

# The influence of loading path on fault reactivation: a laboratory perspective

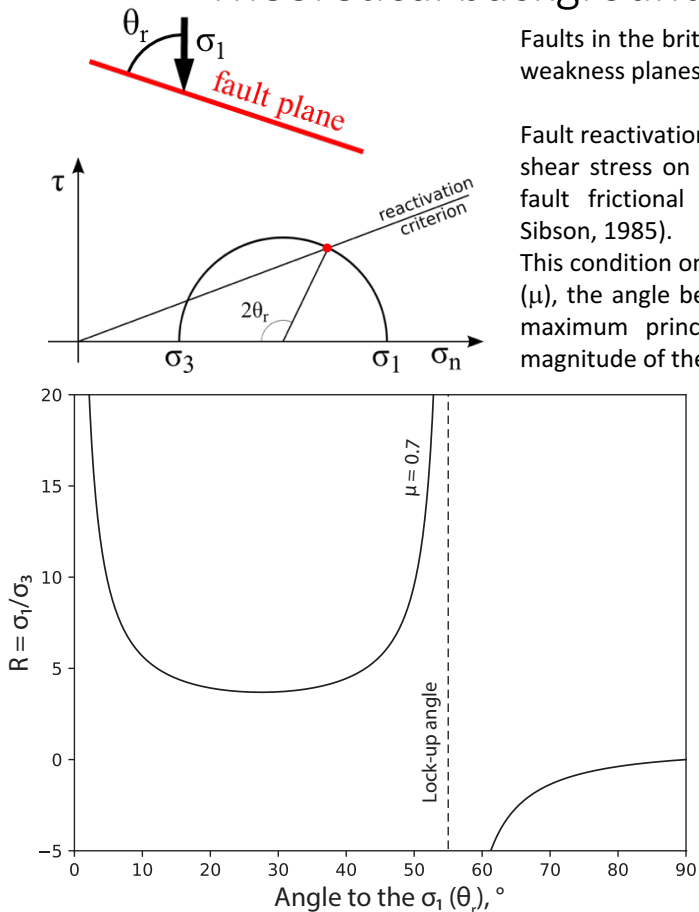


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**Abstract:** Although natural faults are variably oriented to the Earth's surface and to the local stress field, the mechanics of fault reactivation and slip under variable loading paths (sensu Sibson, 1993) is still poorly understood. Nonetheless, different loading paths commonly occur in natural faults, from load-strengthening when the increase in shear stress is coupled with an increase in normal stress (e.g., reverse faults in the absence of fluid pressure increase) to load-weakening when the increase in shear stress is coupled with a decrease in normal stress (e.g., normal faults). According to the Mohr-Coulomb theory, the reactivation of pre-existing faults is only influenced by the fault orientation to the stress field, the fault friction, and the principal stresses magnitude. Therefore, the stress path the fault underwent is often neglected when evaluating the potential for reactivation. Yet, in natural faults characterized by thick, incohesive fault zone and highly fractured damage zone, the loading path could not be ruled out. Here we propose a laboratory approach aimed at reproducing the typical tectonic loading paths for reverse and normal faults. We performed triaxial saw-cut experiments, simulating the reactivation of well-oriented (i.e., 30° to the maximum principal stress) and misoriented (i.e., 50° to the maximum principal stress), normal and reverse gouge-bearing faults under dry and water-saturated conditions. We find that load-strengthening versus load-weakening path results in clearly different hydro-mechanical behaviour. Particularly, prior to reactivation, reverse faults undergo *compaction* even at differential stresses well below the value required for reactivation. Contrarily, normal faults experience *dilation*, most of which occurs only near the differential stress values required for reactivation. Moreover, when reactivating at comparable normal stress, normal faults (load-weakening path) are more prone to slip seismically than reverse fault (load-strengthening path). Indeed, the higher mean stress that normal fault experienced before reactivation compacts more efficiently the gouge layer, thus increasing the fault stiffness and favouring seismic slip. This contrasting fault zone compaction and dilation prior to reactivation may occur in different natural tectonic settings, affecting the fault hydro-mechanical behaviour. Thus, taking into account the loading path the fault underwent is fundamental in evaluating both natural and induced fault reactivation and the related seismic risk assessment.

