

MAR suitability mapping combined with field examination and numerical simulation in the Danube-Tisza Interfluve

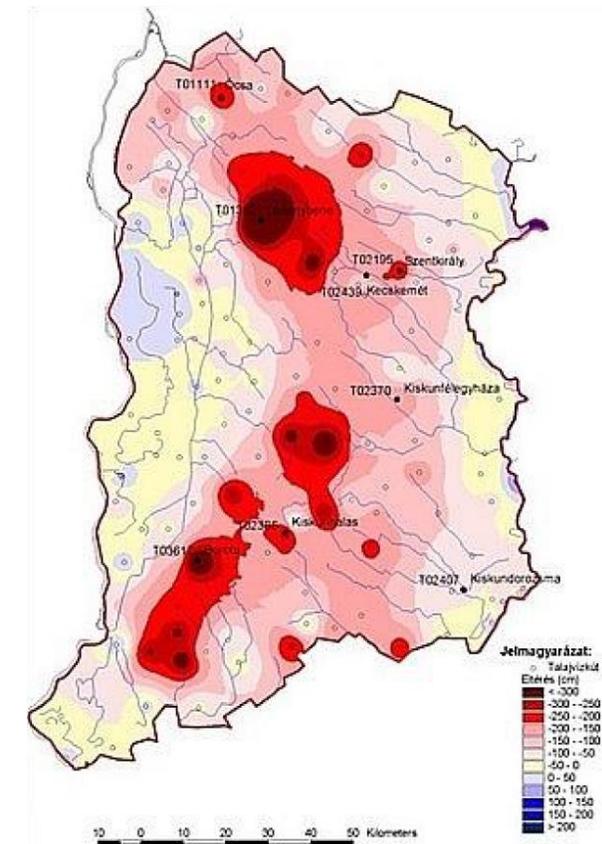
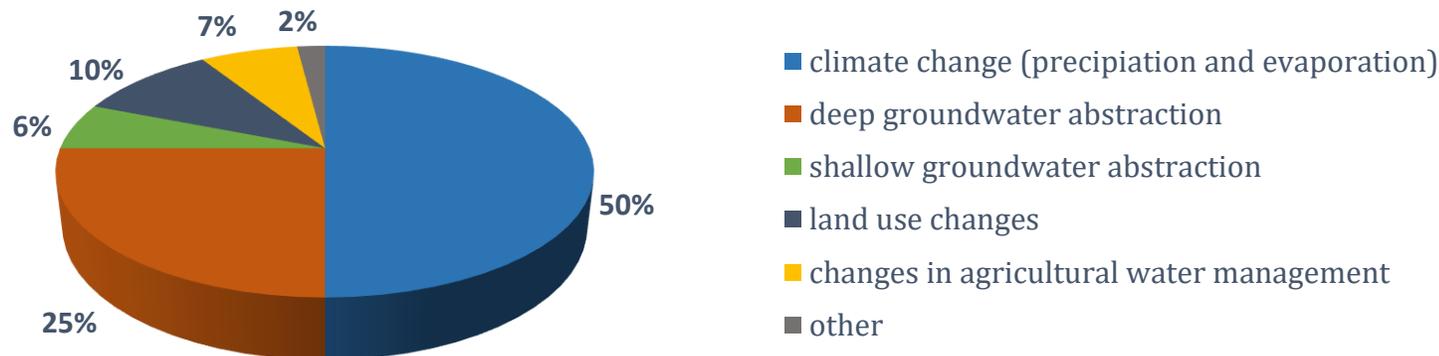
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1. Introduction and aims

- ❖ **Water level drawdown** in the area due to climate change and additional anthropogenic effects
- ❖ Numerous replenishment plans were worked out in the past decades but using **Managed Aquifer Recharge** hasn't been considered yet
- ❖ The aim of this research is to find **suitable areas** for MAR and to assess water recharge possibilities in a local study area

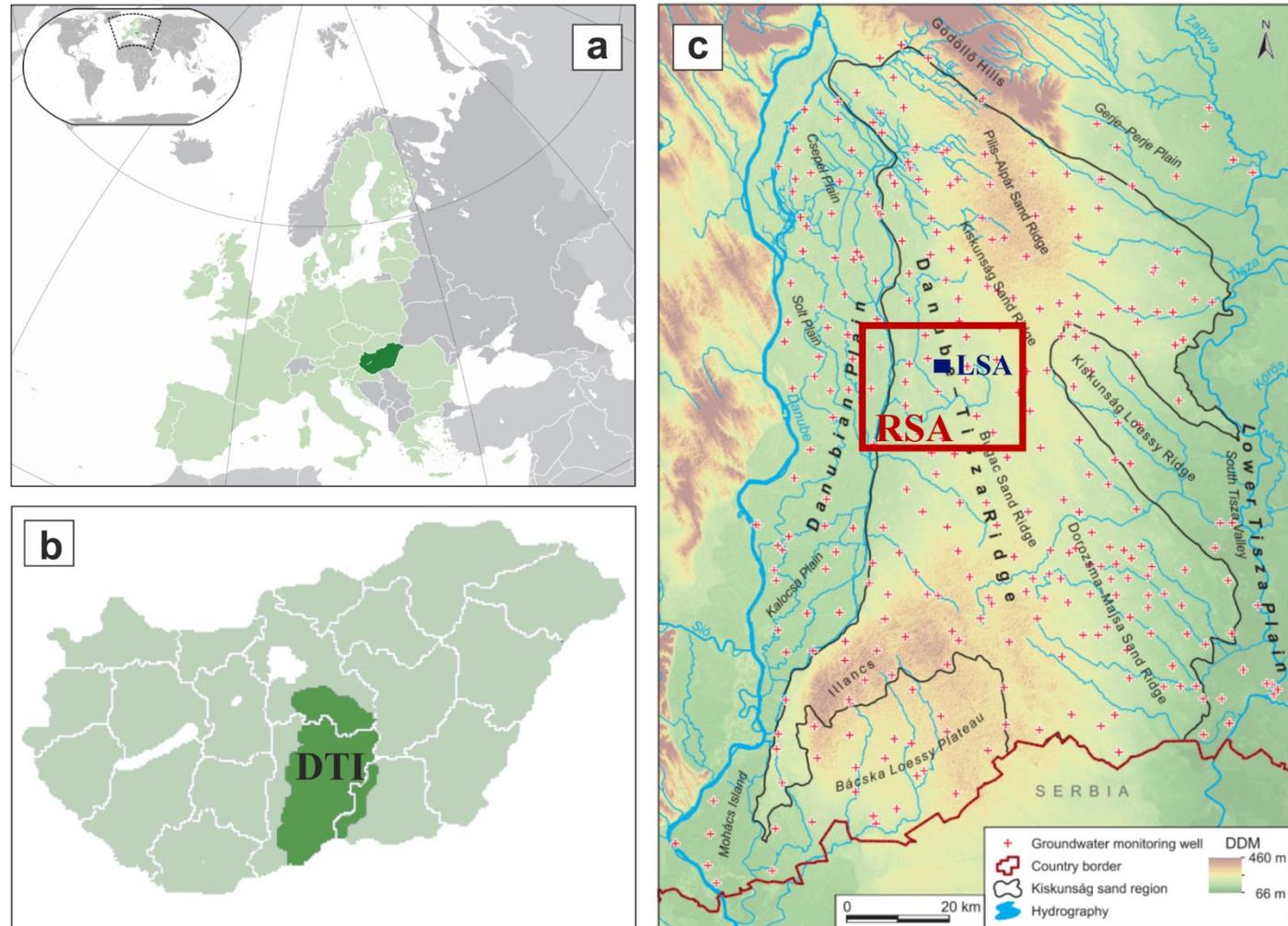


Water level changes between 1956-60 and 2002 (VITUKI, 2002)

