

Stability of subsurface ocean of Pluto

Jun Kimura¹, and Shunichi Kamata²

¹Osaka University, ²Hokkaido University



Introduction and scope: Signatures for subsurface ocean in Pluto

- Mean surface radius=1188.3 km, Bulk density=1.88 g/cc (indicating 30-35 wt% of H₂O)
- No detectable flattening implies warm/deformable interior (ocean?) [1]
- Reorientation indicates that the ocean currently exists [2]
- Extensional tectonics with strong H₂O spectral signature [3, 4]

Astrobiological/geophysical interests to investigate the conditions which Pluto may retain an ocean.

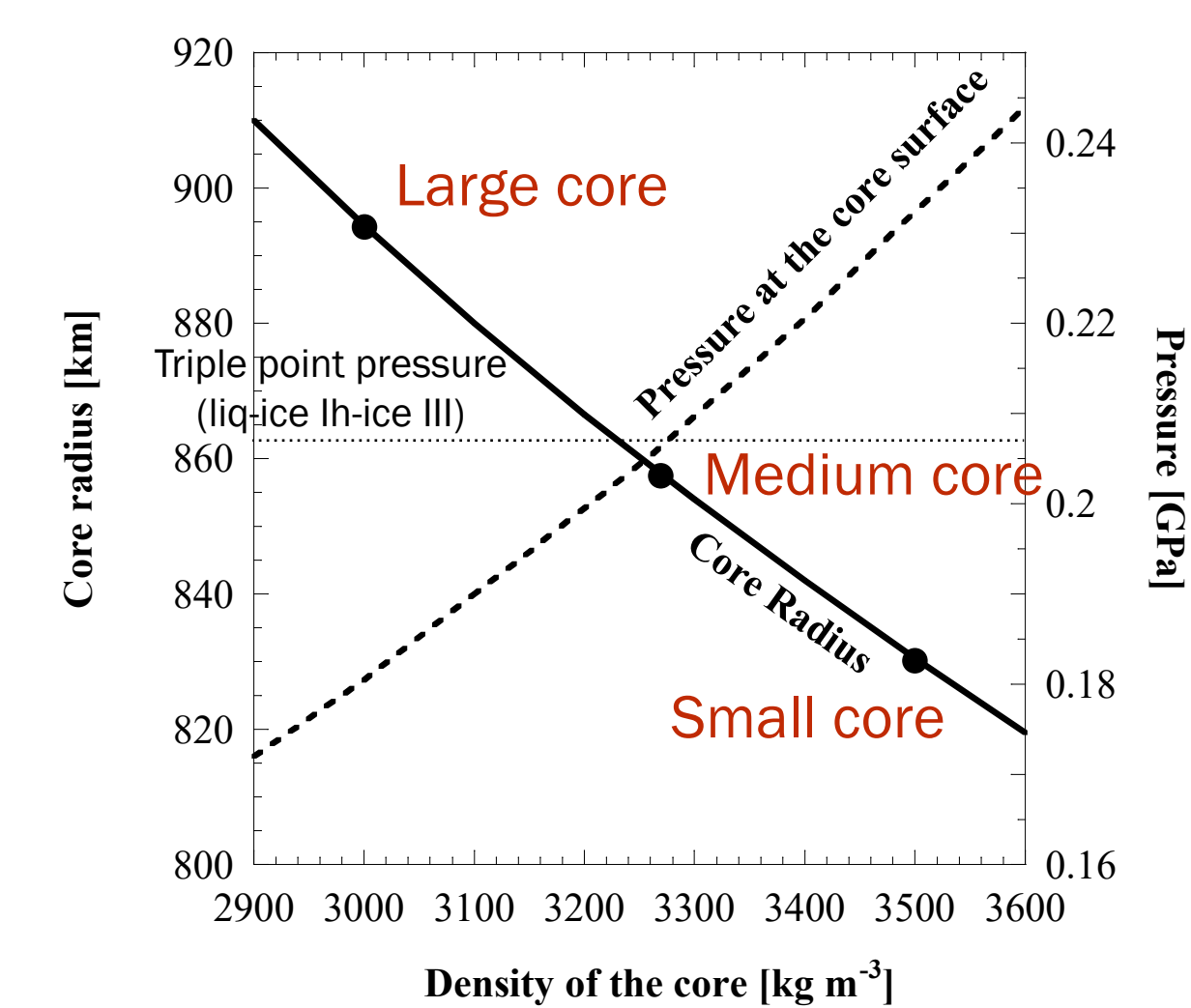
- Former numerical work for the Pluto's ocean [5] considers only a fixed water-rock ratio (only ice-Ih appeared case) and an initially frozen state.

