

# The effect of dust absorption on Sahel precipitation

Session: **Natural Aerosols in Climate Change** (AS3.7/CL3)

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# Plan

- Why are dust particles influencing precipitation?
- Mineralogy and dust particle size
- Analyzing the variables influenced by the presence of dust

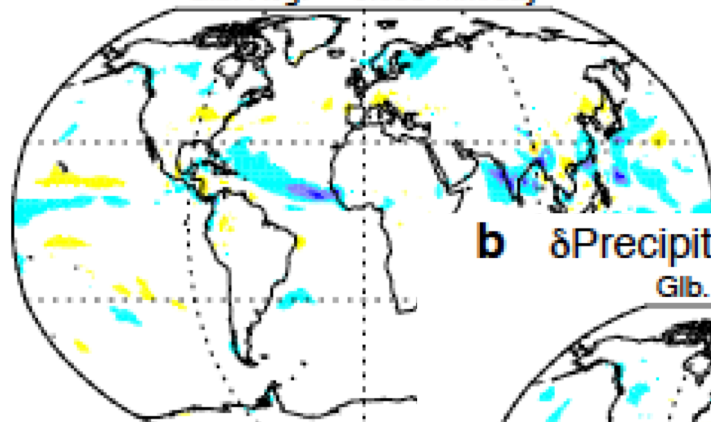
## Existing literature on dust/precipitation connection

- Miller et al. 2004
- Miller et al. 2011
- Yoshioka et al. 2009
- Solomon et al., 2008

## Change in precipitation due to dust (Miller et al., 2014)

**a**  $\delta$ Precipitation (JJAS): Sinyuk

Glb. Avg. = -0.033 mm/day

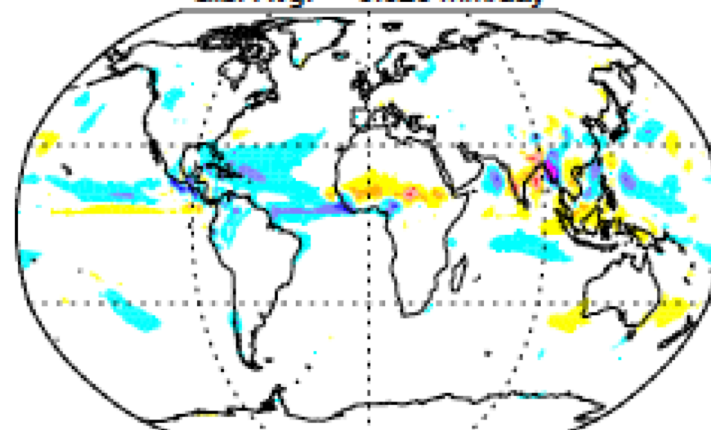


Min = -2.3 ( 91.2E, 23N) M

$w_{0,550nm}=0,95$

**b**  $\delta$ Precipitation (JJAS): Patterson

Glb. Avg. = -0.026 mm/day

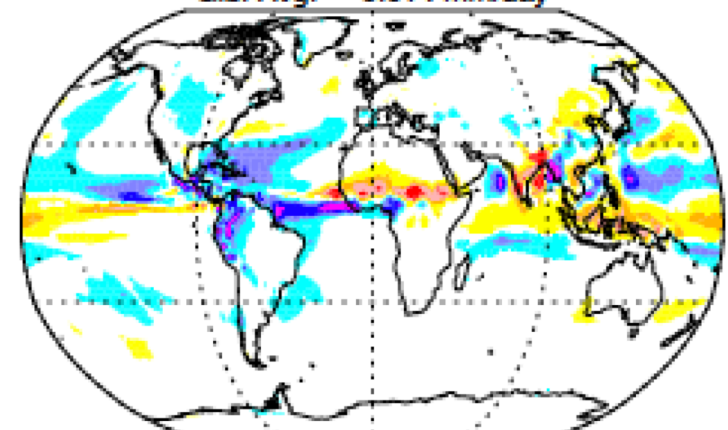


Min = -8.0 ( 93.8E, 21N) Max = 4.4 ( 86.2E, 17N)

$w_{0,550nm}=0,86$

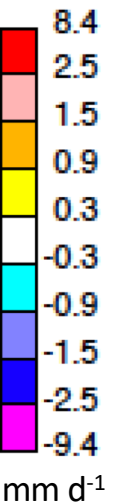
**c**  $\delta$ Precipitation (JJAS):  $0.9 \times \omega_0(\text{Pat})$

Glb. Avg. = -0.014 mm/day



Min = -9.4 ( 93.8E, 21N) Max = 8.4 ( 76.2E, 13N)

