

# The magnetic signatures of oceanic tides in satellite data

A virtual-observatory approach

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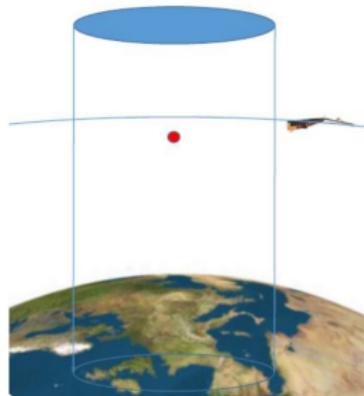
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DTU Space  
National Space Institute  
Kgs. Lyngby

May 2020, EGU2020: Sharing Geoscience Online

# Virtual observatories

- ▶ a robust procedure for estimating high-resolution time series of the secular variation of the core field
- ▶ introduced by Mandea & Olsen (2006), later reused by Olsen, Beggan, Whaler,...
- + local method with local error estimates
- + usually does not imply any regularization in time (unlike global field models)
- problem with external field contamination (ideally should average to zero)
- more strict data selection criteria → strong reduction of data dimension for local inversion



after C. Finlay

# Virtual observatories and tides

Local Laplacian potential field: tidal parameterization

- ▶ local quadratic/cubic parameterization in Cartesian coordinates

$$V(x, y, z; t) = \sum_{a+b+c \leq l} C_{abc}(t) x^a y^b z^c$$

- ▶  $k = 1, \dots, K$ : individual tidal constituents (e.g., M<sub>2</sub>, N<sub>2</sub>, O<sub>1</sub>)

$$C_{abc}(t) = \operatorname{Re} \left\{ \sum_{k=1}^K f_k(t) \hat{c}_{abc,k} \exp [i(\omega_k(t - t_0) + V_{0,k}(t_0) + u_k(t))] \right\}$$

where  $\hat{c}_{abc,k} \in \mathbb{C}$

- ▶ tidal parameters, available from TPXO subroutines (Egbert & Erofeeva 2002)

$\omega_k$  angular frequency

$f_k(t)$  amplitude modulation (seasonal)

$u_k(t)$  phase modulation

$V_{0,k}(t_0)$  Greenwich phase related to  $t_0 = 1992.0$

- ▶ number of free complex parameters  $\hat{c}_{abc,k}$ : 8K/15K



# Virtual observatories and tides

## VO Algorithm for tidal signals

1. select satellite data by quietness criteria
2. subtract a-priori models of main and external fields
3. choose a virtual observatory, a search radius ( $\approx 500$  km) and all times
4. select all residua within the search radius
5. rotate the residua to a local Cartesian coordinate system
6. fit local time-dependent Laplacian potential field with a-priori tidal parameters by Iterative Reweighted Least-Squares with Huber weights
7. repeat from 3 for next VO

# Virtual observatories and tides

## VO Algorithm for tidal signals

1. select satellite data by quietness criteria
  - ▶  $K_p < 3$
  - ▶  $\left| \frac{dRC}{dt} \right| < 3 \text{ nT/hr}$
  - ▶  $E_m \leqslant 0.8 \text{ mV/m}$
  - ▶  $B_z^{\text{IMF}} > 0 \text{ nT}$
  - ▶  $|B_y^{\text{IMF}}| < 10 \text{ nT}$
  - ▶ Sun at least  $10^\circ$  below horizon
2. subtract a-priori models of main and external fields
3. choose a virtual observatory, a search radius ( $\approx 500 \text{ km}$ ) and all times
4. select all residua within the search radius
5. rotate the residua to a local Cartesian coordinate system
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# Virtual observatories and tides