What we did? Further numerical investigations under various parameters for,

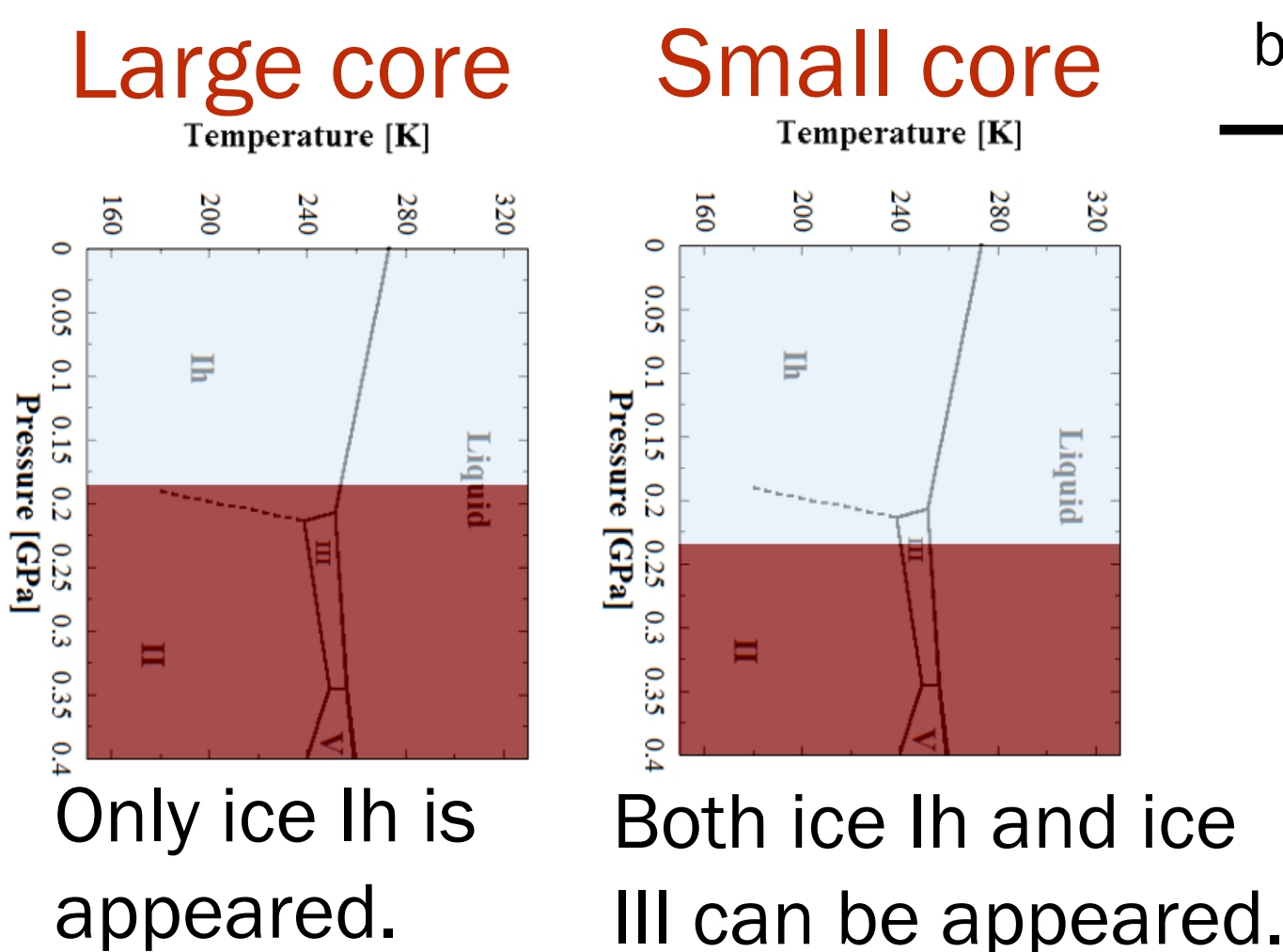
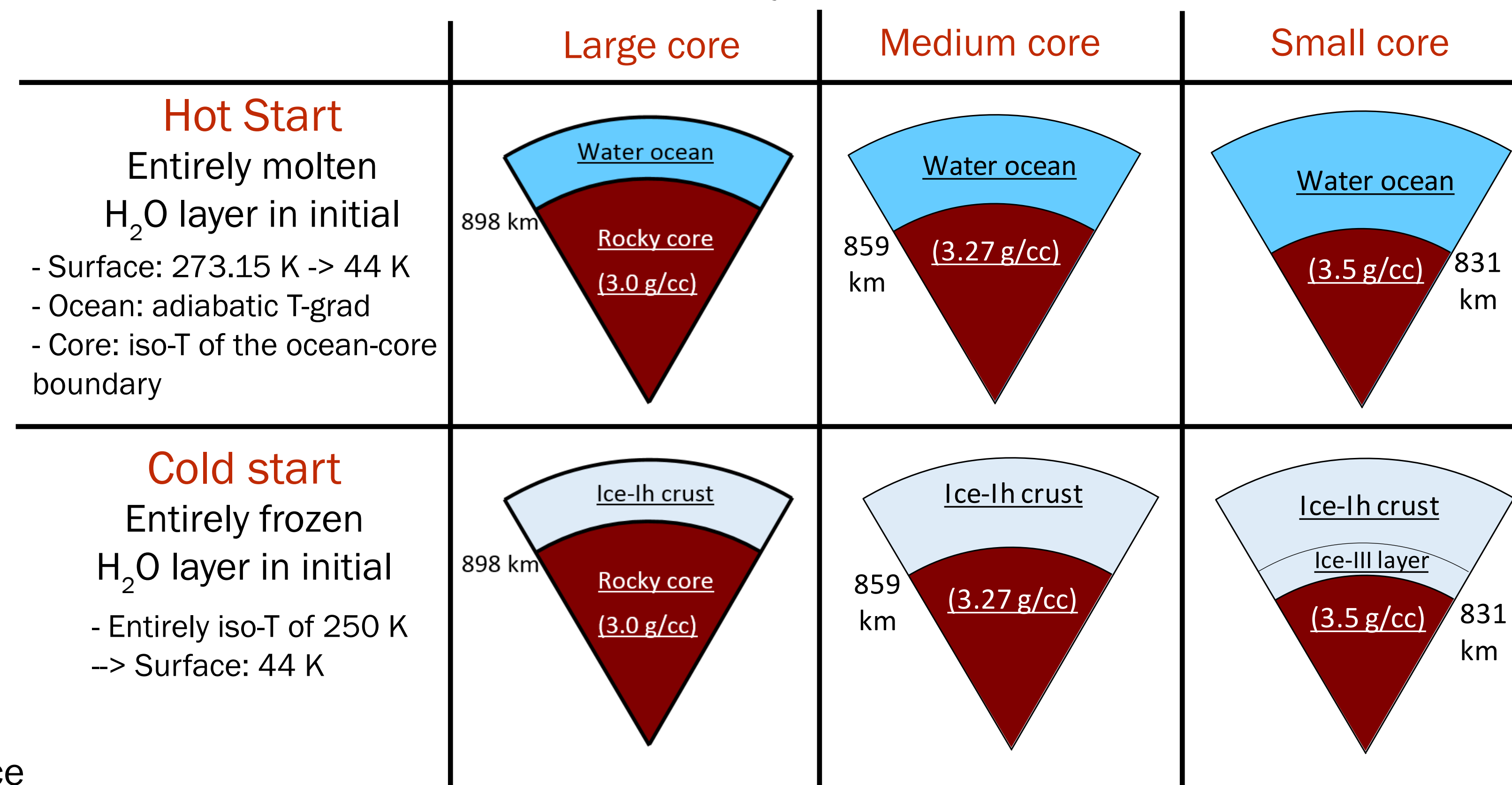
- various water/rock volume ratio
- radioisotope abundances (CI chondritic/ordinary chondritic)
- ice reference viscosity
- initial thermal state (entirely frozen/molten water layer)

Our interior model and case studies

Assumption: 2-components, water (1.0 g/cc) and rock (3.0-3.5 g/cc), are completely differentiated in initial.