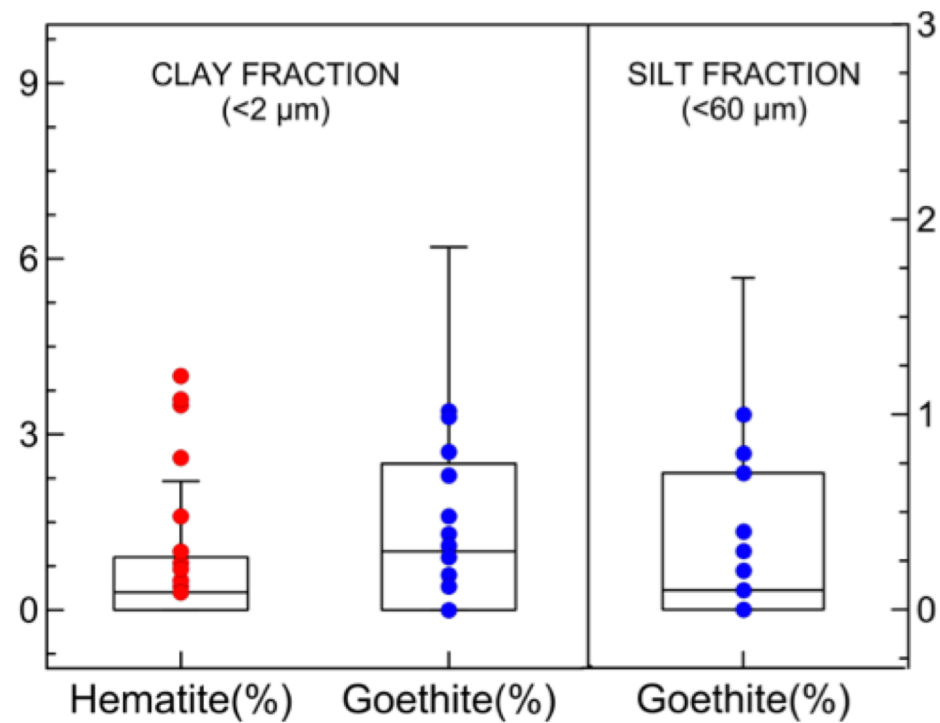
$w_{0,550nm}=0,77$





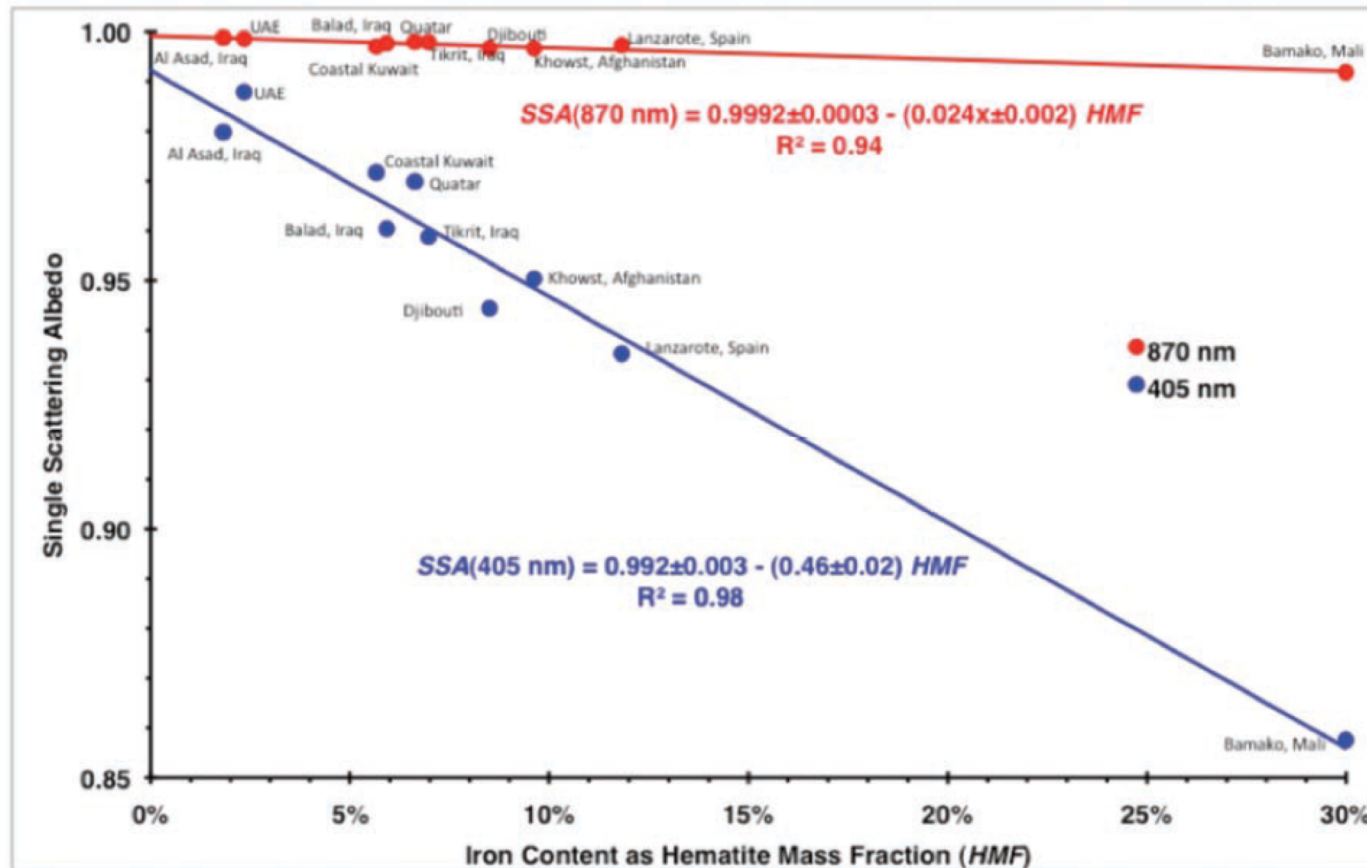
- Over Sahel summer dust absorption reaches  $36 \text{ W.m}^{-2}$  over the Sahel region (see Slide 13)
- Absorption of dust depends on its mineralogy and is mostly driven by the dust iron oxide content
- Large dust particle absorb much more than smaller particles

Goethite and hematite mass fractions (%) measured in the CESAM aerosol chamber for dust samples from major source areas



Di Biagio et al., 2019

# Relationship Between Hematite Content and Single Scattering Albedo



## Computation of the refractive index computed for a mineral assemblage

- We account for six minerals: Kaolinite, Illite, Montmorillonite, Quartz, Calcite et hematite
- An optical model is used in which we vary the iron oxide content to represent respectively: 0.9 1.5 2.7 5.0 10 et 15% of the total volume of the particles.

Step 1: Each mineral is associated with a VOLUME content of respectively: 0.9 1.5 2.7 5.0 10 et 15% using the Maxwell-Bruggeman approximation

Step 2: The combination kaolinite-hematite is associated with illite\_hematite

Step 3: The combinaison kaolinite-illite-hematite is associated to montmorillonite-hematite

Step 4: The combinaison kaol-illi-montmo-hema is associated to quartz\_hematite

Step 5: The combinaison kaol-illi-montmo-hema-quartz is associated to calcite\_hematite

Step 6:: The refractive index of kaol-illi-montmo-hema-quartz-calcite-hematite is obtained

