

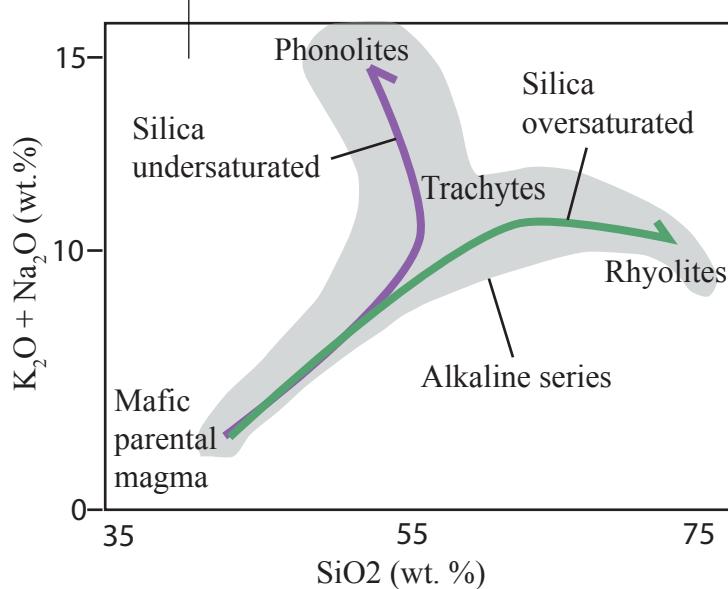
The role of contamination in the tightrope of Gran Canaria felsic magma differentiation



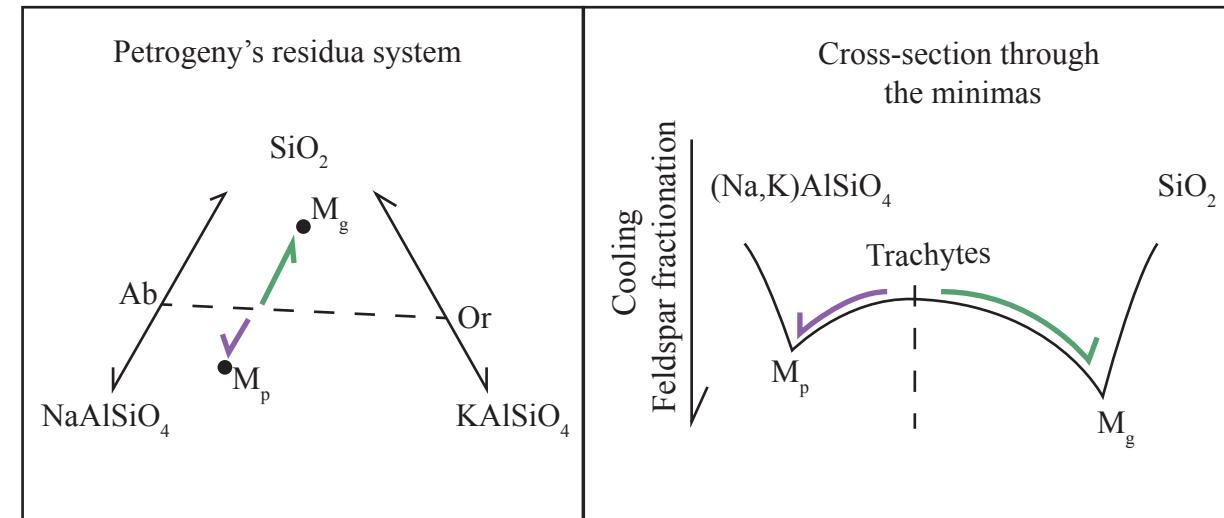
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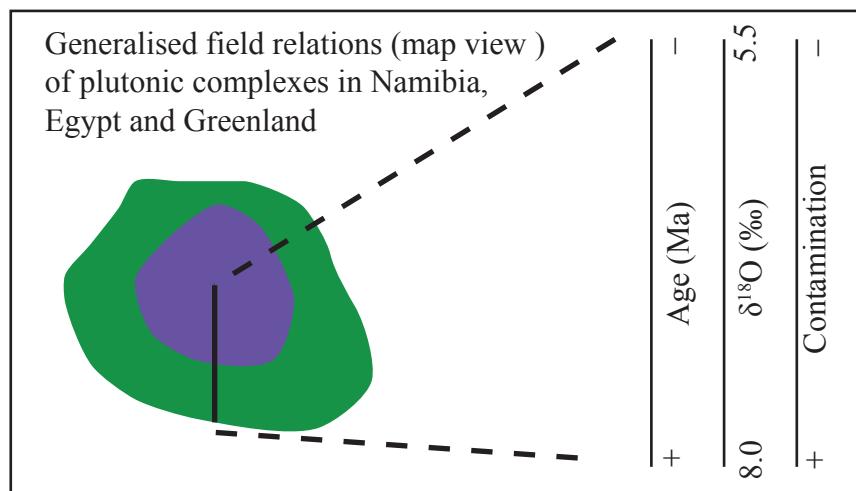
Alkaline Magmas