Reasons of groundwater level reduction (based on Pálfi, 2010 and Nagy et al., 2016)

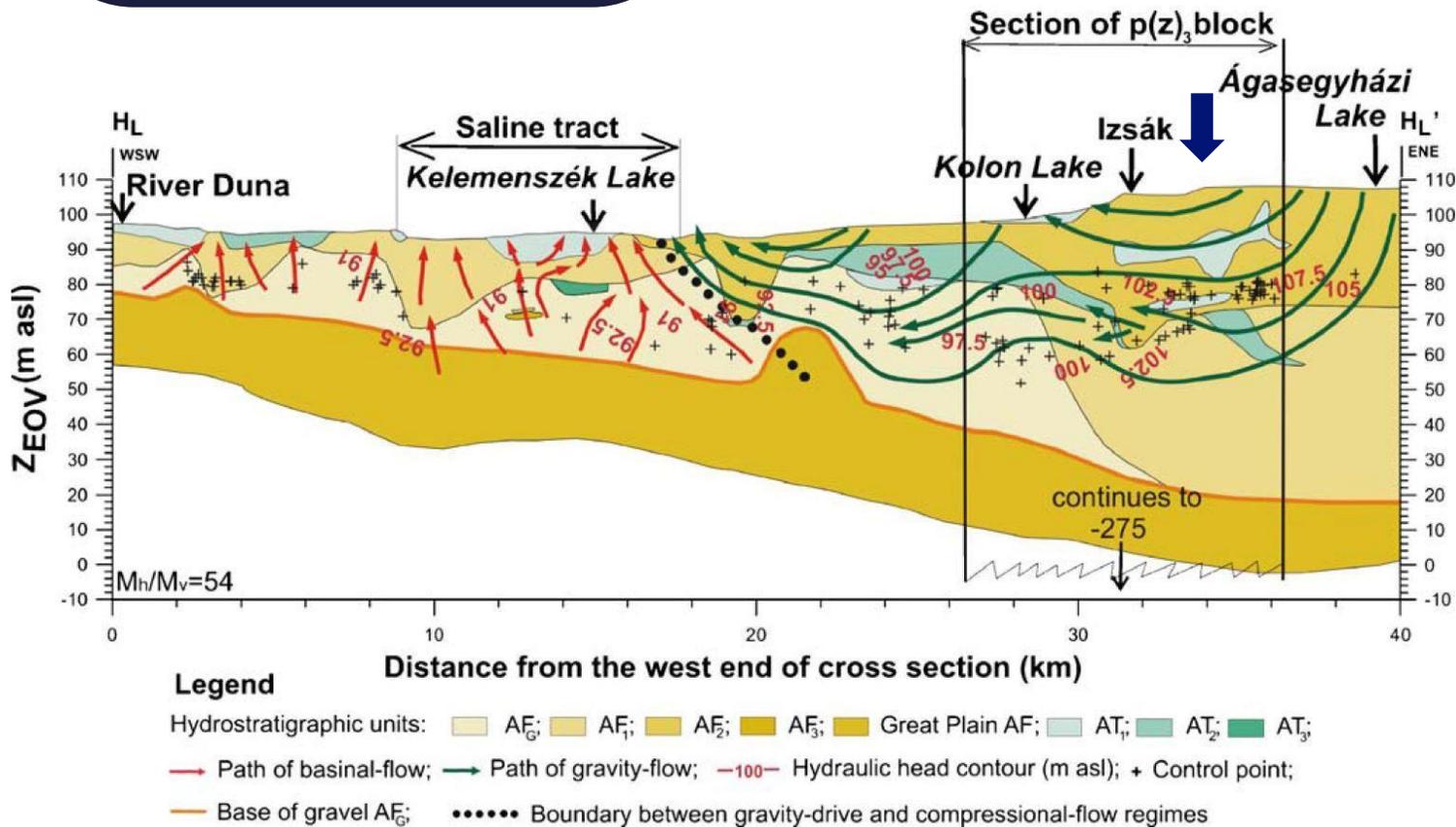
2. Study area

- ❖ Located in Hungary, in the Danube-Tisza Interfluve (DTI)
- ❖ DTI is a ridge region, up to 130 m a.s.l. between Danube and Tisza Rivers
- ❖ The river valleys are situated at 85-90 m above sea level
- ❖ Alluvial sediments of Danube and aeolian sands



Location of Hungary (a), the Danube-Tisza Interfluve – DTI (b), the regional study area – RSA (c, red rectangle) and the local study area – LSA (c, blue rectangle)
(Fig. 1c map modified from Kohán & Szalai, 2014)

2. Study area



Hydrostratigraphic and hydraulic section for the Western part of the Duna-Tisza Interfluve, AF-aquifer, AT-aquitard
 Mádl-Szőnyi & Tóth, 2009

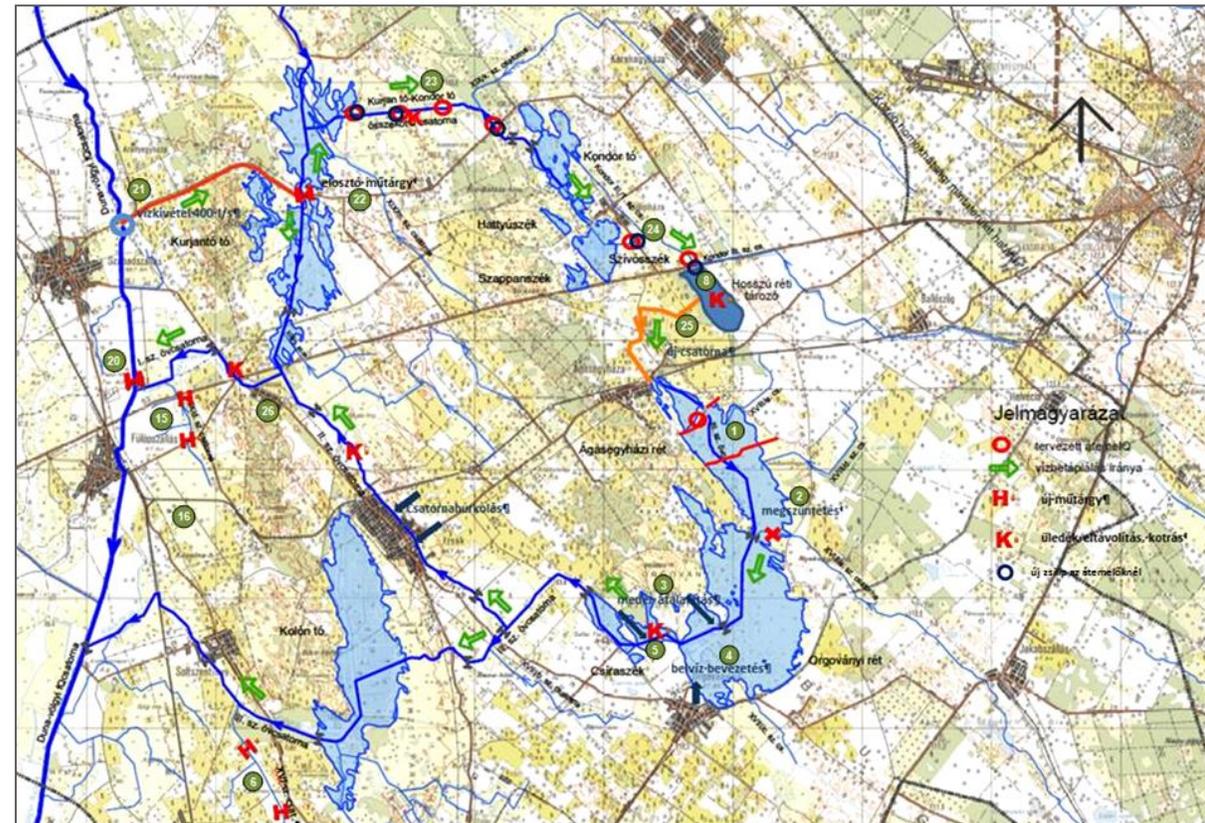
- ❖ The area's groundwater flow systems are characterized by a **gravity driven meteoric water** and an **over-pressured saline water**.
- ❖ The shallow flow systems of the elevated ridge region are under the effect of gravity driven meteoric flow regime.

