

Tracing dew and fog water inputs into temperate grassland using stable water isotopes in the extreme summer 2018

Yafei Li^{1*}, Franziska Aemisegger², Andreas Riedl¹, Nina Buchmann¹, Werner Eugster¹

¹ Institute of Agricultural Sciences, ETH Zurich, Zurich Switzerland * Correspondence to: Yafei Li yafei.li@usys.ethz.ch

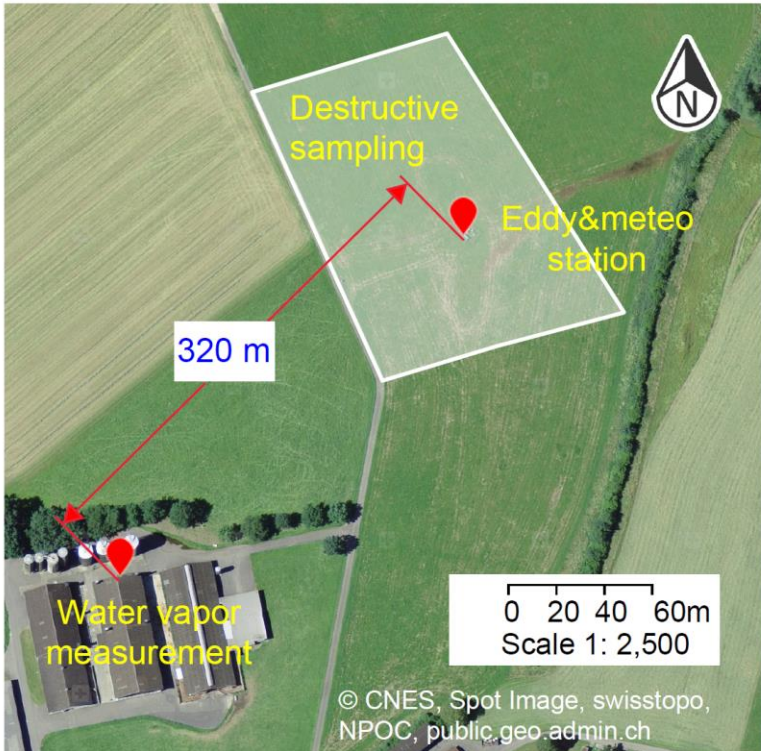
² Institute for Atmospheric and Climate Science, ETH Zurich, Zurich, Switzerland

Using stable isotopes of water to trace the water cycle during dew formation and shallow ground radiation fog deposition