## Theoretical background and motivation



Faults in the brittle crust constitute pre-existing weakness planes.

Fault reactivation occurs whenever the resolved shear stress on the fault plane overcomes the fault frictional strength (e.g., Jaeger, 1960; Sibson, 1985).

This condition only depends on the fault friction ( $\mu$ ), the angle between the fault plane and the maximum principal stress axis ( $\theta$ ), and the magnitude of the principal stresses ( $\sigma_1, \sigma_2, \sigma_3$ ).

The stress ratio required for reactivation is evaluated as follow:

$$R = \frac{\sigma_1}{\sigma_3} = \frac{1 + \mu_s \cot \theta}{1 - \mu_s \tan \theta}$$

This approach assumes that faults are *zero-thickness* planes.

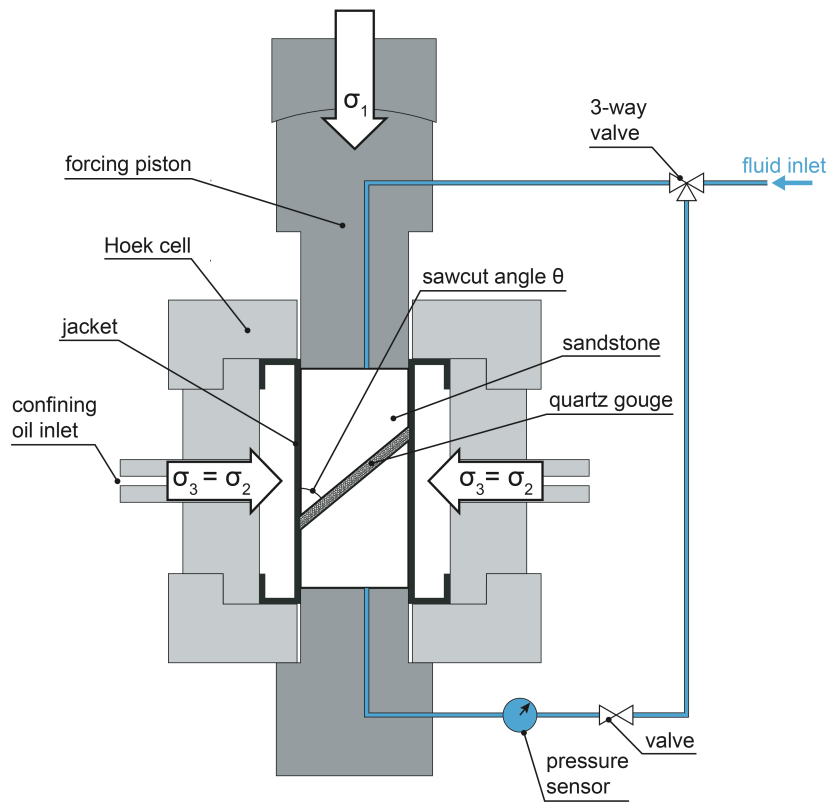
Although this simple approach is reliable, it does not take into account the loading path the fault experienced before reactivation.

Recent studies (e.g., Schorlemmer et al., 2015; Petrucci et al., 2019) show that the faulting type (normal, strike-slip, or reverse), and thus the loading path, influences the seismic behaviour of faults.

In natural faults characterized by thick, incohesive fault zone and highly fractured damage zone, the loading path could not be ruled out.