## Absorption increase (imaginary part of ref. index) as a function of hematite content

Hematite 1,5%

$\omega_{0,550\text{nm}}=0,906$

Hematite 5%

$\omega_{0,550\text{nm}}=0,879$

Hematite 10%

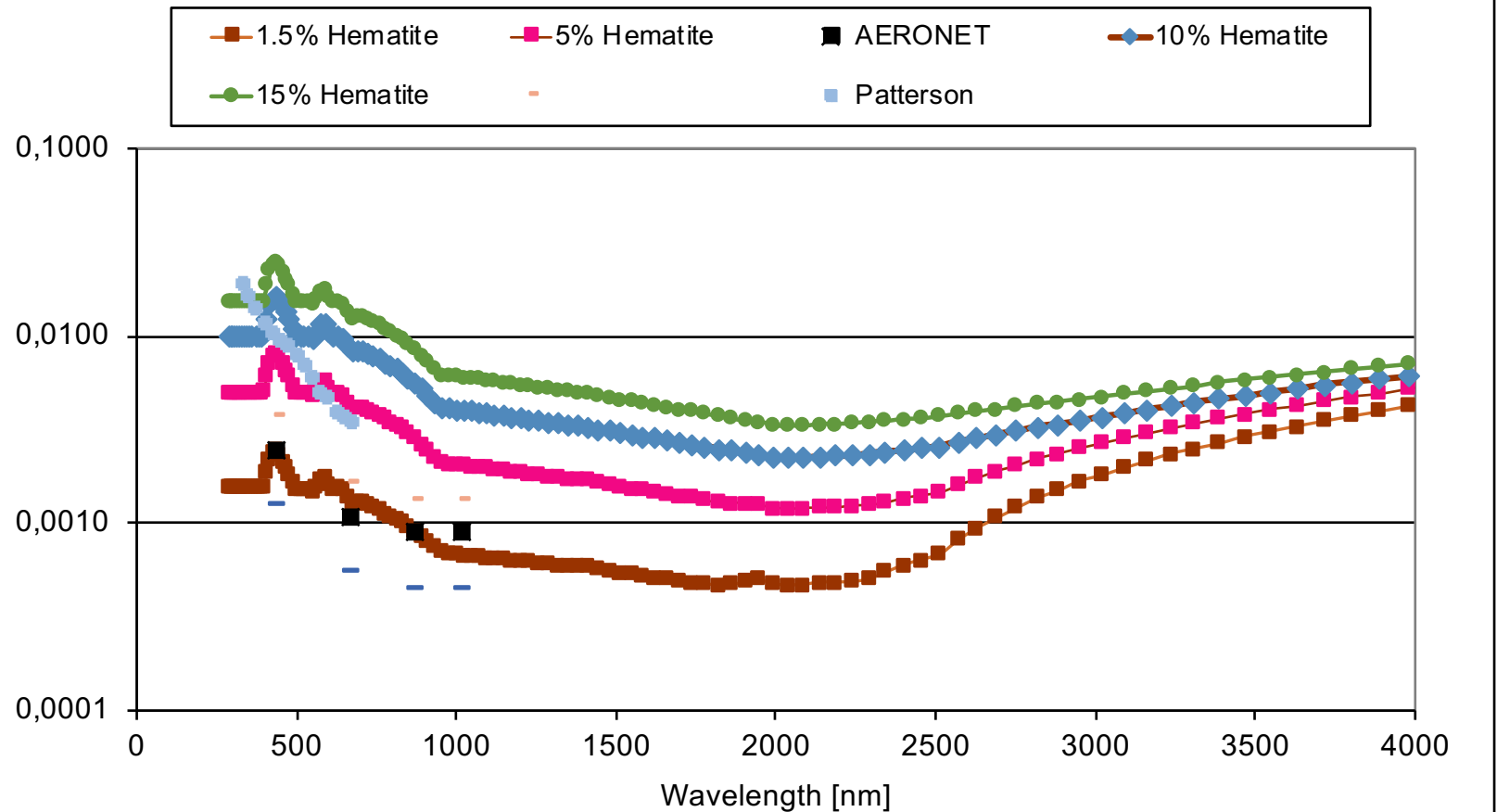
$\omega_{0,550\text{nm}}=0,807$

Hematite 15%

$\omega_{0,550\text{nm}}=0,757$

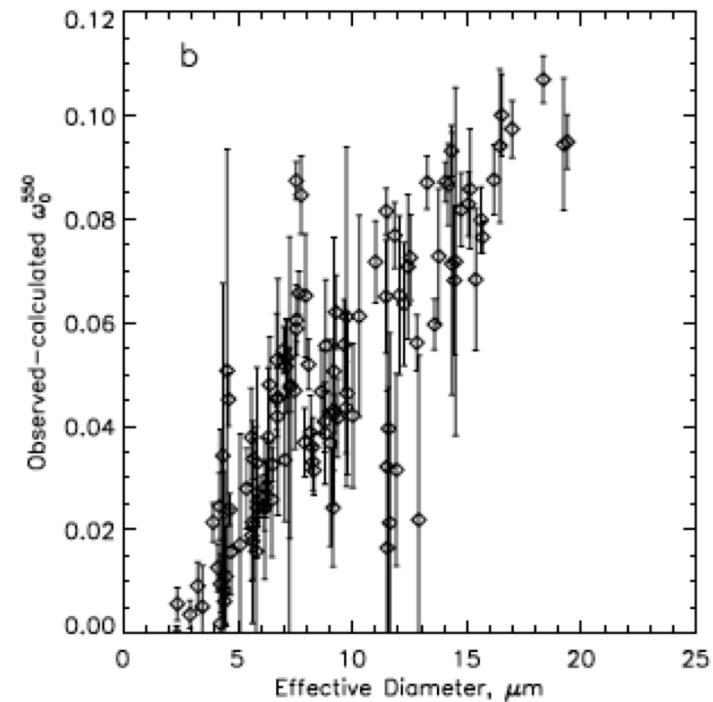
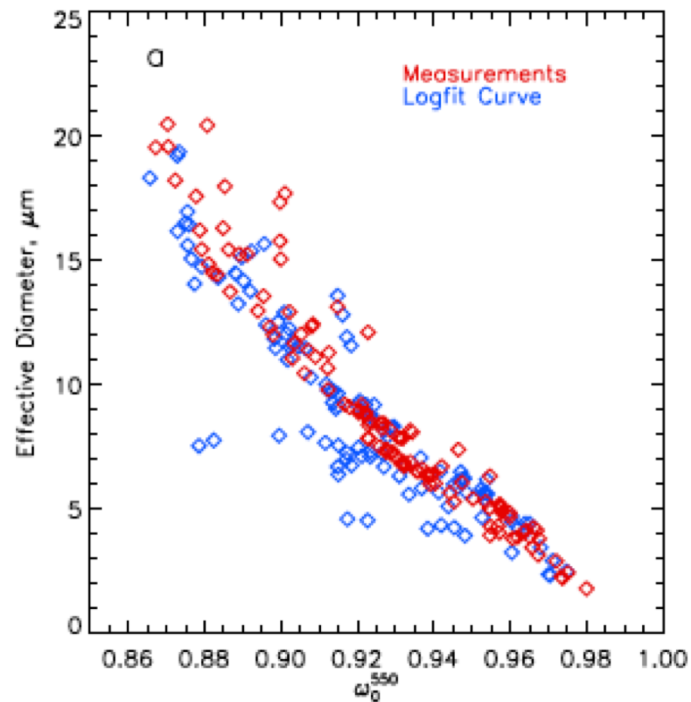
Patterson-Volz

$\omega_{0,550\text{nm}}=0,857$



# Observations of Dust Single Scattering Albedo (SSA) as a function of particle size during the AER-D campaign

*Ryder et al., 2013*

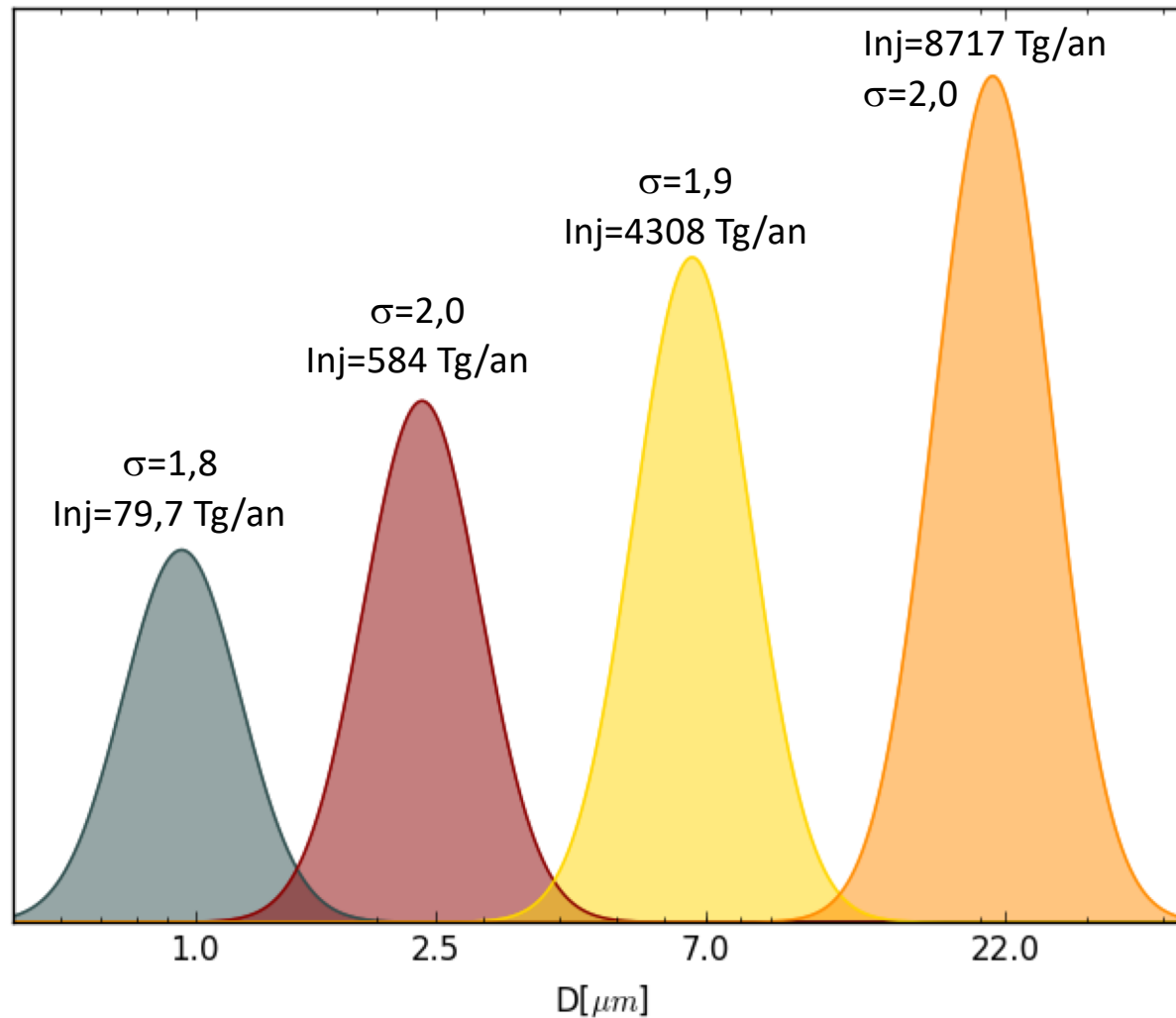


## Description of the simulations

- We modeled the Dust size distribution with either 1 or with 4 modes.
- When 1 mode (MMD=2.5  $\mu\text{m}$ ) is used absorbing dust contains 5% hematite, when 4 modes are used (MMD= 1, 3.5, 7 and 22  $\mu\text{m}$ ) the same absorption is reached with only 3.5% hematite
- We made 2 long simulations:
  - Simulation 1 with fixed SSTs for 18 years (1980 to 1997)
  - Simulation 2 with a fully coupled model for 100 years (1915 to 2014)