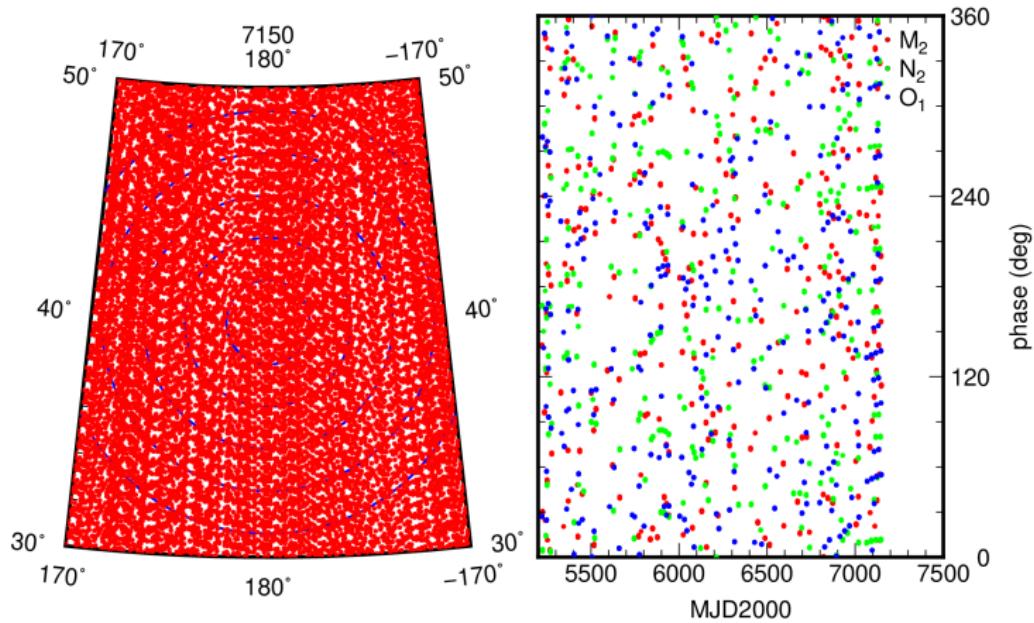
## VO Algorithm for tidal signals

1. select satellite data by quietness criteria
2. subtract a-priori models of main and external fields
  - ▶ core field and lithosphere (CHAOS-6)
  - ▶ magnetospheric external field (CHAOS external model)
  - ▶ ionospheric external and induced field (CIY4 model)
3. choose a virtual observatory, a search radius ( $\approx 500$  km) and all times
4. select all residua within the search radius
5. rotate the residua to a local Cartesian coordinate system
6. fit local time-dependent Laplacian potential field with a-priori tidal parameters by Iterative Reweighted Least-Squares with Huber weights
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# Virtual observatories and tides

## Example of data and phase coverage



left: spatial coverage around a VO in the Northern Pacific; blue circles mark distance from the VO with 200 km steps

right: coverage of the phase of individual tides for 600 km distance

## Numerical modelling

- ▶ elmgFD: frequency-domain spherical harmonic-finite element solver
- ▶ zero external forcing, preconditioned BiCGSTAB(2), OpenMP (Velímský et al. 2018)
- ▶  $j_{\max} = 480$ ,  $K_{3D} = 101$
- ▶ 1-D mantle conductivity profile (Grayver et al. 2017)
- ▶ 3-D ocean conductivity based on collocated temperature and salinity measurements (World Ocean Atlas, Tyler et al. 2017)
- ▶ ocean-bottom sediments (a-priori assigned values and maps of thicknesses, Everett et al. 2003)
- ▶ TPXO9-atlas ocean flows for  $M_2$ ,  $N_2$ ,  $O_1$

# VO analysis setup for Swarm A and C

Parameter study for M<sub>2</sub>, N<sub>2</sub>, O<sub>1</sub>

- ▶ number of VOs in regular grid ( $N_\varphi \times N_\vartheta$ )
- ▶ search radius  $d$
- ▶ fields (A,C):  $\mathbf{B}_i^A, \mathbf{B}_i^C$
- ▶ NS+EW differences (A-C,A+C)

NS along-track differences and sums on both satellites:  $\frac{\mathbf{B}_{i+1}^A - \mathbf{B}_i^A}{2}, \frac{\mathbf{B}_{i+1}^A + \mathbf{B}_i^A}{2}, \frac{\mathbf{B}_{i+1}^C - \mathbf{B}_i^C}{2}, \frac{\mathbf{B}_{i+1}^C + \mathbf{B}_i^C}{2}$