3. Research background

- ❖ Water management problems in the broader area have been known for decades
- ❖ One of the most recent plans was to move water from the Danube Valley Channel to the center of the ridge, through existing channels and lakes (Nagy et al., 2016)
- ❖ Too expensive and not effective enough as the water can easily infiltrate from the channels and it would not reach the higher regions in sufficient amount
- ❖ One of the aim of this research is to find suitable areas for MAR utilisation



Western Water Supply Plan (Nagy et al. 2016) ▼

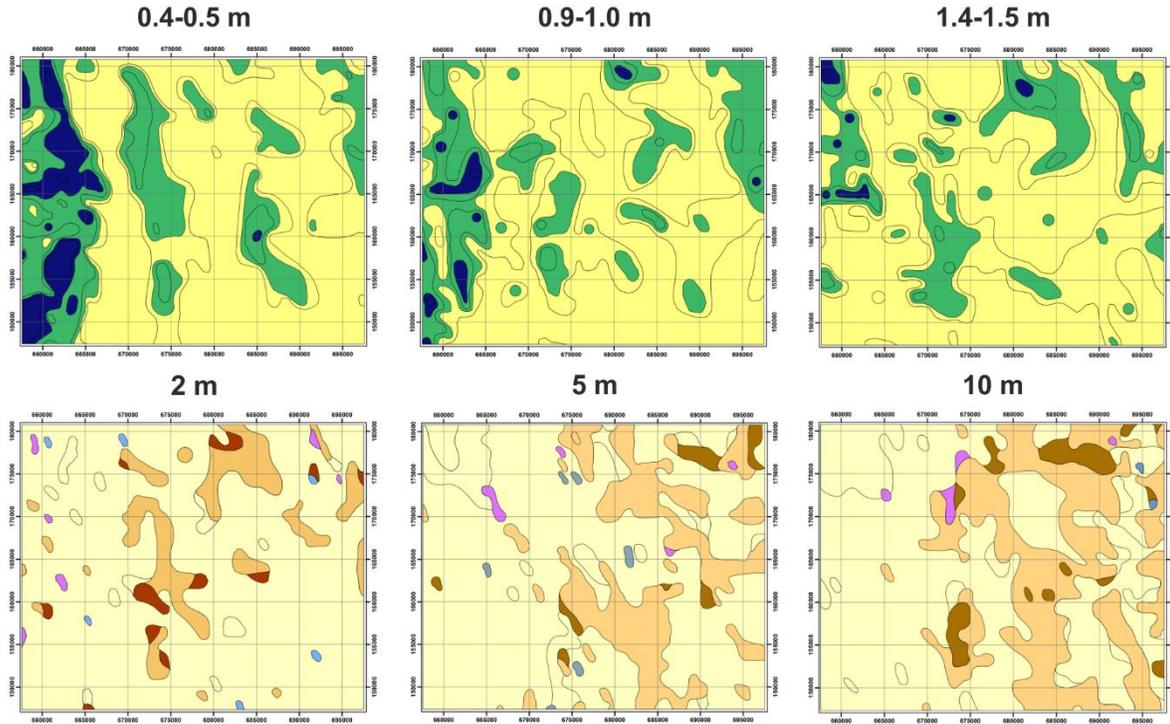


4. Methods

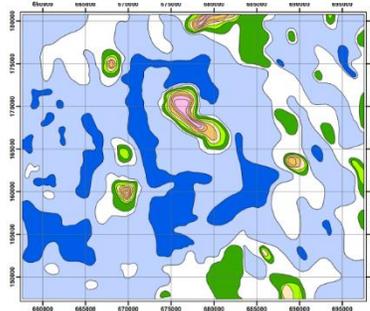
- ❖ **Suitability mapping** (Silva Cisneros, 2019) for the Western Water Supply Area (Regional Study Area – RSA)
- ❖ **Field measurements** (Local Study Area – LSA)
 - ❖ ERT and RMT geophysical measurements
 - ❖ Drilling by hand and soil sampling
 - ❖ Water level and water chemical measurements
- ❖ **Laboratory measurements** (LSA)
 - ❖ Water chemical measurements
 - ❖ Sieving and elutriation of soil samples
- ❖ **Numerical modeling** (LSA, cross section)



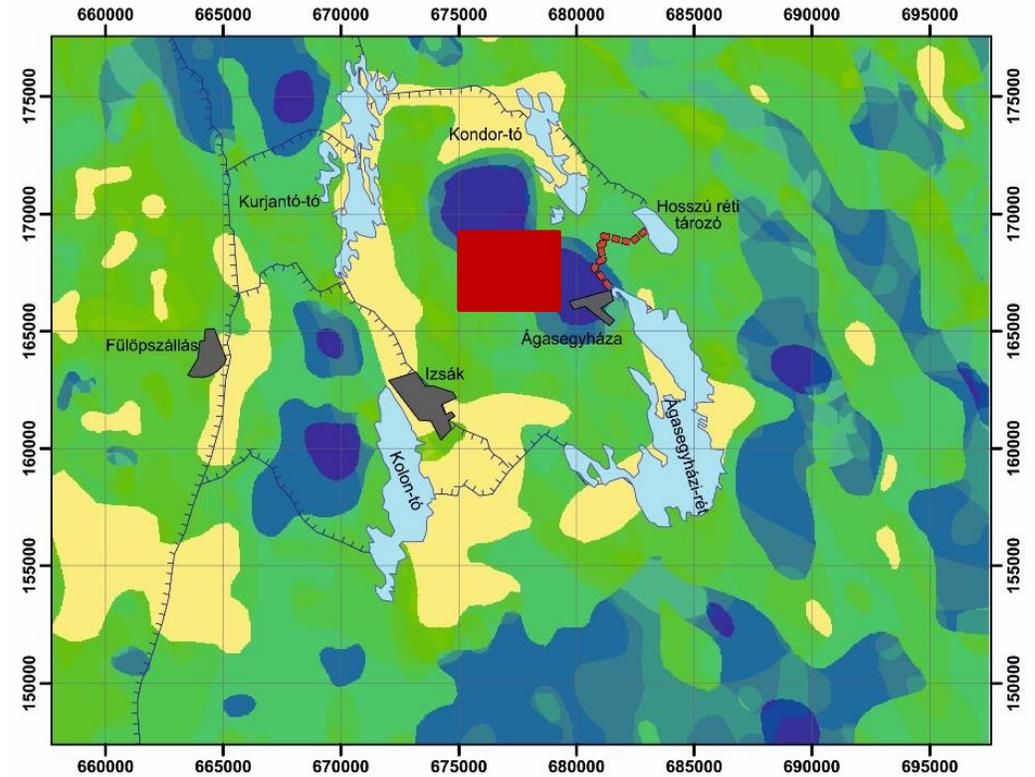
5. Suitability mapping



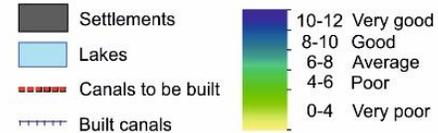
Water depth (1978)



