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Spatial observations of large eddies and cross-canopy coupling with joint fiber-optic distributed sensing and eddy covariance flux measurements

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Intro: Poggi et al. (2004)

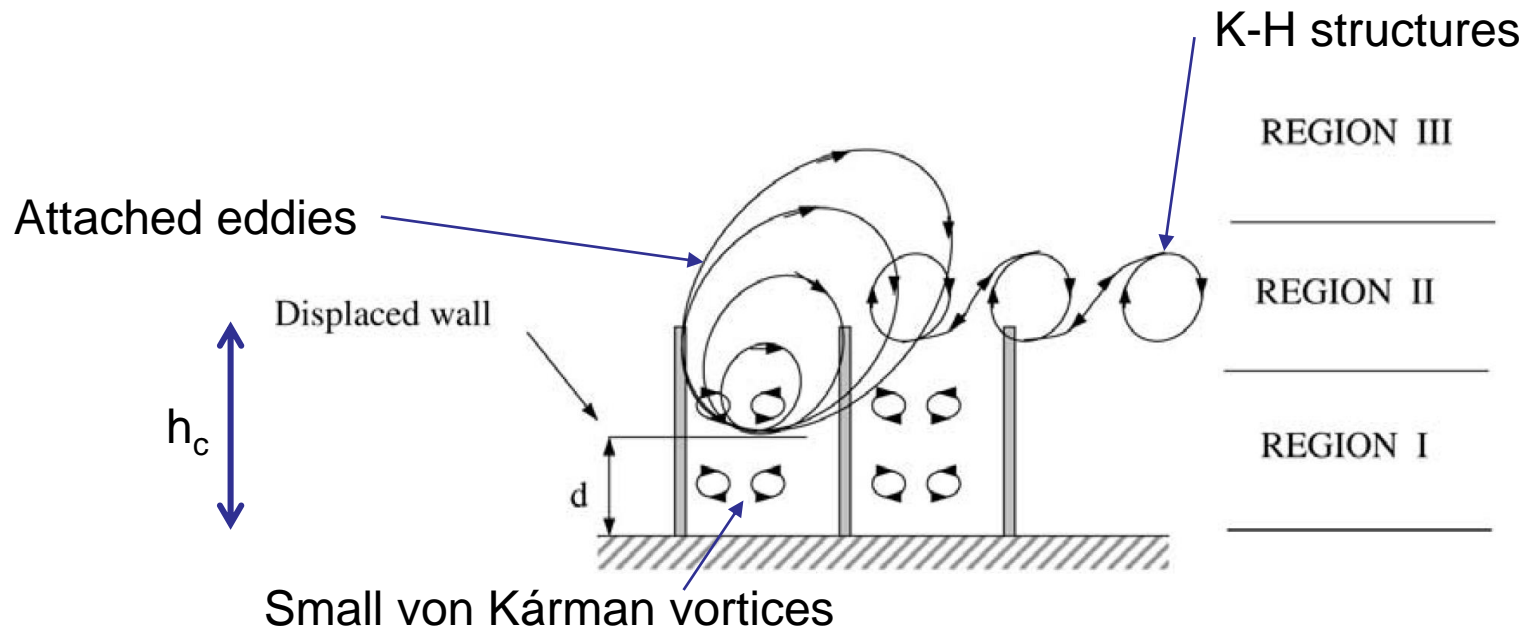
Flows above canopies can be divided into three structures:
attached eddies (scale with $z-d$), K-H structures (scale with h_c) and von Kármán vortices

Three regions

Region I: subcanopy air space, von Kármán vortices

Region II: flows just above canopies, K-H structures dominate (roughness sublayer)

Region III: flows above canopies, attached eddies dominate (inertial sublayer)



Intro: Thomas and Foken (2007)

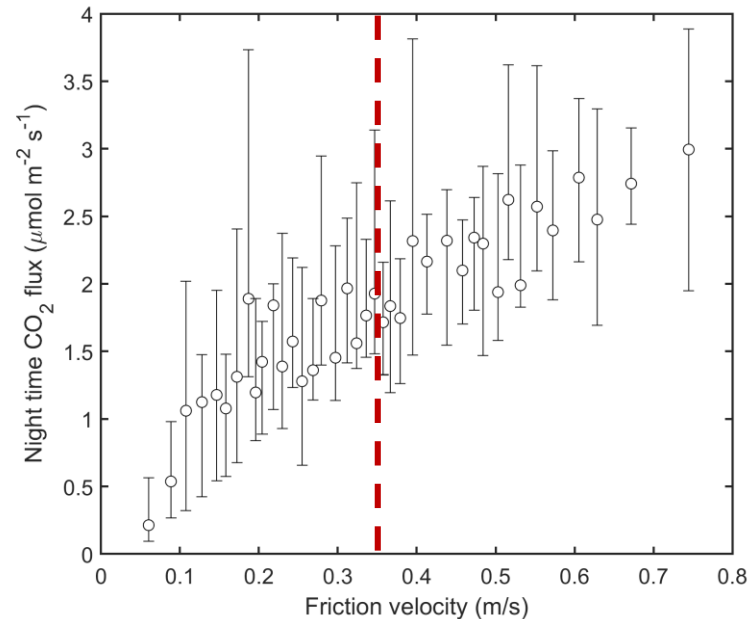
- Four scalar exchange regimes above forests were identified
 - Wa (wave)
 - Waves dominate the above canopy flow. Weak vertical transport
 - Dc (decoupled canopy)
 - Turbulence quiescent. Weak vertical transport
 - Ds (decoupled subcanopy)
 - Turbulent mixing above canopy, but not at subcanopy air space. Turbulent vertical transport misses the subcanopy signal (e.g. forest floor respiration)
 - Cs (Coupled subcanopy by sweeps)
 - Occasional downdrafts (i.e. sweeps) penetrate to the subcanopy air space. Turbulent vertical transport misses part of the subcanopy signal
 - C (fully coupled canopy)
 - Above canopy flow fully coupled with the subcanopy



Intro: CO₂ flux underestimation

- CO₂ fluxes measured above canopy are underestimated at night time when friction velocity (u_*) is low
- Introduces biases to ecosystem carbon balance estimates
 - ⇒ measurements from periods with low u_* are typically discarded
- Clear friction velocity threshold is difficult to find at some sites

Dependence of night time CO₂ fluxes measured above canopy on friction velocity at Hyytiälä pine forest site



