



Impactful Tibetan Plateau Vortices: structure, lifecycle and environmental conditions

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Tibetan Plateau Vortices (TPVs)

- TPVs one of the major precipitation producing/triggering systems on the Tibetan Plateau
- Mainly present at 500 hPa level
- Spatial scales

horizontal: Meso-α-scale, 400-800 km vertical: 2-3 km

- TPVs can trigger extreme precipitation and severe flooding east of the TP
- July 2008: TPV caused heavy rain over large parts of China







Methodology

- 500 hPa relative vorticity field (spectrally filtered: T40-100)
- Objective feature-tracking algorithm TRACK (Hodges et al., 1994, 1995, 1999)
- ERA-Interim and NCEP-CFSR
- UK Met Office model (HadGEM3-GA3; 25 km; 5 ensemble member)
- 6-hourly resolution
- Filtering: originate on Tibetan Plateau (3000 m contour)
 - persistent ≥ 1 day
 - geopotential minimum at one time step (at least)

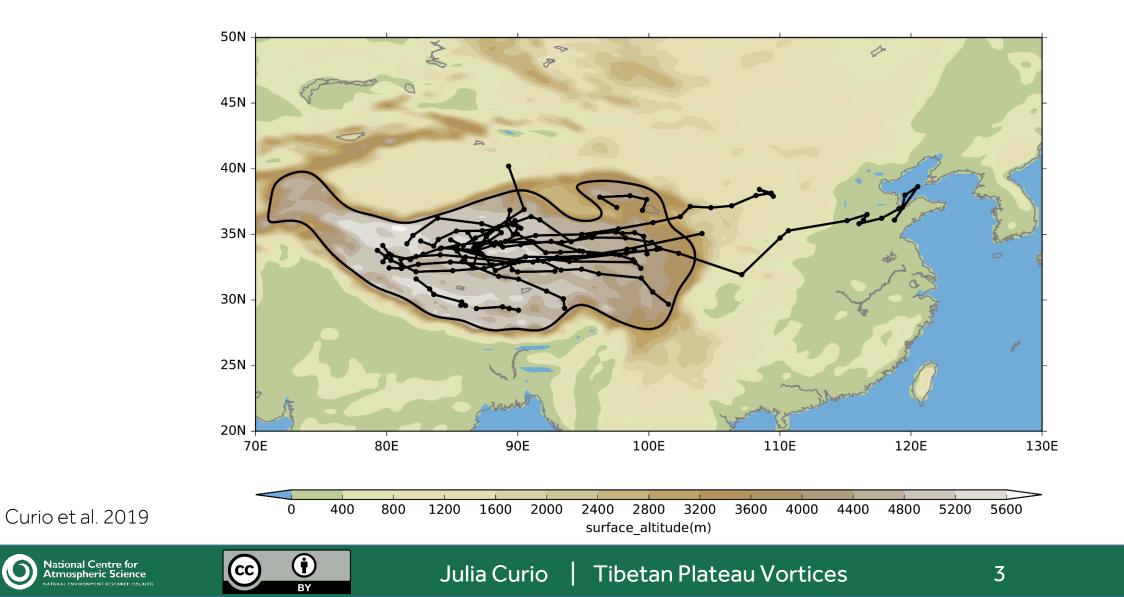
Curio et al. 2019: Climatology of Tibetan Plateau Vortices in reanalysis data and a high-resolution global climate model. Journal of Climate, **32**, 1933–1950, https://doi.org/10.1175/JCLI-D-18-0021.1.