Silva Cisneros, 2019
and Mádl-Szőnyi et al., 2019



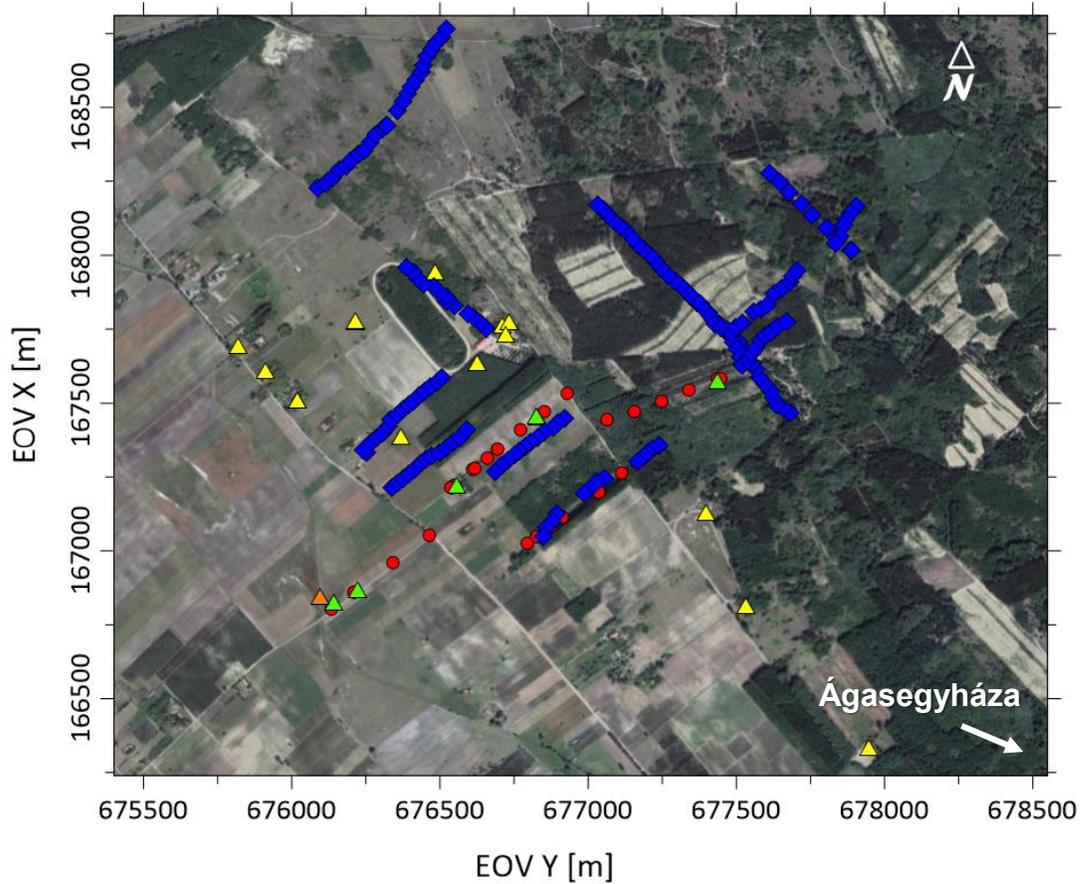
LEGEND



Based on Geological ATLAS of the Great Plain (1978)

- ❖ Based on near surface geology and water table depth (slope is not an important factor in this area)
- ❖ A **local study area** was chosen based on the final suitability map

6. Field measurements



● ERT measurement ◆ RMT measurement ▲ Well ▲ Drilling ▲ Sample from channel

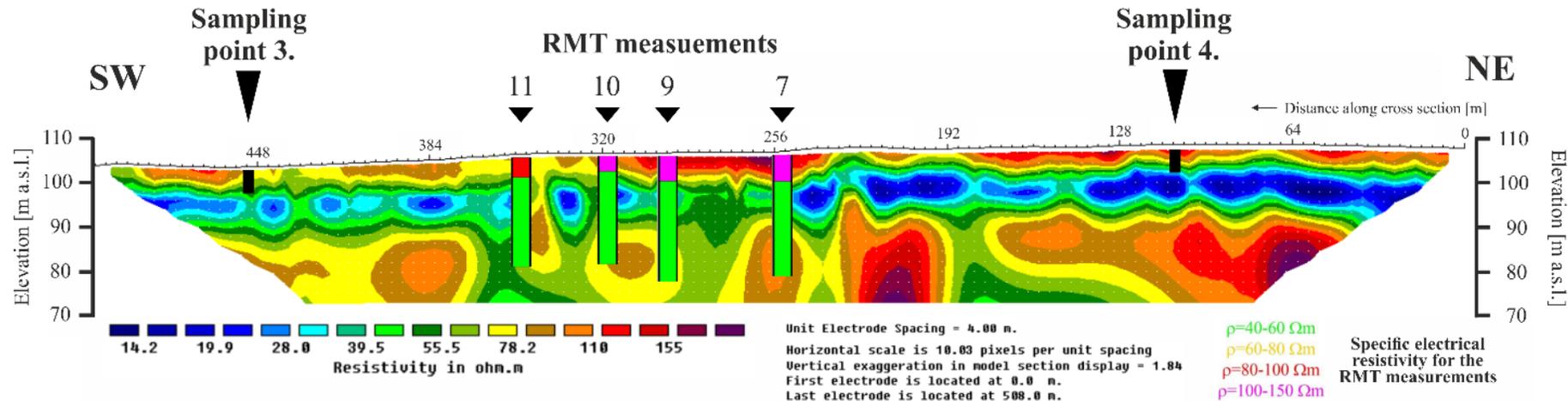
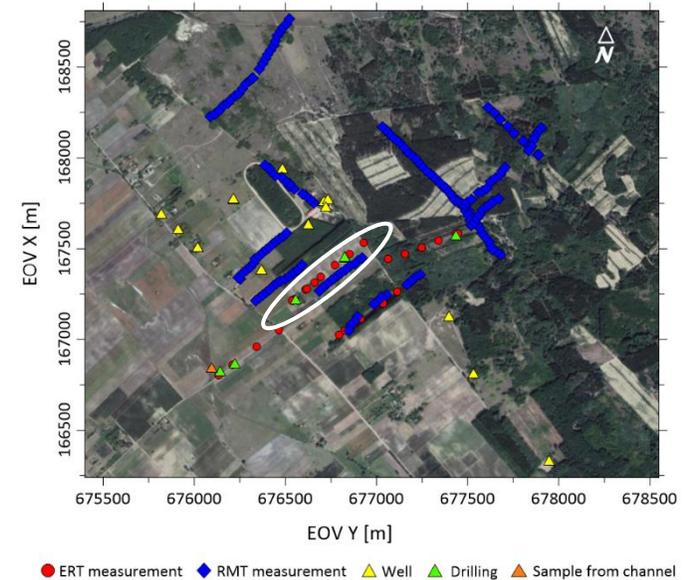
- ❖ Geophysical measurements (RMT, ERT)
- ❖ Drilling and soil sampling
- ❖ Water level and water chemical measurements



7. Results of field measurements

Based on ERT and RMT measurements 3 different layers could be distinguished:

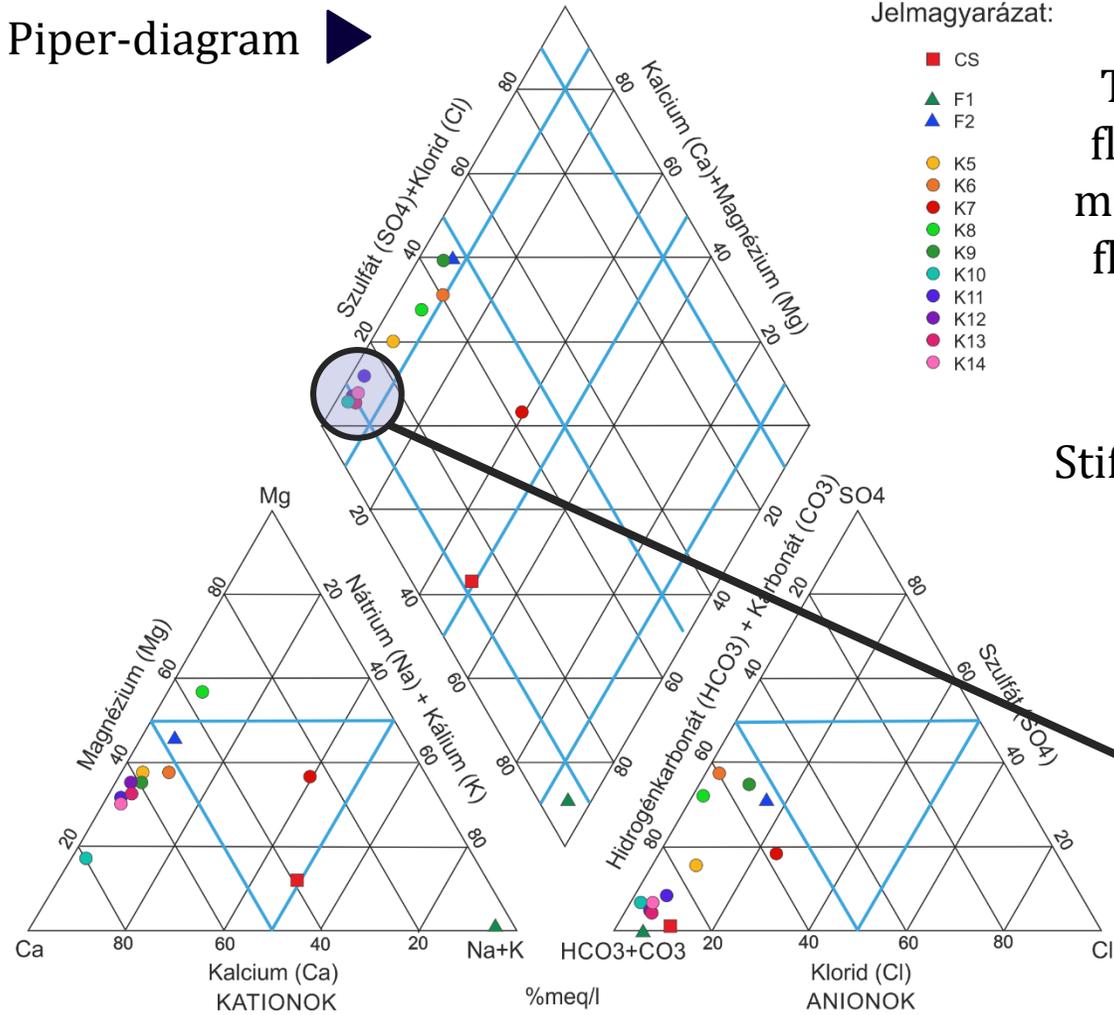
- 1) **Upper aquifer:** a relatively dry upper layer which is approximately coincident with the vadose zone (based on the geophysical measurements it can not be distinguished unequivocally).
- 2) **Aquitard:** a middle layer, with relatively low resistivity, higher clay content.
- 3) **Lower aquifer:** a third layer, which is probably more compact and has a lower hydraulic conductivity, than the upper layer, but still a relatively good aquifer.



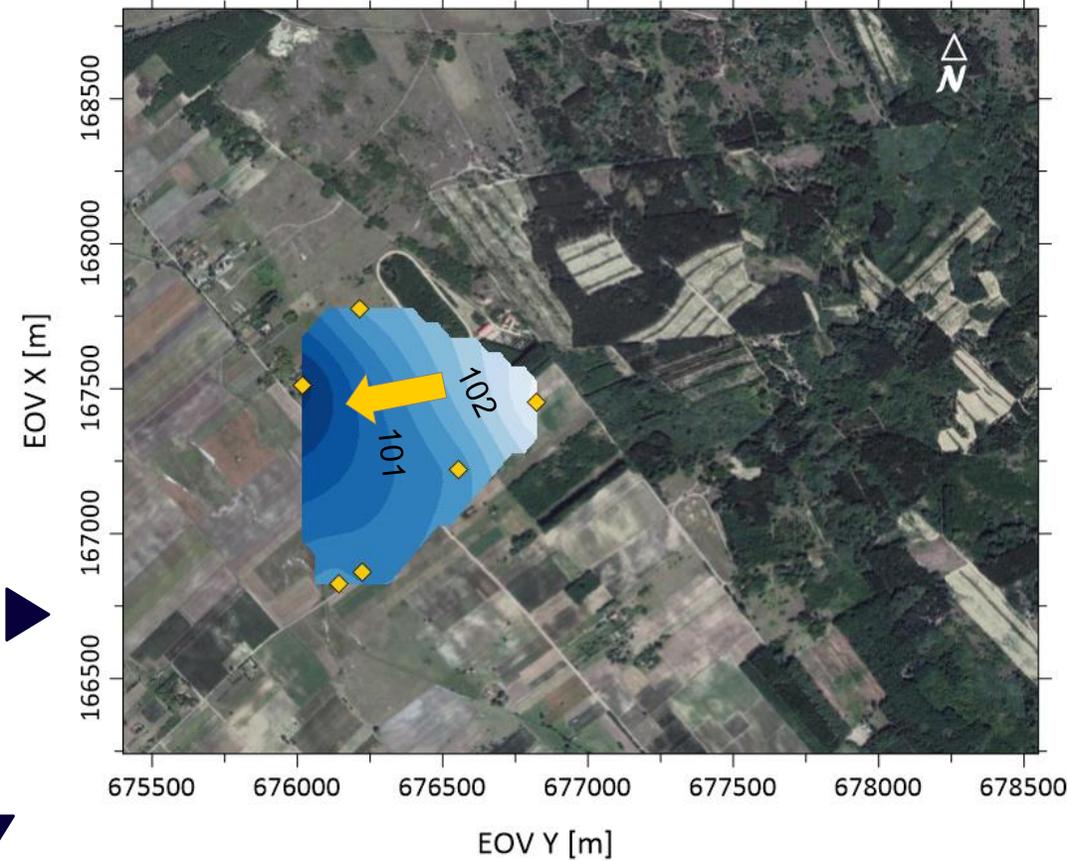
Inverted result of the ERT measurement (ERT 2); RMT measurements (11-7), and the locations of soil samples

