



# Improved predictions for ENSO diversities using a nonlinear forcing singular vector assimilation

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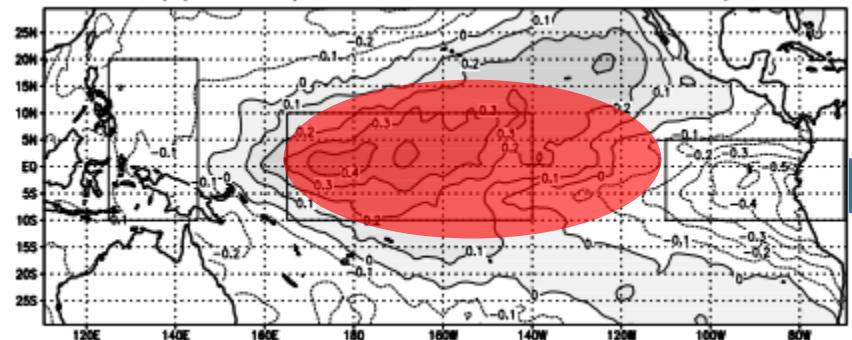
# CONTENT

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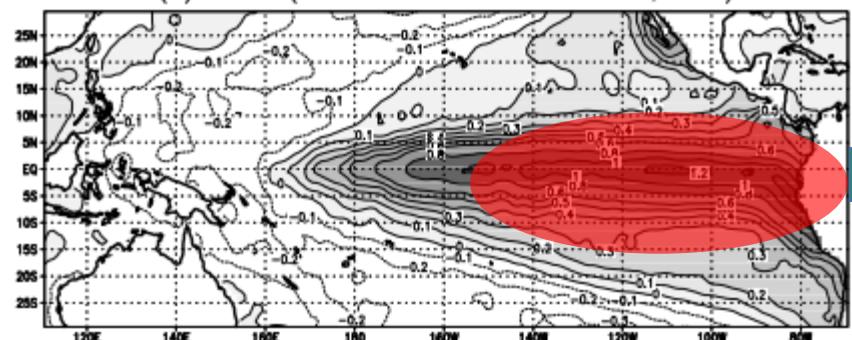
- **1. Motivation**
- **2. NFSV-assimilation and NFSV-ICM**
- **3. Prediction for ENSO diversities**
- **4. The dynamic roles acted by NFSV-perturbation**
- **5. Summary**

# 1. Motivation

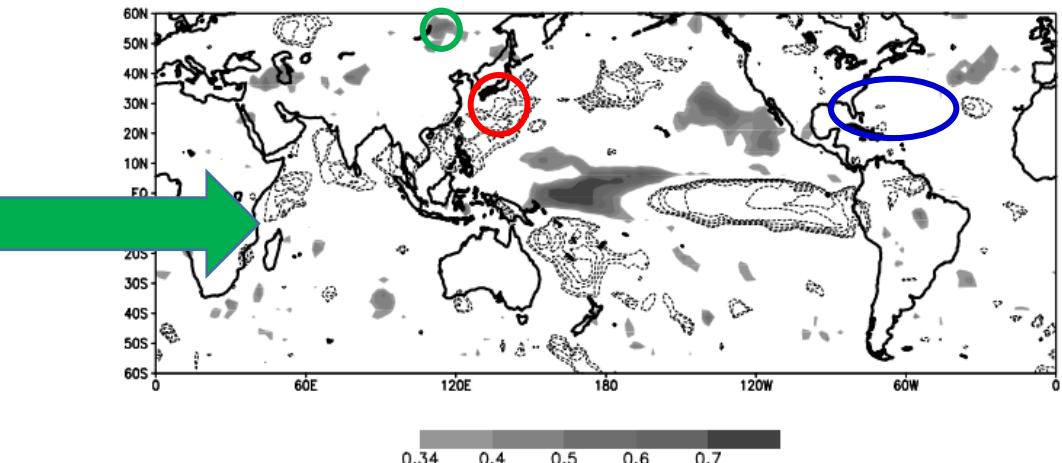
## CP El Niño



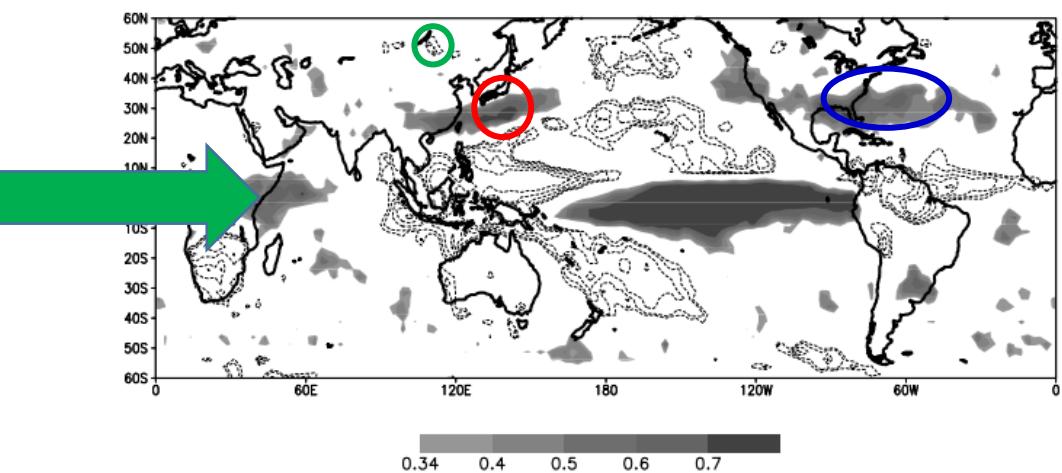
## EP El Niño



Ashork et al., JGR, 2007

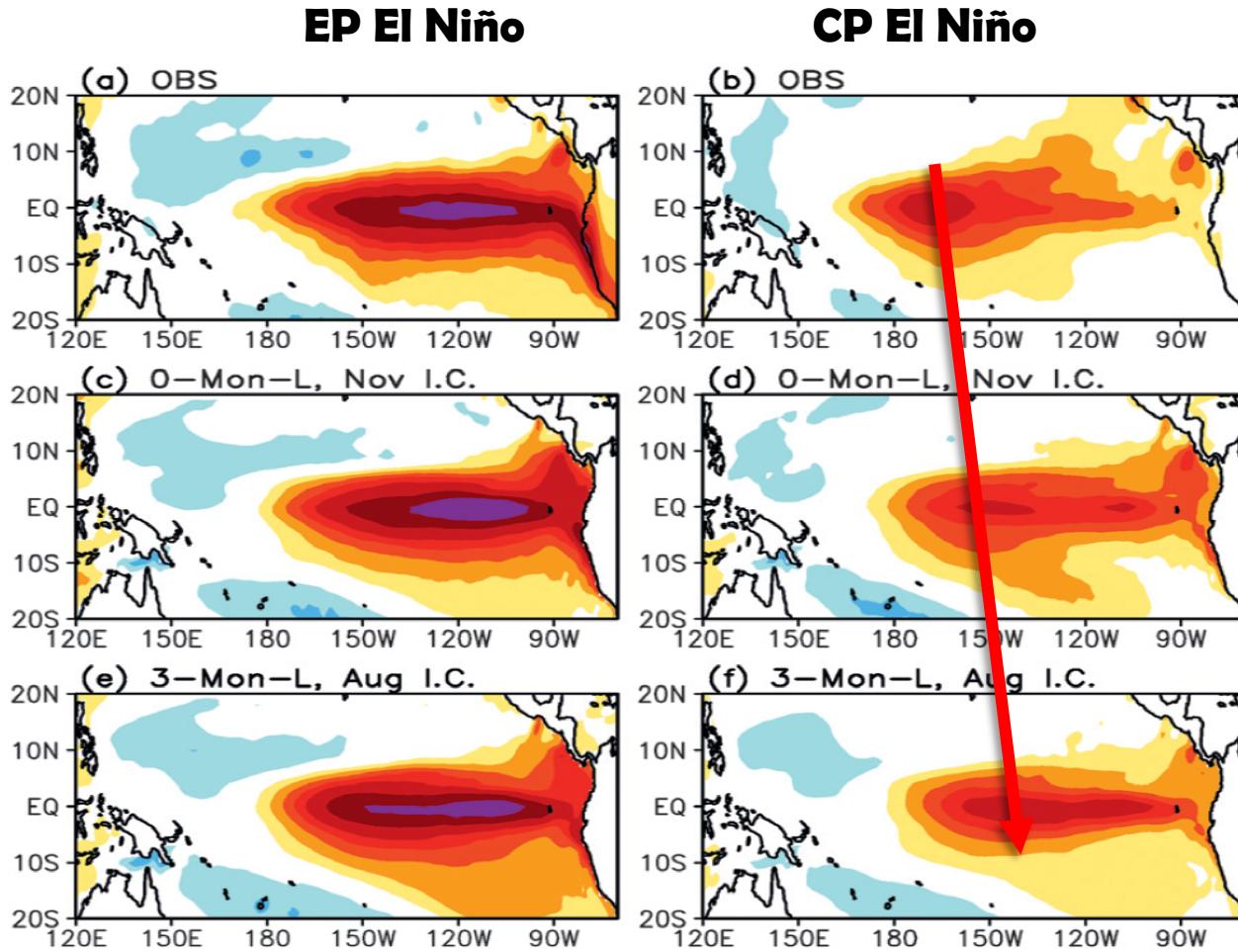


(b)



- The CP- and EP-El Niño induced climate effects are different and even opposite somewhere

# 1. Motivation



Xue et al., 2013, JC

- Most models can only distinguish/predict two types of El Niño at 3-month lead time even 1-month in some models (e.g., Hendon et al., 2009; Jeong et al., 2012; Yang and Jiang, 2014; Zheng and Yu, 2017; Ren et al., 2017)