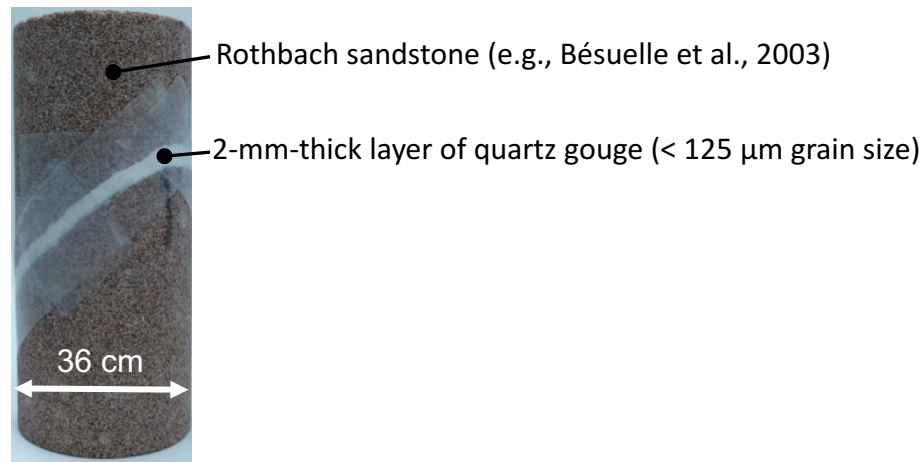
Venere fault (central Italy)  
e.g., Agosta and Aydin, 2006



To investigate the role of the loading path on gouge-bearing fault reactivation, we performed triaxial saw-cut experiments.

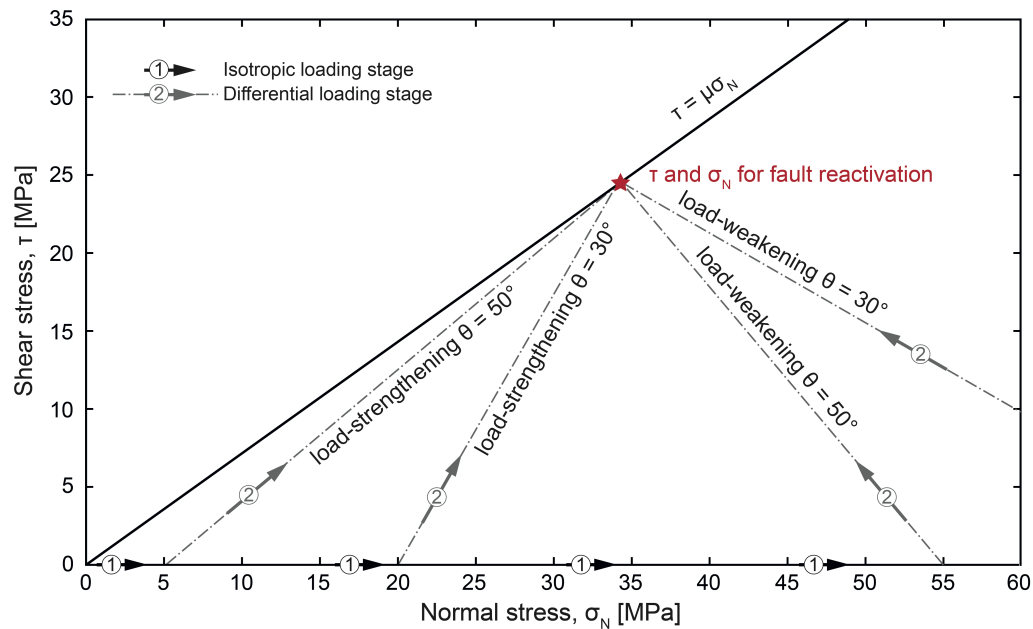
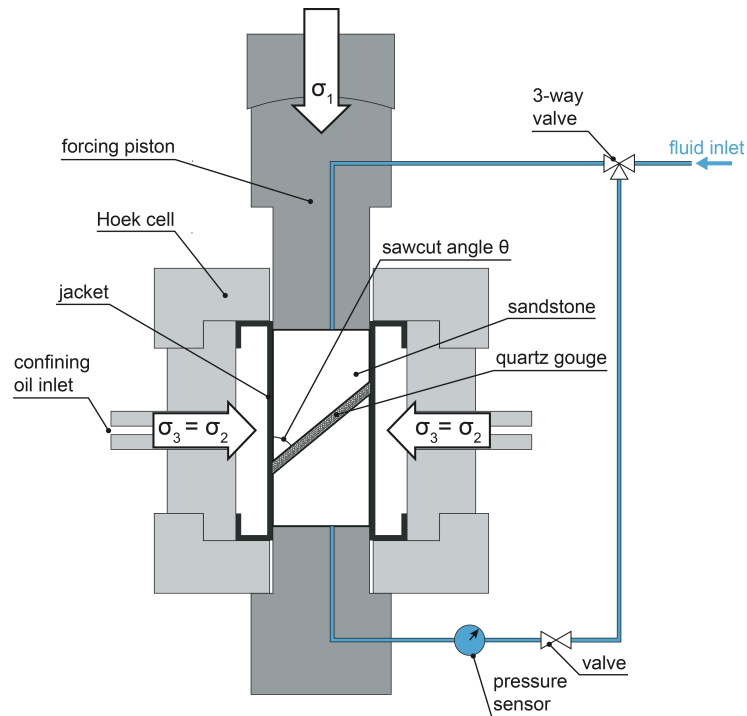
Experimental conditions:

- saw-cut angle: **30°** (well-oriented) and **50°** (misoriented) to the  $\sigma_1$ ;
- **dry** (room humidity) and **water** saturated;
- room temperature.



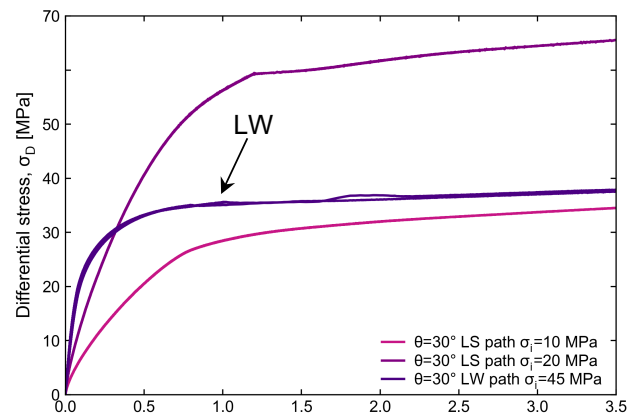
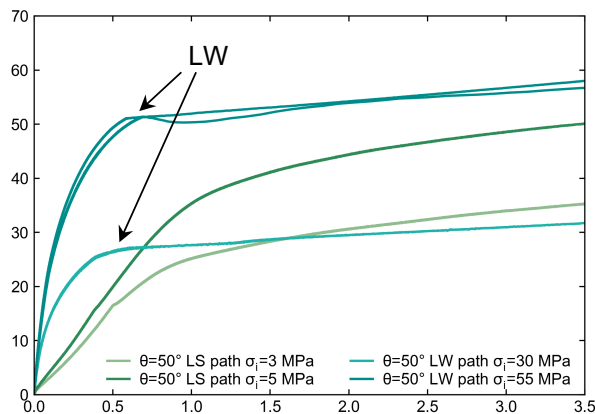
First, the sample is isotropically loaded at a rate of 0.033 MPa/s to the target  $\sigma_3$  and  $P_f$  for load-strengthening experiments and the target  $\sigma_1$  and  $P_f$  for load-weakening experiments.

Once the initial isotropic pressure is achieved, load-strengthening and load-weakening paths were imposed to the sample, respectively increasing the maximum principal stress or decreasing the minimum principal stress at a constant loading/unloading rate  $\delta\sigma_1/\delta t$  or  $\delta\sigma_3/\delta t$ .