EW cross-track differences and sums  $\frac{\mathbf{B}_i^A - \mathbf{B}_i^C}{2}, \frac{\mathbf{B}_i^A + \mathbf{B}_i^C}{2}$

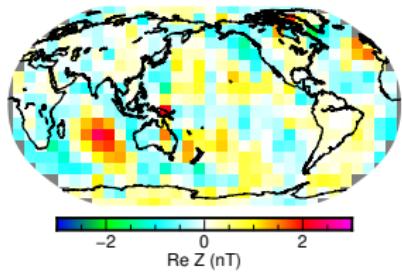


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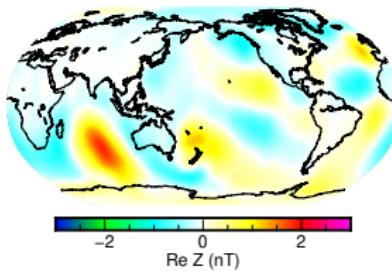
$M_2$

## Swarm A and Swarm C fields

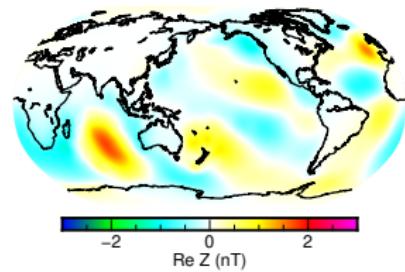
036x018,0500 km,A,C



Grayver & Olsen 2019 (SH28)



Forward model (SH480)



Re  $Z$  (nT)

Re  $Z$  (nT)

Re  $Z$  (nT)

Im  $Z$  (nT)

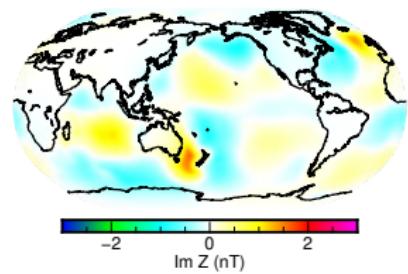
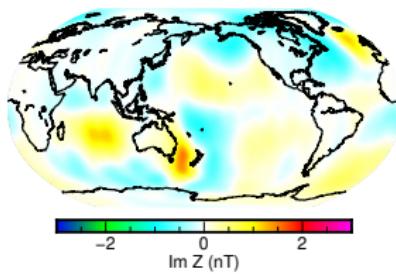
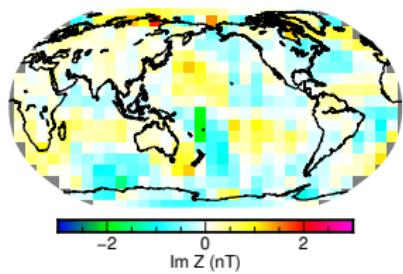
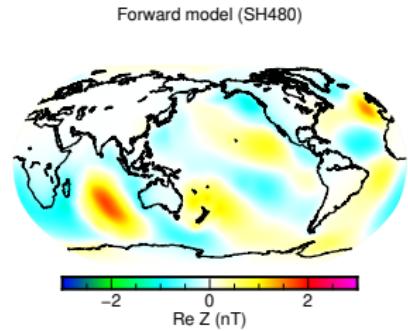
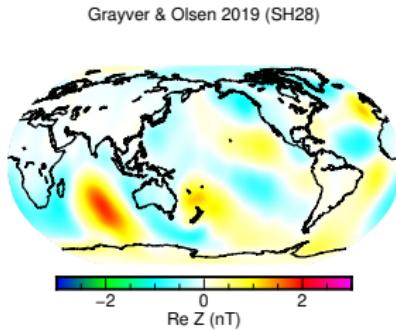
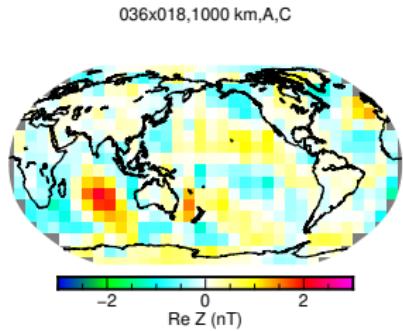
Im  $Z$  (nT)