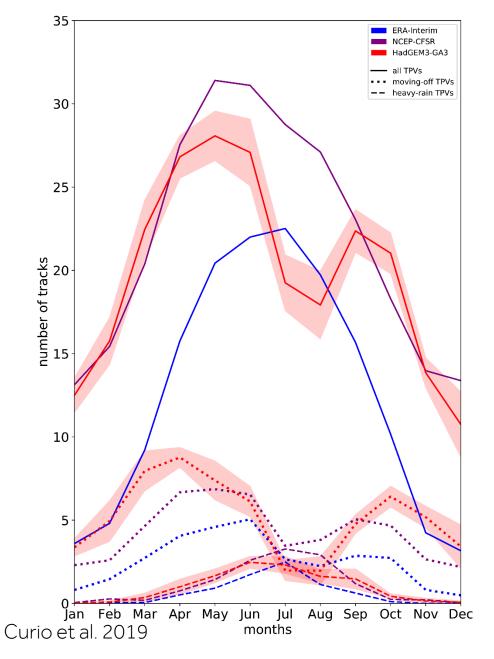




TPV tracks example for July 2008 from ERA-Interim







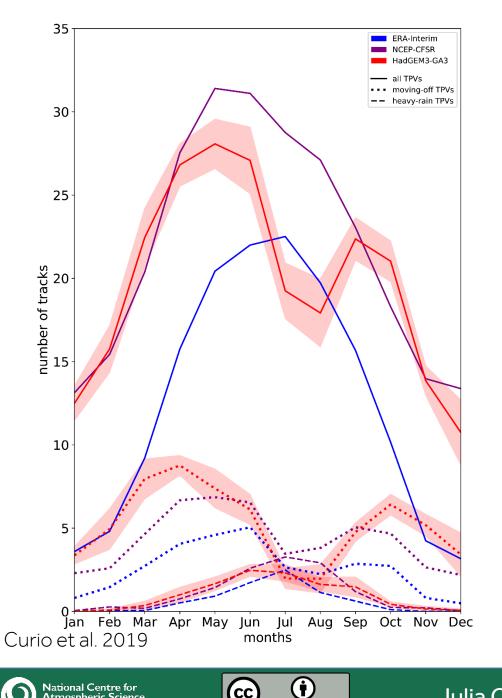
Annual cycle

- Frequent phenomena
- Pronounced annual cycle with maximum occurrence frequency in summer and minimum in winter
- Only a minority of TPVs move off the Tibetan Plateau









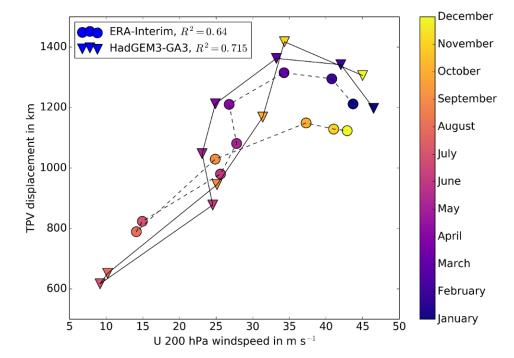
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Annual cycle

• Frequency of moving-off TPVs is highest in spring and autumn



 Close association between jet speed and TPV displacement explains mid summer suppression of moving-off TPVs

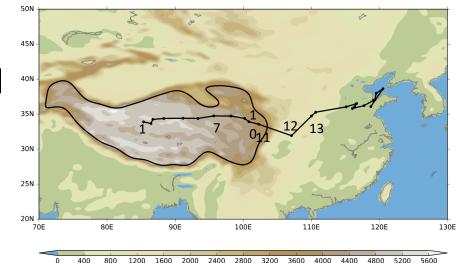
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Objectives

- Under which conditions are TPVs able to move off?
- Connect TPV behaviour and structure with large-scale circulation
- How do TPVs trigger precipitation?

Methodology

- Storm-centred composites of horizontal and vertical structure
- Due to the complex and diverse surface conditions the composites including all TPVs and regions do not provide a complete picture.
- Start with a TPV that caused extreme precipitation in the Sichuan basin in July 2008 as a test case.



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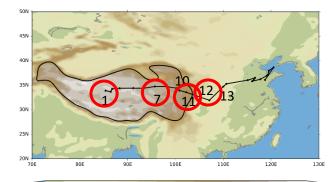




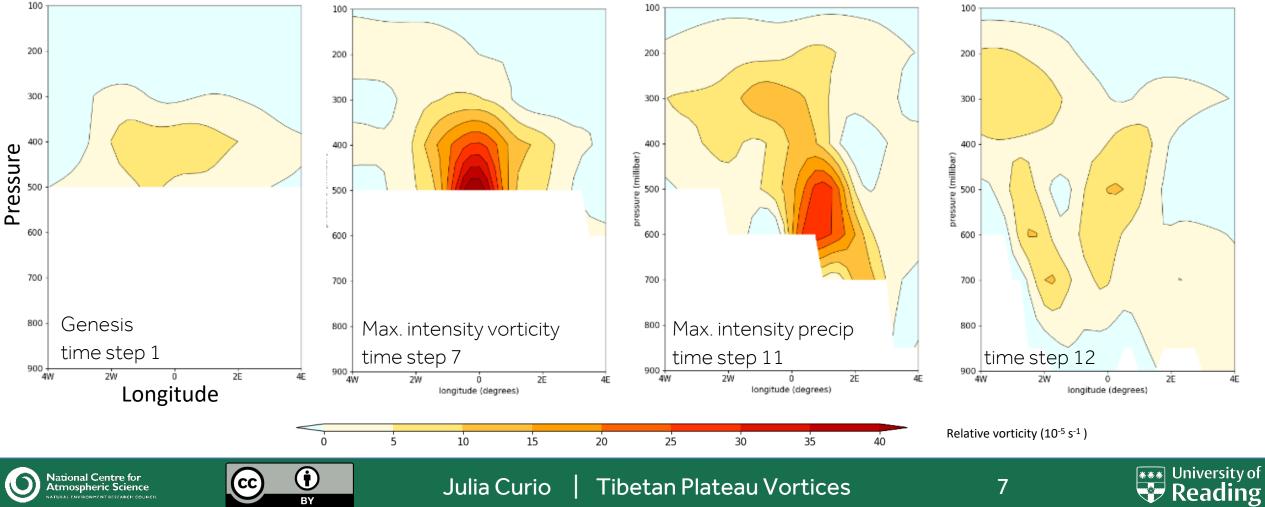
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July 2008 TPV vertical structure: relative vorticity