The analysis is made by comparing these simulations to a control with NO Dust

## Global Emissions derived from observational constraints for 4 Modes of Dust



After Di Biagio et al, 2020



## JJA Average Dust SW Radiative Effect for 5% hematite

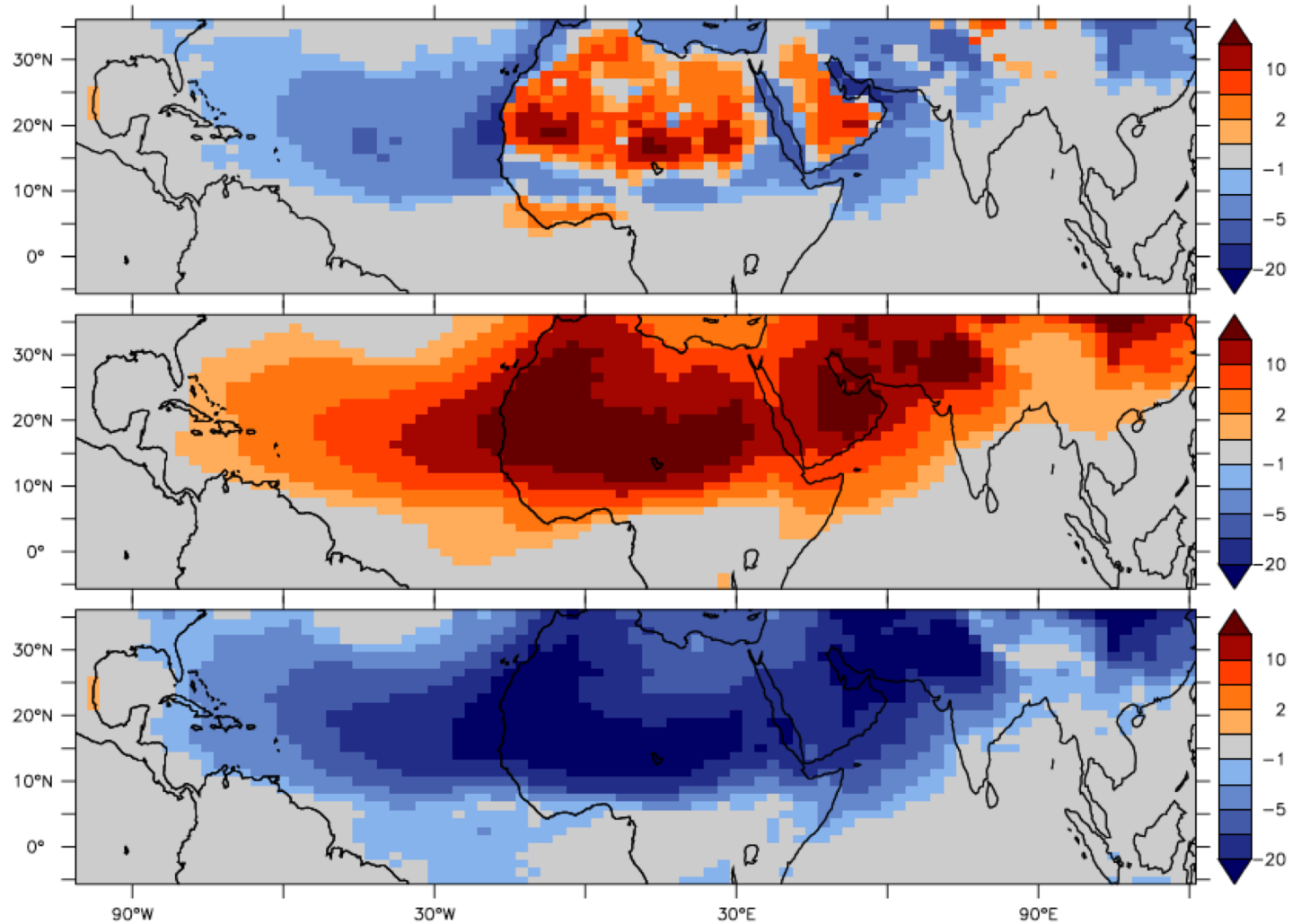
top panel: Top-of-Atmosphere, middle panel: Atmospheric absorption; bottom panel: Surface

Sahel TOA =  $+4.4 \text{ W.m}^{-2}$

SAHEL area: 16W:36E,  
10N:20N

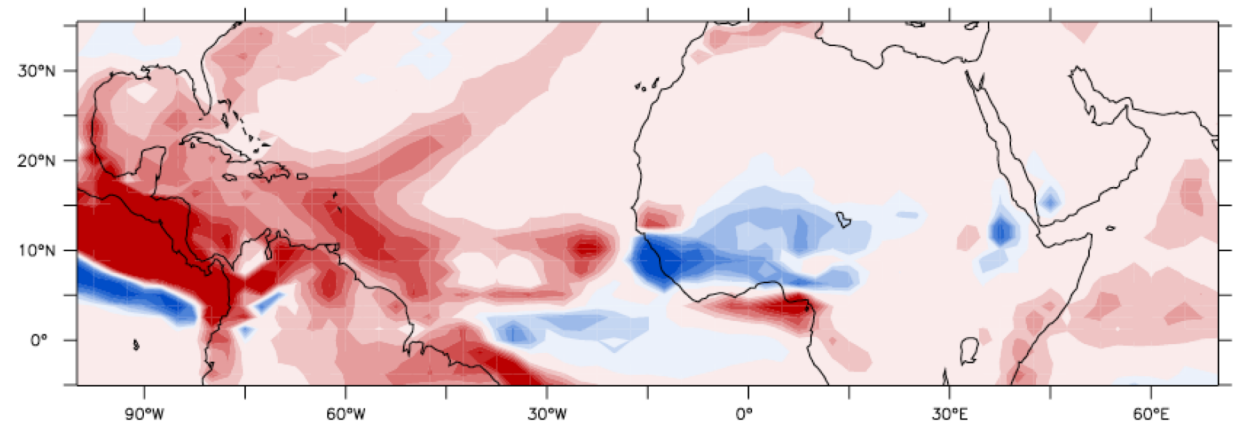
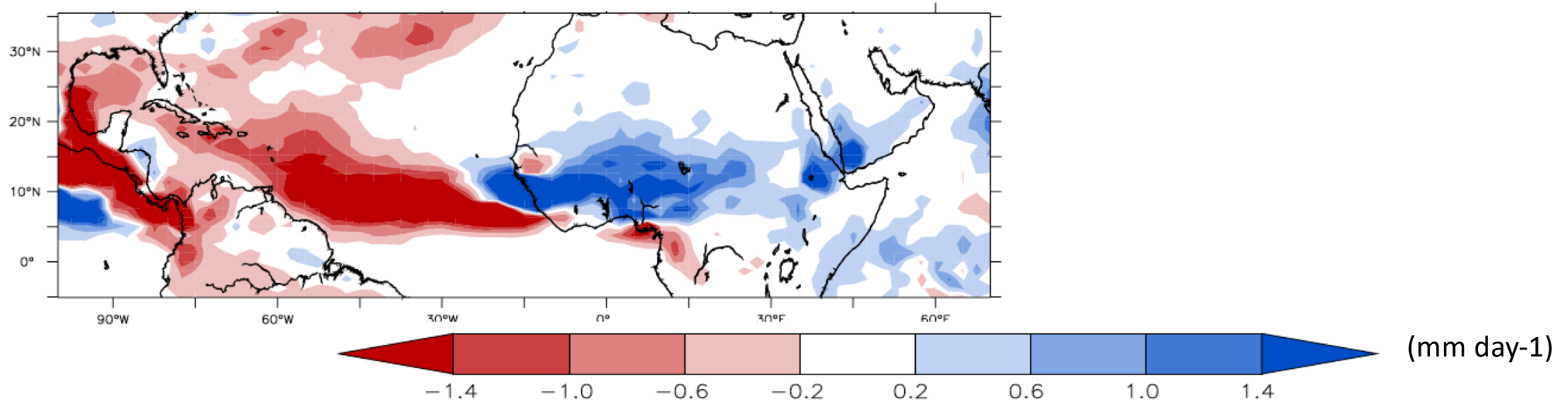
Atm. Absorption =  $+36 \text{ W.m}^{-2}$   
(Sahel area)