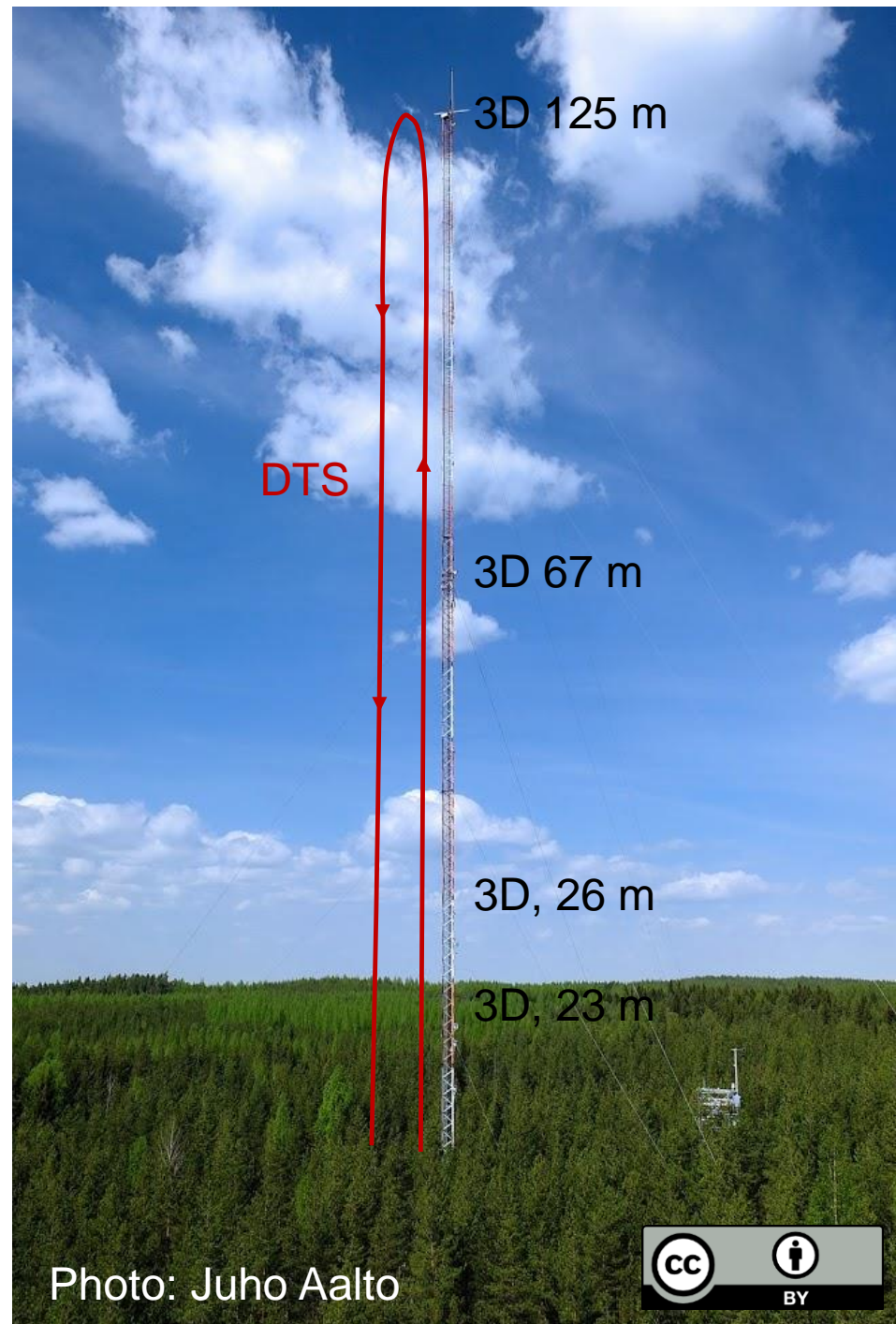
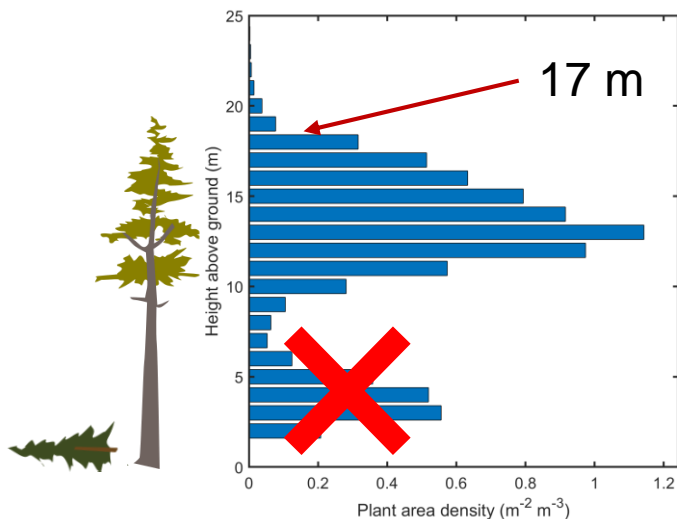
Objective

- Objective of this study is twofold
 - 1) Identify the height-time structure of coherent flow structures above forest canopy from detailed temperature profiles
 - 2) Assess their role in cross-canopy coupling at night when CO₂ fluxes are underestimated



Hyytiälä, SMEAR II

- 125 m tall mast in a boreal pine forest in southern Finland
- Distributed temperature sensing (DTS)
 - Temperature data with 1 Hz and approx. 12 cm resolution along the fibre cable secured to the mast
- => continuous temperature profiles along the mast!
- 3D sonic anemometers at 1 m, 4 m, 23 m, 26 m 67 m and 125 m
- Measurements between 25.5. and 9.7.2019
- All-sided LAI 8 m² m⁻² in year 2015



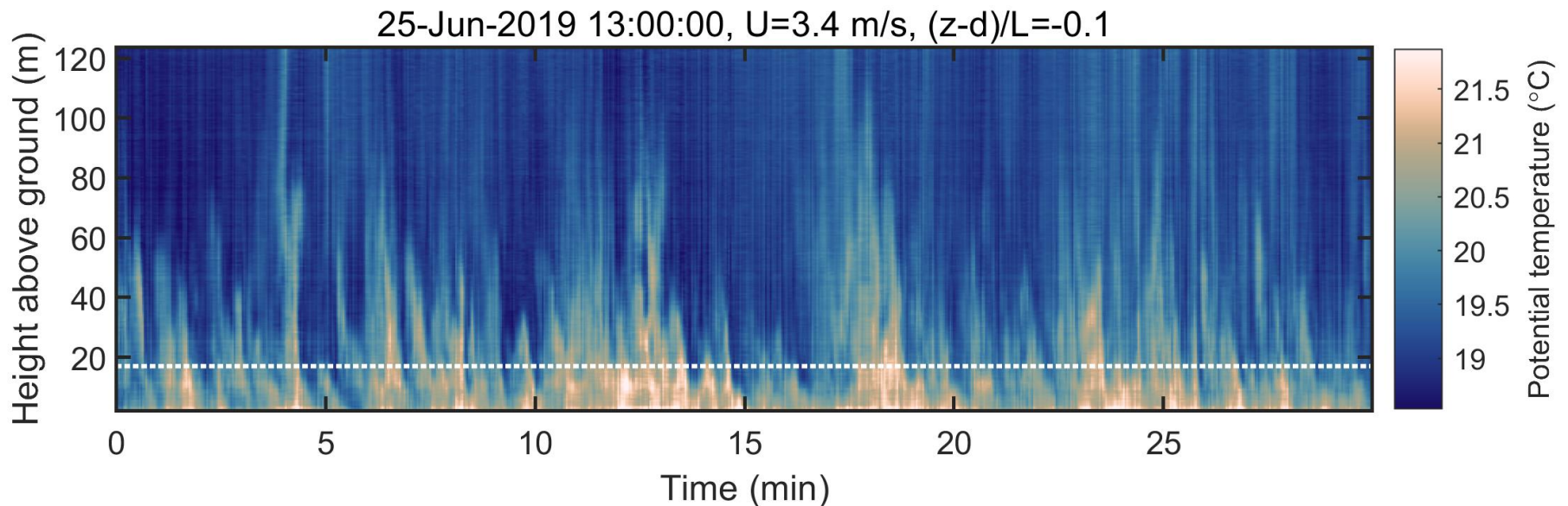
Data processing

- 3D sonic anemometers and EC:
 - Despiking
 - Coordinate rotation following Wilczak et al. (2001)
 - Detrending
 - Cross-covariance maximisation for gas fluxes
- DTS
 - Calibrating single-ended measurements (25.5.-3.6.) following Hausner et al. (2011) and double-ended (3.6.-9.7.) following van de Giesen et al. (2012)
 - DTS measurements contain significant noise component => denoising using proper orthogonal decomposition similarly as Epps and Krivitzky (2019)
- Throughout the study 10-min averaging period was used