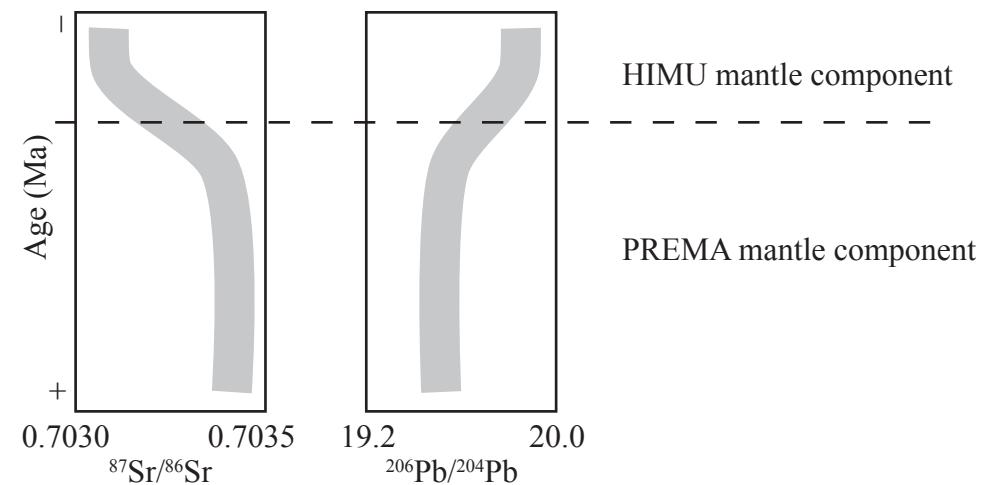
A. Two potential liquid lines of descent (LLD) controlled by fractional crystallisation (FC).
(modified after Conceicao et al., 2007)



B. Thermal barrier hinders the relation of both LLD by FC
(Ab-Albite, Or-Orthoclase, M_g-granite minima, M_p-phonolite minima)
(modified after Tuttle & Bowen, 1958; Wolff, 2017)