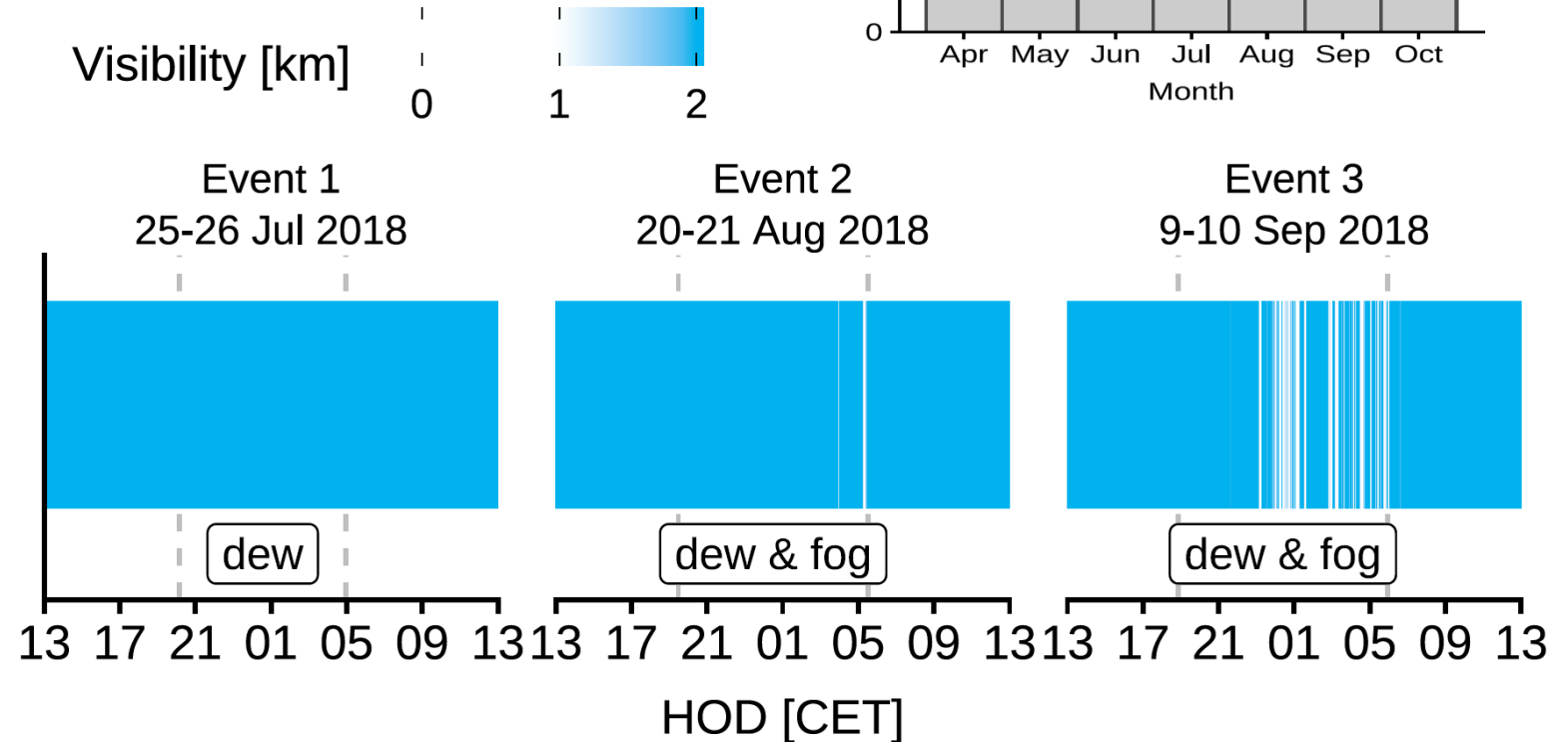
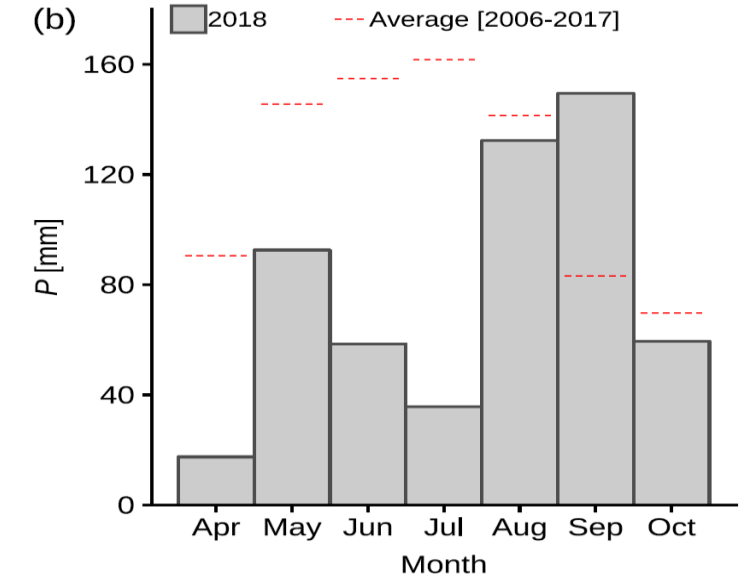
- ✓ Non-rainfall water inputs into grasslands were traced by isotopic dynamics of near-surface atmospheric water vapor
- ✓ Near-surface atmospheric water vapor was the main but not the only pathway of non-rainfall water inputs
- ✓ Internal cycle of water: water vapor from soil condenses back on foliage

Experiment setup

- Study site: Swiss grassland in a valley bottom at 400 m a.s.l.
- Experiment setup
 - ✓ Isotopes of water vapor (6 m a.g.l.)
 - ✓ Isotopes of droplets on leaf
 - ✓ Eddy & meteo (2 – 2.4 m a.g.l.)
- Three observation campaigns in the extreme summer 2018