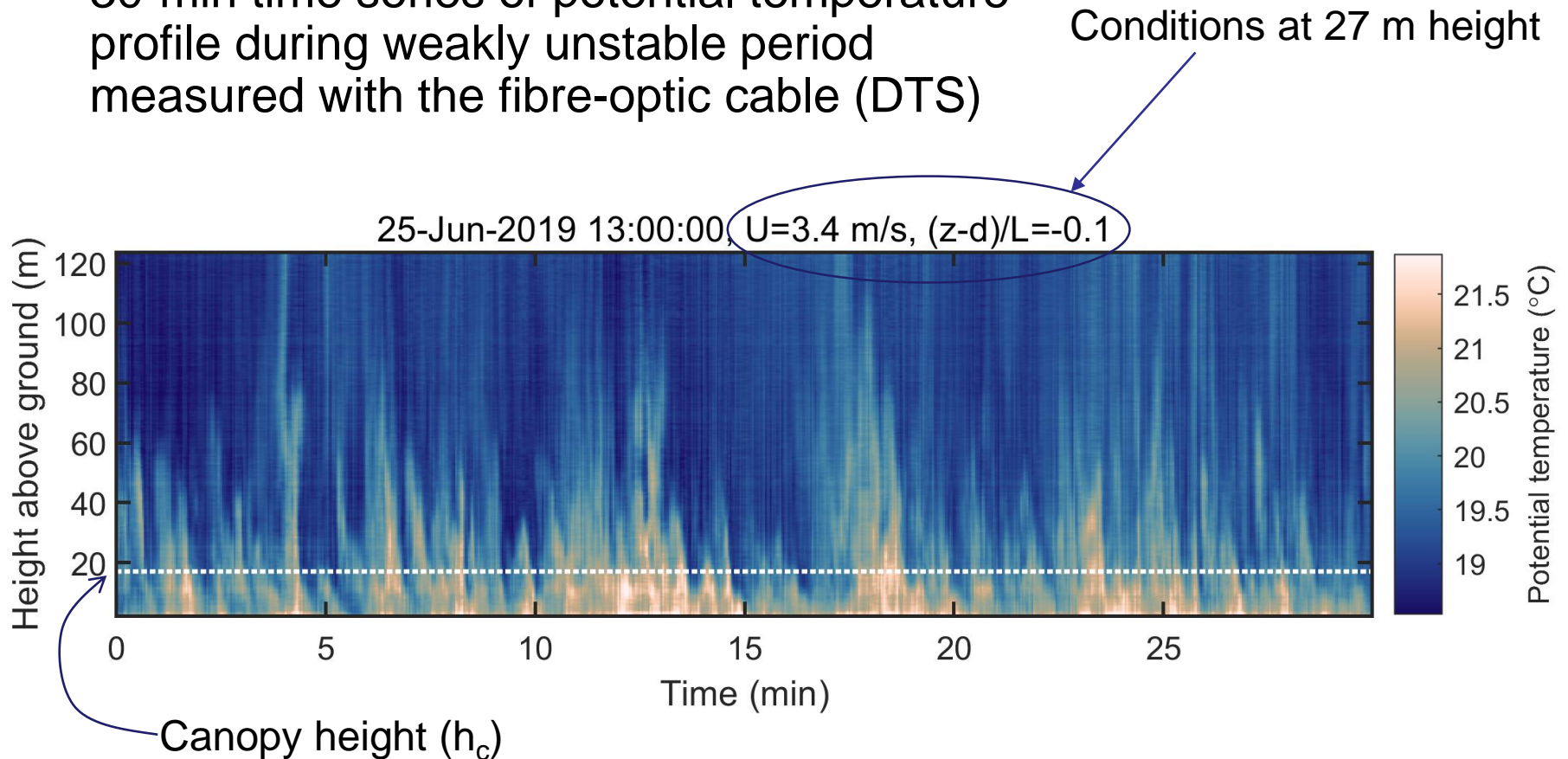
Results: daytime example

- 30-min time series of potential temperature profile during weakly unstable period measured with the fibre-optic cable (DTS)