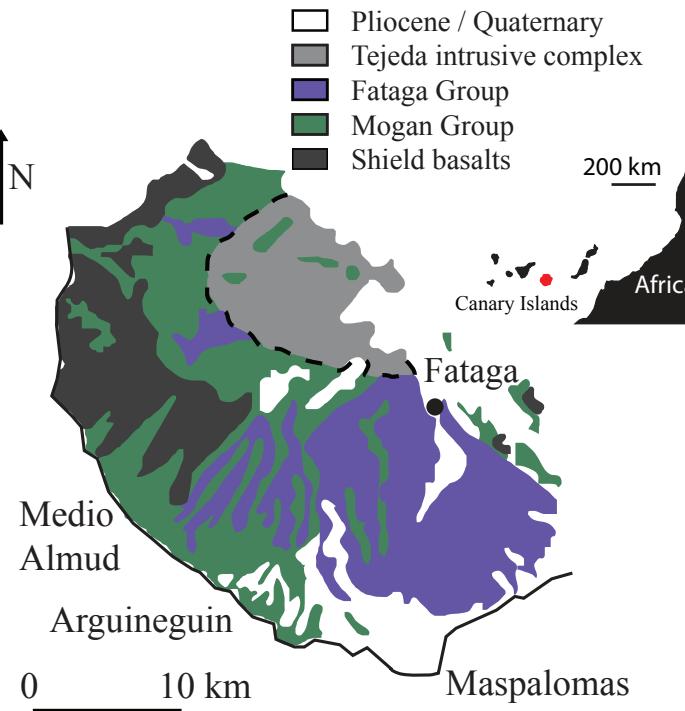


C. Crustal contamination of a silica undersaturated system potentially explains how to cross the thermal barrier
(e.g. Harris, 1999; Riishuus, 2008; Decrée et al., 2019)



D. Identifying contamination is important because isotopic changes in alkaline centers have been linked to changes in mantle source (e.g. Gran Canaria, modified after Cousens et al., 1990)