2400 2800 3200 3600 4000 4400 4800 5200 1200 1600 2000 5600 surface altitude(m

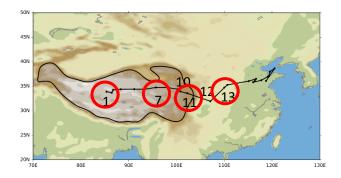


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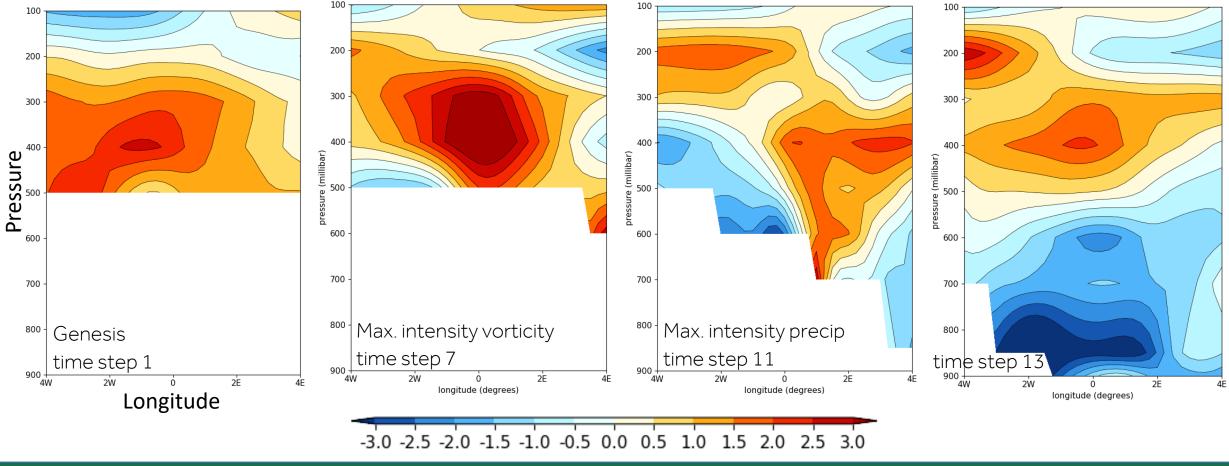
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Tibetan Plateau Vortices

July 2008 TPV vertical structure: temperature anomaly



0 400 800 1200 1600 2000 2400 2800 3200 3600 4000 4400 4800 5200 5600 surface_altitude(m)