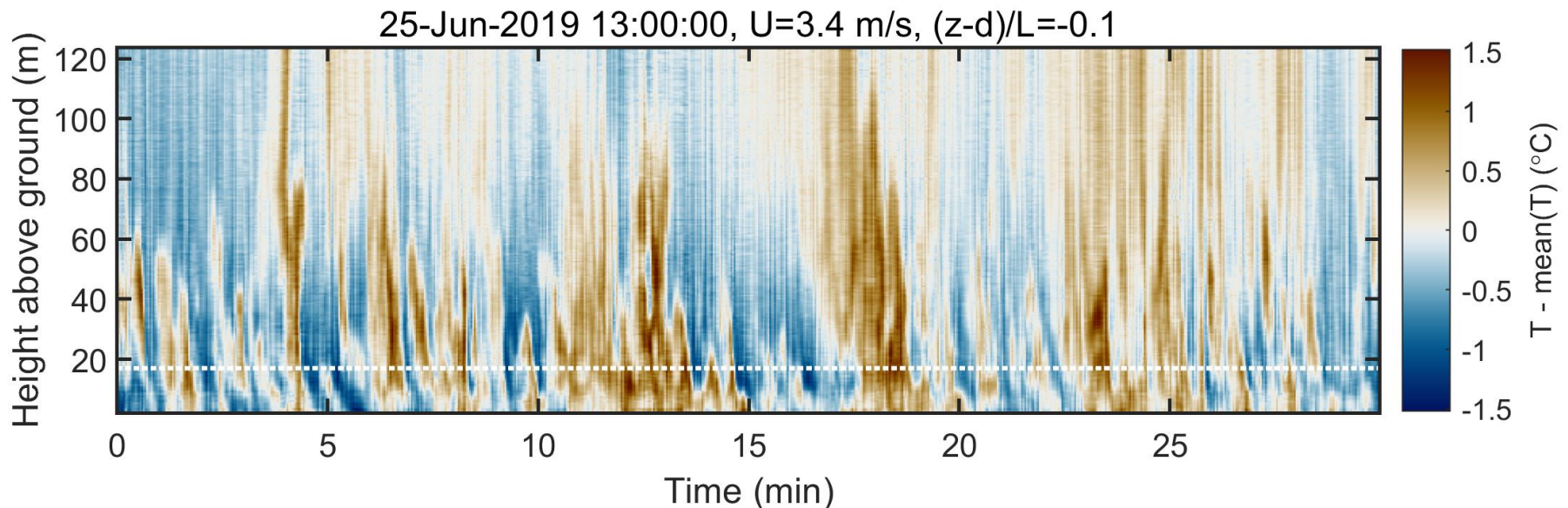
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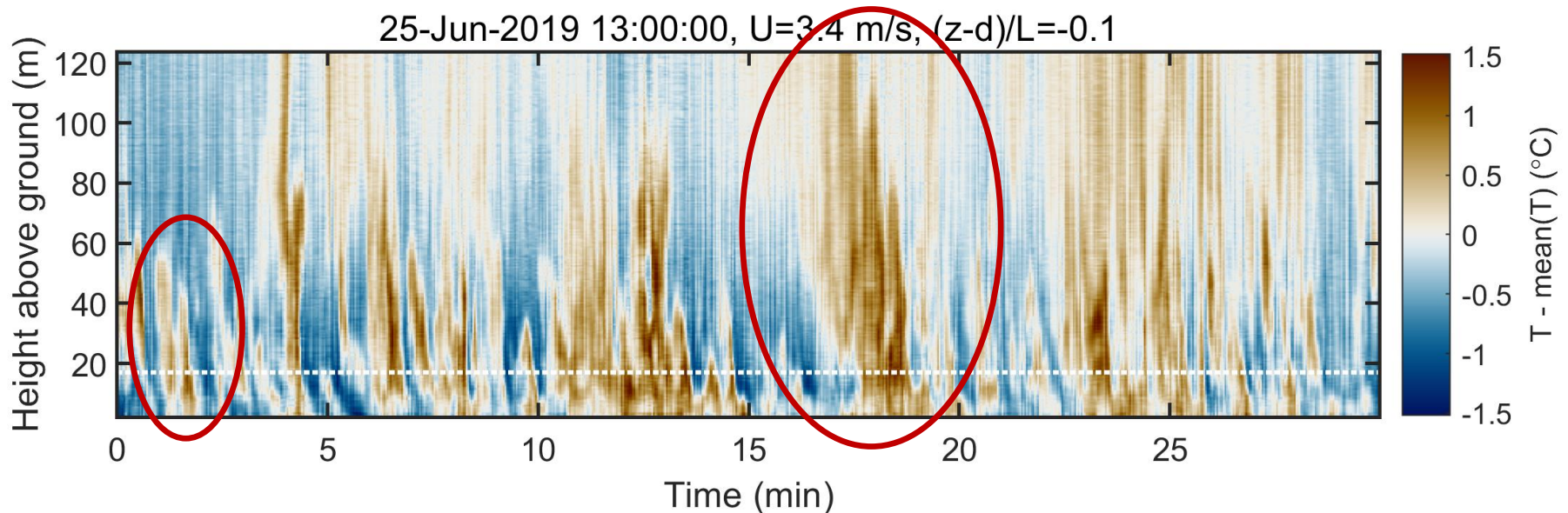
Results: daytime example

- Profile of temperature fluctuations around temporal mean at each height reveal turbulent structures better
- Negative excursions (blue) match with downward and positive (brown) with upward air movement (not shown)



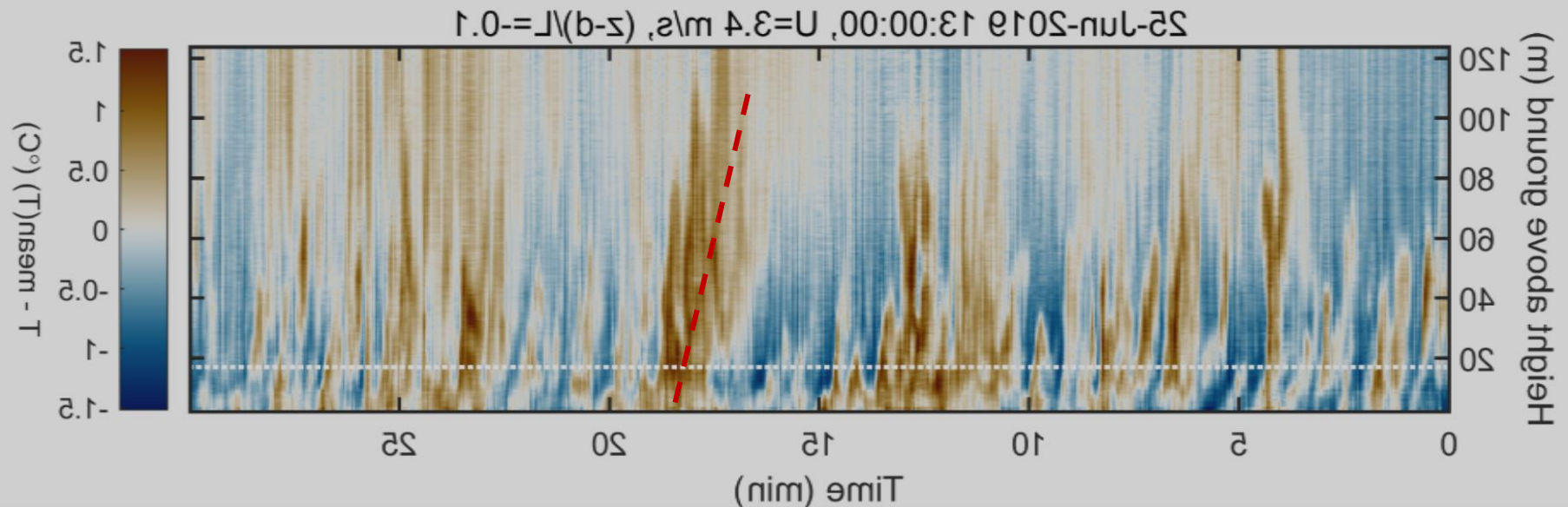
Results: daytime example

- Different sizes of structures can be identified
- By applying Taylor's frozen turbulence hypothesis time can be converted to horizontal distance and hence eddy size