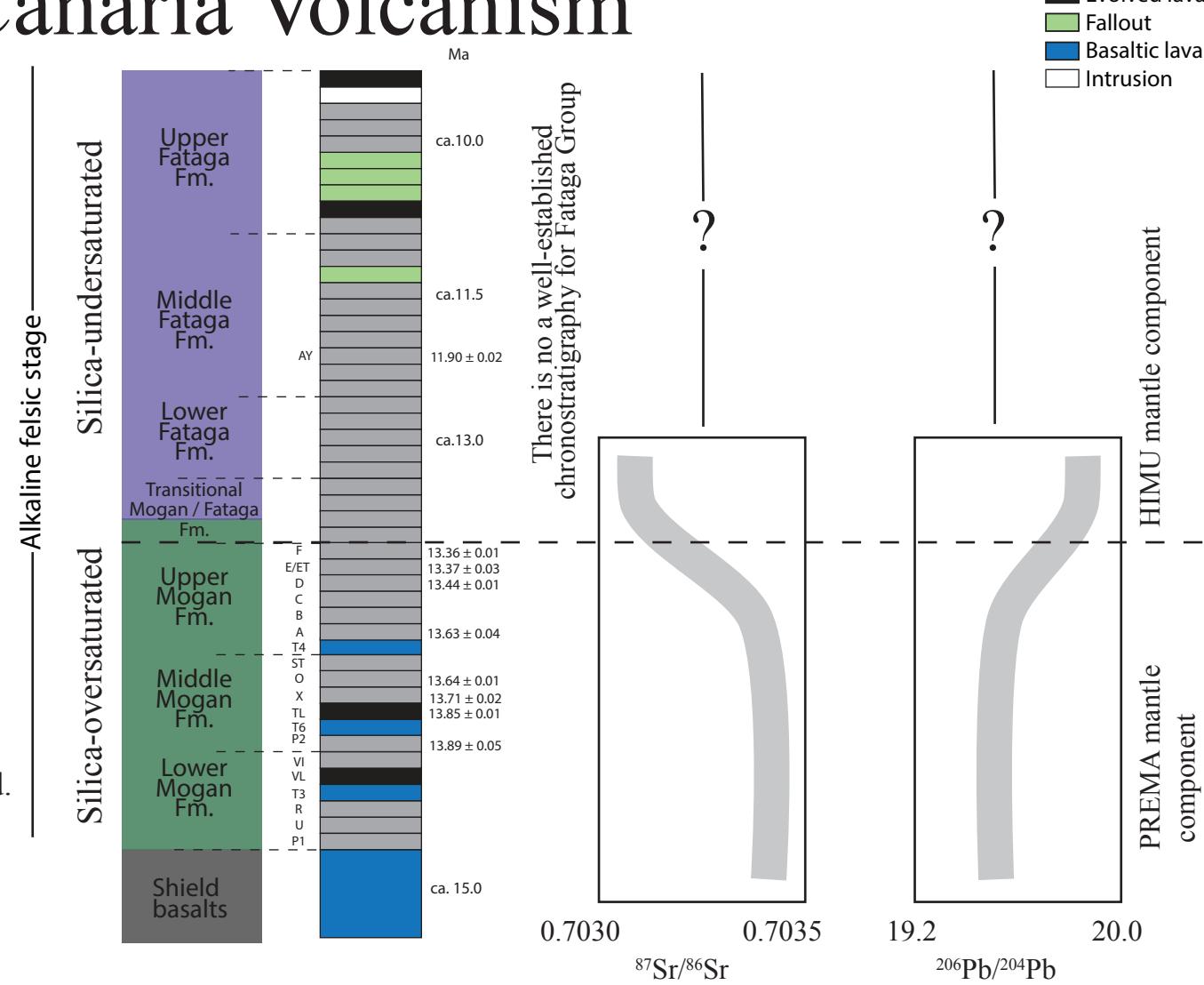
Gran Canaria Volcanism



A. Gran Canaria is an oceanic intraplate volcanic island. Volcanism is divided into: 1) shield stage; 2) alkaline felsic