Sahel Surface =  $-31.6 \text{ W.m}^{-2}$



# Precipitation change – Dust with 5% hematite versus No Dust

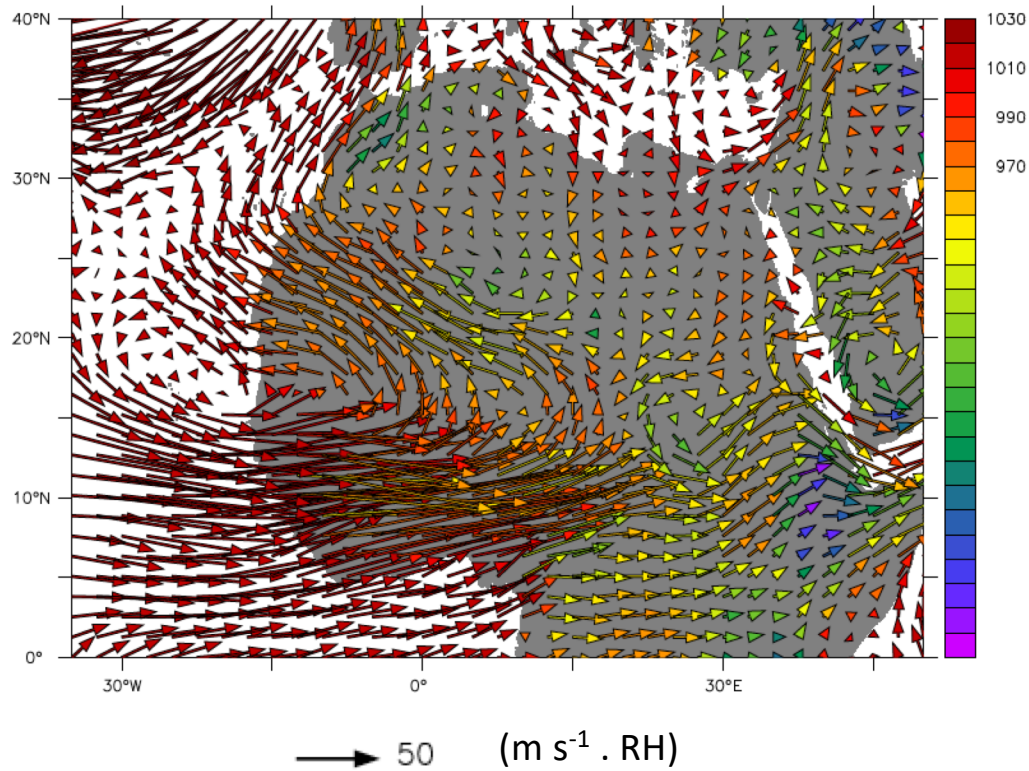
Model with fixed SST 18y mean



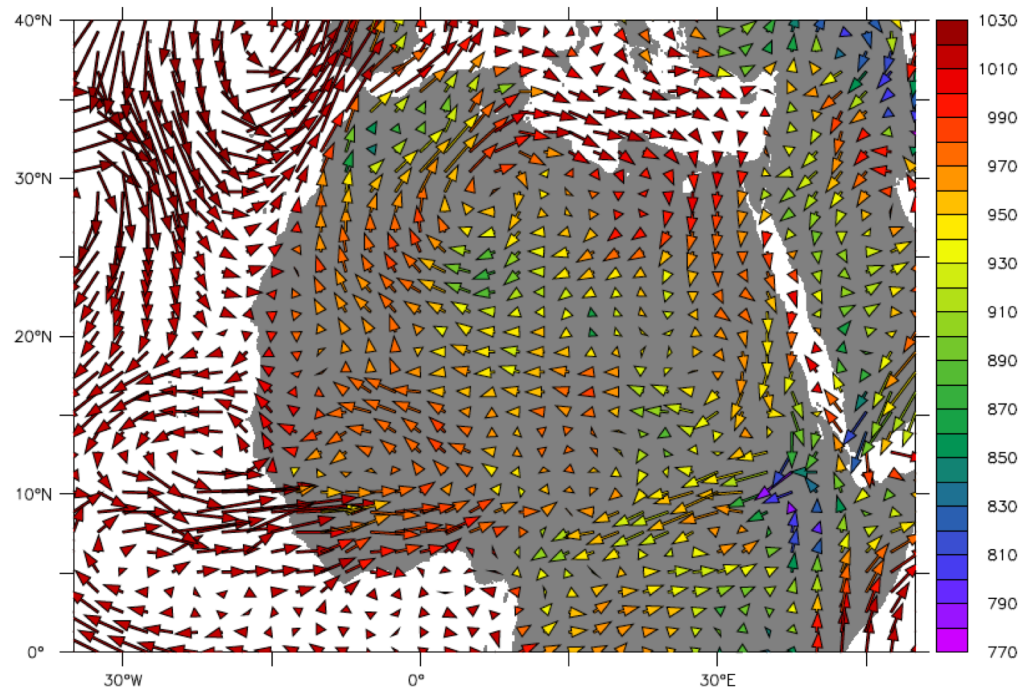
Fully Coupled Model 100y mean