## To improve the prediction skills regarding to ENSO diversities...

### To improve the initial conditions

- More observations
- EnKF
- 4D-Var
- Hybrid data assimilation, e.g., 4DEnVar

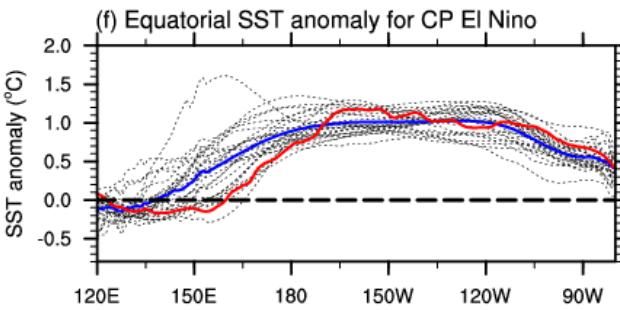
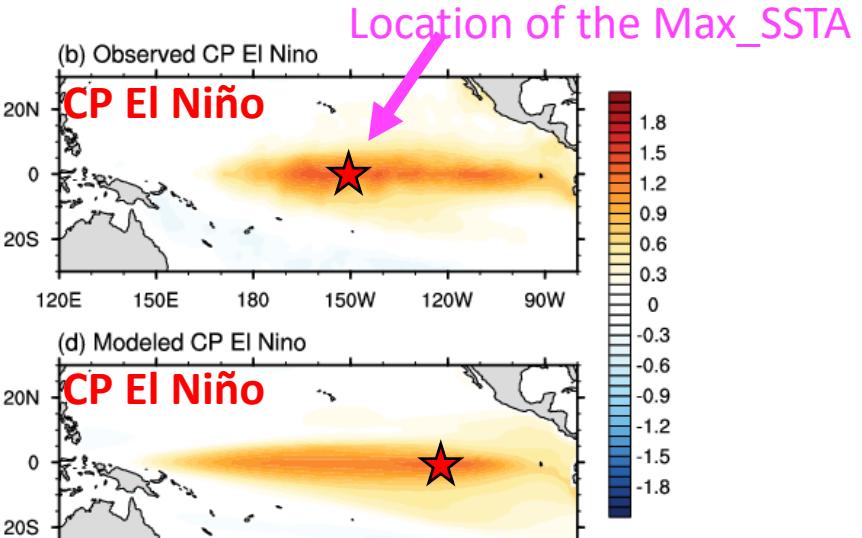
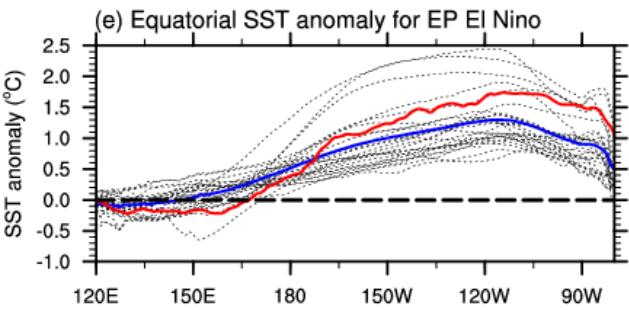
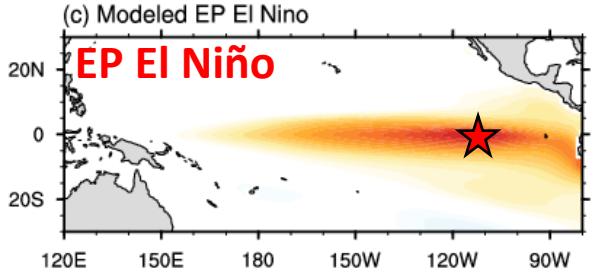
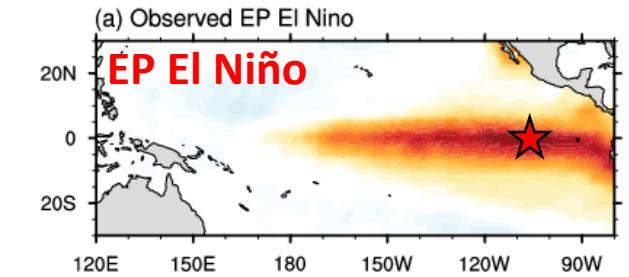
### BUT...

- The persistent barrier for EP El Niño is larger than CP El Niño (Zhang et al., 2014; Ren et al., 2016)
- Initial error-induced error growth during CP El Niño prediction is less than during EP El Niño prediction (Tian and Duan, 2016).

The CP El Niño is not as sensitive to initial condition as EP El Niño

# 1. Motivation

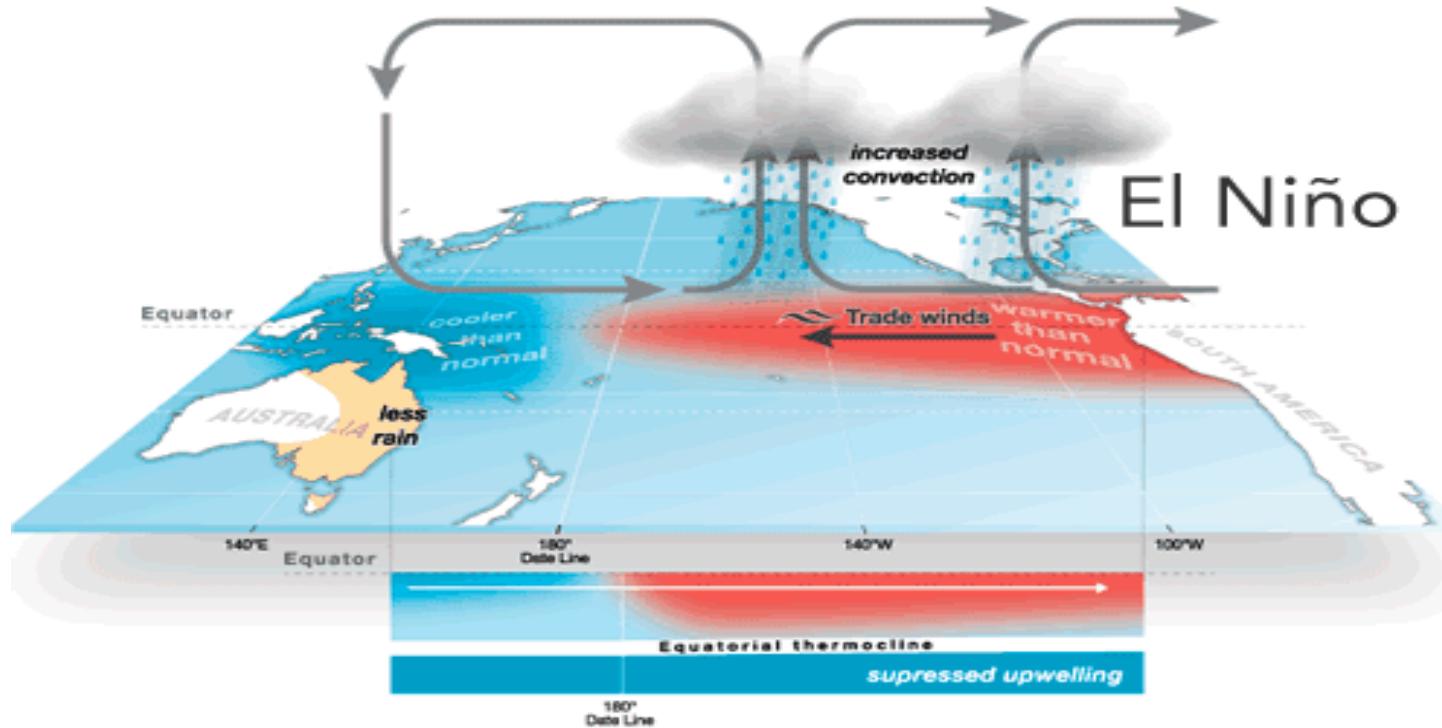
**Observed**



- **CMIP3 and CMIP5 models still have large bias in CP El Niño simulation** (e.g., Kug et al., 2010; Yu and Kim, 2011; Ham and Kug, 2002; Capotondi, 2013; Ham and Kug, 2012; Taschetto et al., 2014, Xu et al., 2017)
- **Improving model will effectively improve the CP El Niño prediction**

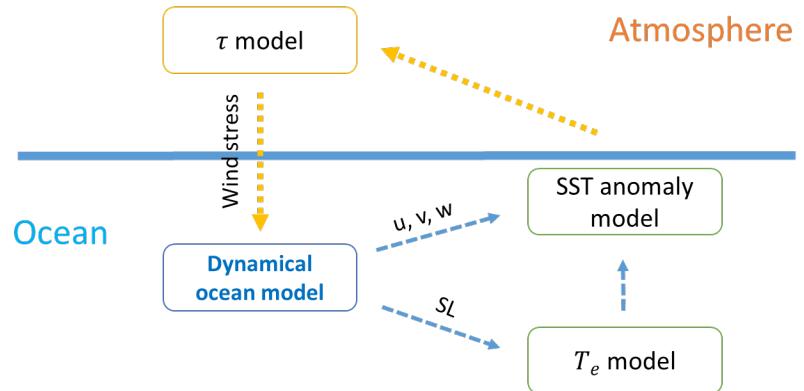
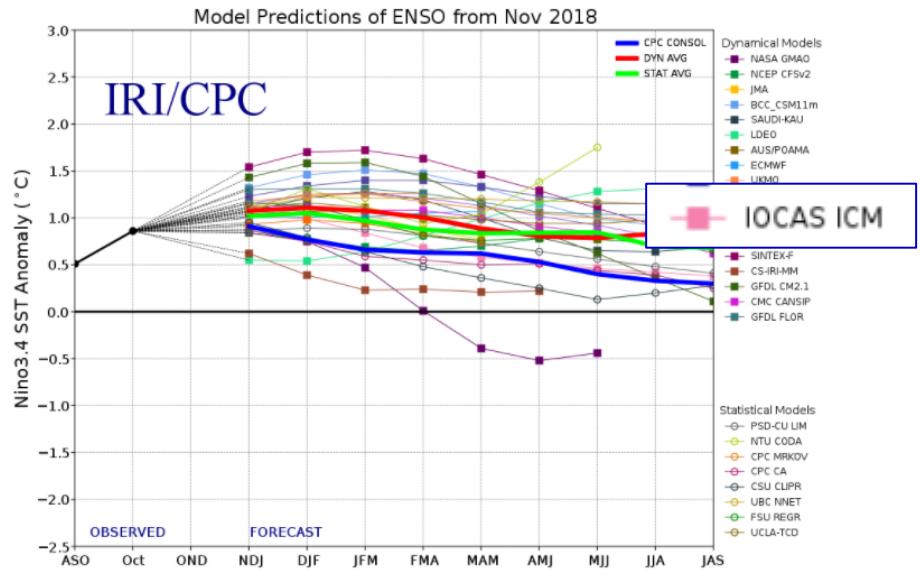
# 1. Motivation

## Source of model errors



- Air-sea coupling
- Convection and cloud
- Ocean mixing
- Extra-tropical forcing
- Subgrid scale parameterization
- ...