Im  $Z$  (nT)

# Selected results

$M_2$

## Swarm A and Swarm C fields

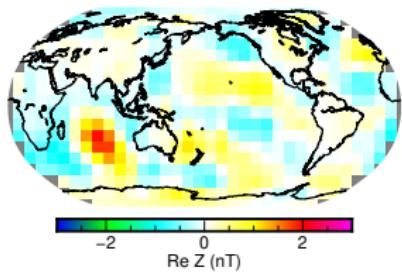


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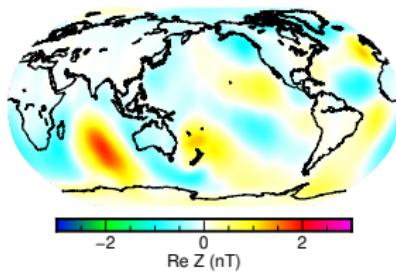
$M_2$

## Swarm A and Swarm C fields

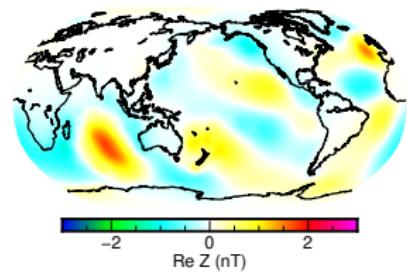
036x018,2000 km,A,C



Grayver & Olsen 2019 (SH28)



Forward model (SH480)



Re  $Z$  (nT)

Re  $Z$  (nT)

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Im  $Z$  (nT)

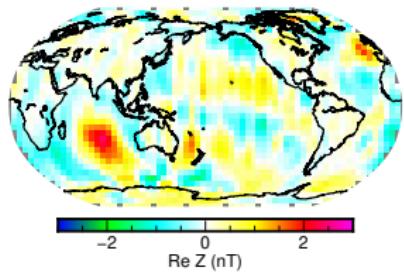
Im  $Z$  (nT)

# Selected results

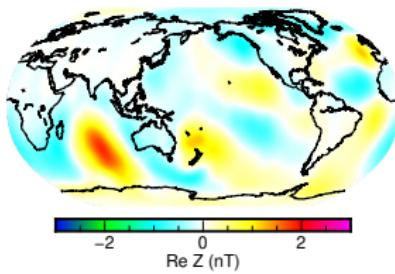
$M_2$

## Swarm A and Swarm C fields

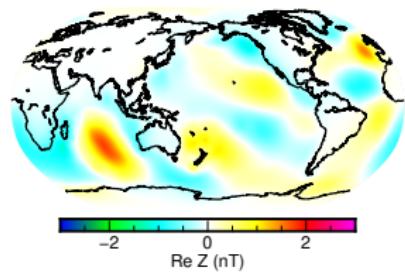
072x036,1000 km,A,C



Grayver & Olsen 2019 (SH28)



Forward model (SH480)



$Re Z$  (nT)

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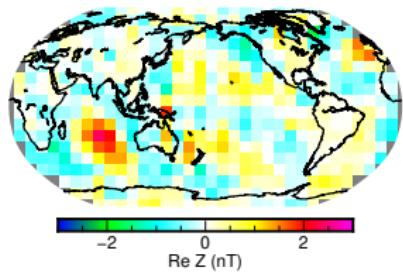
$Im Z$  (nT)

# Selected results

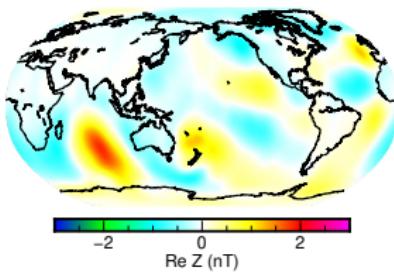
$M_2$

## Swarm A and C NS and EW sums and differences

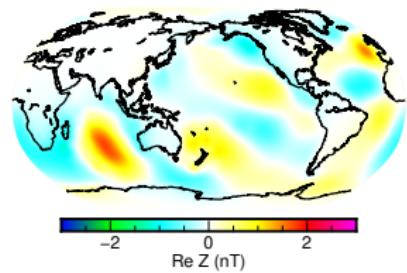
036x018,0500 km,A+C,A-C



Grayver & Olsen 2019 (SH28)