# Change in Humidity transport at the surface – vector ( $uq, vq$ )

Forced Model JAS 1989



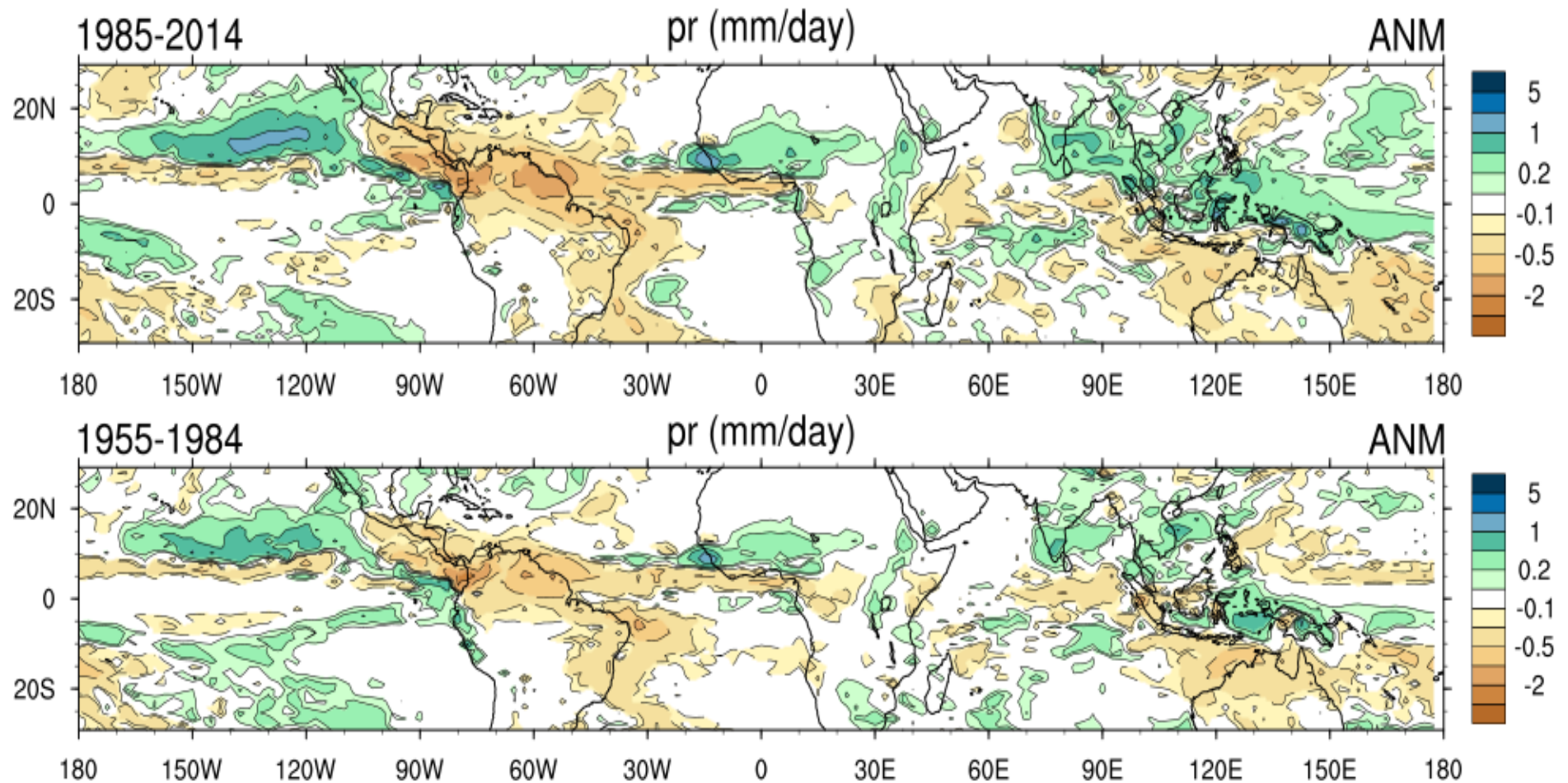
Coupled Model JAS 1924





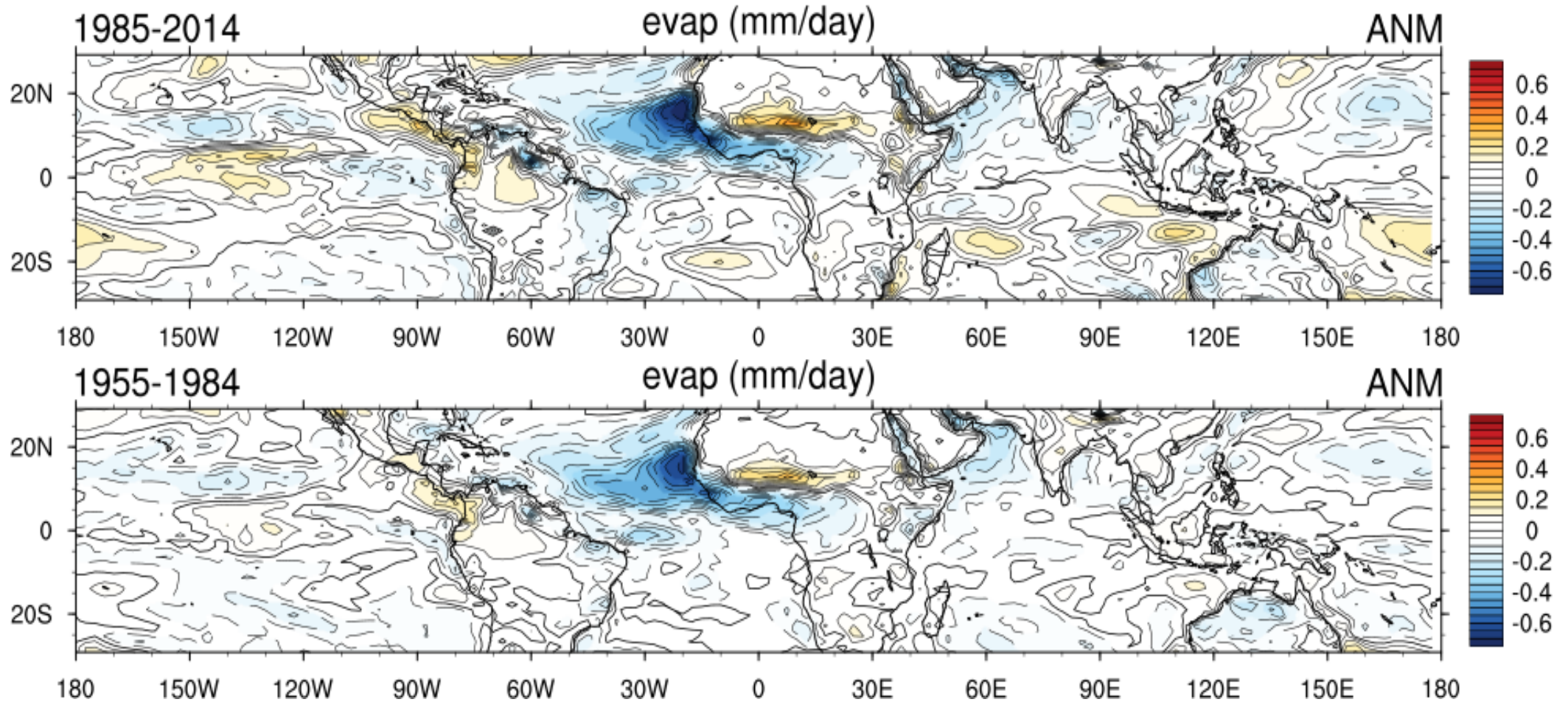
# Precipitation change for the periods 1985-2014 and 1955-1984

