Radiative feedbacks in 1D-RCE

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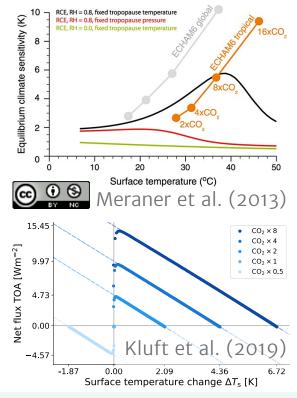




The state

State-dependent climate feedback

- Quantifying the temperature dependence of the climate feedback parameter
- Meraner et al. (2013) find a strong dependence for RCE-like model (black curve)
- Romps (2020) confirms the qualitative behaviour in high-resolution simulations
- Only a small state-dependence of the climate feedback in Kluft et al. (2019)

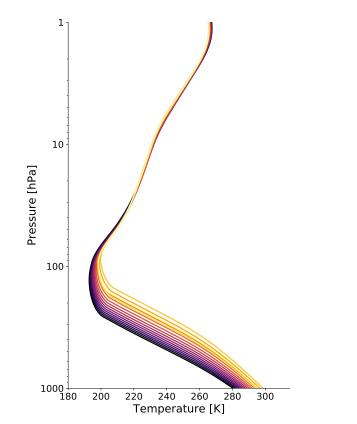




How to change the climate?

How to adjust the state?

- Adjusting the boundary conditions to sample surface temperatures:
 - Solar constant
 Affects stratospheric temperature
 - Relative humidity
 Reduces water vapour feedback
 - Fixed surface temperature
 Unable to quantify radiative forcing
- Introduce a surface heat sink



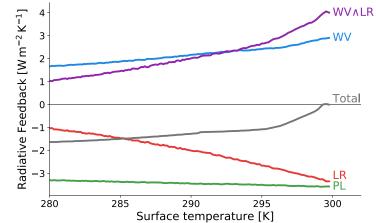
Temperature profiles for different values of the surface heat sink





Decomposed feedbacks

- The 1D-RCE model Konrad is run with T_s between 280 and 300 K
- We find a robust state-dependence of the total climate feedback
- Decompose feedbacks by turning their mechanism on/off
- The increase is driven by a strengthening of the water-vapour feedback [in agreement with Meraner et al. (2013)]

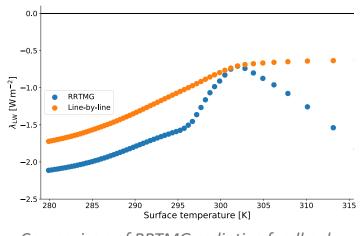






Is it a feature or a bug?

- RRTMG is optimised for a reasonable set of atmospheric conditions
- The tropical model configuration reaches these boundaries
 - Water vapour amount (80% RH)
 - Tropopause height above 100 hPa
- Koll and Cronin (2018) find a closing of the atmospheric window for surface temperatures around 300 K



Comparison of RRTMG radiative feedbacks with offline line-by-line simulations



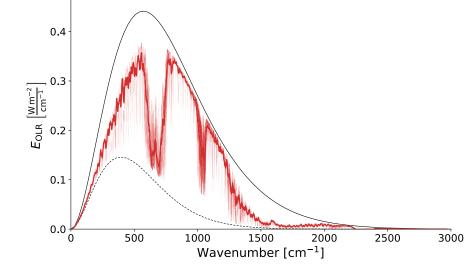


High-resolution* modelling

*radiative transfer

Line-by-line longwave radiation

- Replace RRTMG LW with the line-by-line model ARTS
- 30,000 frequencies between
 1 and 3000 cm⁻¹
- 128 pressure levels
- Shift the ozone profile together with the expanding troposphere
- Keep RRTMG for the shortwave component (possible inconsistencies)





Preliminary results

- 1D-RCE with on-the-fly line-by-line radiative transfer
- Robust state-dependence of the total climate feedback
- Weaker (or no) temperature dependence above 300 K
 surface temperature [in agreement with a conceptual by Ingram (2010)]
- Be careful when using fast radiation schemes in extreme climates

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