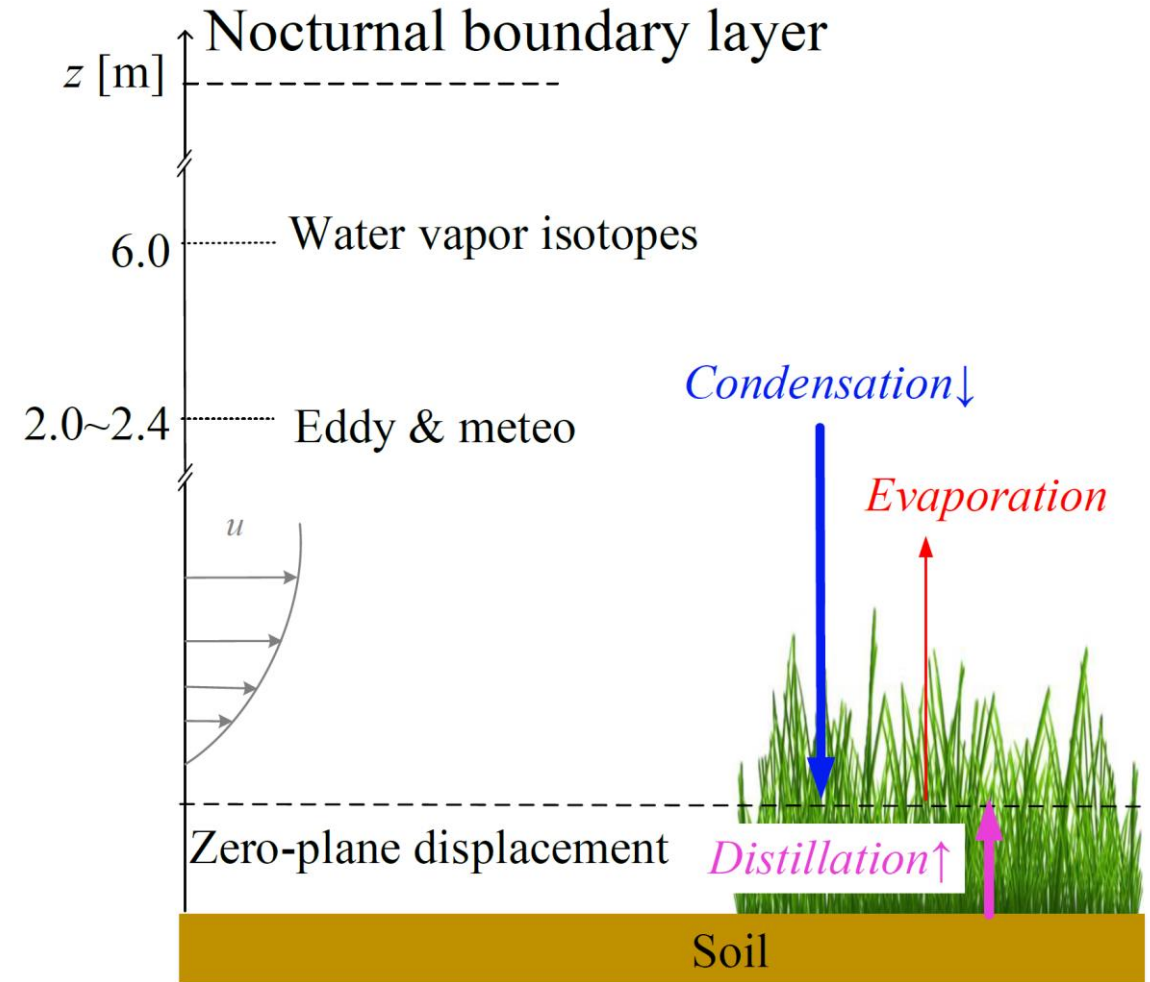
■ Precipitation



Water cycle during non-rainfall water inputs

During dew formation and shallow ground radiation fog deposition:

- Above zero-plane displacement, downward vapor flux from near-surface atmospheric water vapor condenses on leaf surfaces or close to ground atmosphere (*condensation*↓)
- Below zero-plane displacement, upward vapor flux from soil condenses on leaf surfaces (*distillation*↑)
- Condensation was accompanied by *evaporation* due to humidity gradient between surface and atmosphere