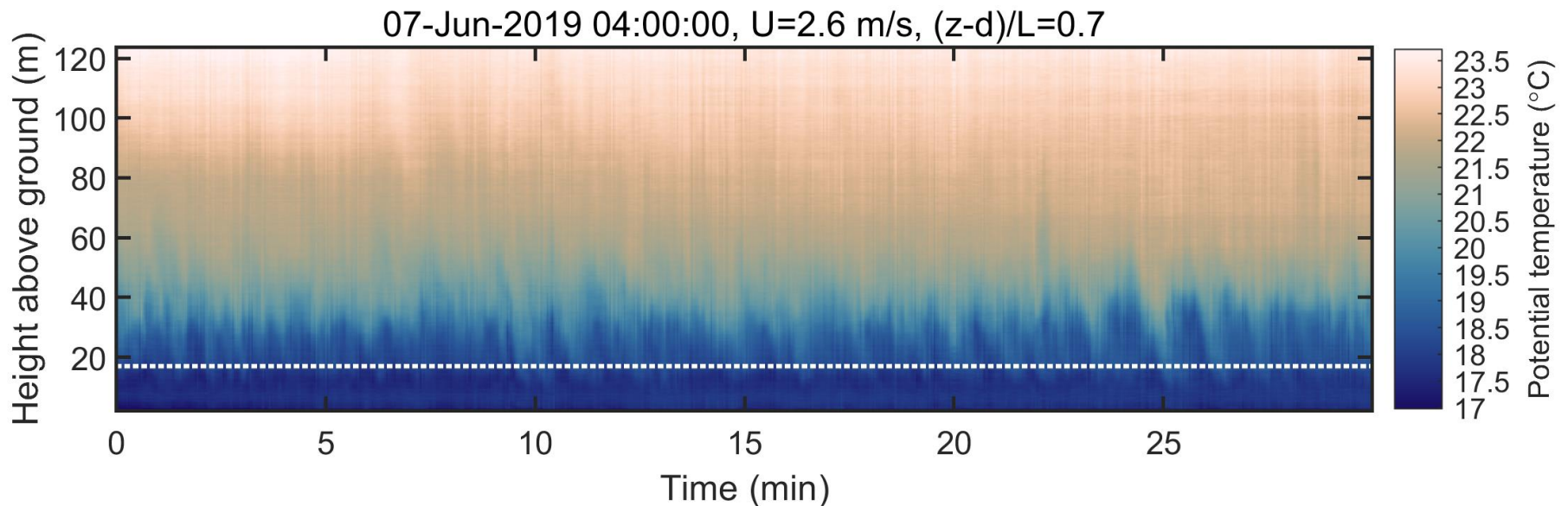
Results: daytime example

- By inverting the time axis we observe the eddies as they are advected past the measurement mast
- The eddies are inclined as they are advected by the mean flow and hence they are observed first at the highest levels



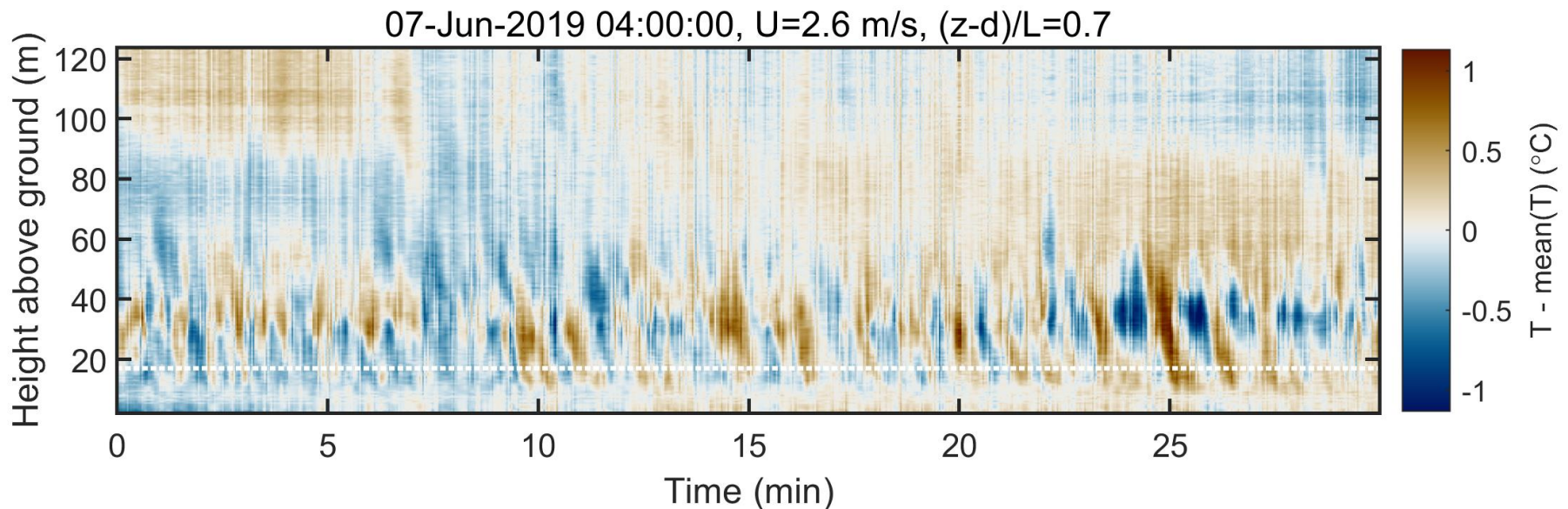
Results: night time example

- Stably stratified situation
- Coherent eddies between 10 m and 60 m
- No strong fluctuations above 60 m or below 10 m => negligible mixing



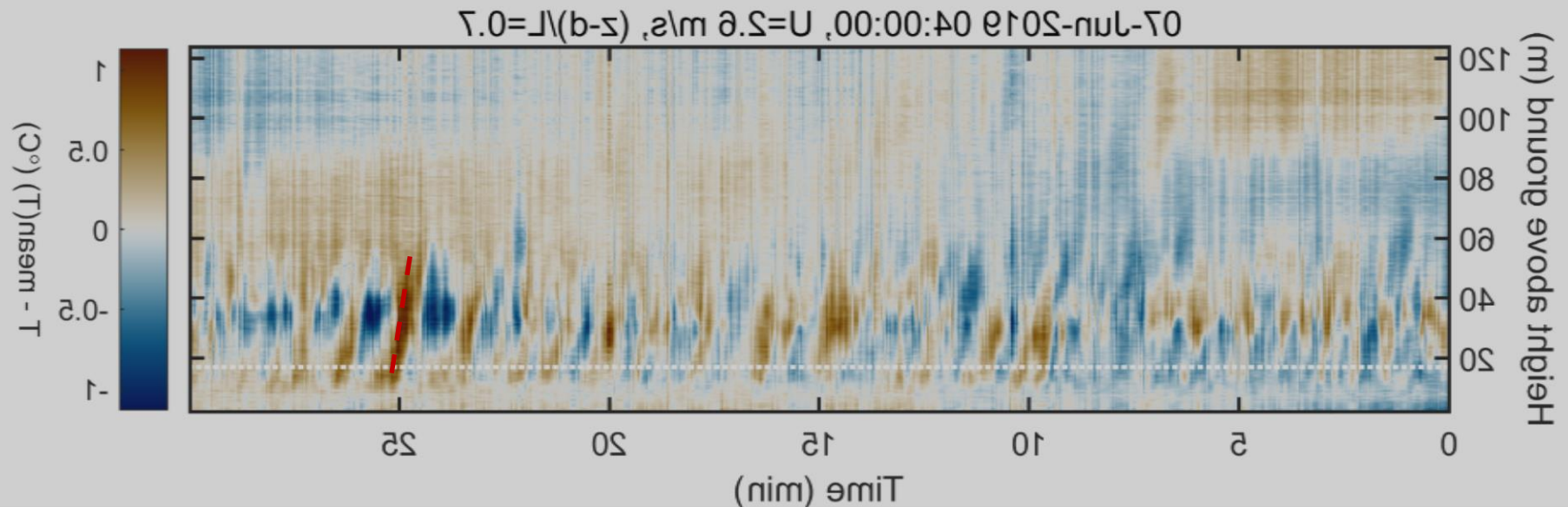
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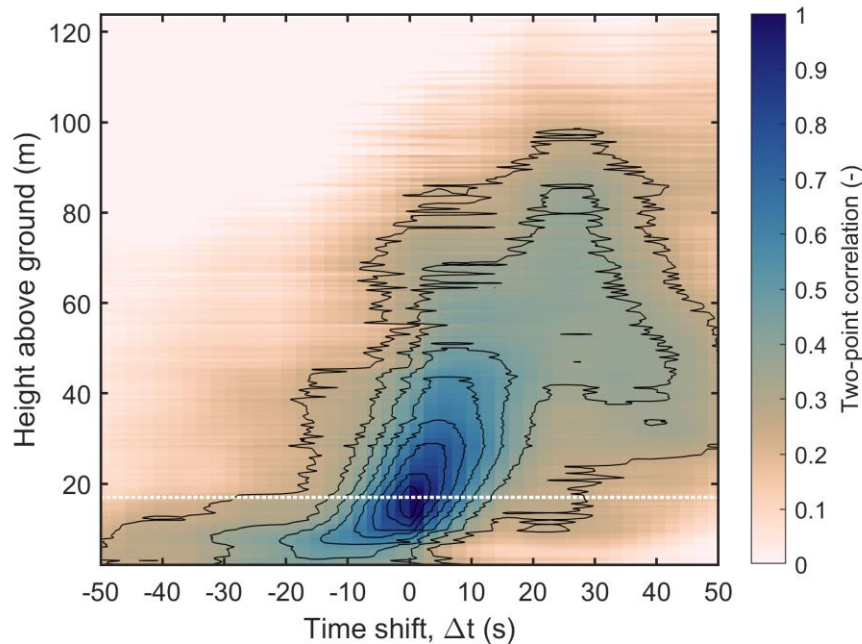