Forward model (SH480)



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$Re Z$  (nT)

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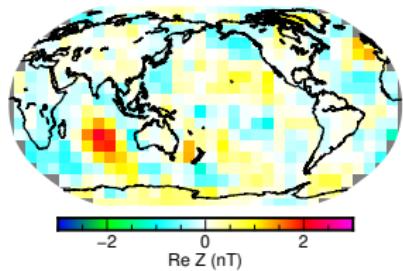
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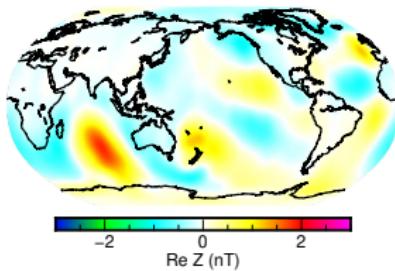
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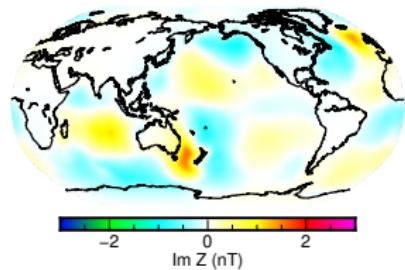
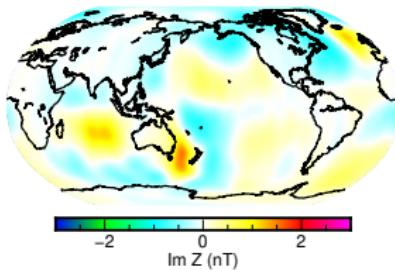
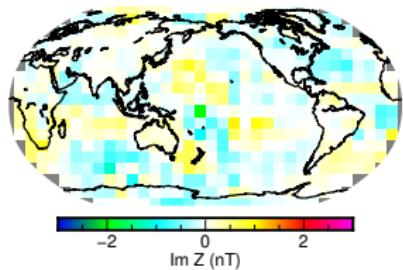
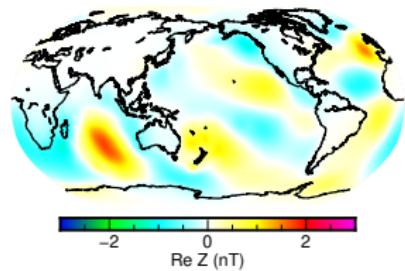
036x018, 1000 km, A+C, A-C



Grayver & Olsen 2019 (SH28)



Forward model (SH480)

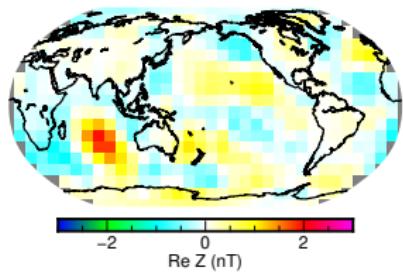


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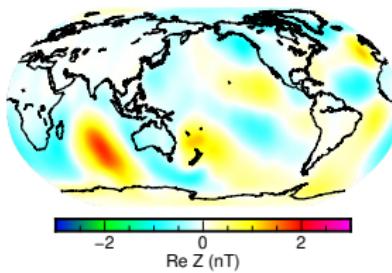
$M_2$