**CM61-LR-5perc-pi-AER-01 (vs IPSLCM6 CM61-LR-NoDust-hist-AER-01)**



# Evaporation change for 1985-2014 and 1955-1984 periods

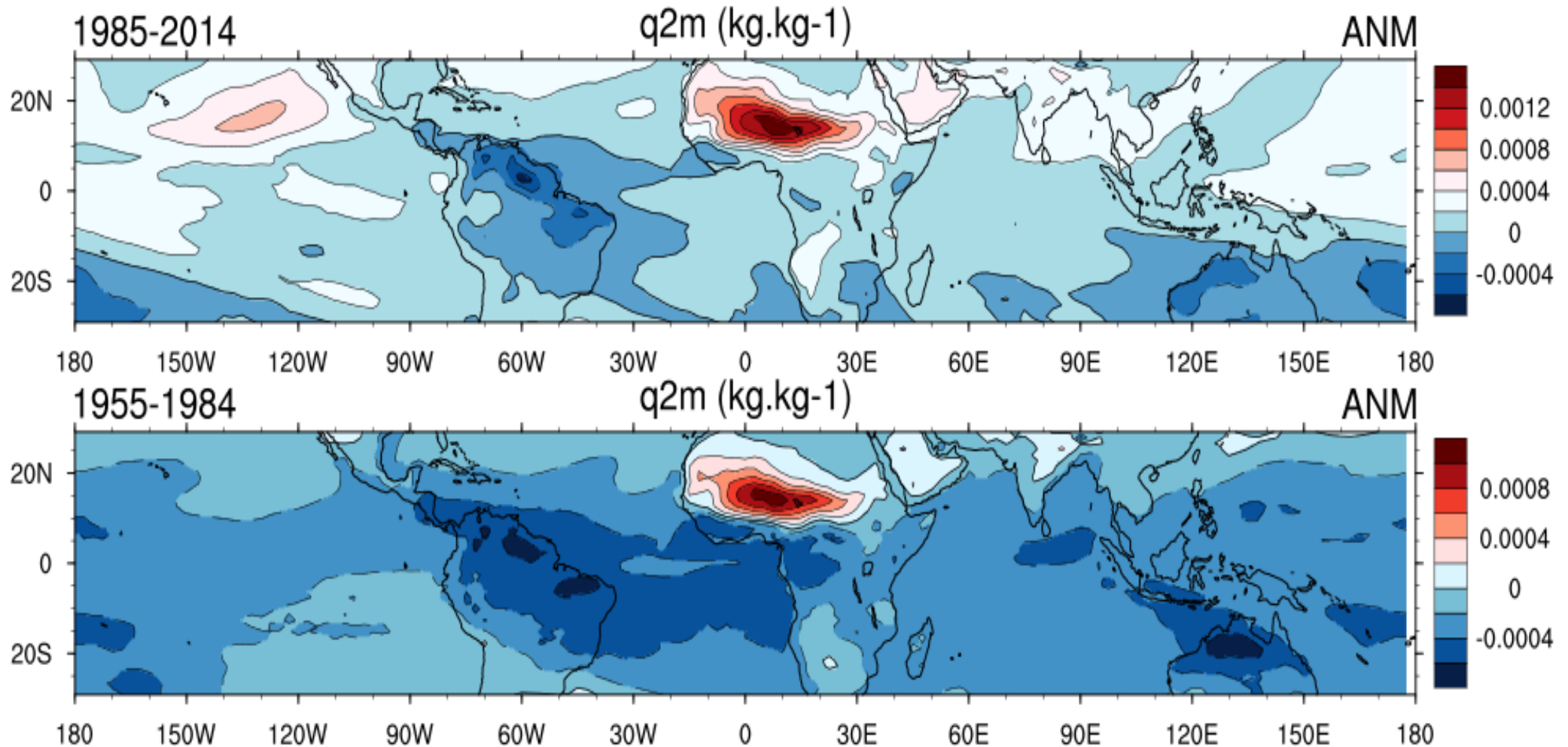
**CM61-LR-5perc-pi-AER-01 (vs IPSLCM6 CM61-LR-NoDust-hist-AER-01)**





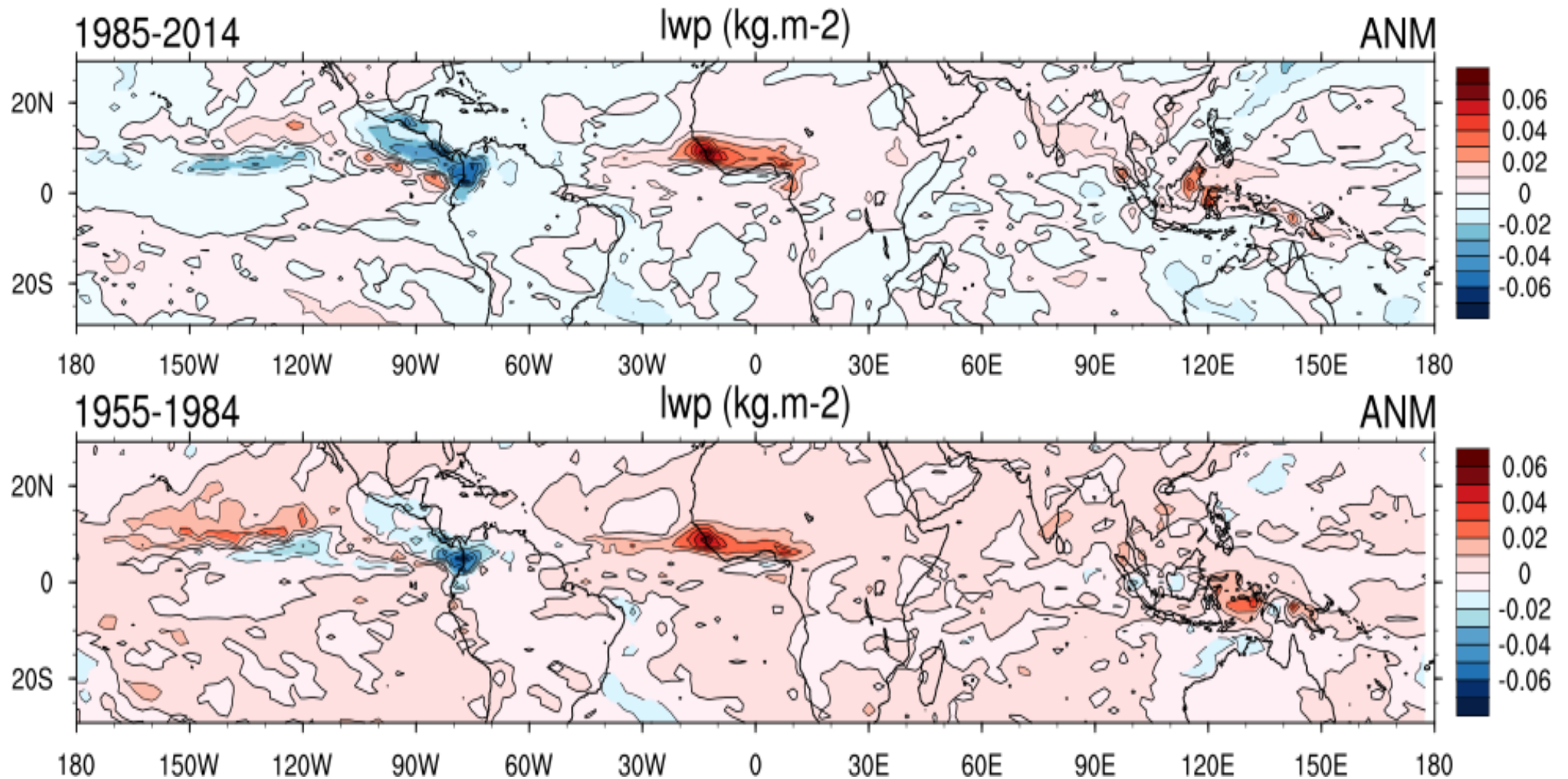
Relative humidity (at 2m) change for 1985-2014 and 1955-1984 periods