Two-point correlation

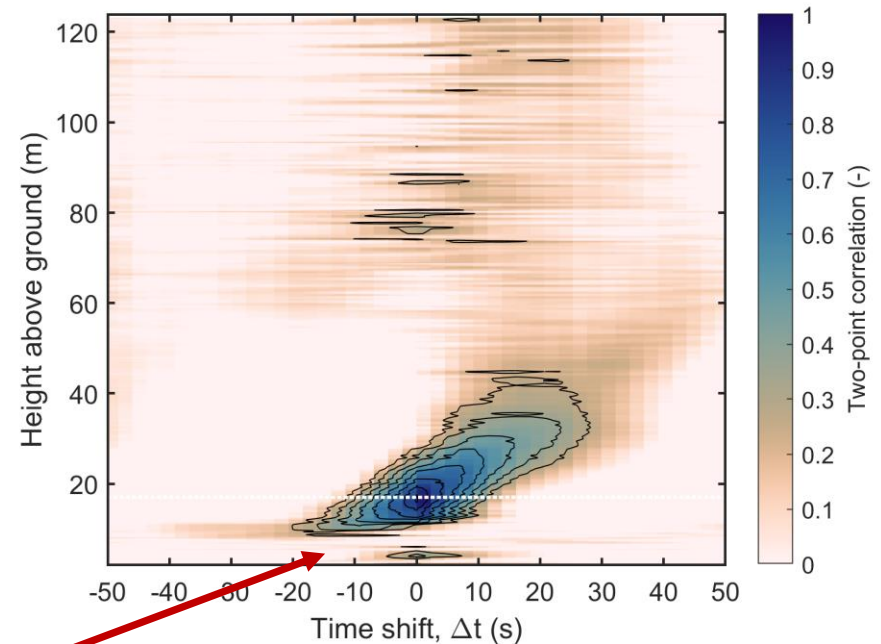
- Calculating two-point correlations for each averaging period with the following equation:

$$R(z, \Delta t) = \frac{T(h_c, t)T(z, t + \Delta t)}{\sigma_T(h_c)\sigma_T(z)}$$

Daytime example



Night time example



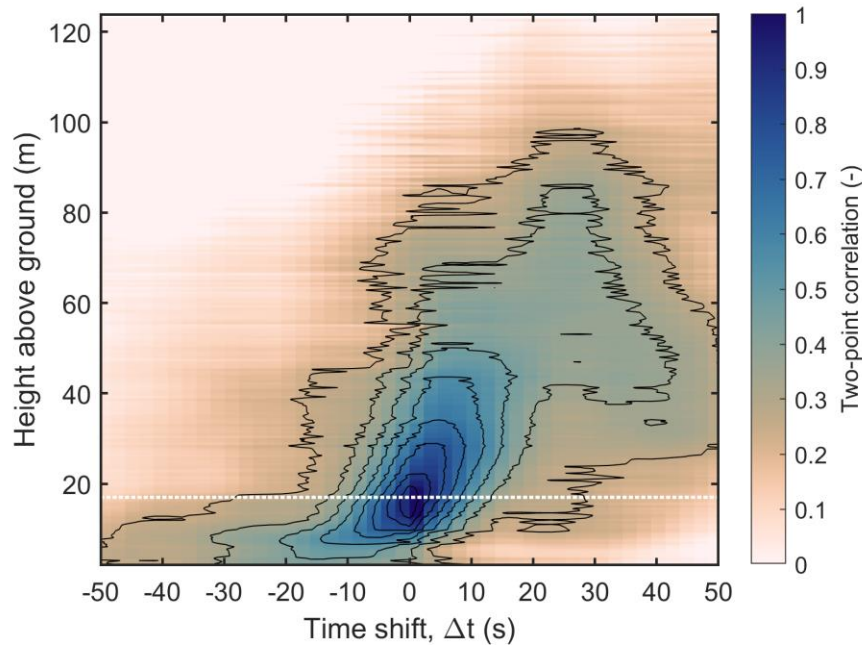
Note that the large eddies are decoupled from the forest floor



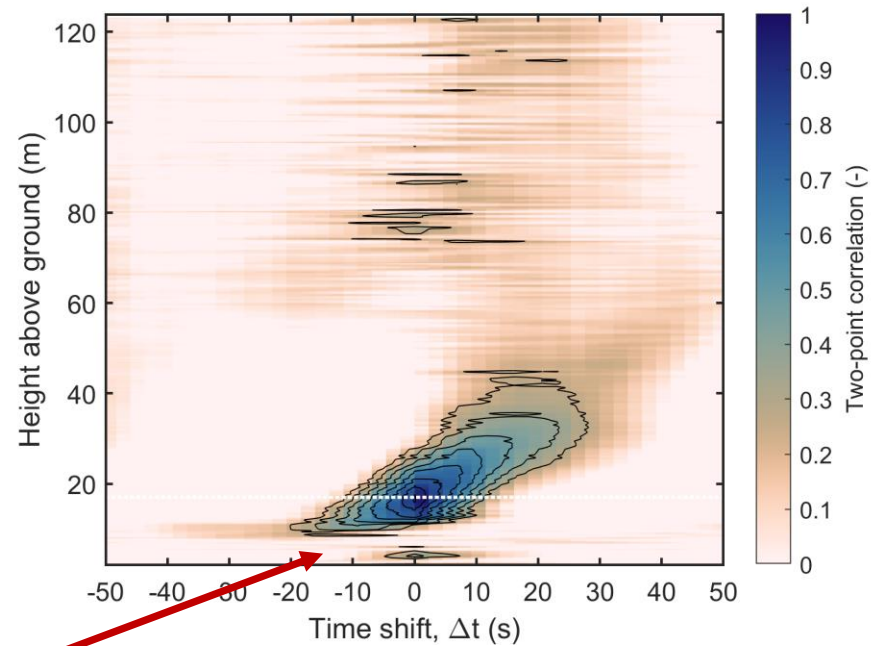
Two-point correlation

- Two-point correlations can be approximated to show the vertical and temporal extend of those flow structures that dominate the T variability at the reference height (here canopy height h_c)

