The dynamical composition of the Madden-Julian Oscillation

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Problem

What are the modes more excited by the convection associated with the MJO, and in which modes is the energy generated?

The energy balance equation

The balance of the total energy (the kinetic plus the available potential energy), *E_{nlk}*, of the atmospheric circulation is given by

$$\frac{\mathrm{d}}{\mathrm{d}t}E_{nlk} = C_{nlk} + S_{nlk},\tag{1}$$

n, l, k stand for zonal wavenumber, meridional and vertical mode indexes, respectively. Terms C_{nlk} and S_{nlk} represent the nonlinear energy interactions and the energy source/sink due to diabatic heating and dissipation.

(See, Marques et al, 2020,

https://doi.org/10.5194/gmd-2019-340, for more details).

Data & Method

 Horizontal wind field and geopotential from ERA-Interim reanalysis (1979-2016), and the outgoing longwave radiation (OLR)-based MJO index (OMI – https://www.esrl.noaa.gov/psd/mjo/mjoindex/).

Energy spectra for the composite differences between strong and weak convection were calculated.

The composite were calculated finding the days when the MJO index strength falls in the first quartil or forth quartil of each phase.

Index Convention for Figures

In the next figures the following convention for the meridional index, *l*, was followed. Negative indexes denote Rossby waves with *l* = −1 representing the first equatorial Rossby mode. The index *l* = 0 denotes the mixed Rossby-gravity mode.

Positive indexes denote gravity mode with l = 1 denoting the Kelvin waves.

Generation & Dissipation



composite difference (strong - weak convection)

Nonlinear Interaction



composite difference (strong - weak convection)

Total Energy



composite difference (strong - weak convection)

Horizontal Circulation and Geopotential Fields



strong - weak convection: shading represents the OLR anomalies; the thick blue vector in lower right corner of which panel represents a wind speed difference of $5 \text{ m} \cdot \text{s}^{-1}$.

Conclusions

► The MJO is characterised by:

- 1. Strong generation of energy in the equatorial Rossby waves of wavenumbers 1, 2 and 3;
- 2. Energy transfer from the Rossby waves to the Kelvin waves of wavenumber 1;
- 3. Strong dissipation of energy in the Kelvin waves of wavenumber 1;
- 4. Compared to the case of weak convection, during strong convection the energy in the equatorial Rossby waves and in the Kelvin waves of wavenumber 1 is 1.5 and 1.9 times higher, respectively.