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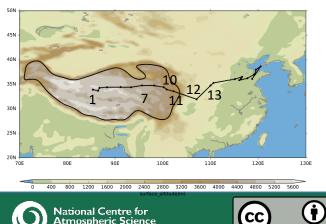
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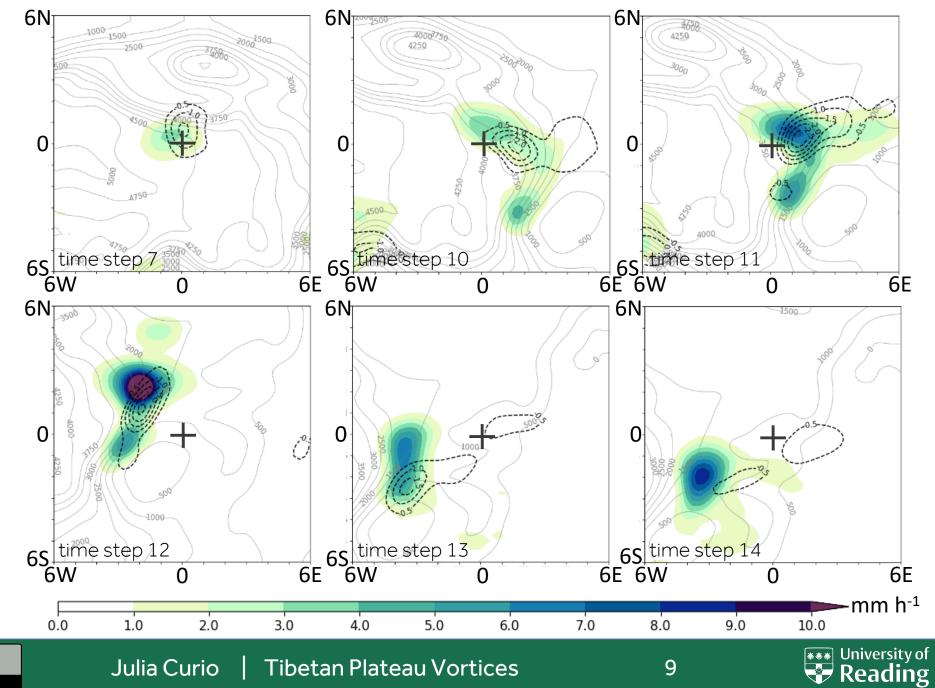
July 2008 TPV precipitation

Precipitation starts to form at the edge of the TP and stays terrain locked at the slopes of the TP when the TPV continues to moves further eastwards.

Colour shading: precipitation rate Grey contours: orography Dotted contours: updraft

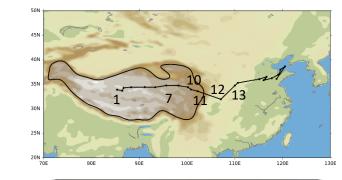


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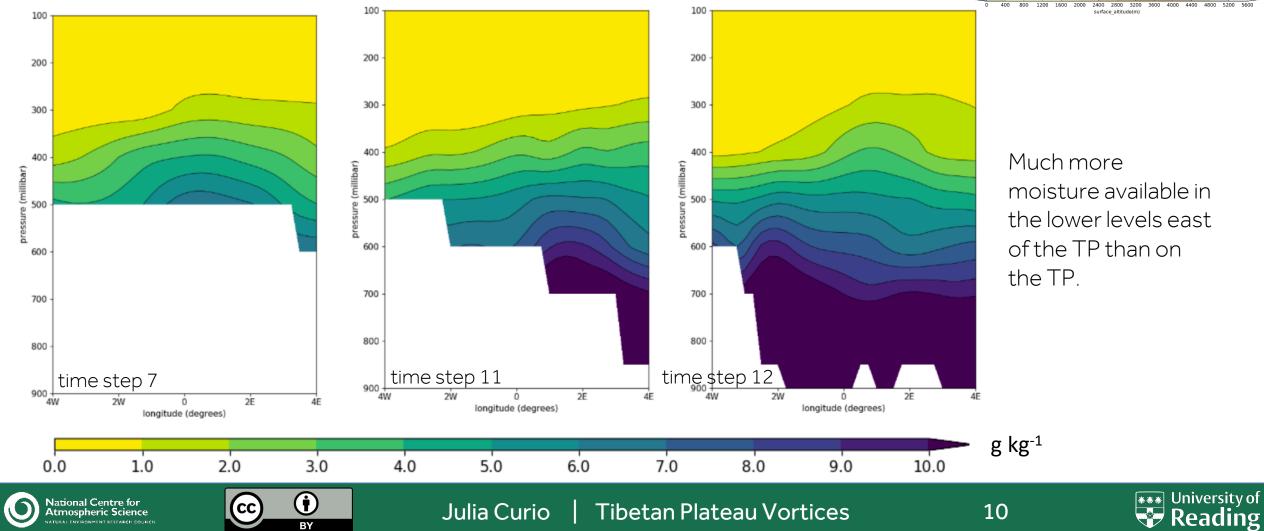


July 2008 TPV vertical structure: moisture

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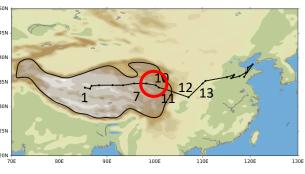
4800 5200



moisture available in the lower levels east of the TP than on

- Westerly winds in the Sichuan basin at time step 9, when the TPV is still located over the eastern TP.
- The next slides show how the local circulation in the Sichuan basin changes to a easterly winds with the TPV moving off the TP.
- \rightarrow The TPV appears to change the local circulation, thereby directing a flow of moist air towards the eastern slope of the TP.