## 2.1 IOCAS ICM



ICM (Zhang et al., 2003; Zheng et al., 2010)

### Simplified or missing processes in the ICM

- Air-sea heat interaction
- The stochastic atmospheric forcing (i.e., MJO)
- The fresh water flux (FWF) and salinity
- Tropical instability waves (TIW)
- Subtropical-tropical teleconnection
- Interaction between multi oceanic basins
- Ocean ecological/biological feedback



Model uncertainties are complicated and diverse

- How to give a comprehensive consideration to model errors and correct the model?
- Nonlinear forcing singular vector-assimilation (Duan et al., 2014)

## 2.2 NFSV-assimilation

### Control prediction

Forecast model:

$$\begin{cases} \frac{\partial \mathbf{u}}{\partial t} = \mathbf{F}(\mathbf{u}, t) \\ \mathbf{u}|_{t=0} = \mathbf{u}_0 \end{cases}$$

Prediction:

$$\mathbf{u}(\mathbf{x}, \tau) = \mathbf{M}_\tau(\mathbf{u}_0)$$

Prediction error:

$$E = \sum_{t=t_0}^{\tau} \|\mathbf{u}^p(\mathbf{x}, t) - \mathbf{u}^{obs}(\mathbf{x}, t)\| \gg \varepsilon$$

### NFSV-assimilation

Forecast model with forcing vector  $\mathbf{f}$ :

$$\begin{cases} \frac{\partial \mathbf{u}}{\partial t} = \mathbf{F}(\mathbf{u}, t) + \mathbf{f}(\mathbf{x}, t) \\ \mathbf{u}|_{t=0} = \mathbf{u}_0 \end{cases}$$

Prediction:

$$\mathbf{u}(\mathbf{x}, \tau) = \mathbf{M}_\tau(\mathbf{f}_\tau)(\mathbf{u}_0)$$

Duan et al., 2014, *Clim Dyn*

Prediction error:

$$E(\mathbf{f}_t^*) = \min \sum_{t=t_0}^{\tau} \|\mathbf{M}_t(\mathbf{f}_t, \mathbf{u}_0) - \mathbf{u}^{obs}(\mathbf{x}, t)\| \approx 0$$

$\mathbf{f}_t^*$  is the NFSV-tendency error, which represents the combined effect of kinds of model errors

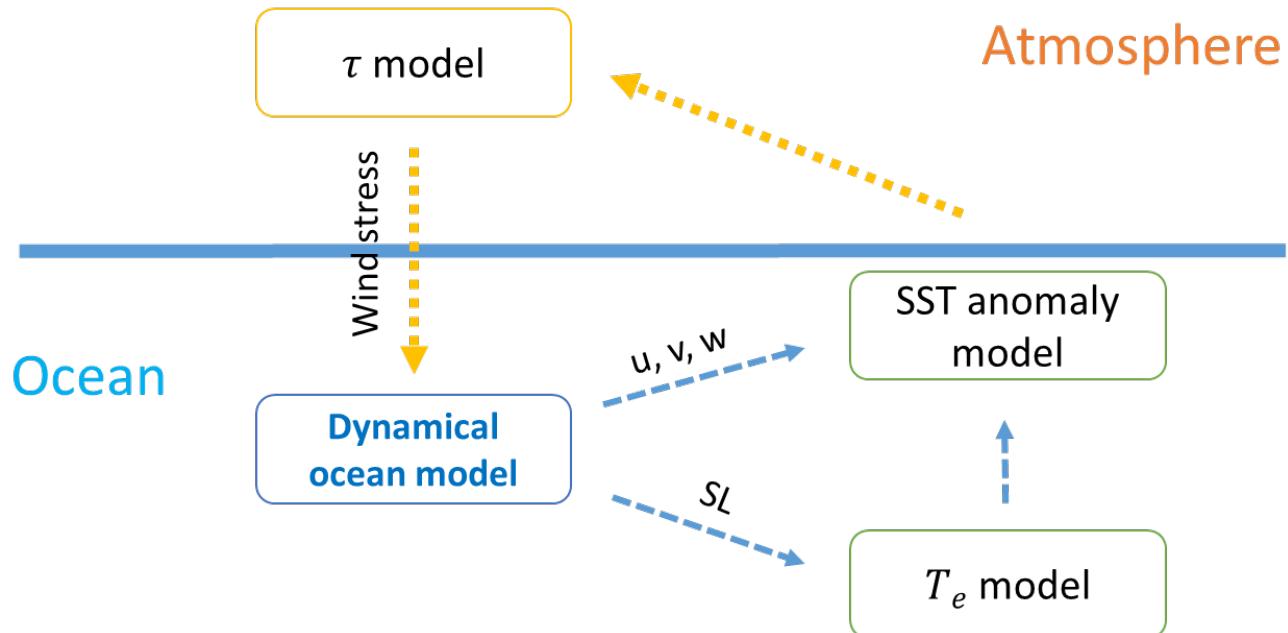
## 2.2 NFSV-assimilation

$$E(\mathbf{f}_t^*) = \min \sum_{t=t_0}^{\tau} \| \mathbf{M}_t(\mathbf{f}_t, \mathbf{u}_0) - \mathbf{u}^{obs}(\mathbf{x}, t) \|$$

- **SST data:** ERSSTv3b (1854 - now)

**SST model**

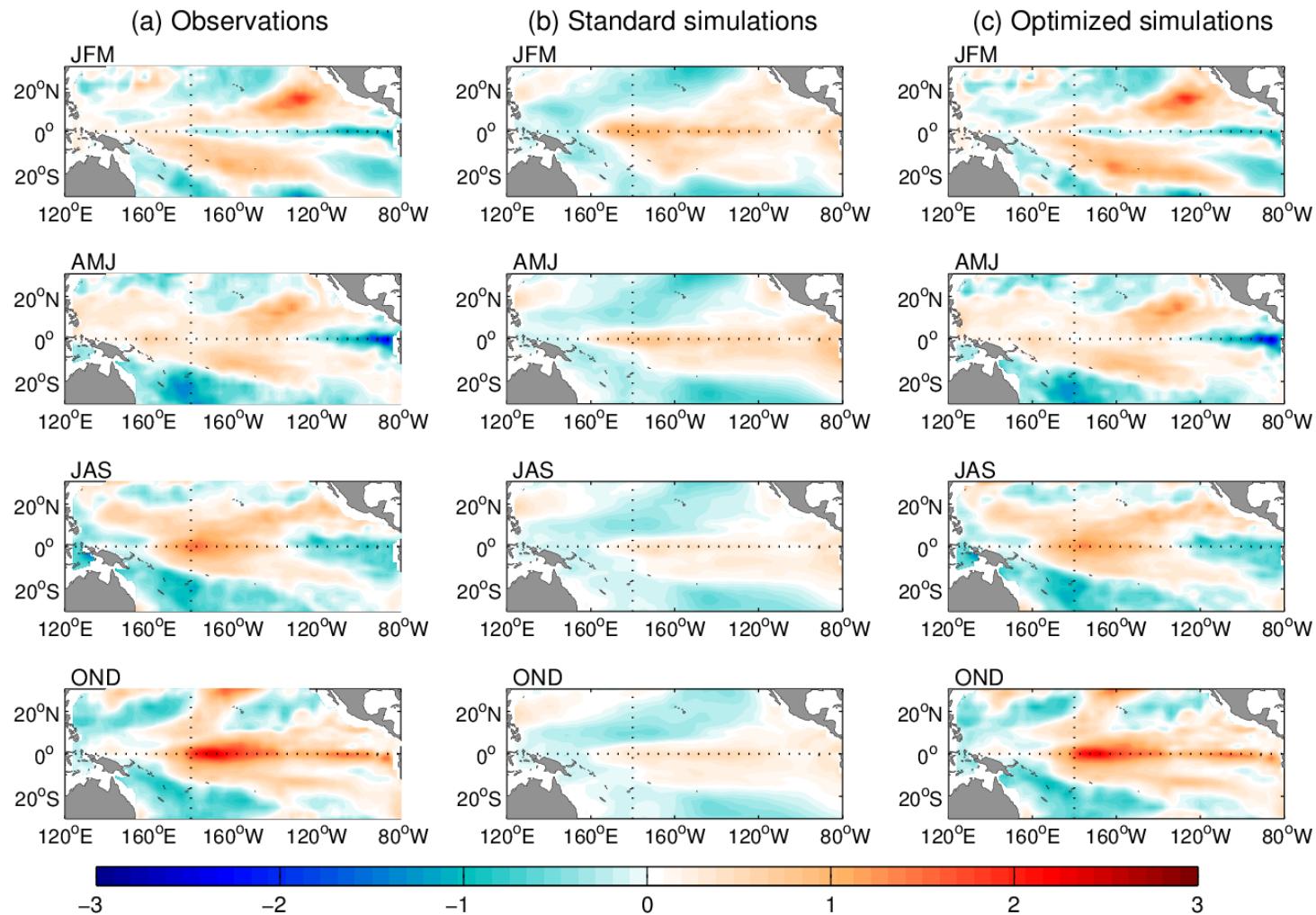
$$\frac{\partial T}{\partial t} = G(T, u, v, h, p) + \mathbf{F}$$