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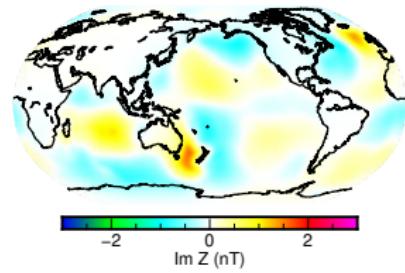
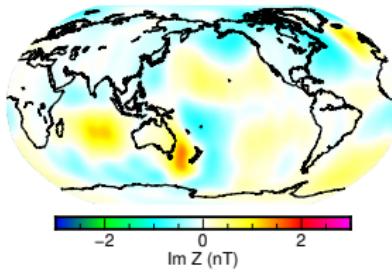
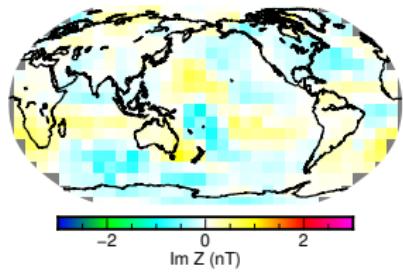
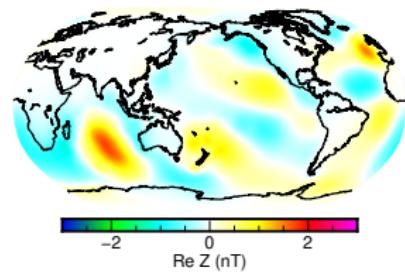
036x018,2000 km,A+C,A-C



Grayver & Olsen 2019 (SH28)



Forward model (SH480)

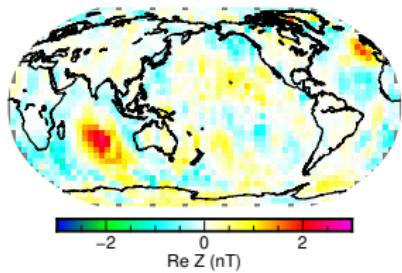


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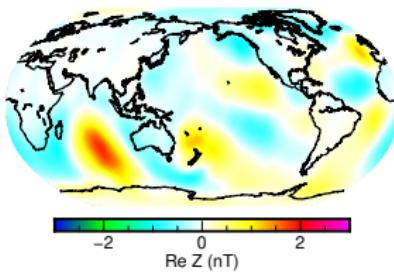
$M_2$

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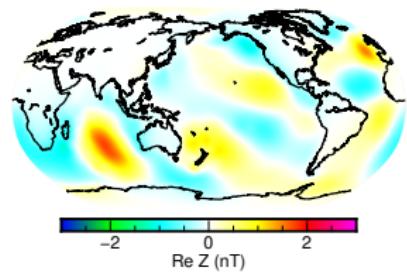
072x036, 1000 km, A+C, A-C



Grayver & Olsen 2019 (SH28)



Forward model (SH480)



-2 0 2

-2 0 2

-2 0 2

Im Z (nT)

Im Z (nT)

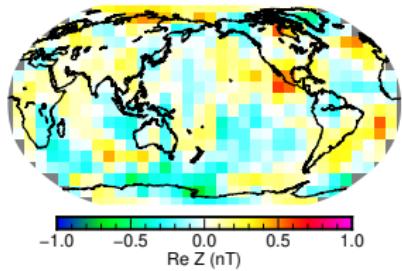
Im Z (nT)

# Selected results

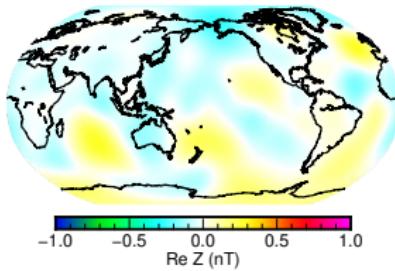
$N_2$

## Swarm A and C NS and EW sums and differences

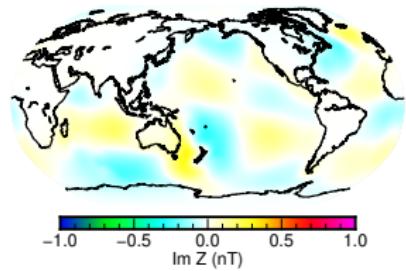
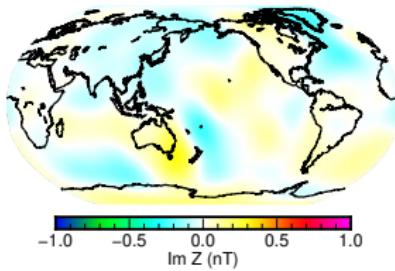
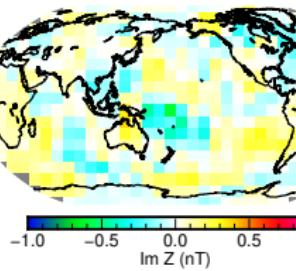
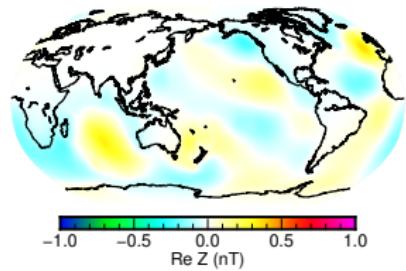
036x018,2000 km,A+C,A-C