Daytime example



Night time example



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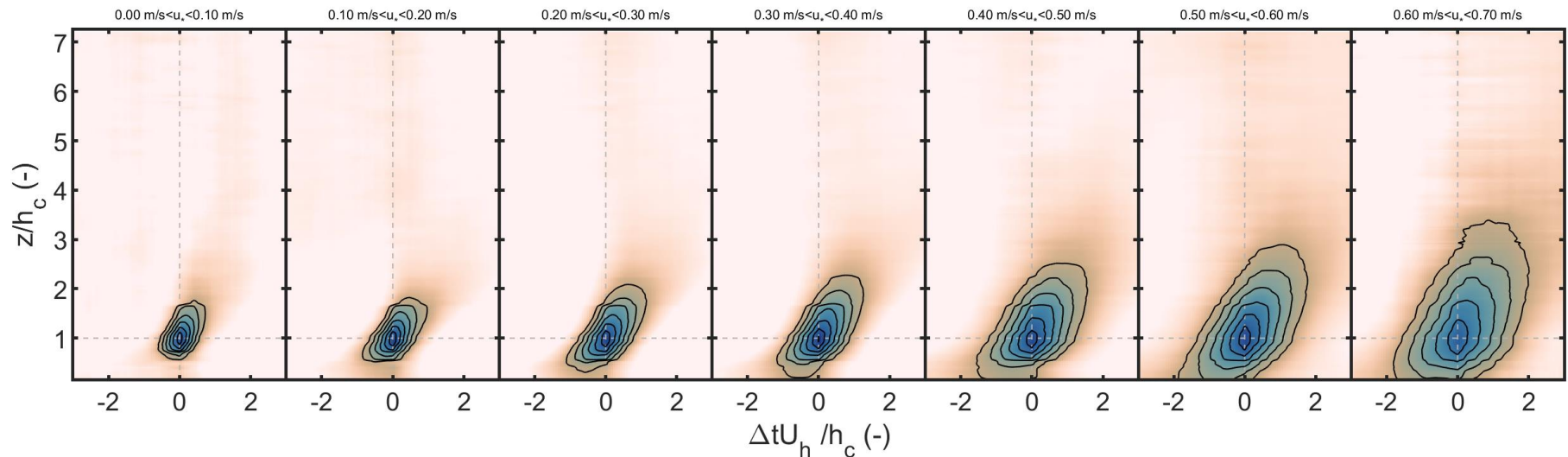
Canopy eddies and decoupling

- Calculation of two-point correlations from the DTS data for each averaging period during the 1.5 month long measurement campaign
- Converting the x-axis from temporal (time shift) to spatial (distance) scale using Taylor's frozen turbulence hypothesis.
- Ensemble averaging the calculated two-point correlation matrices for different conditions



Canopy eddies and decoupling

- Ensemble averaged two-point correlations in friction velocity bins
- Only night time data



Friction velocity increases

U_h = eddy advection velocity (wind speed at canopy height)

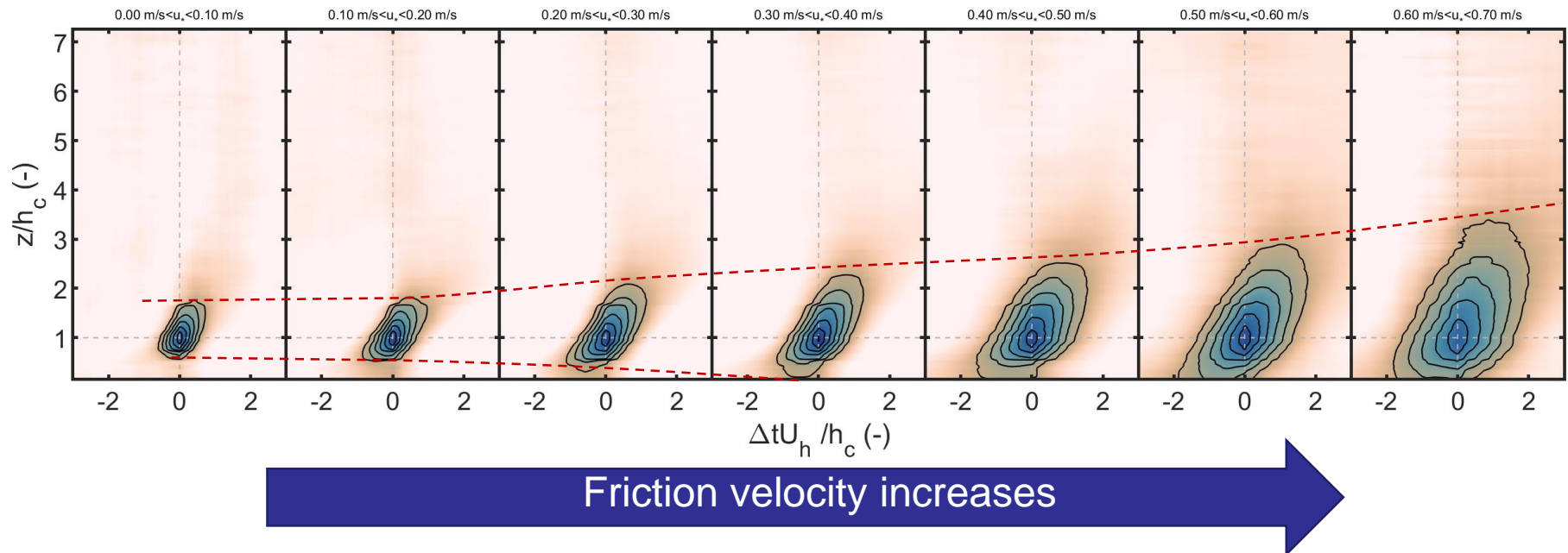


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Canopy eddies and decoupling