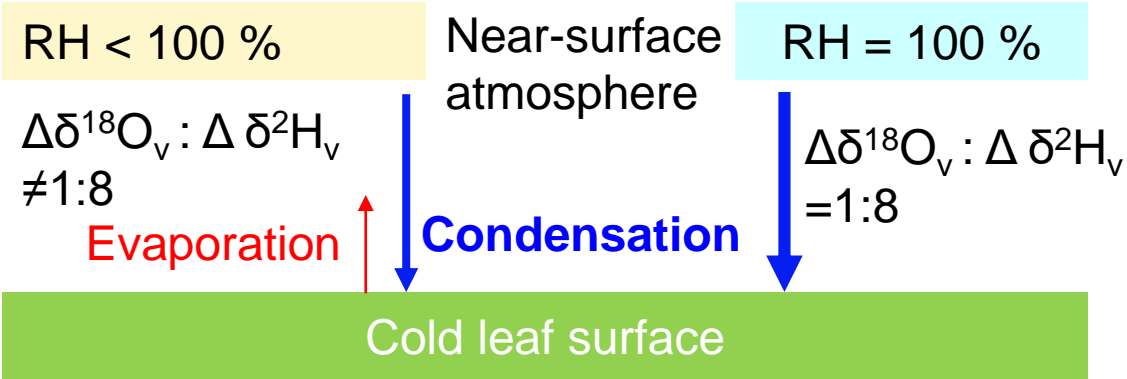


Isotopic dynamics of water vapor during non-rainfall water inputs

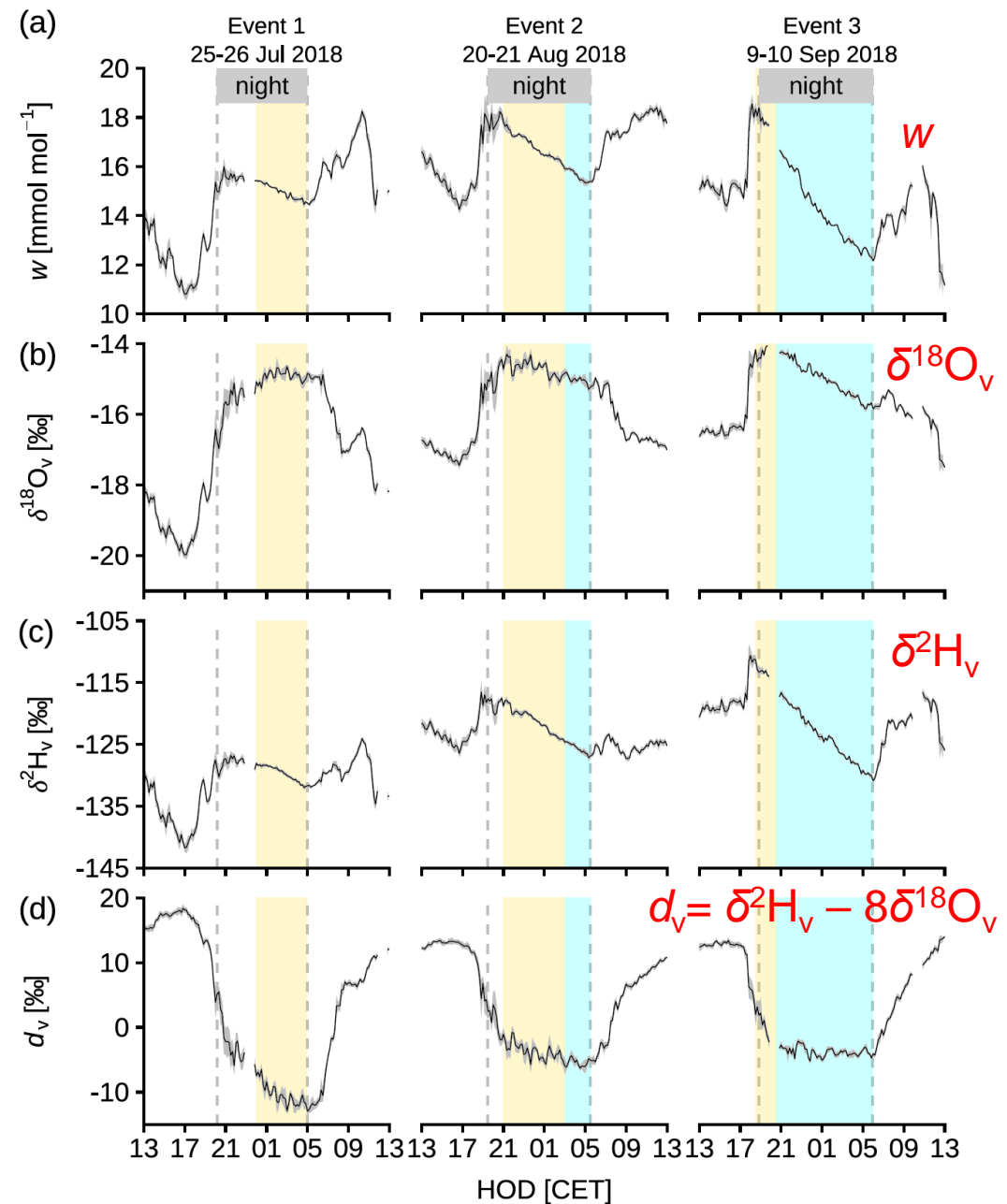
Non-rainfall water inputs into grasslands were traced by water vapor isotopes:

| Water vapor | nsaturated | saturated |
|--------------------------------|-------------|-----------|
| w (water vapor mixing ratio) | ↓ | ↓ |
| $\delta^{18}\text{O}_v$ * | Fluctuating | ↓ |
| $\delta^2\text{H}_v$ * | Slight ↓ | Faster ↓ |
| d_v | ↓ | Constant |

*In the roughly 2 h before sunset, decreased $\delta^{18}\text{O}_v$ indicated that evaporated vapor $\delta^{18}\text{O}_E$ was more enriched than $\delta^{18}\text{O}_v$, but evaporated vapor $\delta^2\text{H}_E$ was not significantly different from $\delta^2\text{H}_v$, thus under unsaturation, evaporation caused $\delta^{18}\text{O}_v$ fluctuation, but weak disturbance on $\delta^2\text{H}_v$.



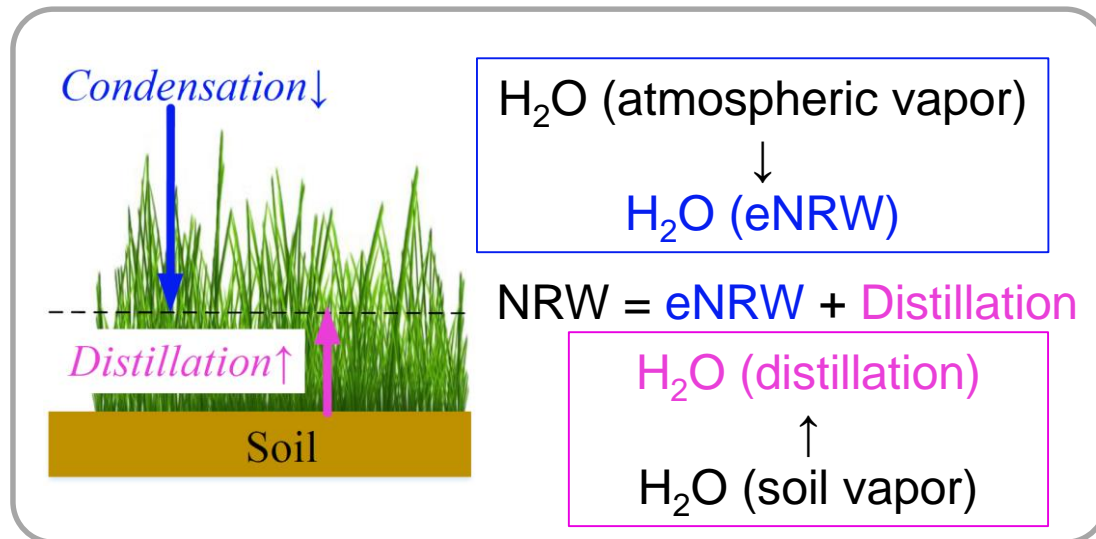
(Horita and Wesolowski, 1994; Majoube, 1971; Dansgaard, 1964)