Grayver & Olsen 2019 (SH12)

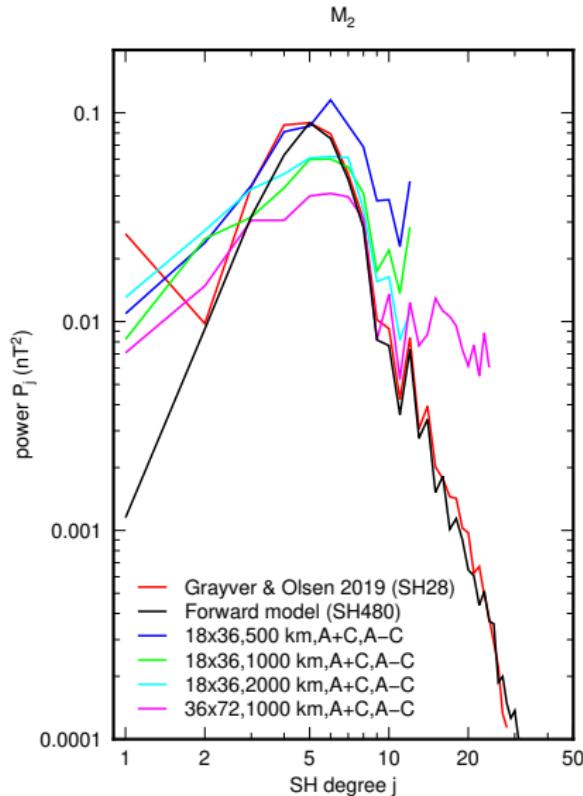
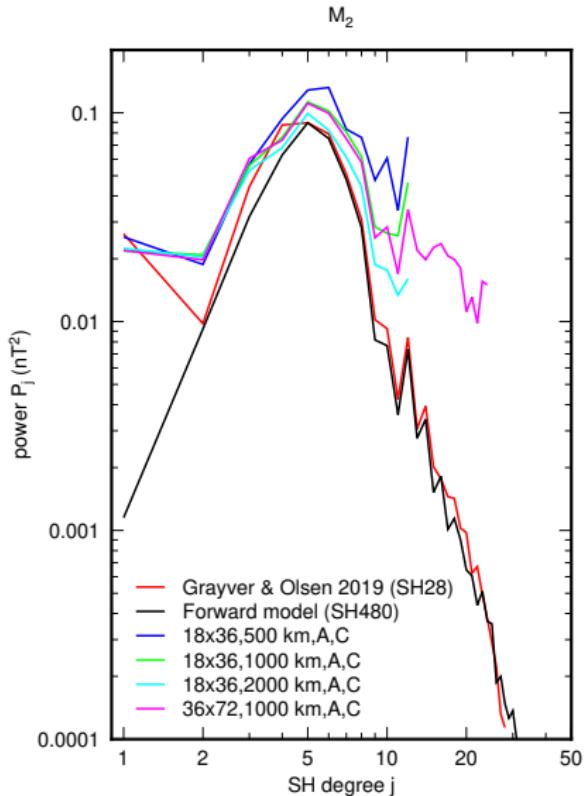


Forward model (SH480)



# Selected results

## Comparison of power spectra for $M_2$



# Conclusions

$M_2$  successfully recovered from Swarm A,C data by VO approach

- ▶ significant dependence on the choice of search radius  $d$   
(smoother solution with suppressed higher harmonics for large  $d$ )
- ▶ use of NS and EW differences does not introduce any particular advantage
- ▶ alternative corrections for external field to be exploited

$N_2$  poorly recovered

$O_1$  not recovered