stage; 3) rejuvenation stage (modified after Cousens et al., 1990; Troll & Carracedo, 2016)

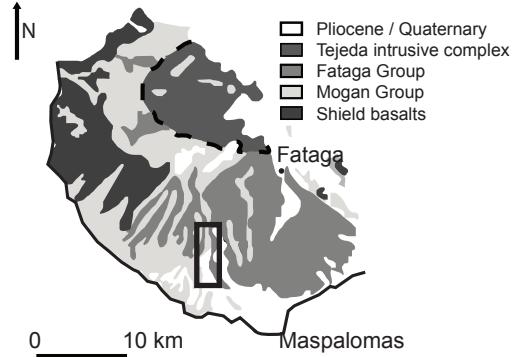
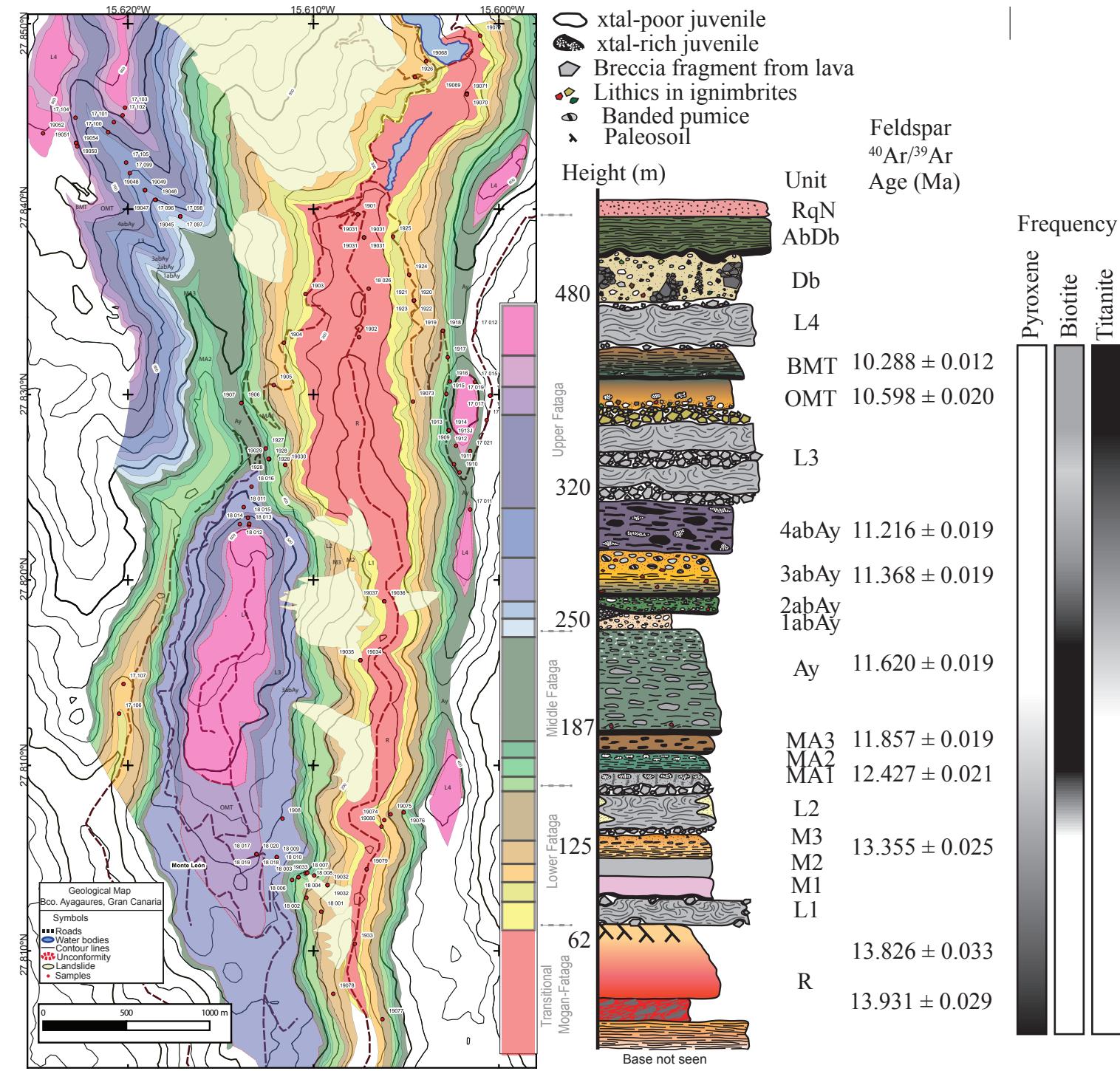
Research questions