## 2.2 NFSV-assimilation: ENSO simulation

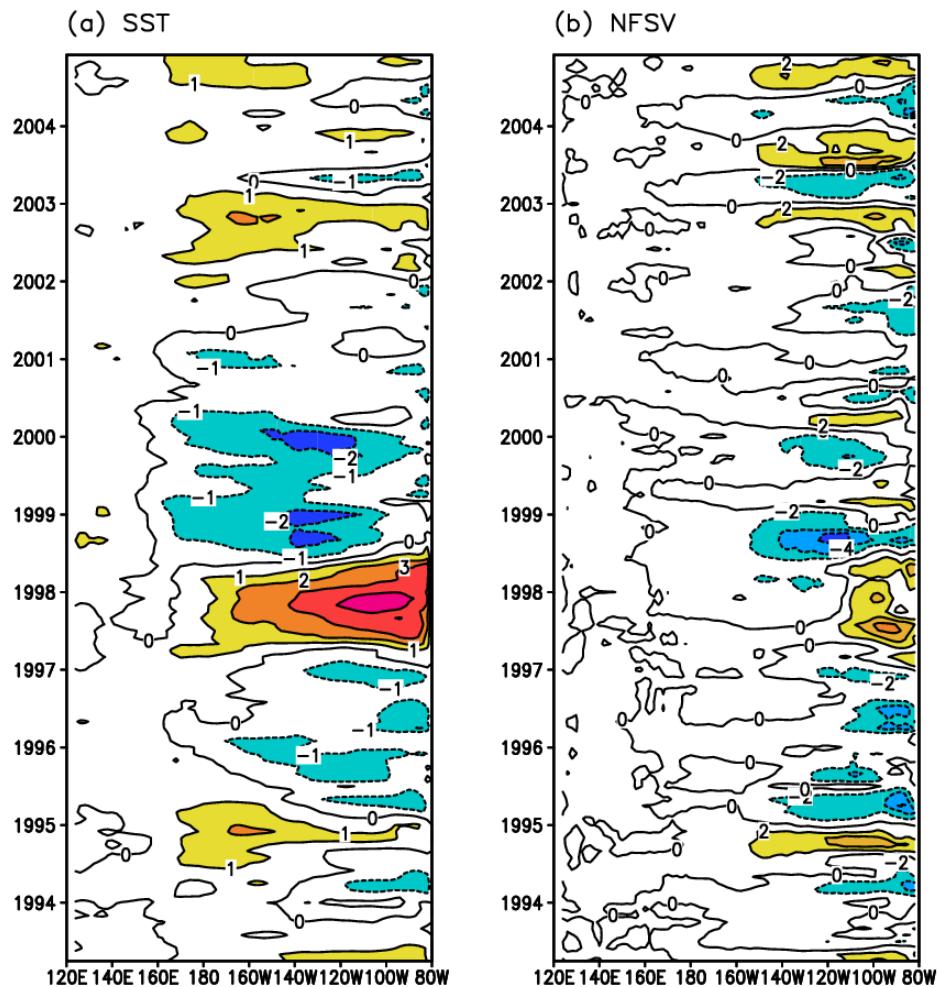


# CP El Niño: 1994/1995



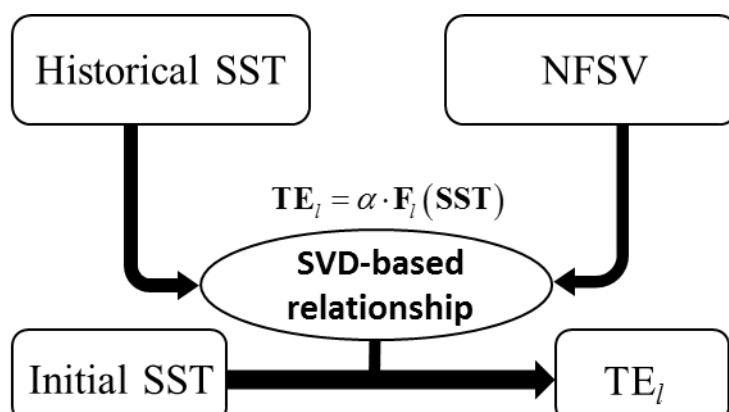
## 2.3 NFSV-tendency error forecast model

1960 Training period 1996 Prediction period 2018



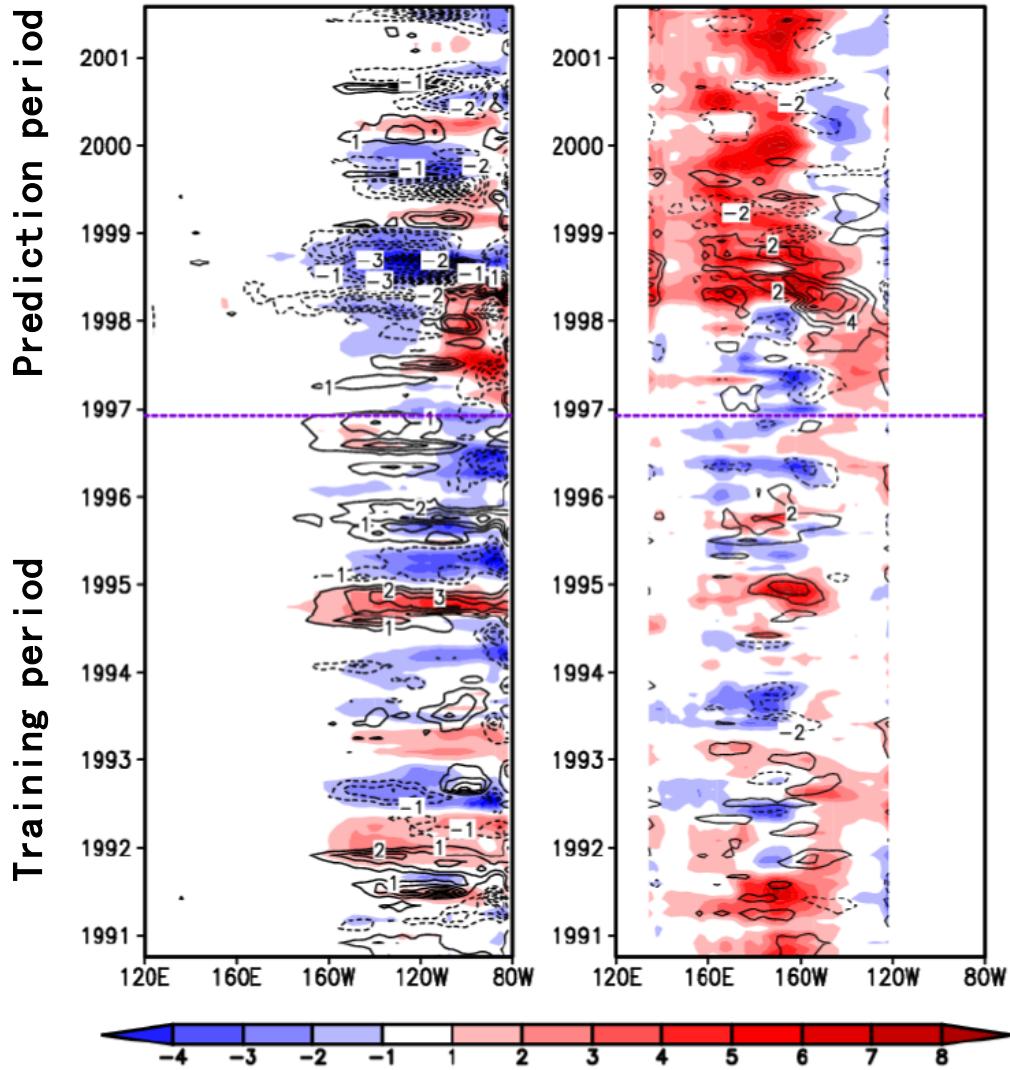
$$E(\mathbf{f}_t^*) = \min \sum_{t=t_0}^{\tau} \| \mathbf{M}_t(\mathbf{f}_t, \mathbf{u}_0) - \mathbf{u}^{obs}(\mathbf{x}, t) \|$$

- SST data:** ERSSTv3b (1854 - now)
- Training period:** 1960-1996
- Relationship:** SST-NFSV