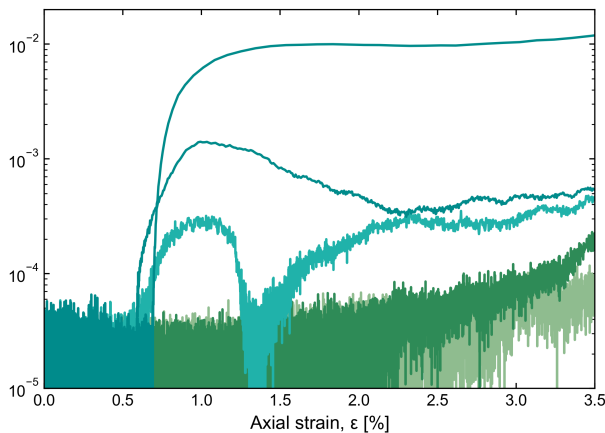
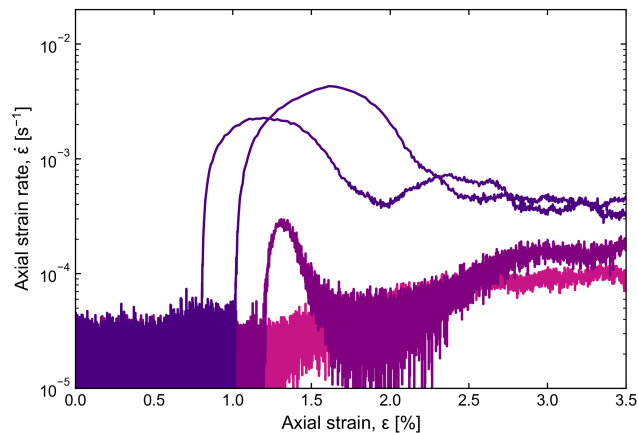


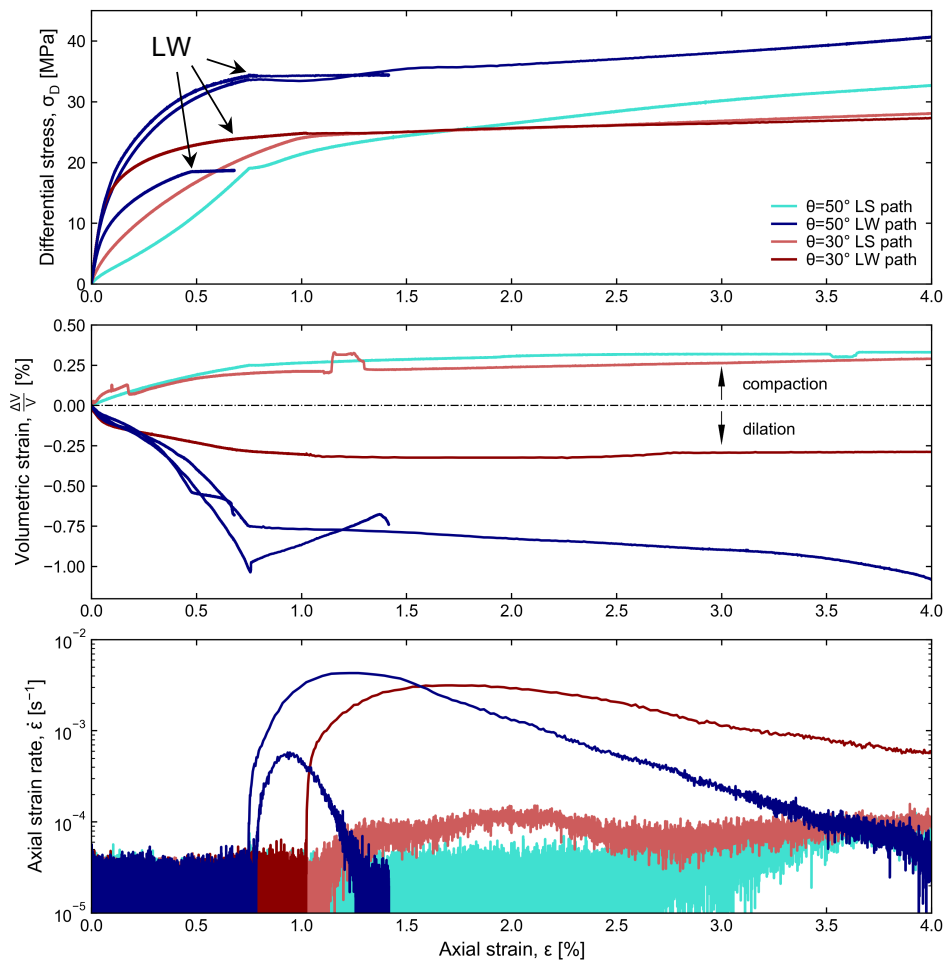
The imposed loading/unloading rates  $\delta\sigma_1/\delta t$  and  $\delta\sigma_3/\delta t$  range between 0.0035 MPa/s and 0.0100 MPa/s to achieve reactivation after a comparable amount of time following different loading paths. The starting isotropic pressure values were chosen in order to achieve reactivation at the same normal stress for faults undergoing load-strengthening (LS) and load-weakening (LW) path. The tested normal stresses at reactivation are  $\sigma_n \approx 12\text{-}35$  MPa. In water-saturated experiments the pore pressure is maintained constant during the experiments in order to result in a pore fluid factor  $\lambda=0.3$ , equal to  $Pf/\sigma_3$  in load-strengthening path and to  $Pf/\sigma_1$  in load-weakening path (e.g., Sibson, 1993).

