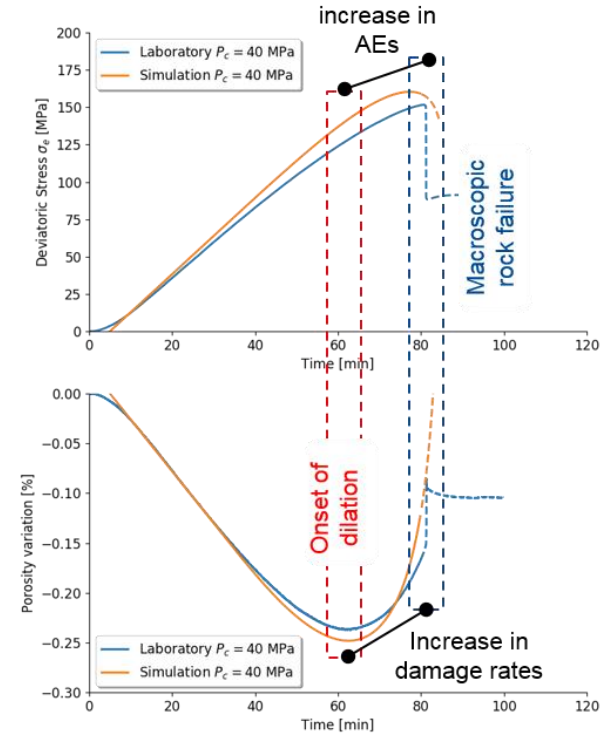
- High confinement - ductile onset at inelastic stage, power law distribution (b-value ~ 0.3)
- Low confinement - brittle power law, but macroscopic peak @ failure (b-value ~ 0.62 -0.8)



$$\log[N(n)] = a - b \log[n]$$

similarities to Omori's law and rate and state
frictional behaviour

- Thermodynamic consistent framework for multiphysics at BDT conditions
 - Frictional damage rheology == function to the state (p , T) and rates of deformation
 - Self-promoting (porosity dependent) damage weakening on friction and capped yield
 - Porosity evolution as precursor to run-away instabilities (though look at current work on brittle creep!)
-
- Across scale studies required linking to existing approaches and community build up:
 - Upscaling from strain rate dependency
 - Rate and state friction (see Lyhakovsky et al., 2005)
 - Long-term geodynamic approaches



Cacace and Jacquey, 2019