(i)

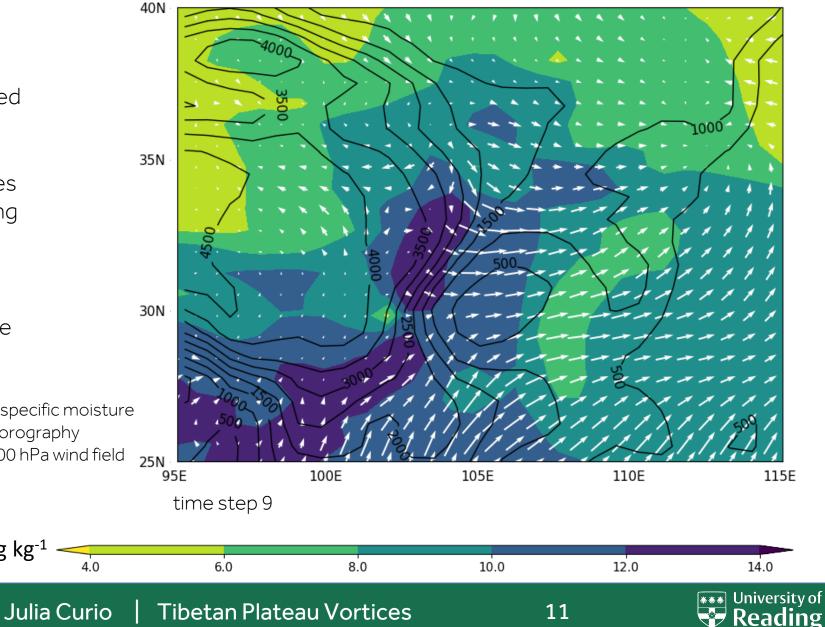


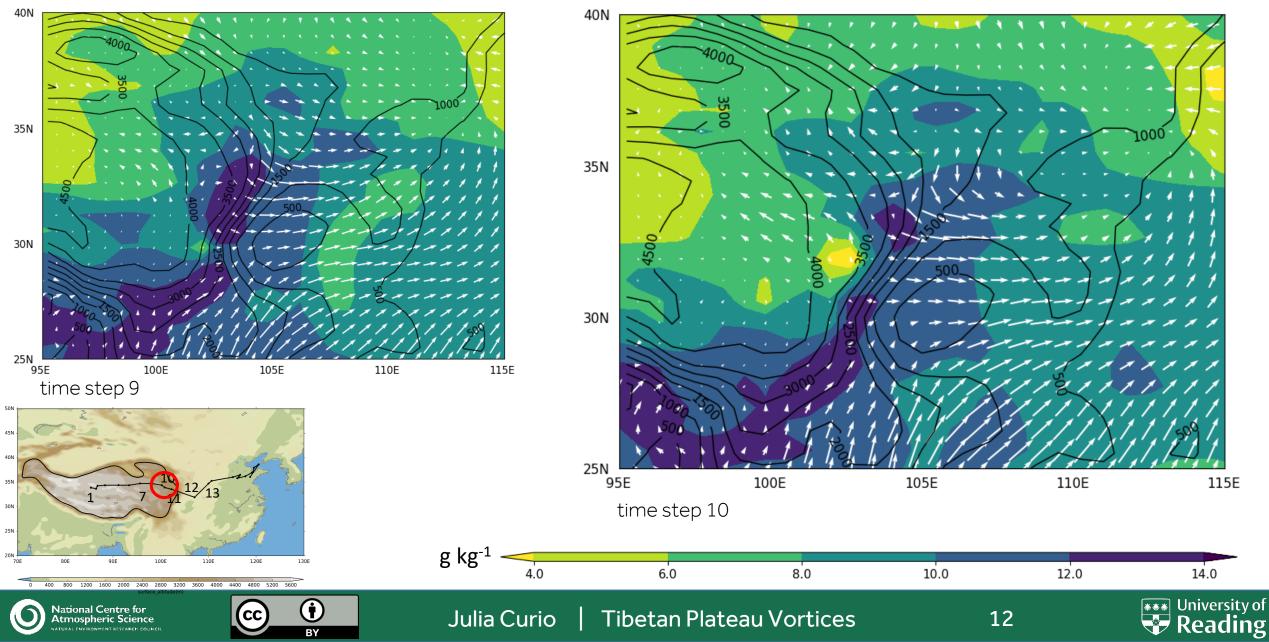
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