$\theta$	hydrostatic $P_c$ (MPa)	$Pf$ (MPa)	loading path	rate	$\sigma_n$ at reactivation
50°	5	0	LS	0.010 MPa/s	$\approx 35$ MPa
50°	55	0	LW	0.010 MPa/s	$\approx 35$ MPa
30°	20	0	LS	0.009 MPa/s	$\approx 35$ MPa
50°	30	0	LW	0.006 MPa/s	$\approx 20$ MPa
30°	10	0	LS	0.007 MPa/s	$\approx 20$ MPa
30°	45	0	LW	0.007 MPa/s	$\approx 20$ MPa
30°	42.9	12.9	LW	0.0046 MPa/s	$\approx 12$ MPa
30°	10	3	LS	0.0046 MPa/s	$\approx 12$ MPa
50°	28	8.4	LS	0.0034 MPa/s	$\approx 12$ MPa
50°	48.6	14.6	LW	0.0062 MPa/s	$\approx 22$ MPa
50°	4.3	1.3	LS	0.0062 MPa/s	$\approx 22$ MPa

well-oriented faults  $\theta = 30^\circ$ misoriented faults  $\theta = 50^\circ$ 

- Under load-weakening (LW) path faults appear stiffer than under load-strengthening (LS) path.
- Under LW path, once the differential stress for reactivation is achieved, the fault suddenly accelerates.
- Under LS path, once the differential stress for reactivation is achieved, the fault creeps continuously and smoothly accelerates.