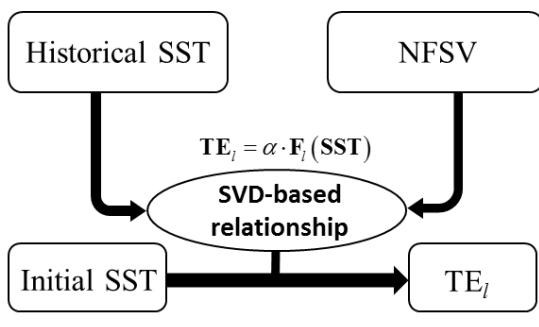
NFSV-tendency error forecast model

## Estimated NFSV-error at 6-month lead time



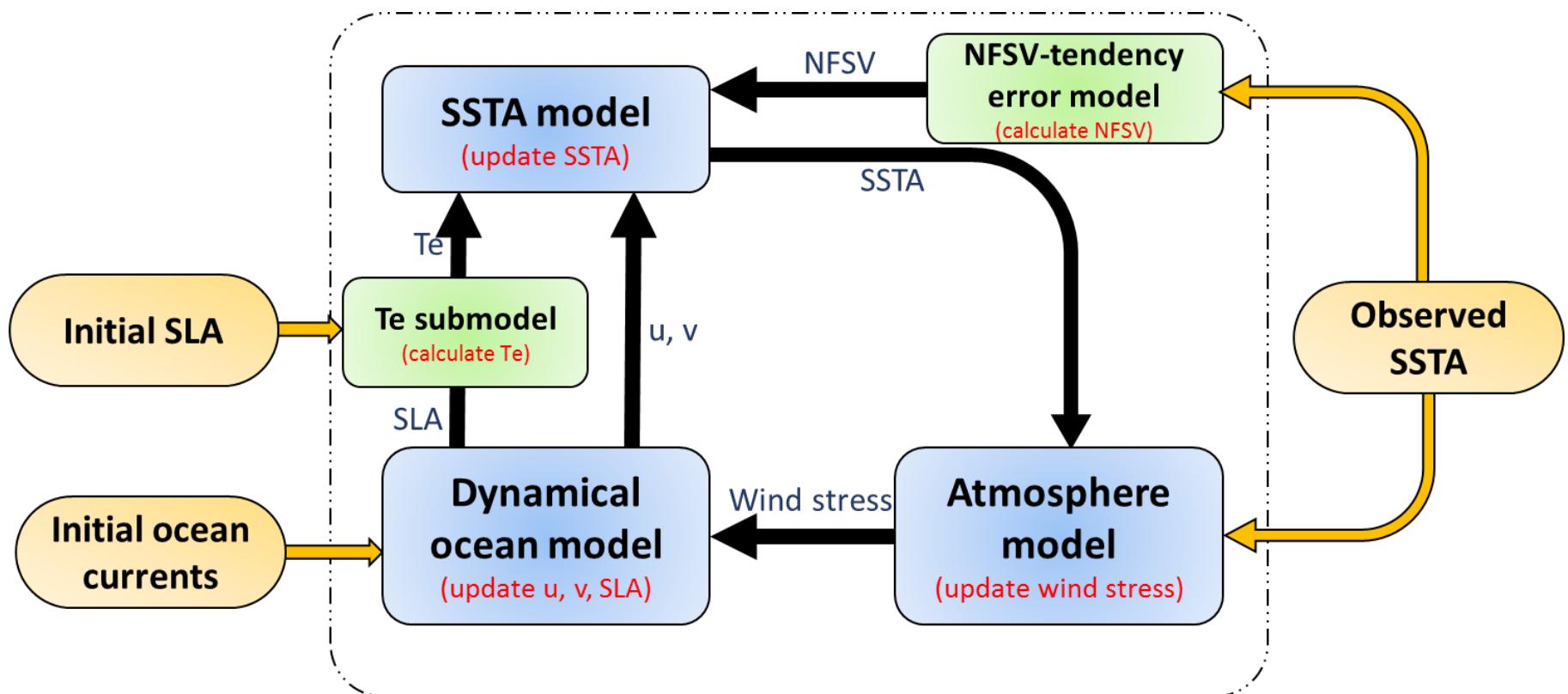
**Shaded:** NFSV-tendency error determined by the NFSV-assimilation

**Contours:** NFSV-tendency error determined by the constructed error model



Error forecast model

## 2.4 NFSV-ICM



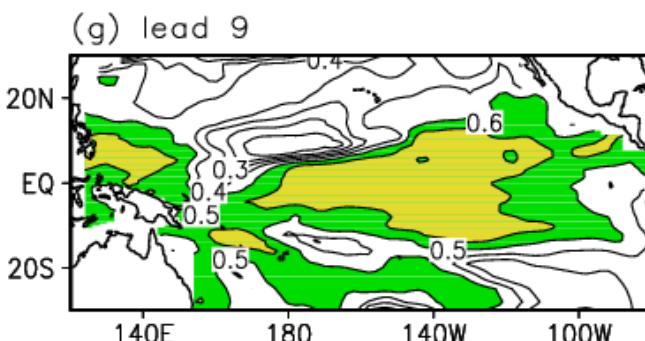
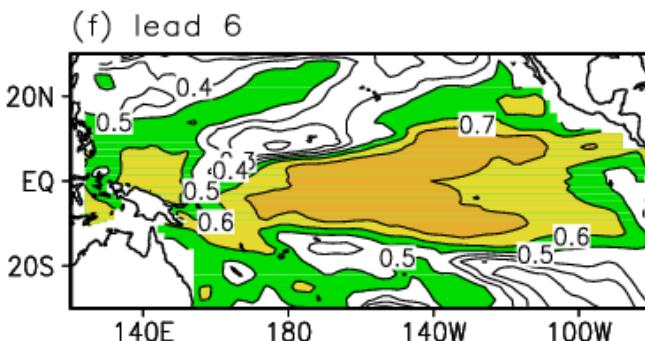
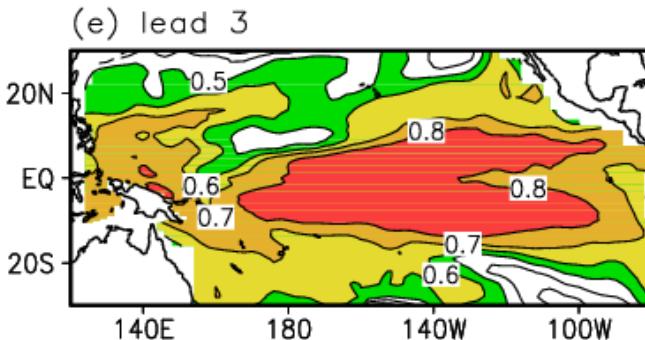
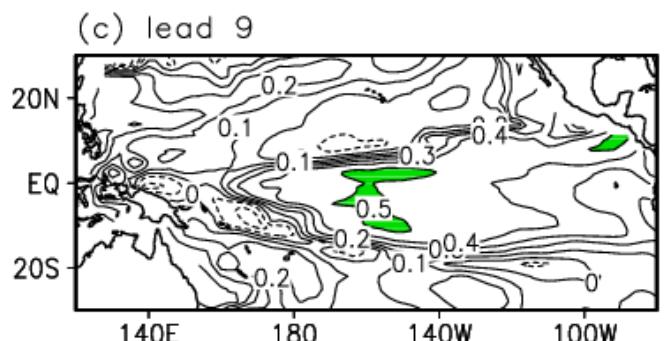
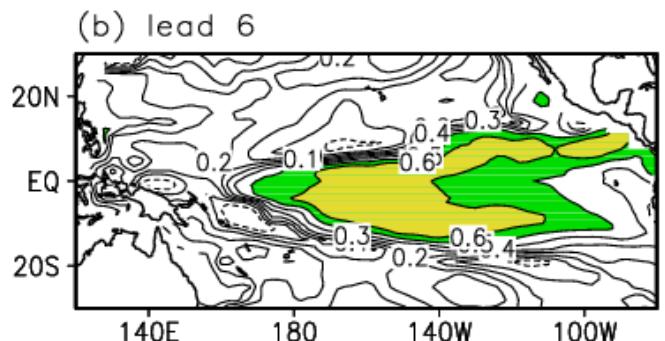
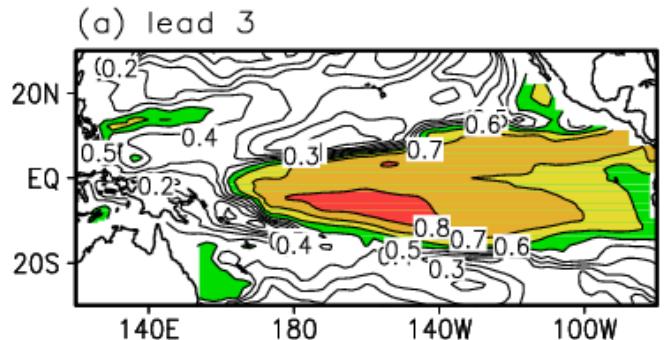
A new ENSO forecast system (NFSV-ICM)

### 3.1 Mean prediction skill

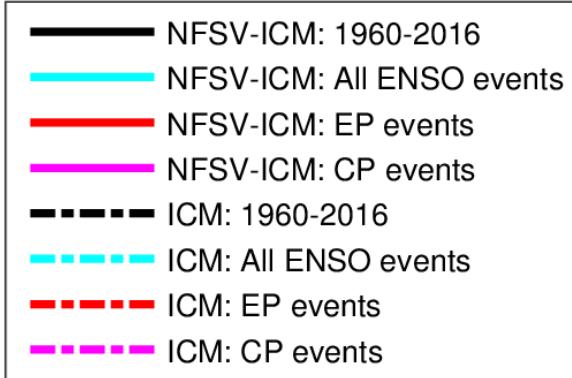
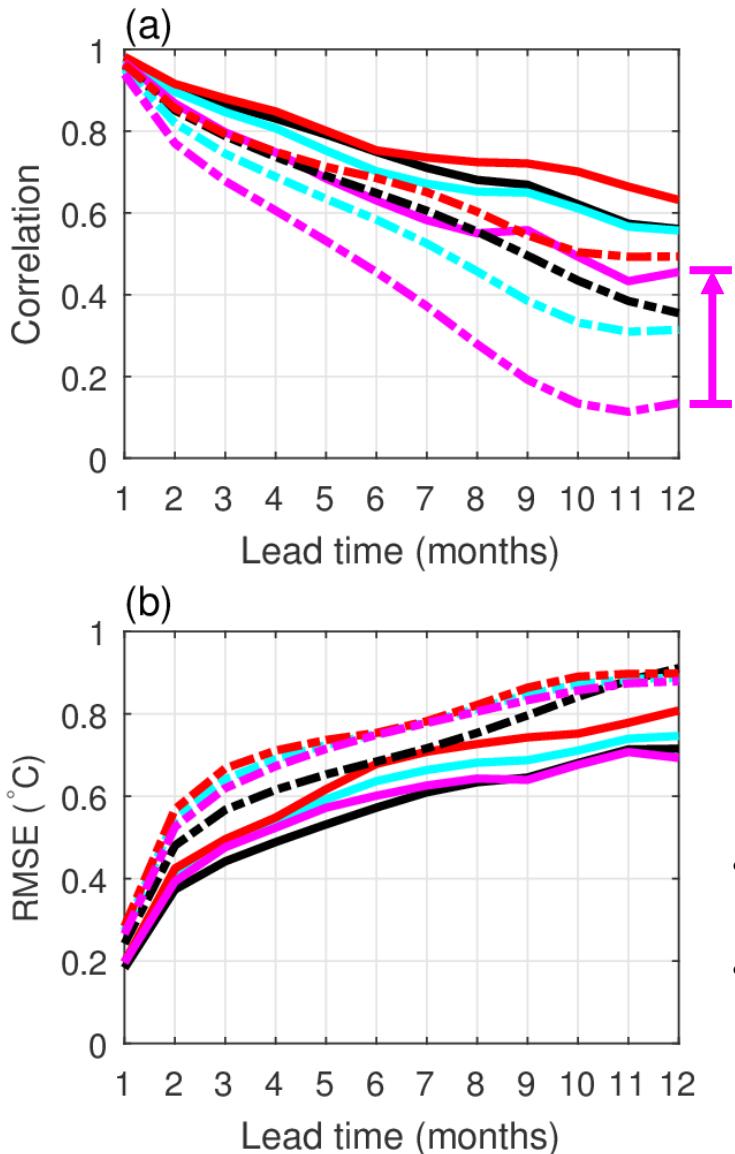
**ICM**