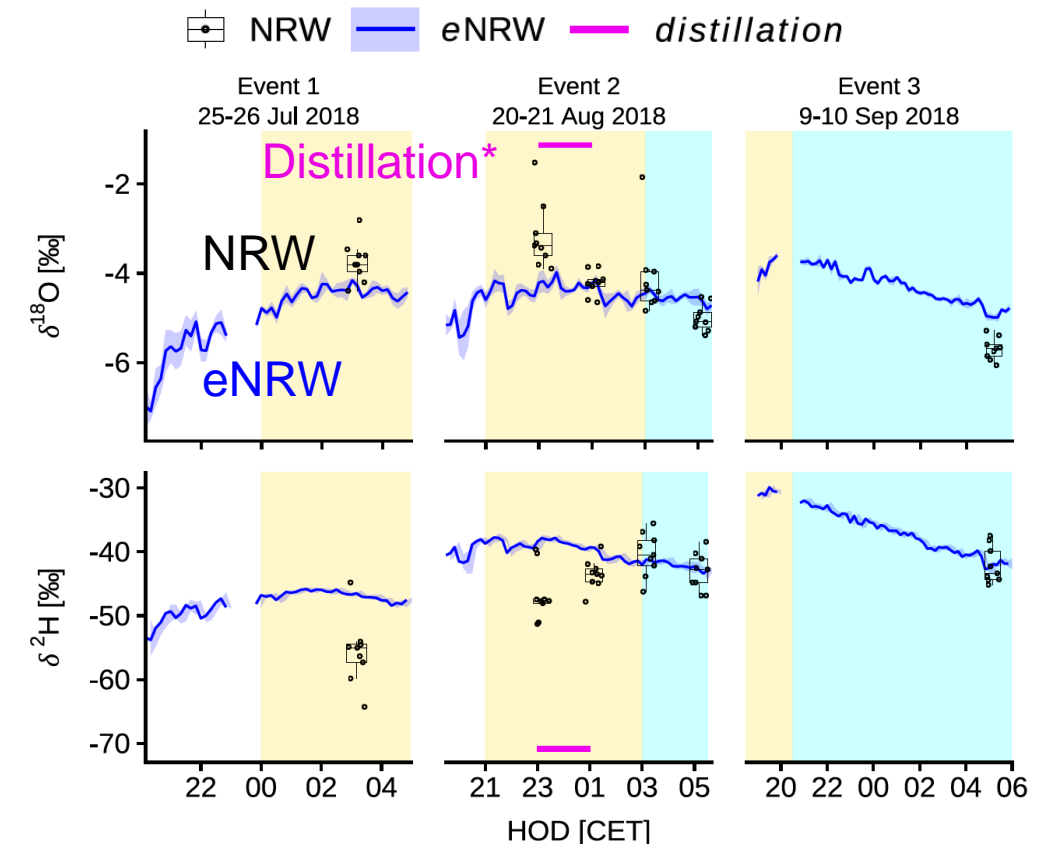
Two pathways of non-rainfall water inputs

The difference of δ between NRW and eNRW indicated **two pathways of non-rainfall water inputs**:

- **Condensation**↓: flux from atmospheric water vapor (as main pathway)
- **Distillation**↑: vapor flux from soil (internal recycling)
- With increased RH, **condensation**↓ became stronger, thus **Distillation**↑ contribution became smaller due to stronger **condensation**↓



- NRW: non-rainfall water taken from leaf surfaces
- eNRW: equilibrium liquids of atmospheric water vapor
- Distillation: vapor from soil condenses on foliage



* δ of distillation was estimated from the δ of NRW and eNRW under unsaturation, and was roughly assumed constant.

Conclusion

- **Dew formation and shallow ground radiation fog deposition** in temperate grassland under summer drought were confirmed by liquid and vapor isotopes
- Non-rainfall water inputs transform water source that are sparingly accessible by plants into water accessible to plants
 - ✓ Near-surface atmospheric water vapor as the main source
 - ✓ Water vapor from soil (distillation) — thus internal recycling of water — is expected to be more important than previously thought in dew formation
- Future study should investigate whether this internal recycling (distillation) is sufficiently important to make deeper soil water available to plants that would otherwise not be accessible via plant roots