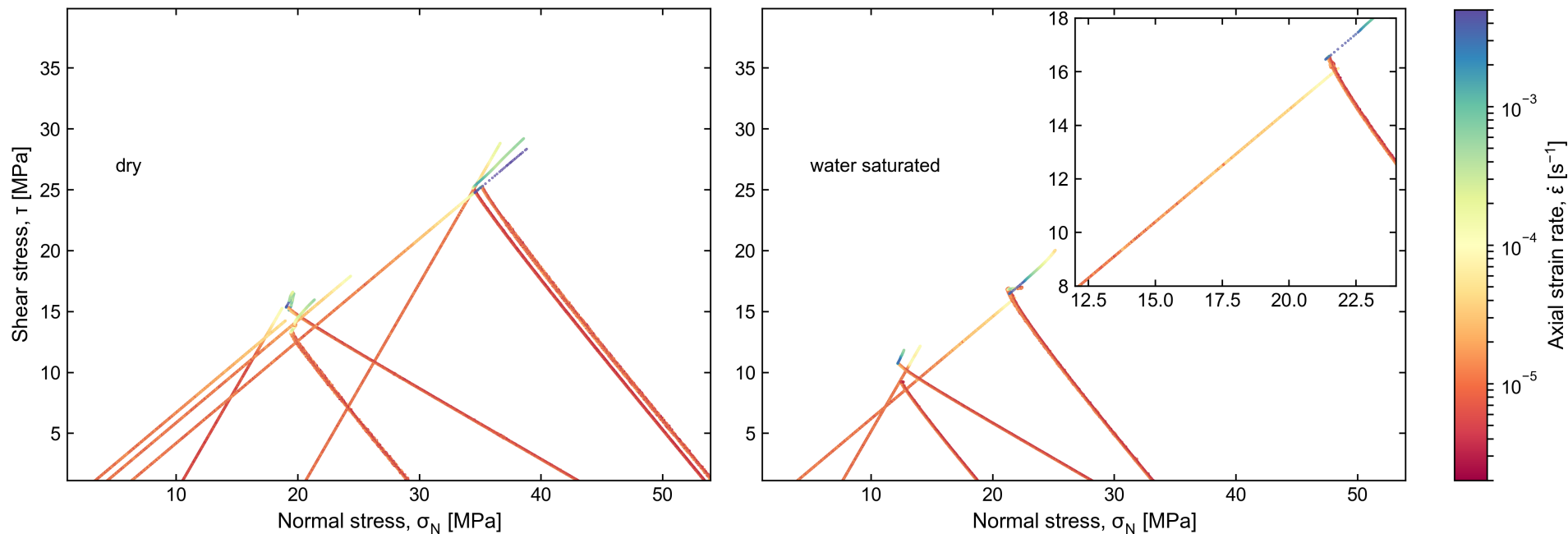




- The results are consistent with the observations collected from dry experiments.
- The fault gouges experiencing load-strengthening path undergo *compaction* until the steady-state differential stress is attained.
- Load-weakening paths result in *dilation* of the fault gouge during the differential stress increase before the abrupt fault acceleration.

Under LS path, the deformation starts to slowly accelerate at lower shear stress, especially for misoriented faults, with the only exception of the experiment conducted at 20 MPa of confining pressure ( $\sigma_3$ ) under dry conditions.

Under LW path, the deformation abruptly accelerates only when the stresses for reactivation are reached, with the only exception of the experiment conducted at  $\sim 19$  MPa of axial stress ( $\sigma_1$ ) under wet conditions.