7. Results of field measurements

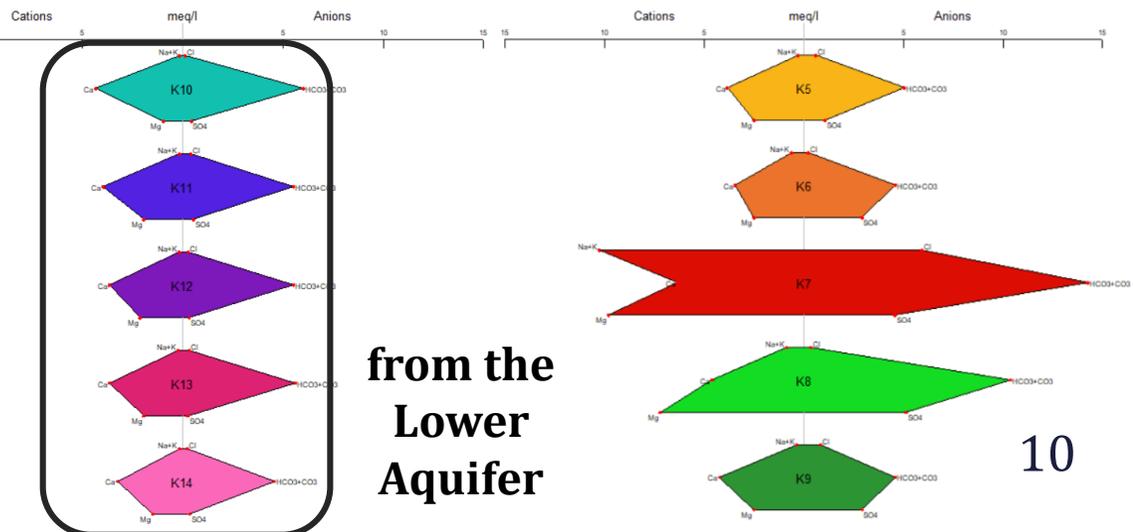
Piper-diagram



Tomographic fluid potential map, horizontal flow direction

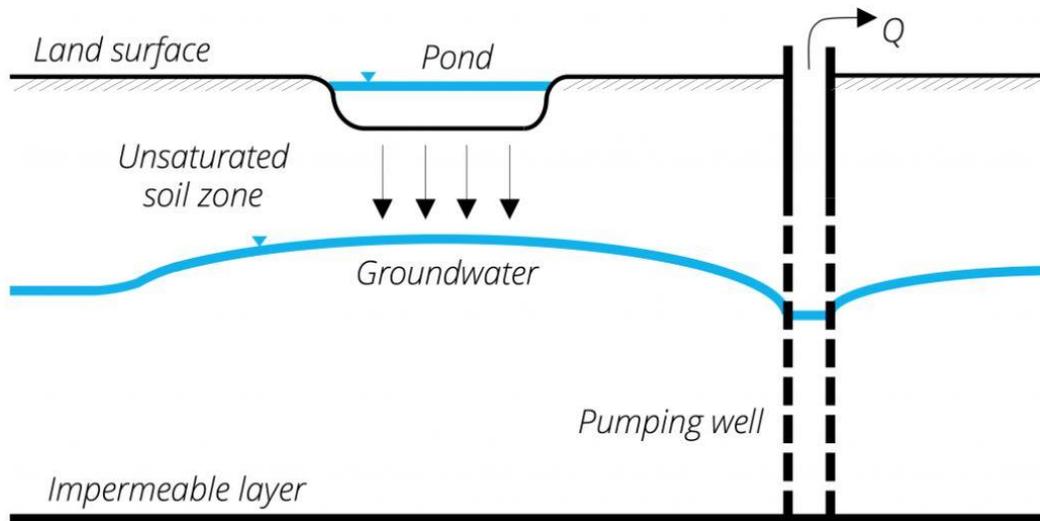


Stiff-diagrams



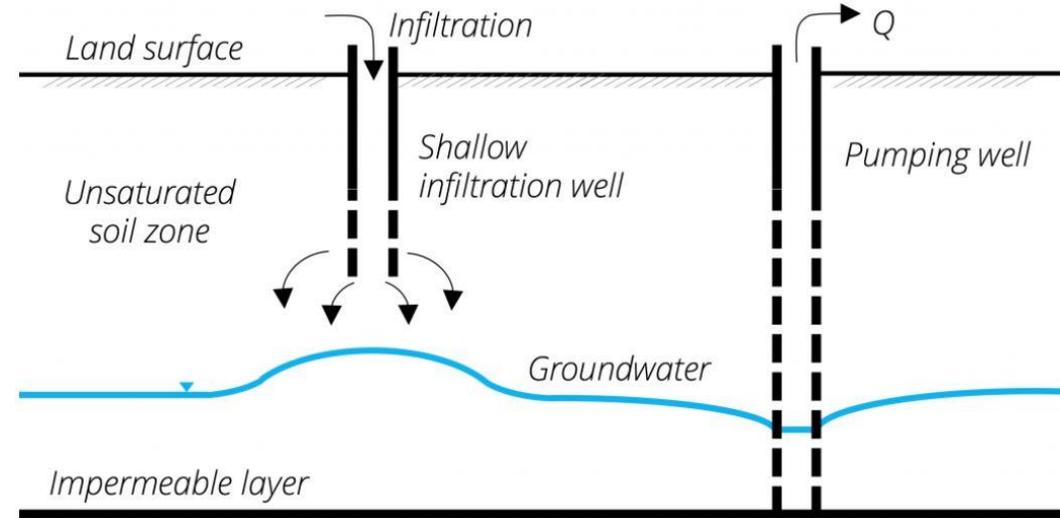
Possible Managed Aquifer Recharge methods