**NFSV-ICM**

Correlations between observed and predicted SSTA



### 3.2 ENSO diversity predictions



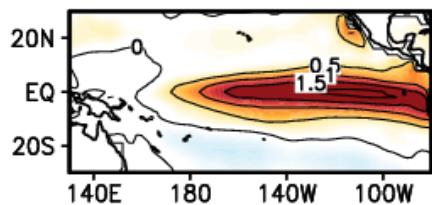
**SSTA:** 7 months  $\Rightarrow$  10 months  
**ENSO:** 5 months  $\Rightarrow$  10 months  
**EP ENSO:** 8 months  $\Rightarrow$  12 months  
**CP ENSO:** 4 months  $\Rightarrow$  7 months

- The prediction skills for ENSO-related SSTA is significantly improved.
- The improvement of ENSO prediction skills is attributed to the high performance for the CP ENSO: lower RMSE and higher correlation coefficient

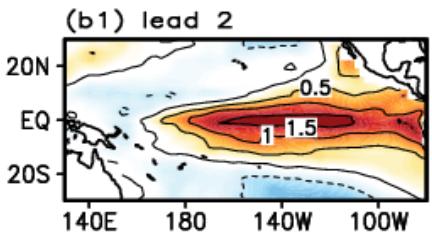
### 3.2 ENSO diversity predictions

#### Composite EP El Niño event

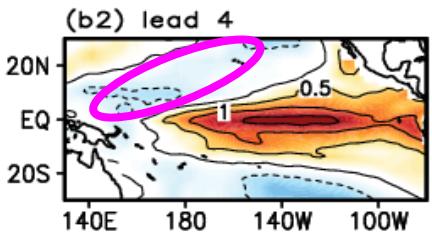
(a) Obs.



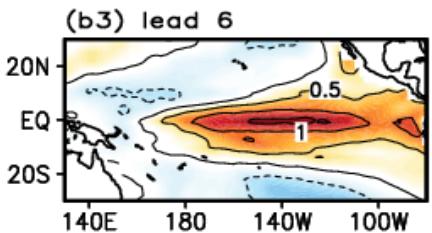
(b) ICM



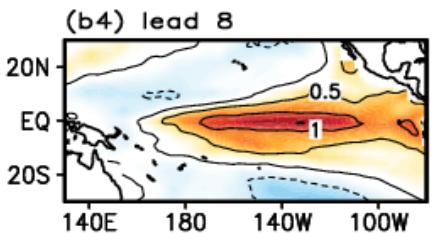
(b1) lead 2



(b2) lead 4

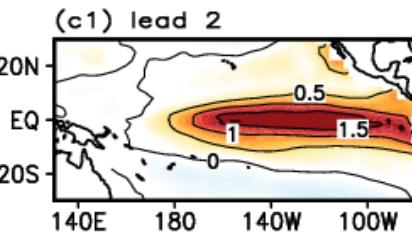


(b3) lead 6

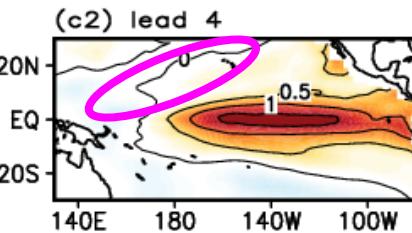


(b4) lead 8

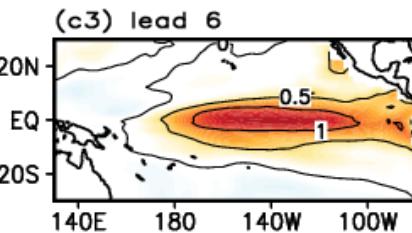
(c) NFSV–ICM



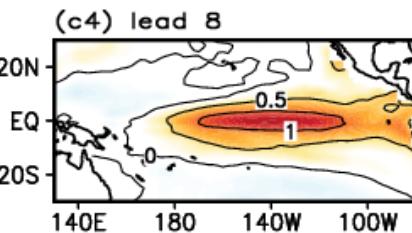
(c1) lead 2



(c2) lead 4

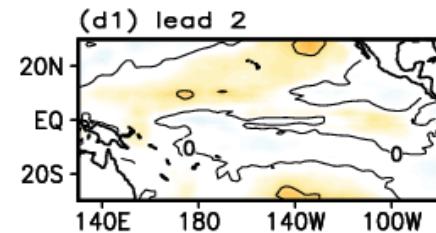


(c3) lead 6

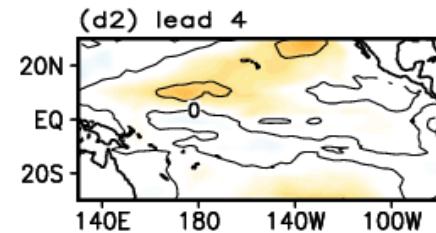


(c4) lead 8

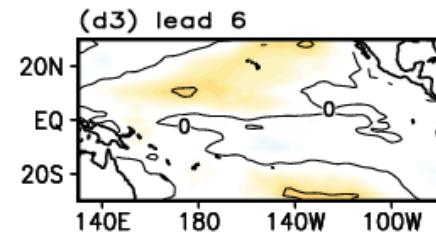
(d) Diff.