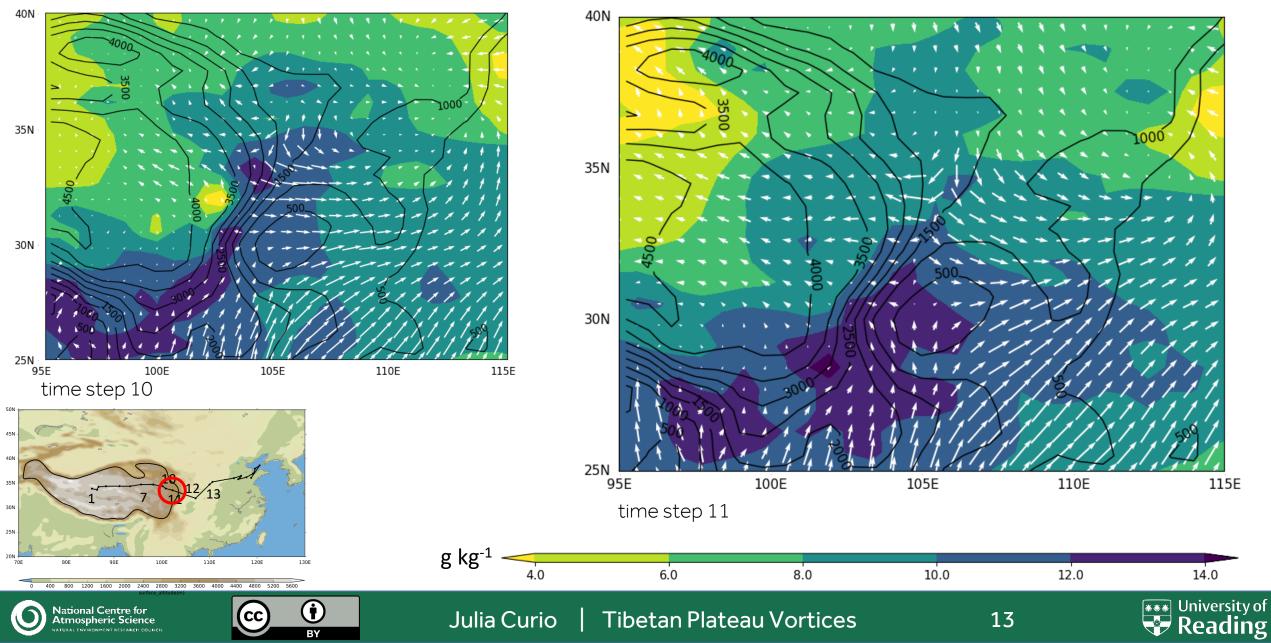
Colour shading: specific moisture Black contours: orography White arrows: 700 hPa wind field