Infiltration pond

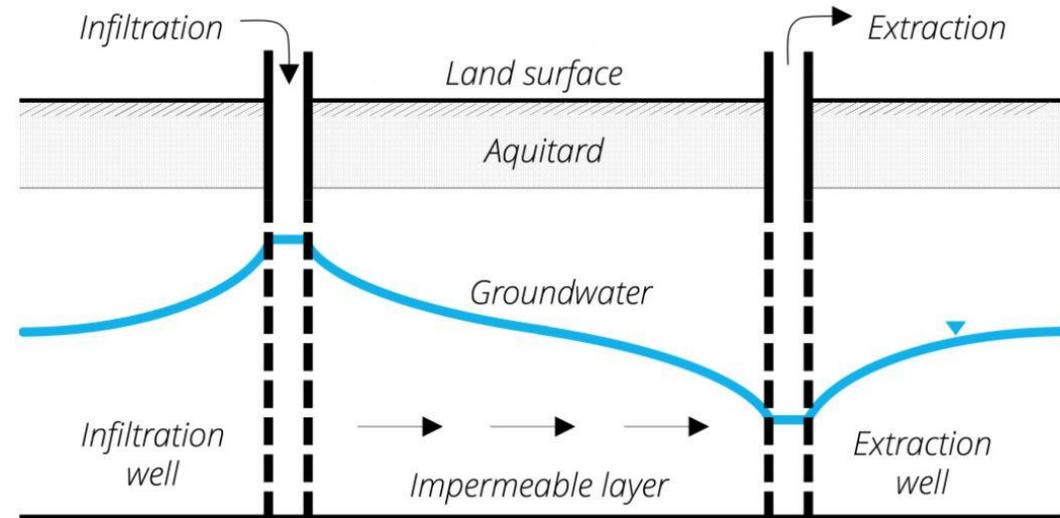


<https://inowas.com/managed-aquifer-recharge/>

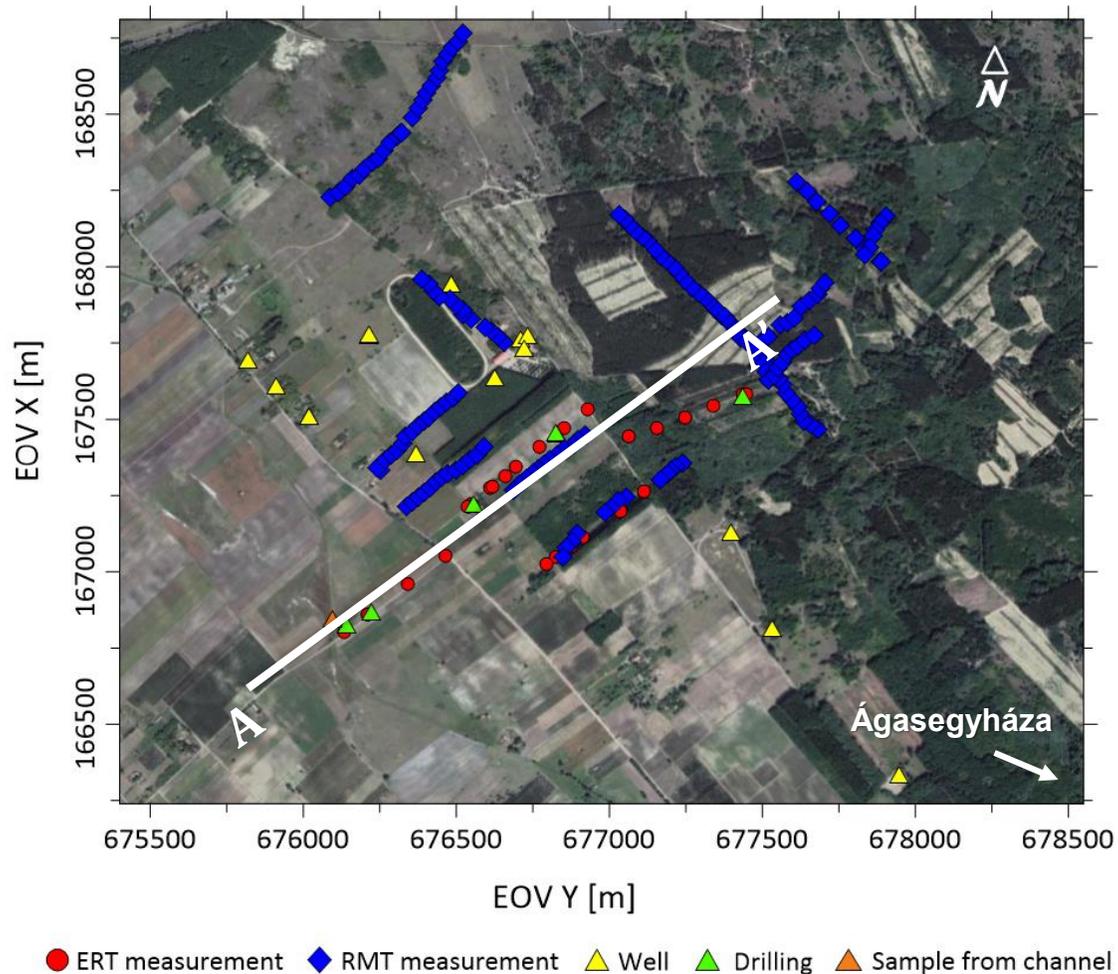
Shallow well



ASTR



8. Numerical modeling



❖ Possible modeling scenarios regarding geological build-up

1. Homogeneous, 1-layer model for only the Upper Aquifer
2. 3-layer model (continuous aquitard in between)
3. 3-layer model (aquitard in between, with discontinuities)

❖ Possible modeling scenarios regarding MAR methods

1. Infiltration basin
2. Shallow well
3. Deeper well recharging the Lower Aquifer

❖ **Material properties** based on laboratory measurements and geophysical measurements

❖ Upper Aquifer: $K \sim 1e-05$ m/s

❖ Aquitard: $K \sim 1e-06$ m/s (Müller et al. 2008)

❖ Lower Aquifer: $K \sim 5e-06$ m/s (Müller et al. 2008)

Vertical/Horizontal anisotropy: 0.1

❖ **Initial water table** specified based on Great Plain Atlas of Hungary, 1978 and field measurements

❖ **Boundary conditions:**

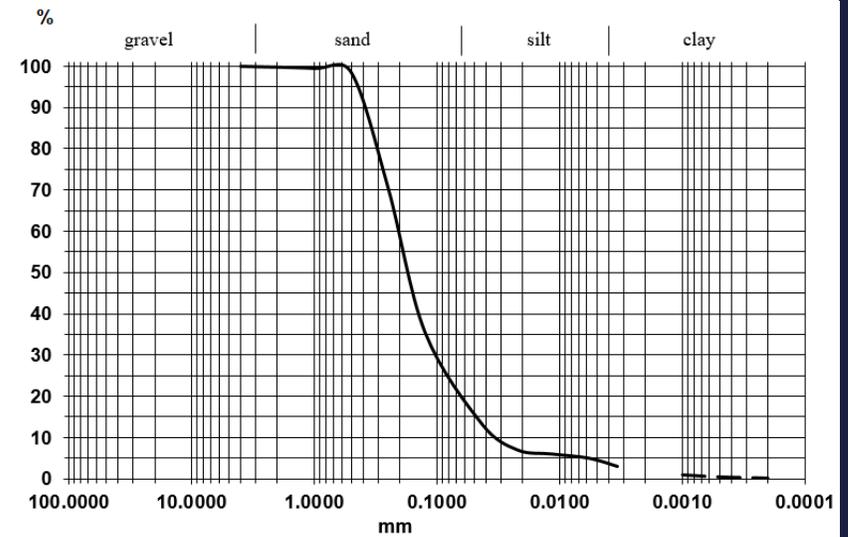
❖ Top boundary: annual recharge of 100 mm/year (Szilágyi et al. 2012)

❖ Right boundary: no flow (highest elevation)

❖ Left boundary: outflow based on natural hydraulic gradient: $3.4e-08$ m³/s/m², $3.4e-09$ m³/s/m² and $1.8e-09$ m³/s/m² for the different layers, respectively

❖ Bottom boundary: no flow (*moderate recharge area, adequate outflow rate is under assessment*)

❖ **Transient model:** 1 year (300 exponentially increasing time steps)