1. Are the LLDs in the alkaline felsic stage of Gran Canaria akin to those in alkaline plutonic complexes showing changes in crustal contamination?
2. If there is crustal contamination, which lithologies are involved and to which extent?



B. Studies based on radiogenic isotopes (Sr, Pb) in the Mogan Group and Lower Fataga formation suggest that the changes in isotopes are due to changes in the mantle source (Cousens et al., 1990). The outputs of any isotopic study rely upon chronostratigraphy, but this is not the case for the silica-undersaturated volcanics in Gran Canaria. Furthermore, all this studies extrapolate the results of the lower Fataga formation to the whole Group. (stratigraphic column modified after Van den Bogaard & Schmincke, 1998; radiogenic logs modified after Cousens et al., 1990)

We carried out sampling, mapping and logging in Bco. la Data.



Our general stratigraphy consists of 21 members (lava flows and tuffs), which span ca. 3 Myr. The main mineral phase in these deposits is alkali feldspar. Deposits often show crystal-poor juveniles at the bottom and crystal-rich at the top of each member. Mafic minerals represent less than 20% of the crystallising phases (biotite, pyroxene, Fe-Ti oxides, titanite).

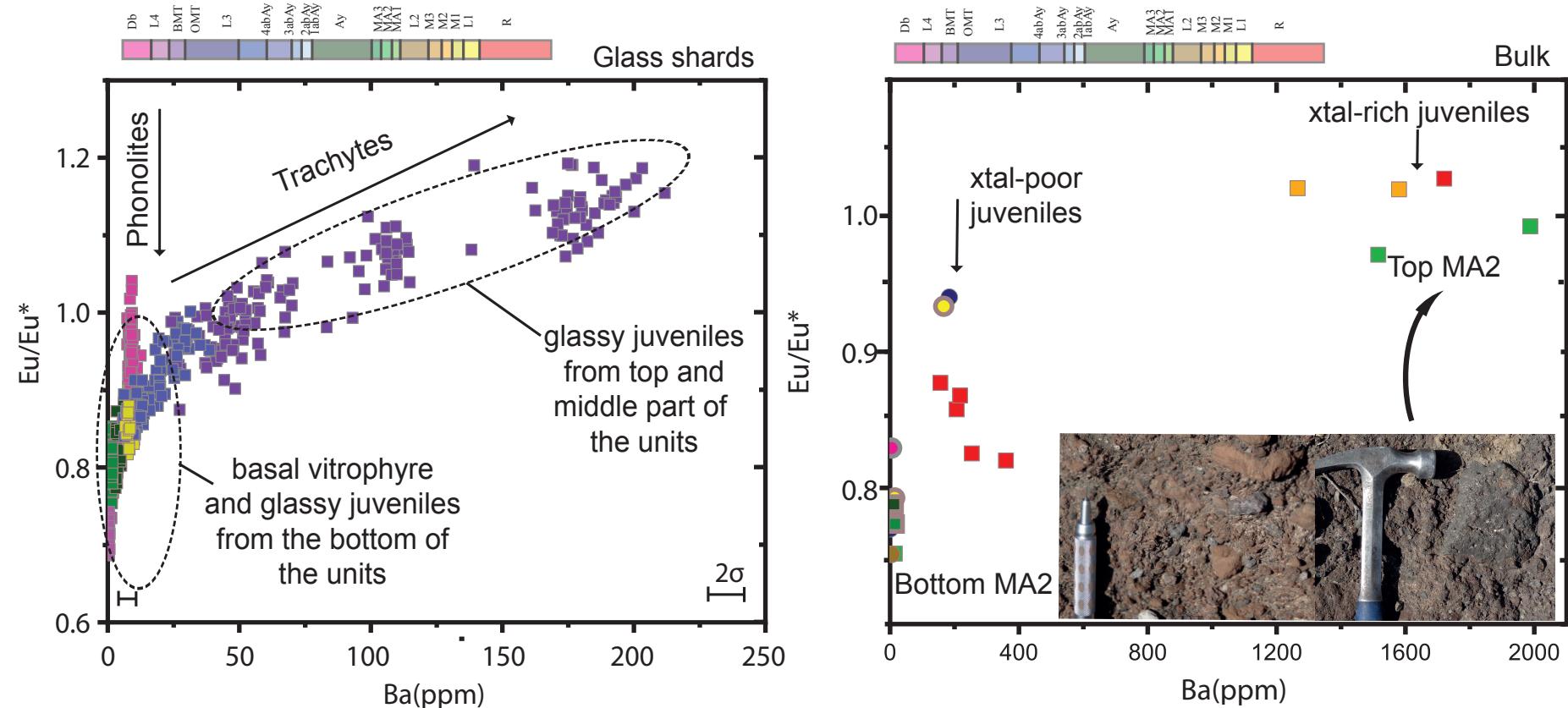
Mineral and glass chemistry by EPMA, LA-ICP-MS

Bulk chemistry of juvenile clasts by XRF, LA-ICP-MS

Oxygen isotopes in feldspar, biotite and pyroxene by laserfluorination

Eruption ages by high-precision single crystal $^{40}\text{Ar}/^{39}\text{Ar}$ in feldspars

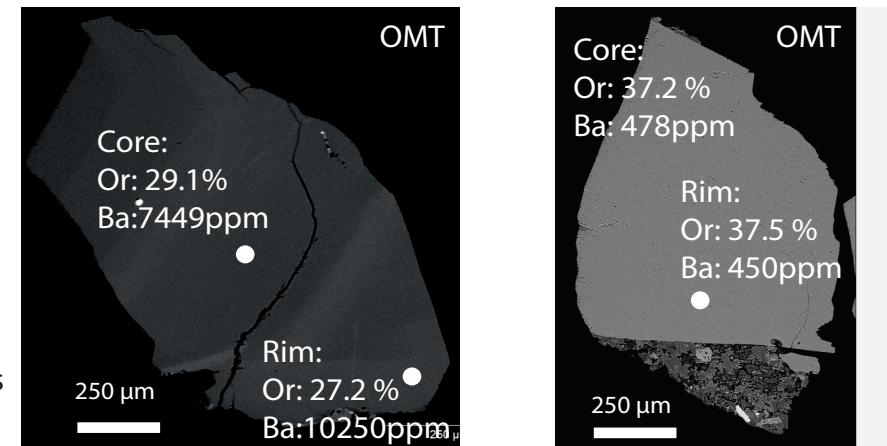
Crustal contamination: Petrological cannibalism