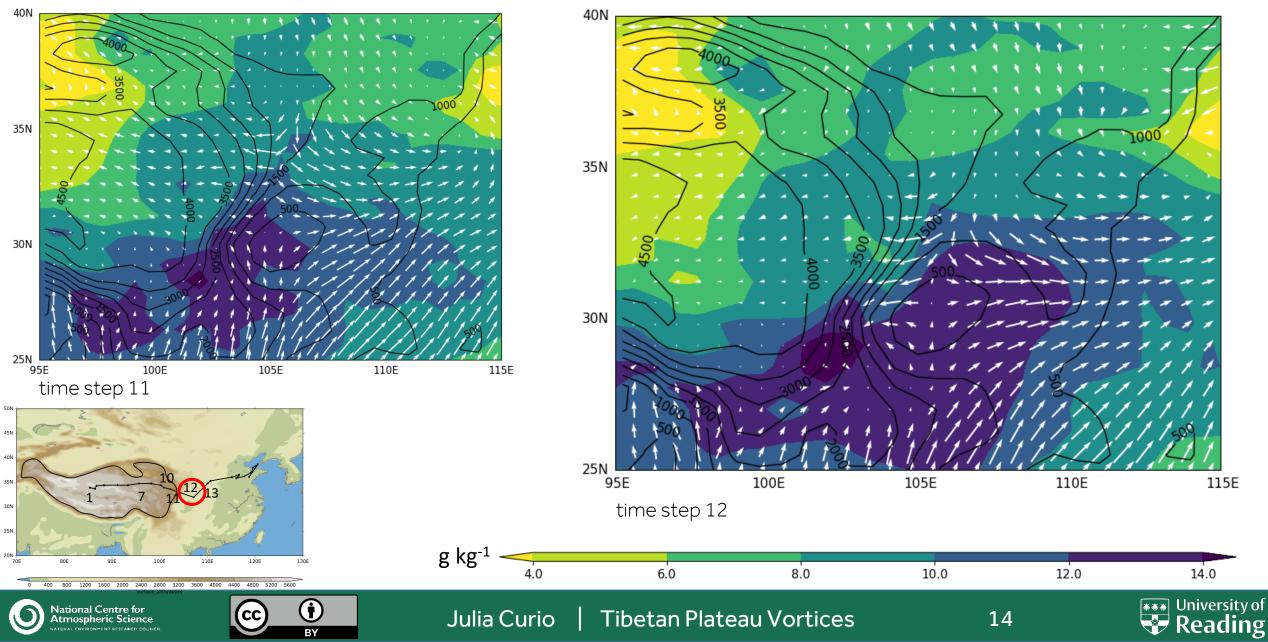
g kg⁻

4.0



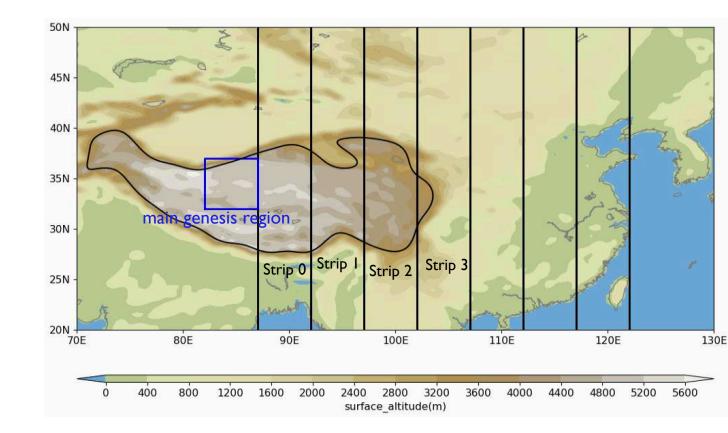






Composites for different regions

- subsample all moving-off and not moving-off TPVs generated in the main genesis region (all months)
- generating relatively homogeneous groups of TPVs
- composites over 5° wide longitudinal strips





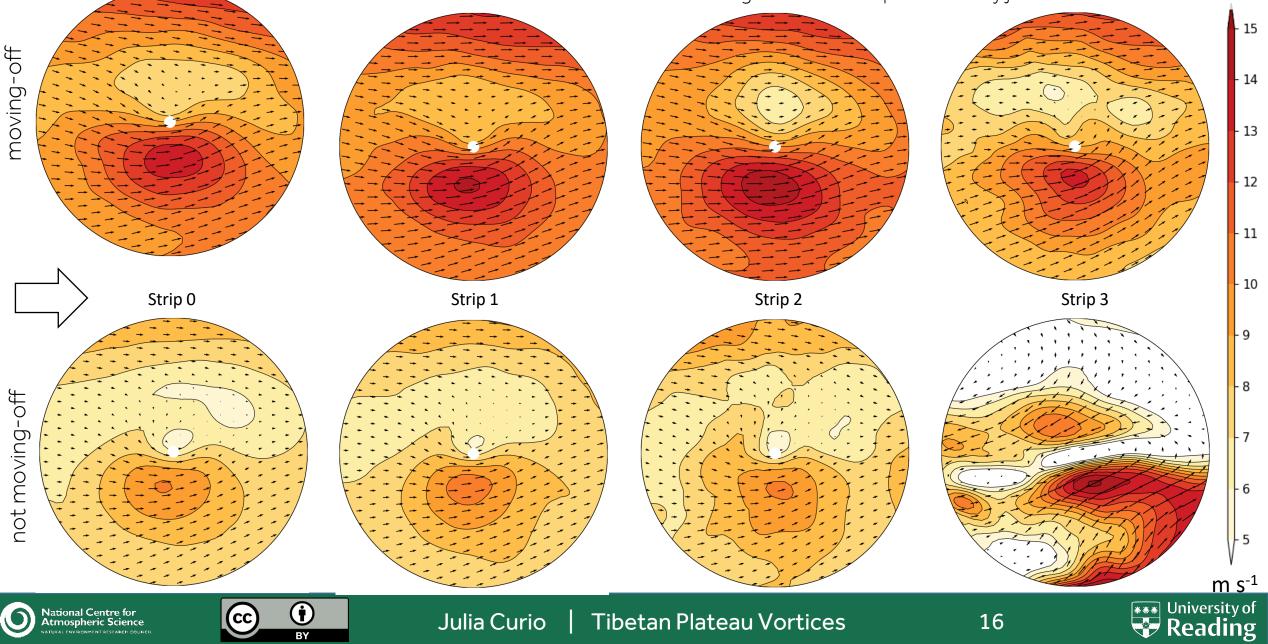


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Earth relative winds 500 hPa