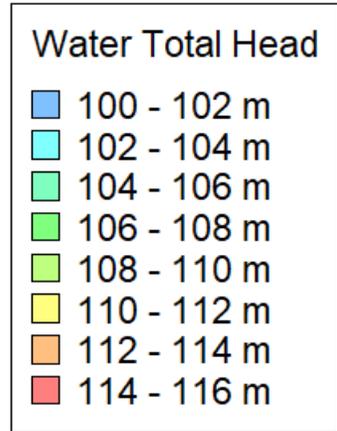
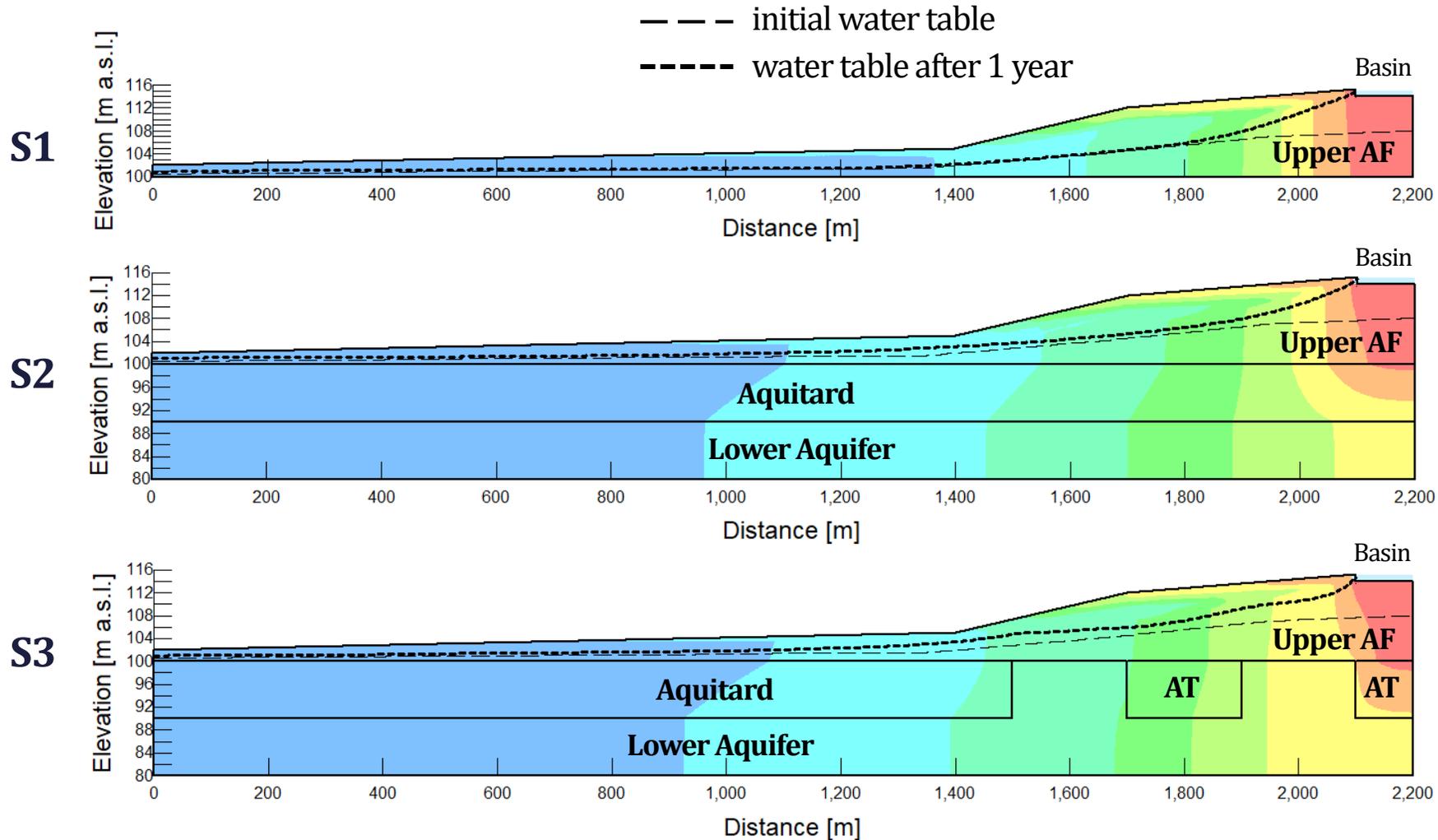


Preliminary results of numerical modeling

using  GeoStudio 2019
 SEEP/W

Infiltration basin

→ width: 100 m, water level: 1 m (constant)



Water level increase after 1 year at selected points

	400 m	1000 m
S1	35 cm	24 cm
S2	41 cm	55 cm
S3	41 cm	60 cm

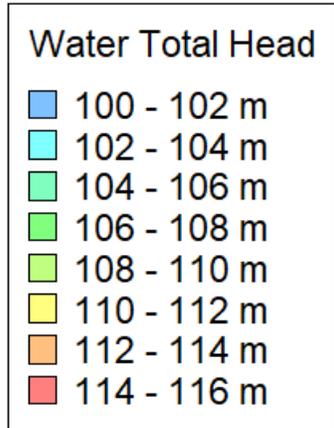
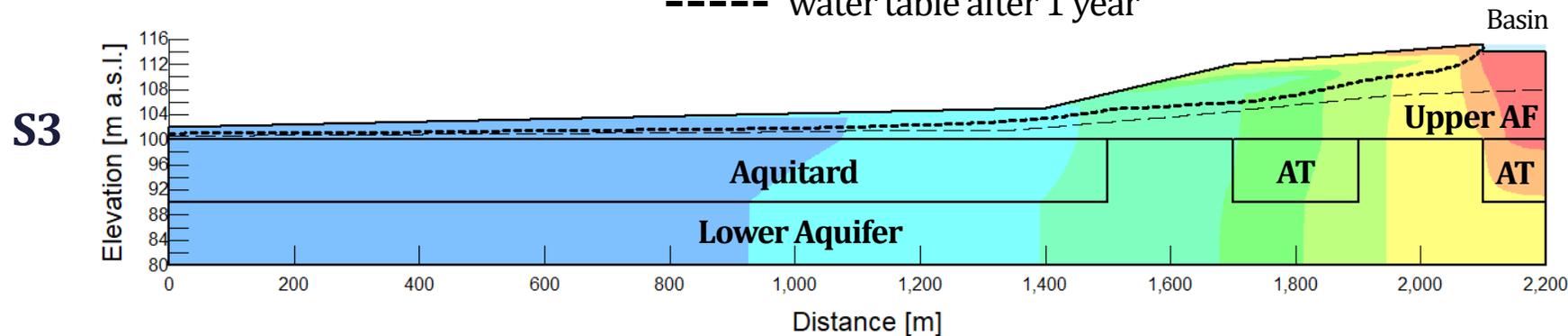
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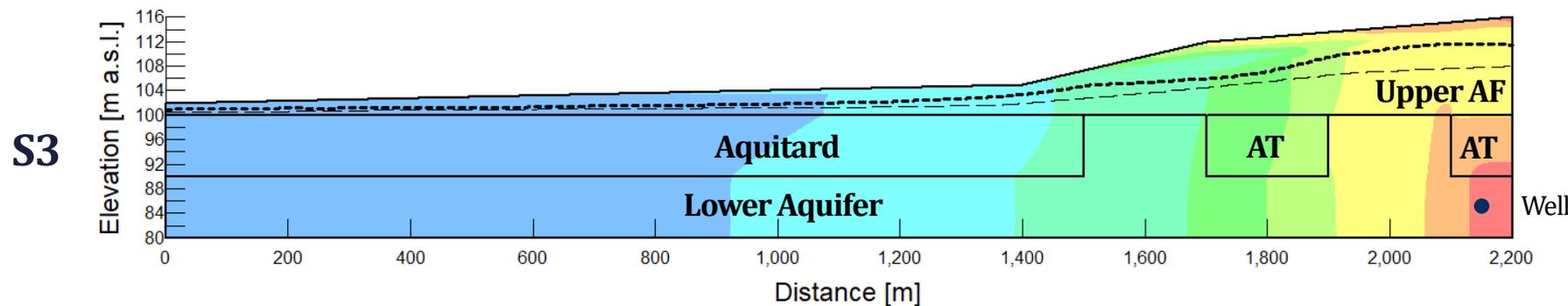
--- initial water table
 - - - water table after 1 year



Water level increase after 1 year at selected points

Well

→ equivalent amount of water ($Q \sim 2e-05 \text{ m}^3/\text{s}$)



S3	400 m	1000 m
Basin	41 cm	60 cm
Well	41 cm	59 cm

→ Similar results (due to no flow boundary at the bottom + sloping terrain)

9. Conclusions

- ❖ The research area **can be suitable** for using Managed Aquifer Recharge methods
- ❖ Possible methods: surface infiltration, shallow wells and deeper wells
- ❖ With the modeled infiltration basin, water level can be increased by 0.5 m in 1 year
- ❖ Groundwater flow regime can influence MAR possibilities, thus it **must be considered**
- ❖ **Local scale solutions** could ease the water shortage of this area

Further research aims:

- ❖ Scenario models for different MAR solutions
- ❖ A more detailed geological and hydrogeological study in the area → validation of modeling results
- ❖ (Rain)water infiltration experiments

10. References

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Thank you for your time!