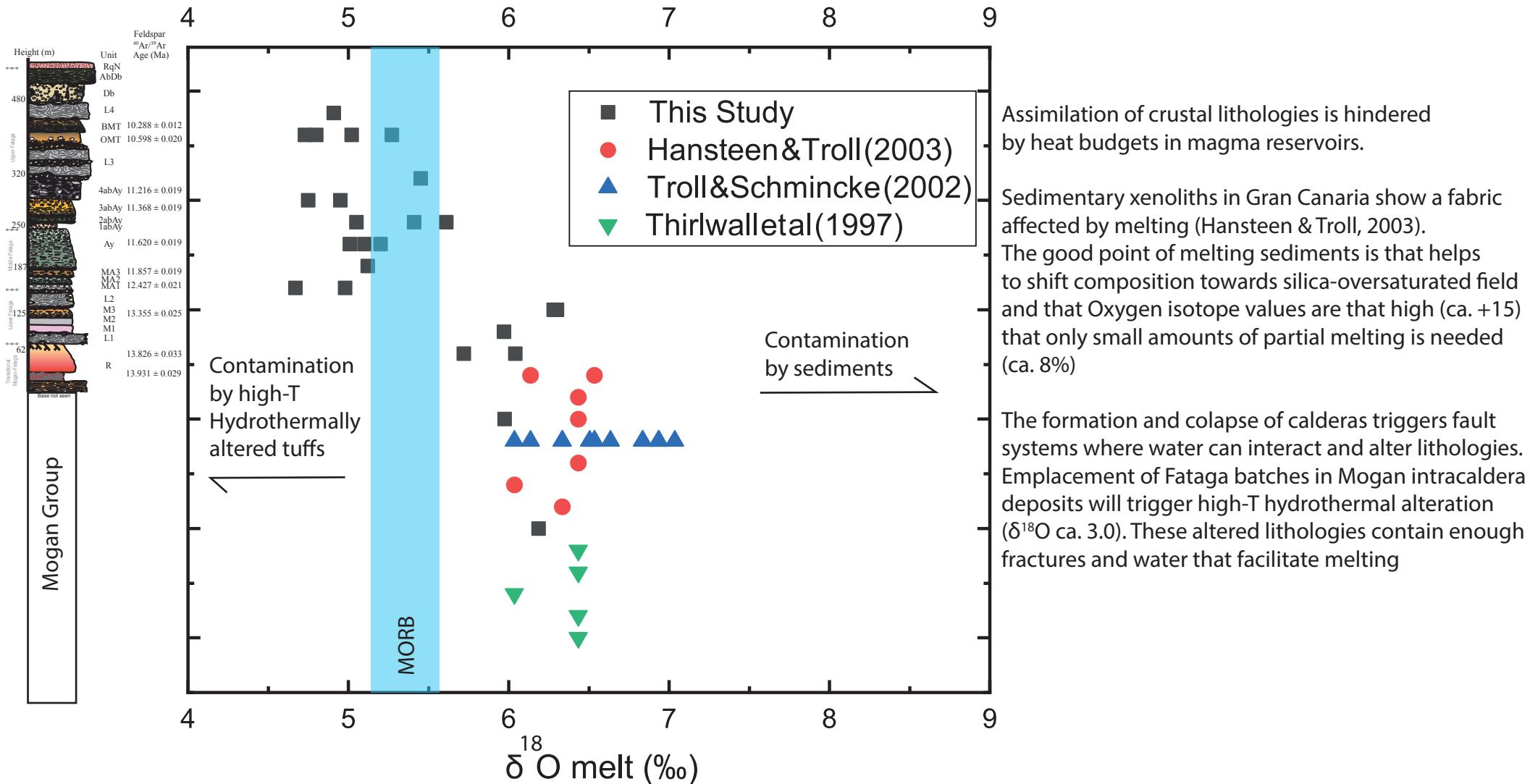
Two alkali-feldspar compositions are found throughout the volcanic sequence. Anorthoclases show strong chemical zonation in Ba and Or (orthoclase) content (rims richer than cores). Sanidines are unzoned and with lower Ba content than anorthoclases. Or content in sanidines increase towards younger deposits as well as the frequency

Cumulate signatures (feldspar accumulation) in juvenile clast are characterised by an increase in Eu/Eu* and Ba.