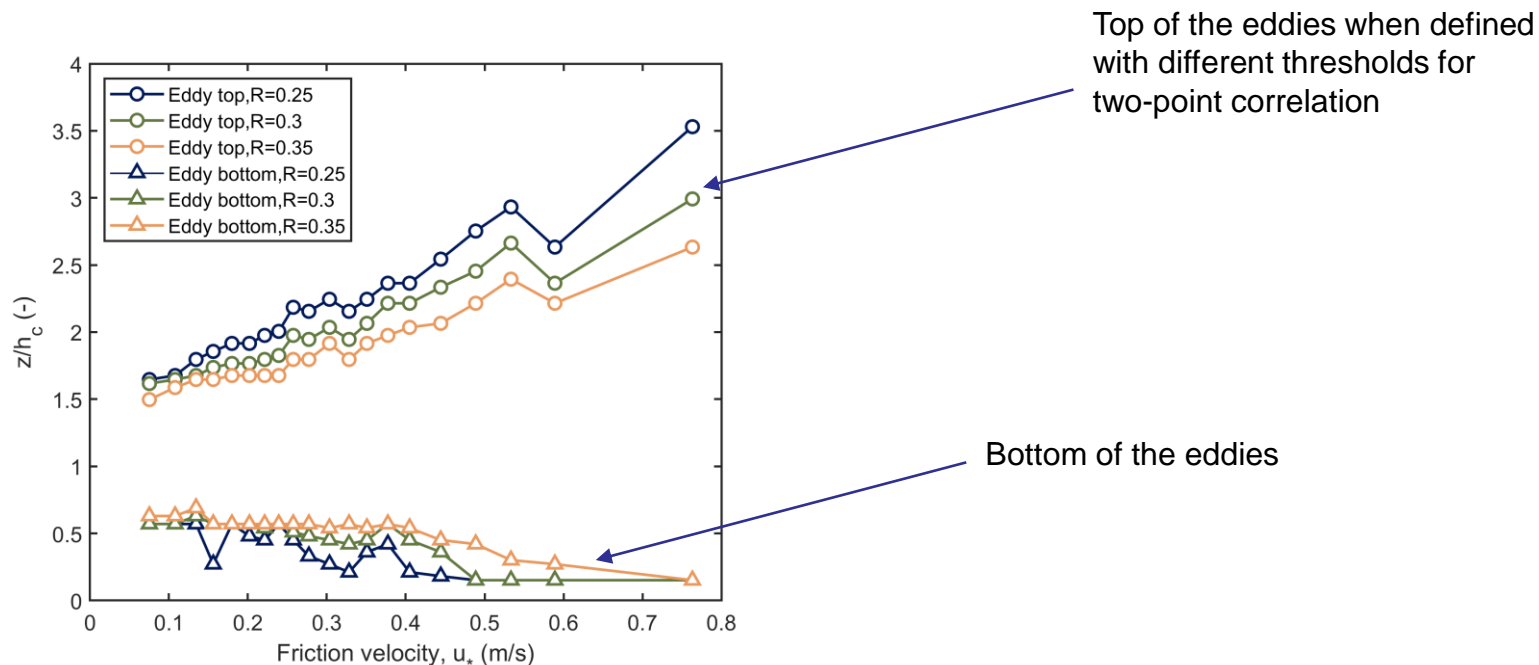
- The height of the eddies dominating T variability at canopy height increase from $2h_c$ to $3h_c$ as u_* increases
- At low u_* the eddies are decoupled from the surface, but couple with the surface as u_* increases



U_h = eddy advection velocity (wind speed at canopy height)

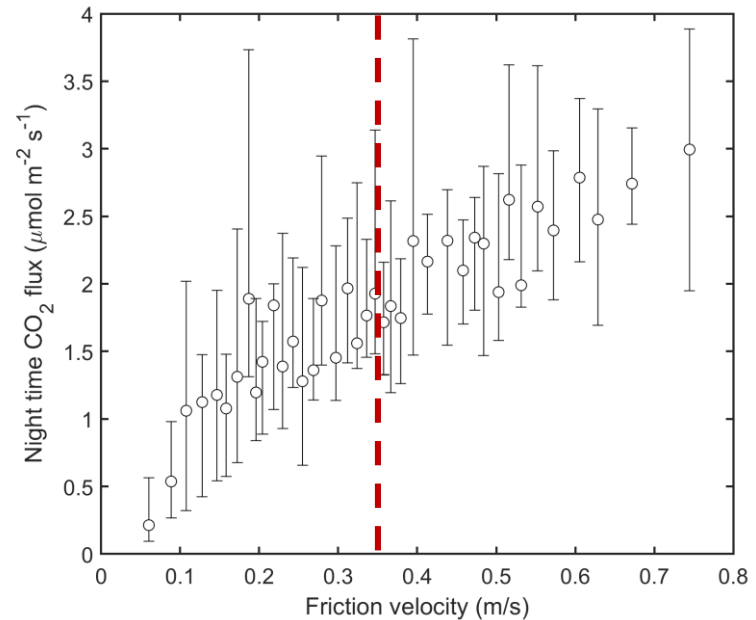
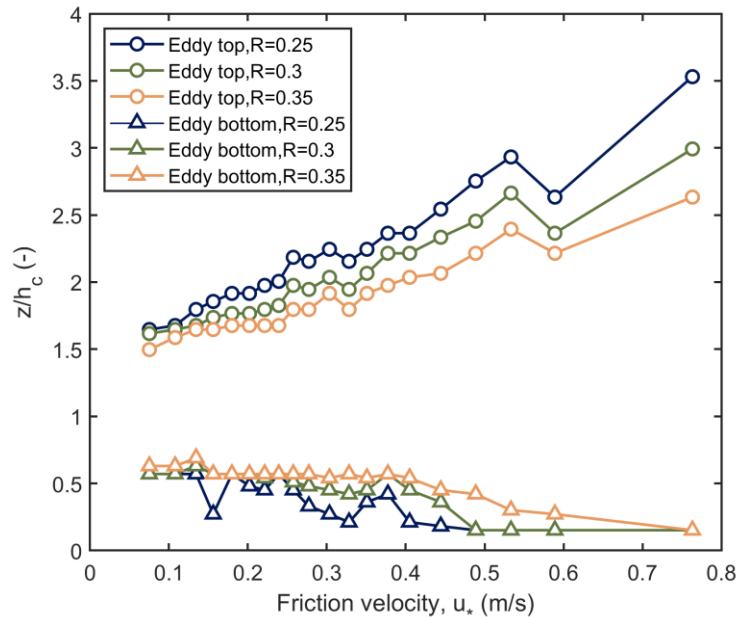
Canopy eddies and decoupling

- Analysing the bottom and top of the canopy eddies at night in different friction velocity bins
- Note that the eddies couple with the surface when u_* around 0.3...0.5 m/s



Canopy eddies and decoupling

- Note that when canopy eddies couple with the surface, CO₂ fluxes are almost independent of u_*



Conclusions

- Coherent eddies were identified in the vertical DTS measurements
- Different eddy sizes were observed in different mixing conditions
 - Large thermal eddies in daytime unstable situation
 - Eddies that scale with h_c and produced by wind shear at the canopy top at night time
- At night time canopy eddies were decoupled from the forest floor in low u^* conditions, but coupled with the surface once u^* increased
- The u^* limit at which the eddies coupled with the surface was similar to the u^* threshold used in filtering CO_2 fluxes
 - ⇒ Is the coupling of large coherent eddies with the ground a necessary condition for observing unbiased CO_2 fluxes above canopies?
- Here fluctuations in T were used to estimate large canopy eddies and to study the cross-canopy coupling. This approach circumvents the critique of Freundorfer et al. (2019) on using above- and below-canopy σ_w on analyzing cross-canopy coupling
- This analysis will be complemented in the future with more in-depth analysis of coupling between below- and above-canopy wind data





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