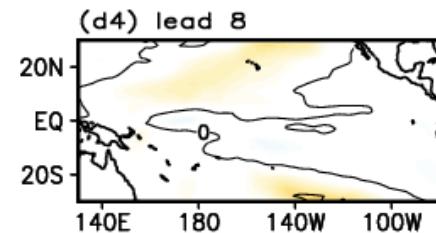
(d1) lead 2



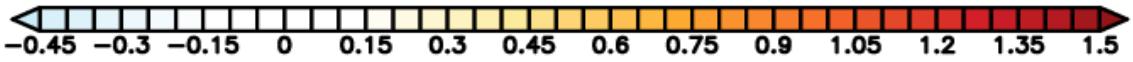
(d2) lead 4



(d3) lead 6

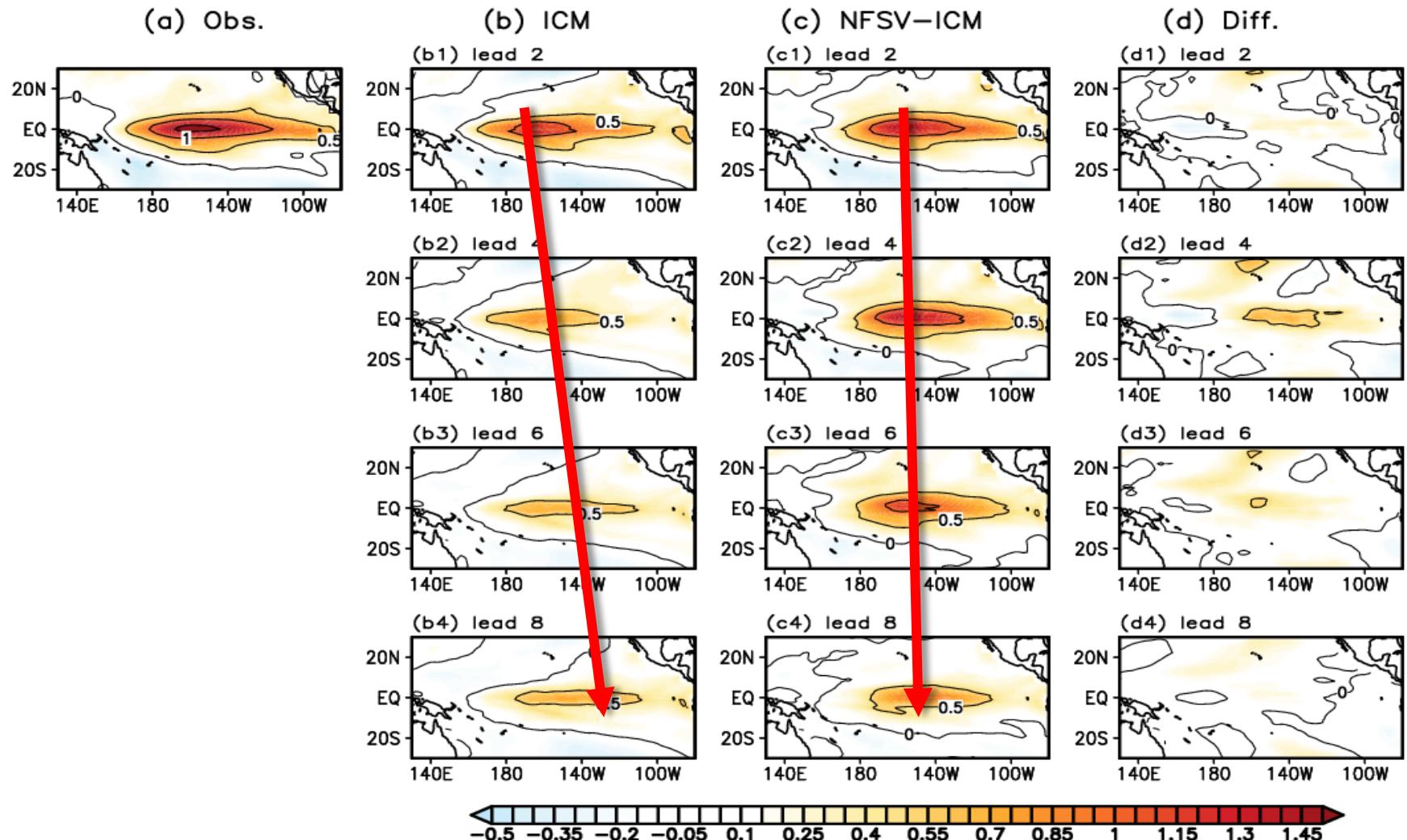


(d4) lead 8



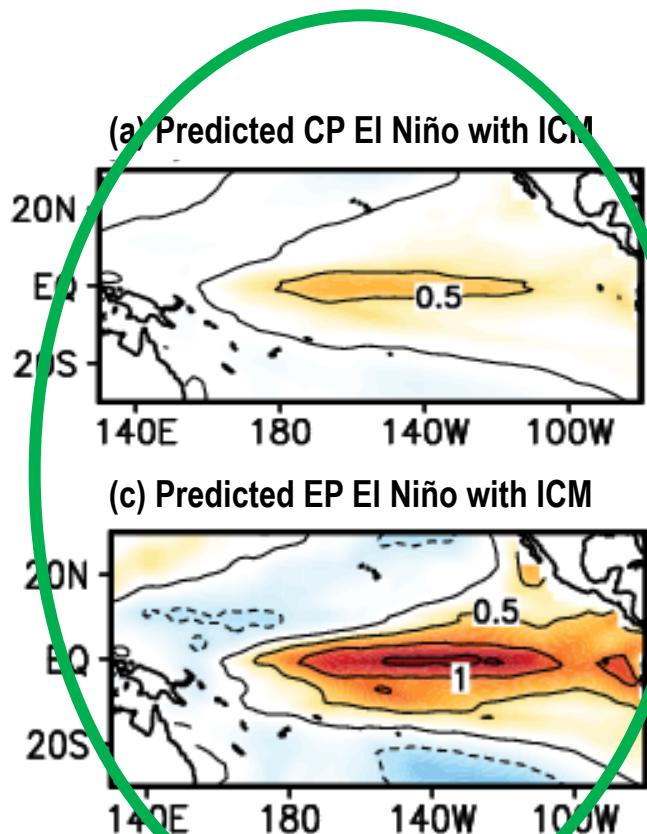
### 3.2 ENSO diversity predictions

#### Composite CP El Niño event

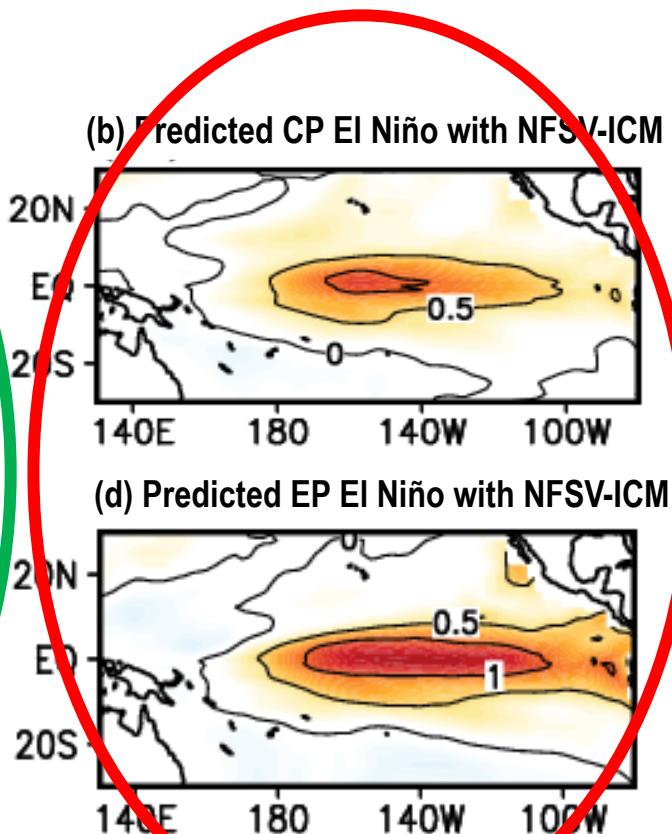


### 3.2 ENSO diversity predictions

**ICM**



**NFSV-ICM**

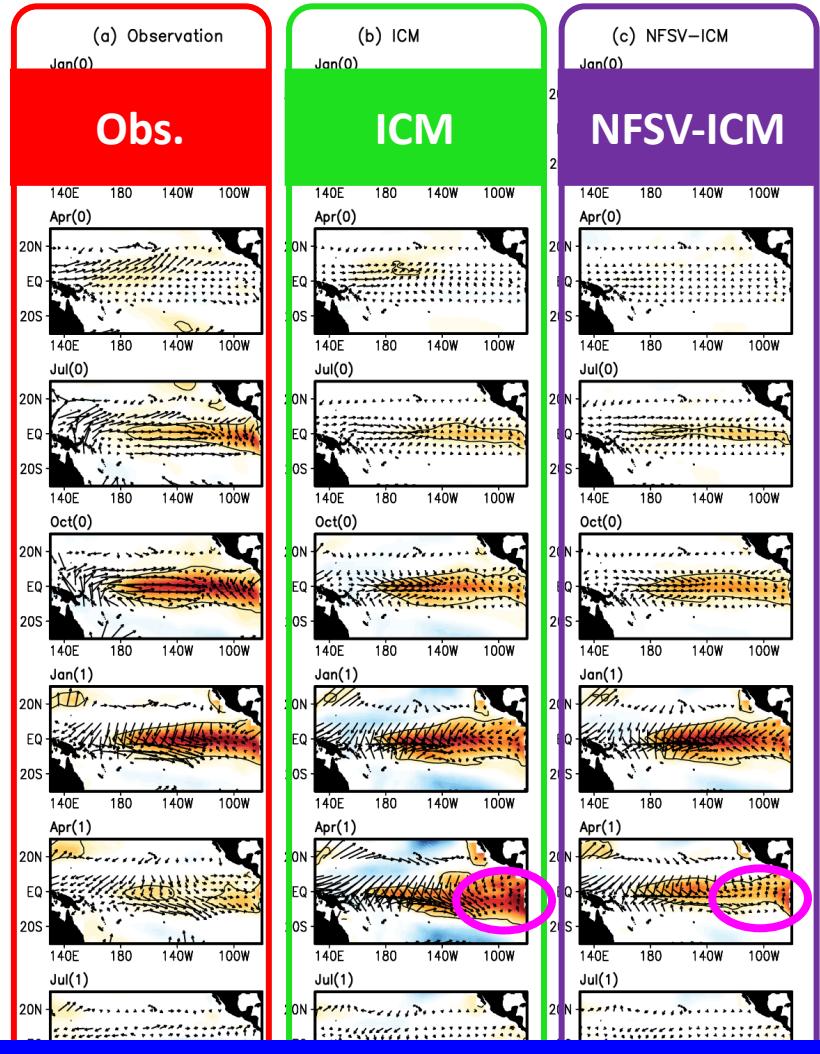


- Prediction of the major differences in pattern of SSTA between two types of El Niño is extended to larger than 2-season lead time.