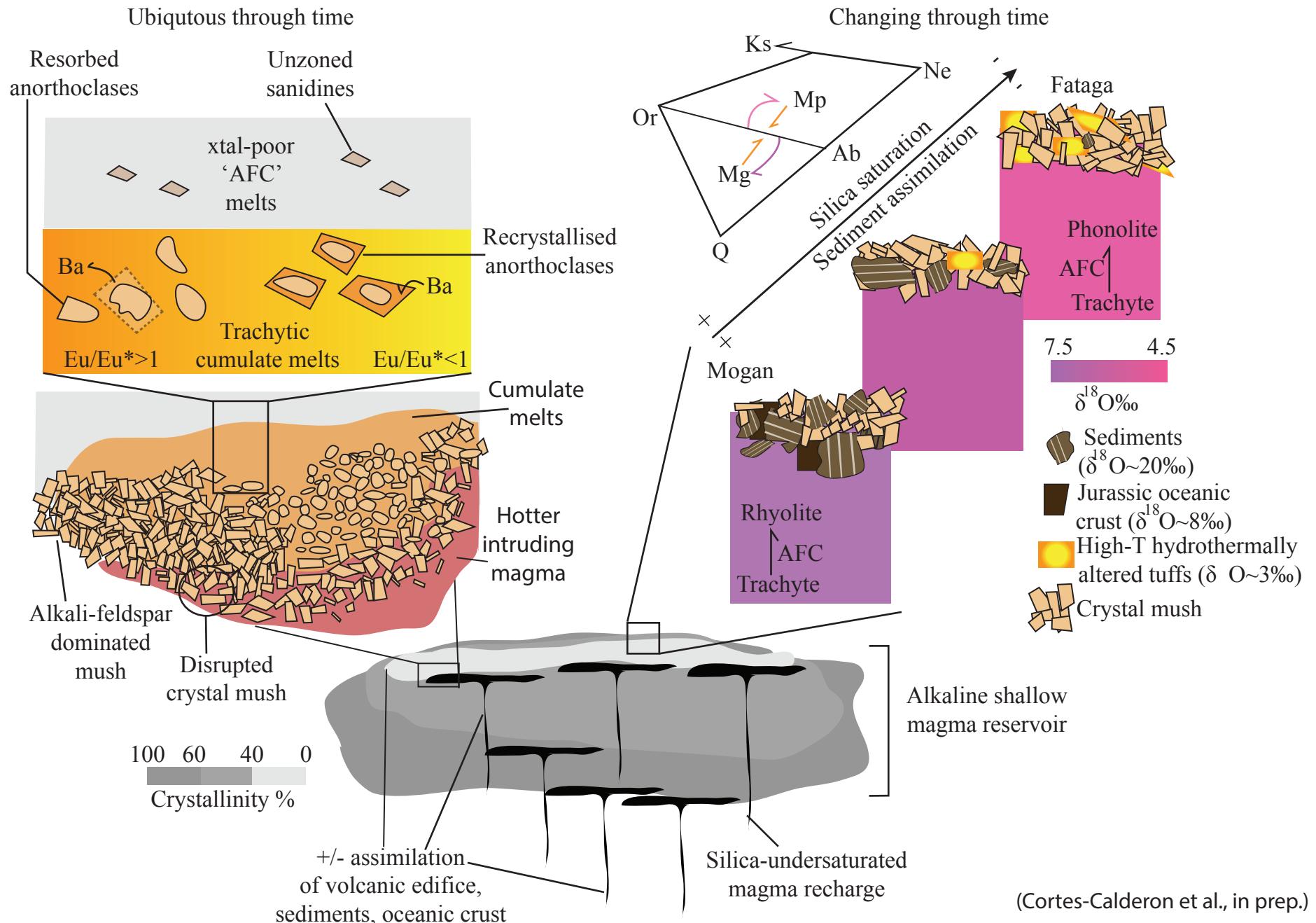
Remelting of cumulates during rejuvenation can be traced in glasses with positive Eu anomalies and increase in Ba.



Crustal contamination: Assimilation of sediments and hydrothermally altered crust



Petrological model for the Miocene felsic stage



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