**CM61-LR-5perc-pi-AER-01 (vs IPSLCM6 CM61-LR-NoDust-hist-AER-01)**



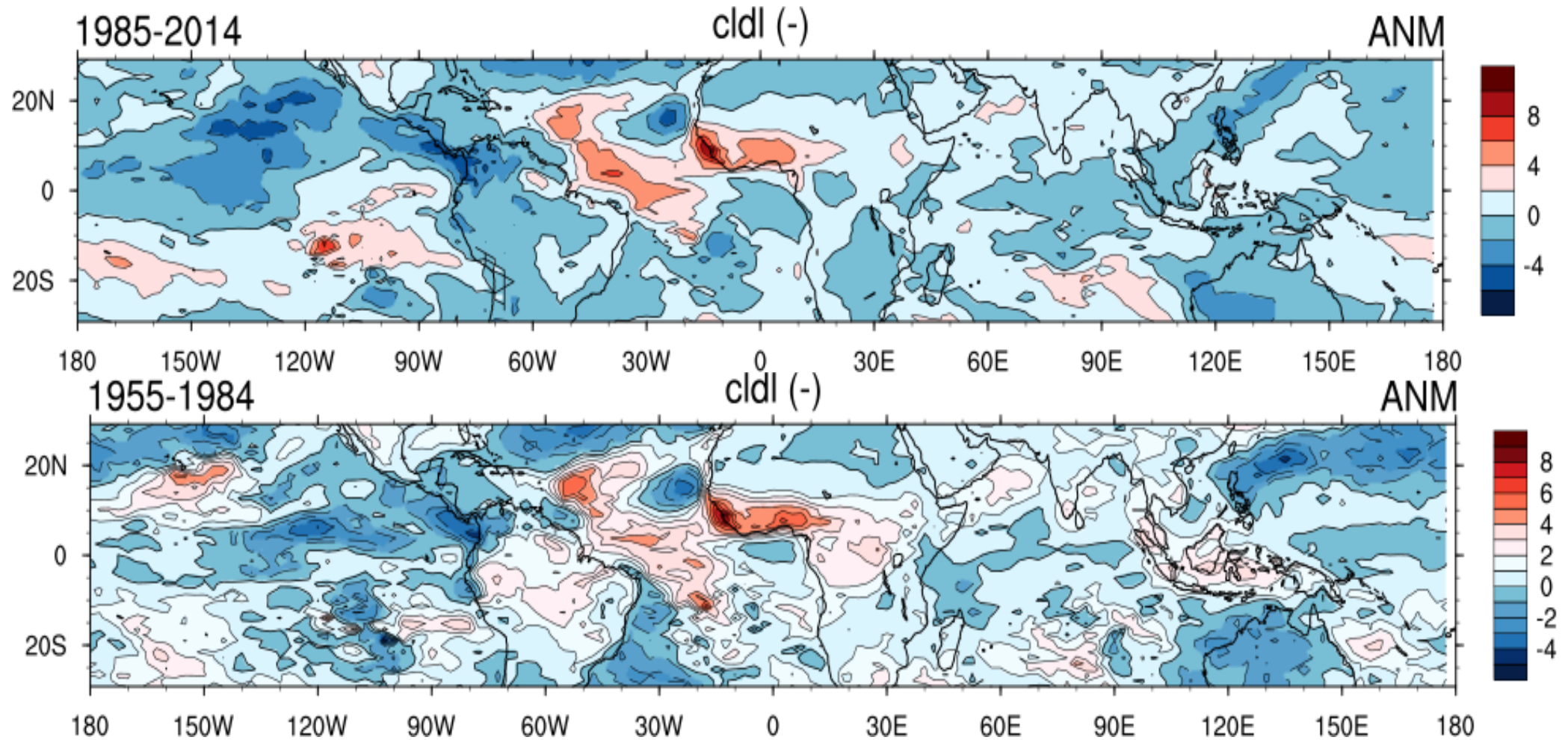
# Liquid Water Path change for 1985-2014 and 1955-1984 periods

**CM61-LR-5perc-pi-AER-01 (vs IPSLCM6 CM61-LR-NoDust-hist-AER-01)**



# Low Clouds change (%) for 1985-2014 and 1955-1984 periods

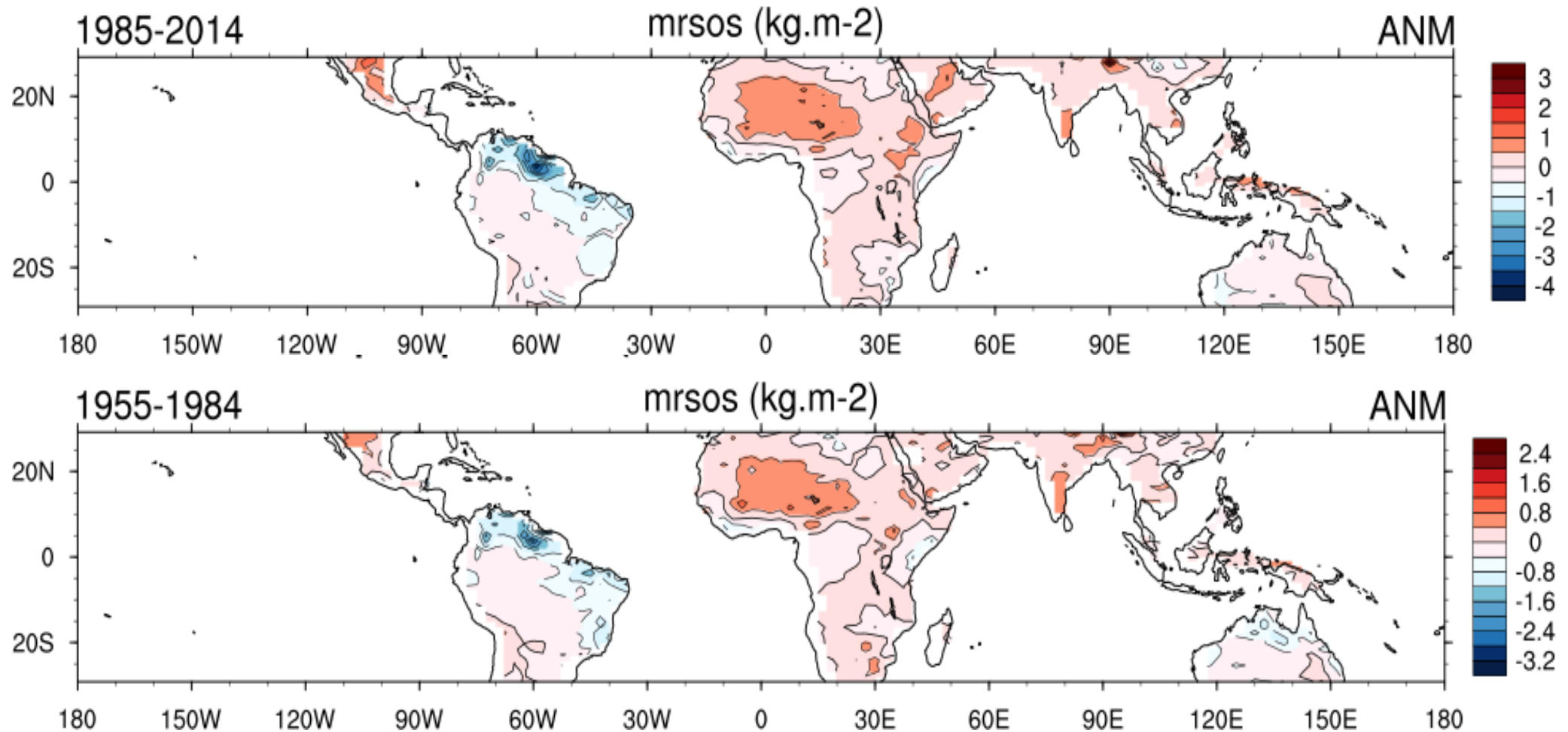
**CM61-LR-5perc-pi-AER-01 (vs IPSLCM6 CM61-LR-NoDust-hist-AER-01)**





# Total soil water content change for 1985-2014 and 1955-1984 periods

**CM61-LR-5perc-pi-AER-01 (vs IPSLCM6 CM61-LR-NoDust-hist-AER-01)**



## Conclusions

- Absorbing dust strongly influences Sahel summer precipitations
- Two factors control this absorption: the iron oxide content and the particle size (absorption increases with particle size)
- We ran a fully coupled Earth System Model with and without dust for 100 years over the period 1915 to 2014
- Over Sahel, the energy absorbed can be as high as  $+36 \text{ W.m}^{-2}$  for the JJA period
- We analyzed changes in water vapor budget in response to changes in this energy
- This analysis will be prolonged by the analysis of the moist static energy