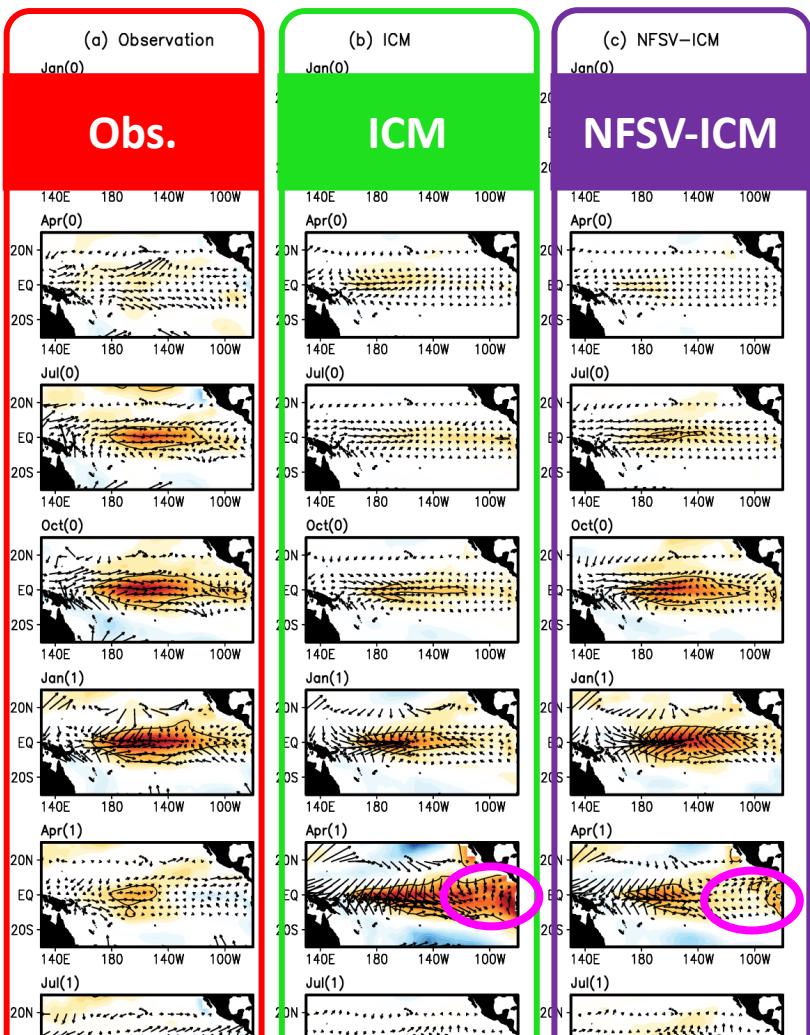
## 3.2 ENSO diversity predictions



# EP El Niño



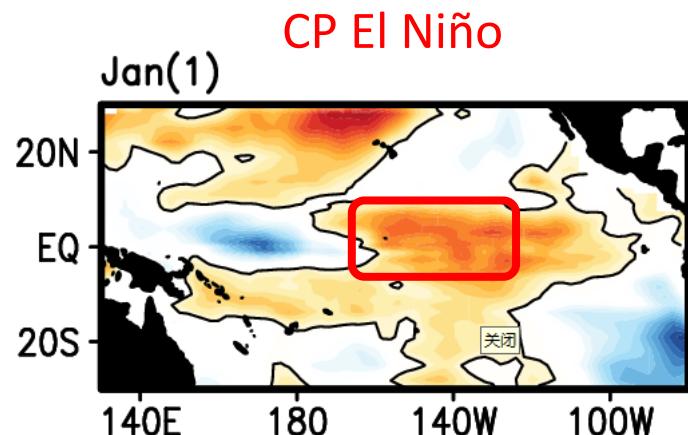
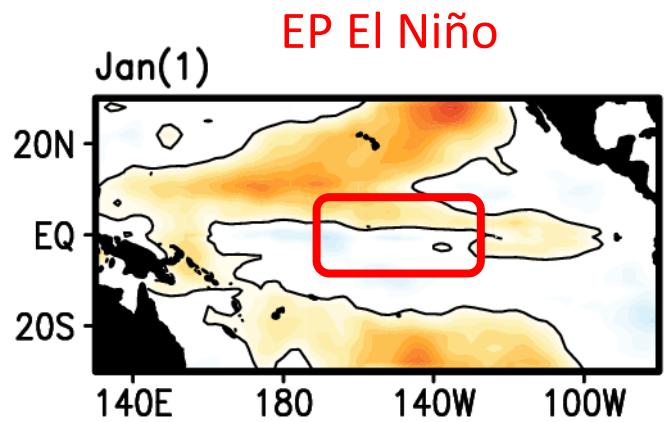
# CP El Niño



- NFSV-ICM can not only improve the identification to the El Niño diversities but also their evolutions

## 4. Dynamics of NFSV-tendency error: CP El Niño

### NFSV-tendency perturbation effect on SST predictions



$$\frac{\partial T'}{\partial t} = -u' \frac{\partial \bar{T}}{\partial x} - \bar{u} \frac{\partial T'}{\partial y} - u' \frac{\partial T'}{\partial y}$$

**Zonal advection terms**

$$-v' \frac{\partial \bar{T}}{\partial y} - \bar{v} \frac{\partial T'}{\partial y} - v' \frac{\partial T'}{\partial y}$$

**Meridional advection terms**

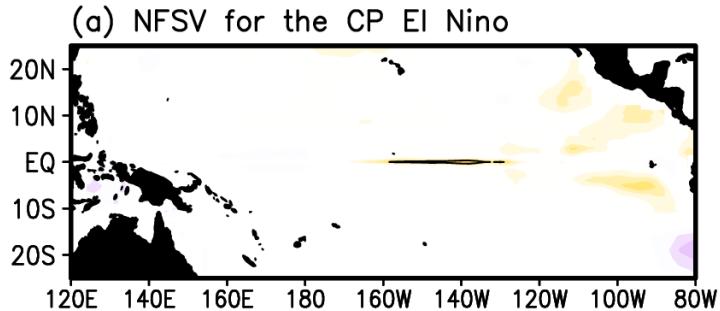
$$-w' \frac{\partial \bar{T}}{\partial z} - \bar{w} \frac{\partial T'}{\partial z} - w' \frac{\partial T'}{\partial z}$$

**Vertical advection terms**

+Q

# 4. Dynamics of NFSV-tendency error: CP El Niño

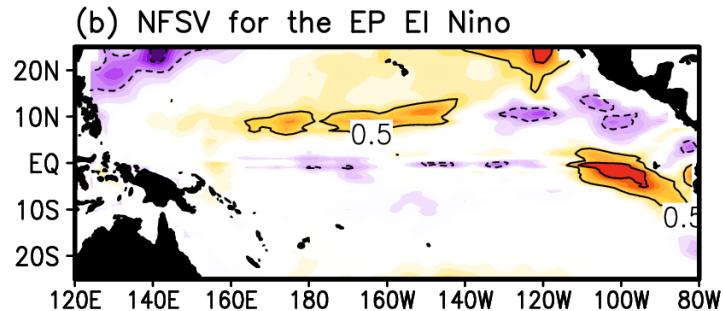
## NFSV in CP El Niño prediction



$-\bar{v} \cdot \frac{\partial T'}{\partial y}$ : Meridional heat transport

$-\bar{u} \cdot \frac{\partial T'}{\partial x}$ : SSTA transport by SEC

## NFSV in EP El Niño prediction



$-\bar{v} \cdot \frac{\partial T'}{\partial y}$ : Meridional heat transport

$-\bar{u} \cdot \frac{\partial T'}{\partial x}$ : SSTA transport by SEC

## Changes of the heat budget

|                            | $-\bar{u} \cdot \frac{\partial T'}{\partial x}$ | $-u' \cdot \frac{\partial \bar{T}}{\partial x}$ | $-u' \cdot \frac{\partial T'}{\partial x}$ | total  |
|----------------------------|---|---|--|--------|
| Zonal advection terms      | <b>1.40</b>                                     | <b>3.17</b>                                     | -1.32                                      | 3.25   |
| Meridional advection terms | <b>5.09</b>                                     | 0.19  | 0.01                                       | 5.10   |
| Vertical advection terms   | <b>-10.56</b>                                   | 0.13  | -0.11                                      | -10.54 |
| dT/dt                      |   |   |  | 9.5    |



$-\mathbf{u}' \cdot \frac{\partial \bar{T}}{\partial x}$ : Zonal advection feedback



$-\bar{w} \cdot \frac{\partial T'}{\partial z}$ : Thermocline feedback



## Changes of the heat budget

|                            | $-\bar{u} \cdot \frac{\partial T'}{\partial x}$ | $-u' \cdot \frac{\partial \bar{T}}{\partial x}$ | $-u' \cdot \frac{\partial T'}{\partial x}$ | total |
|----------------------------|---|---|--|-------|
| Zonal advection terms      | <b>1.25</b>                                     | <b>-2.67</b>                                    | 0.80                                       | -0.61 |
| Meridional advection terms | <b>-5.50</b>                                    | -0.20   | 3.04                                       | -2.75 |
| Vertical advection terms   | <b>1.81</b>                                     | 0.03  | -0.17                                      | 1.61  |
| dT/dt                      |   |   |  | -2.60 |



$-\mathbf{u}' \cdot \frac{\partial \bar{T}}{\partial x}$ : Zonal advection feedback

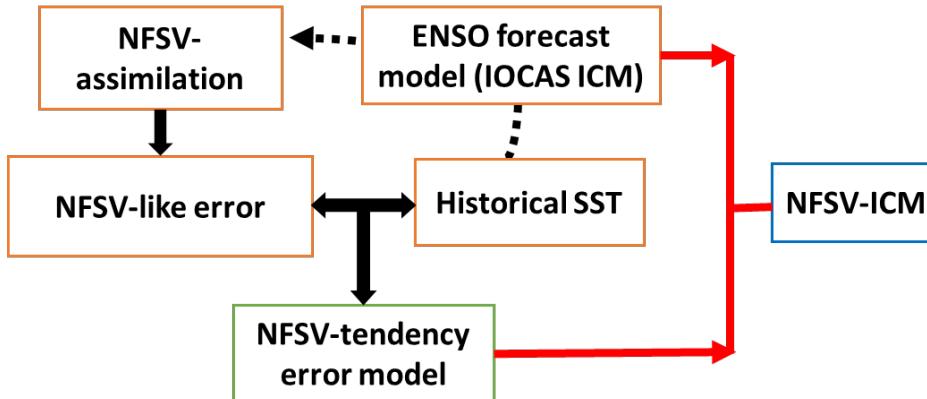


$-\bar{w} \cdot \frac{\partial T'}{\partial z}$ : Thermocline feedback



## 5. Summary

### ➤ How to use NFSV in correcting the ENSO model during predictions?



- The NFSV-ICM can successfully identify CP- and EP- El Niño at least 6-month lead time
- Additional to the local effect by the NFSV-tendency perturbation, it can enhance the thermocline feedback in the EP El Niño prediction but enhance the zonal advection feedback in the CP El Niño prediction.

- Reference:**
1. Tao, L. J., and W. S., Duan, 2019: Using a Nonlinear Forcing Singular Vector Approach to Reduce Model Error Effects in ENSO Forecasting. *Weather and Forecasting*, **34**, 1321-1342.
  2. Tao, L. J., W. S., Duan, and S. Vannitsem, 2019: Improving the forecast skill of El Niño diversity: A nonlinear forcing singular vector approach. *Clim Dynm*, Under Review.

# **Improved predictions for ENSO diversities using a nonlinear forcing singular vector assimilation**

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# Thank you!

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