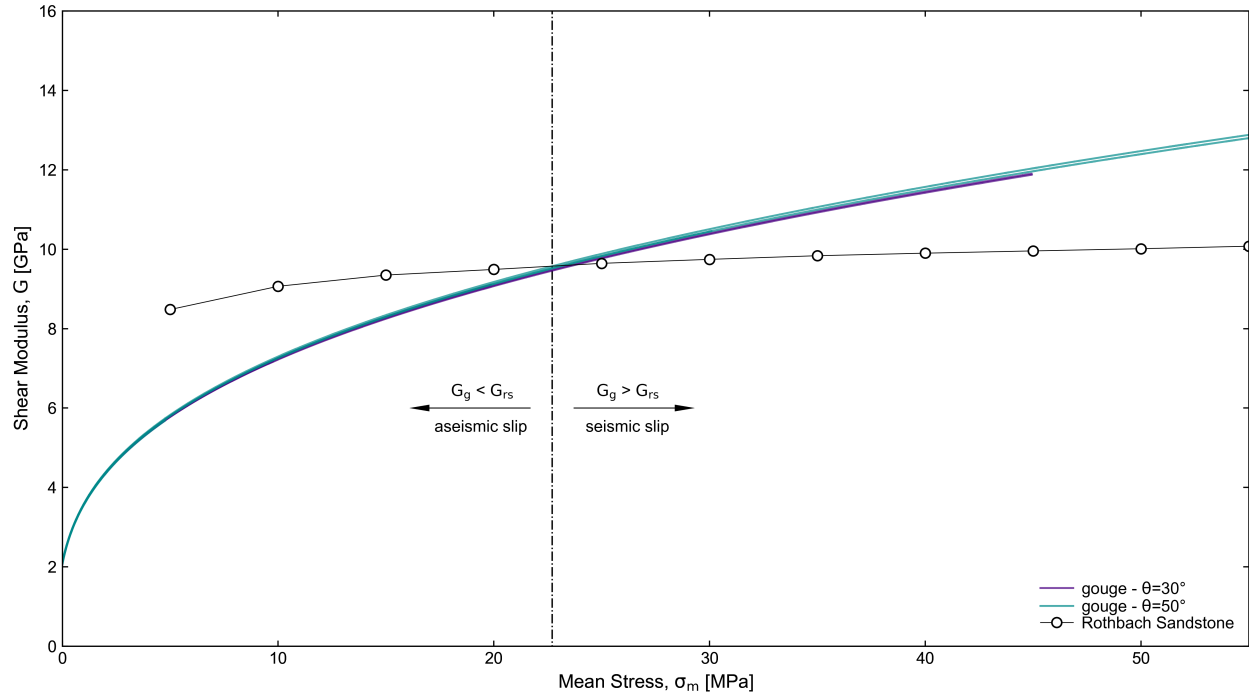




The observed behaviour suggests that faults undergoing LW path (normal faults) tend to slip seismically, whereas faults undergoing LS path (reverse faults) tend to slip aseismically. To understand this behaviour, we investigated the interplay between the elastic properties of the fault gouge and the elastic properties of the surrounding sandstone.

We experimentally measured the shear modulus of the intact sandstone at different confining pressure under dry conditions, through the measurement of ultrasonic wave velocities.

We inferred the evolution of the shear modulus of the fault gouge, from the change in porosity estimated during the isotropic pressure increase under dry conditions.



The change in porosity is calculated from the axial deformation measured during the experiments, assuming that all deformation measured originates from porosity loss in the gouge.

The shear modulus is evaluated as follow (Walton, 1987):

$$G_{gouge} = \frac{3}{5} \sqrt{\frac{C^2(1-\varphi)^2 G^2}{18\pi^2(1-\nu)^2}} \sigma_m$$

- C = 6. coordination number
- G = 44.0 GPa shear modulus of quartz mineral grains
- ν = 0.07 Poisson's ratio of quartz mineral grains
- σ<sub>m</sub> = mean stress

There is a critical value of mean stress σ<sub>m</sub> for higher values of which the fault gouge becomes stiffer than the surrounding sandstone. Consistently, the limit separates seismic to aseismic behaviour.

- Under load-strengthening conditions, faults start to creep at differential stress values that are lower than the values required for the abrupt acceleration of the fault under load weakening conditions.
- Before reactivation, reverse faults undergo *compaction*, whereas normal faults experience *dilation*.
- When reactivating at comparable normal stress, normal faults (load-weakening path) are more prone to slip seismically than reverse fault (load-strengthening path).
- The higher mean stress that normal faults (load-weakening path) experienced before reactivation compacts more efficiently the gouge layer, thus increasing fault stiffness and favouring seismic slip.
- The stress path history of a fault may influence its seismic behaviour, especially if in the presence of a thick gouge-bearing fault zone.

- Thank you for reading! -

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