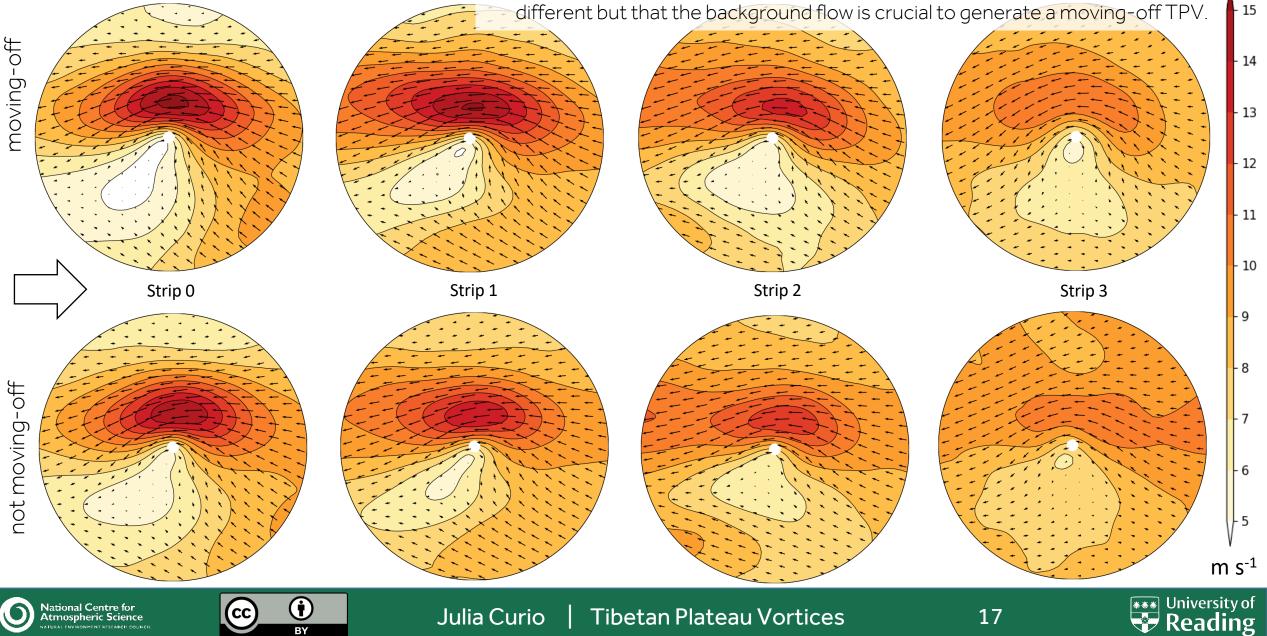
- Strongest winds south of the TPV centre for both groups
- Winds stronger for moving-off TPVs, a second maximum north of the TPV centre -> southern edge of the subtropical westerly jet



System relative winds 500 hPa

BY

- When looking at the system relative winds, propagation speed removed, the wind fields of both groups look very similar.
- This suggests that moving-off and not moving-off TPVs are not structurally different but that the background flow is crucial to generate a moving-off TPV.



Summary

- Only a small fraction of TPVs can move off the TP to the east.
- Impact of subtropical westerly jet on moving-off TPVs visible.
- Background flow appears to be crucial for moving-off
- TPV triggers precipitation on the eastern slopes of the TP when moving off: change in circulation and higher moisture availability in region with conditions already favourable for convection.







Next steps

- Where do impactful TPVs form?
- Which ingredients necessary to produce an impactful TPV?
- Where is the moisture for TPV-associated precipitation coming from?
- Understanding how a combination of the (right) large-scale atmospheric conditions and a TPV-induced change in the local circulation downstream of the TP can create an impactful TPV may enable improved forecasts of TPVs and their impacts.











Please let me know if you have any questions

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Curio, J., R. Schiemann, K. I. Hodges, and A. G. Turner, 2019: Climatology of Tibetan Plateau Vortices in reanalysis data and a high-resolution global climate model. J. Climate, **32**, 1933–1950, https://doi.org/10.1175/JCLI-D-18-0021.1.

Curio, J., Y. Chen, R. Schiemann, A. G. Turner, K. C. Wong, K. I. Hodges, and Y. Li, 2018: Comparison of a manual and an automated tracking method for Tibetan Plateau vortices. Advances in Atmospheric Sciences, 35 (8), 965-980, https://doi.org/10.1007/s00376-018-7278-4.