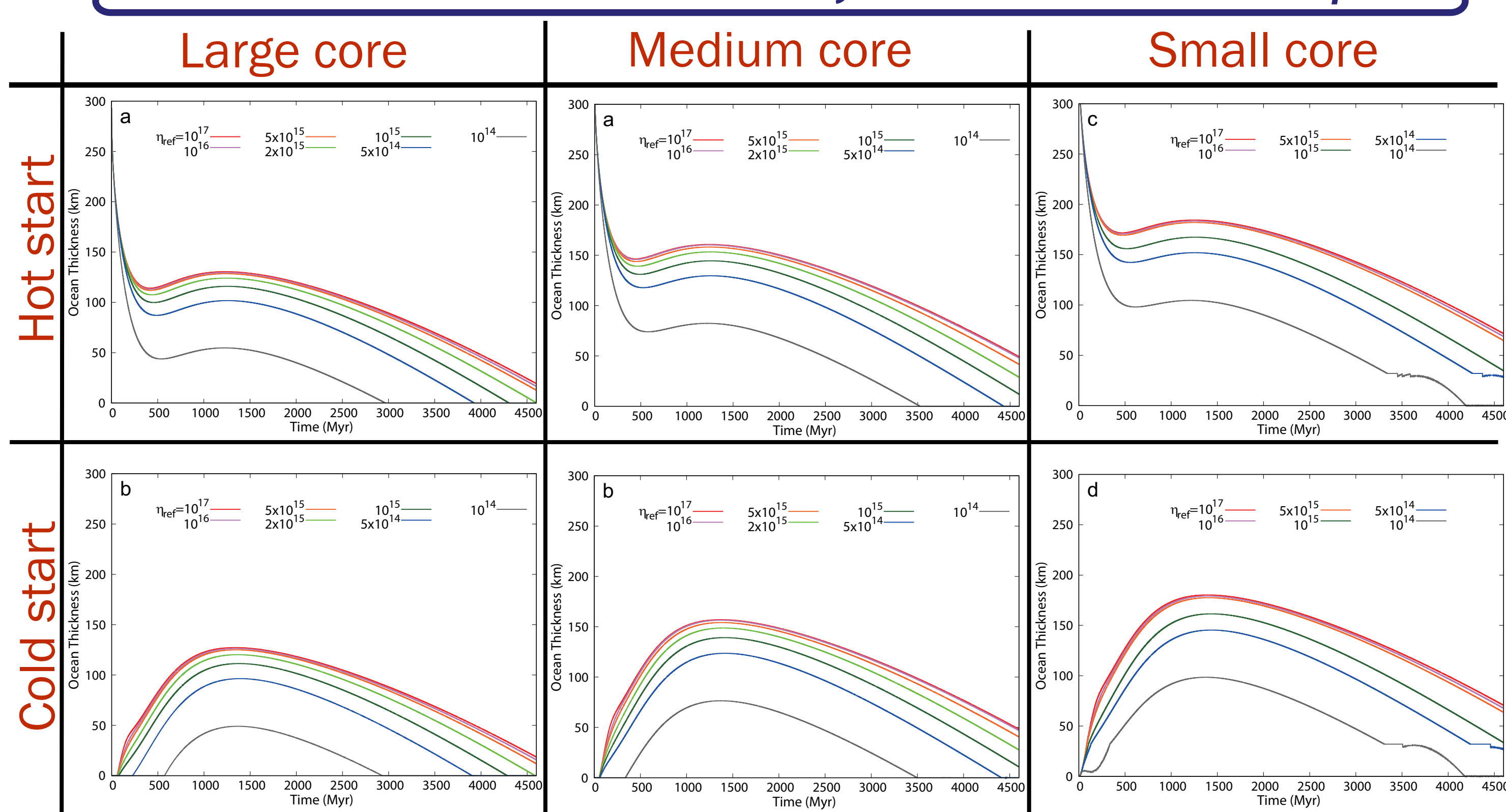


Possible range of Pluto's core radius (solid line) composed of the outer H₂O layer and inner rocky core. HP-ice can be appeared only in the small core case above the core surface when the temperature is sufficiently low.

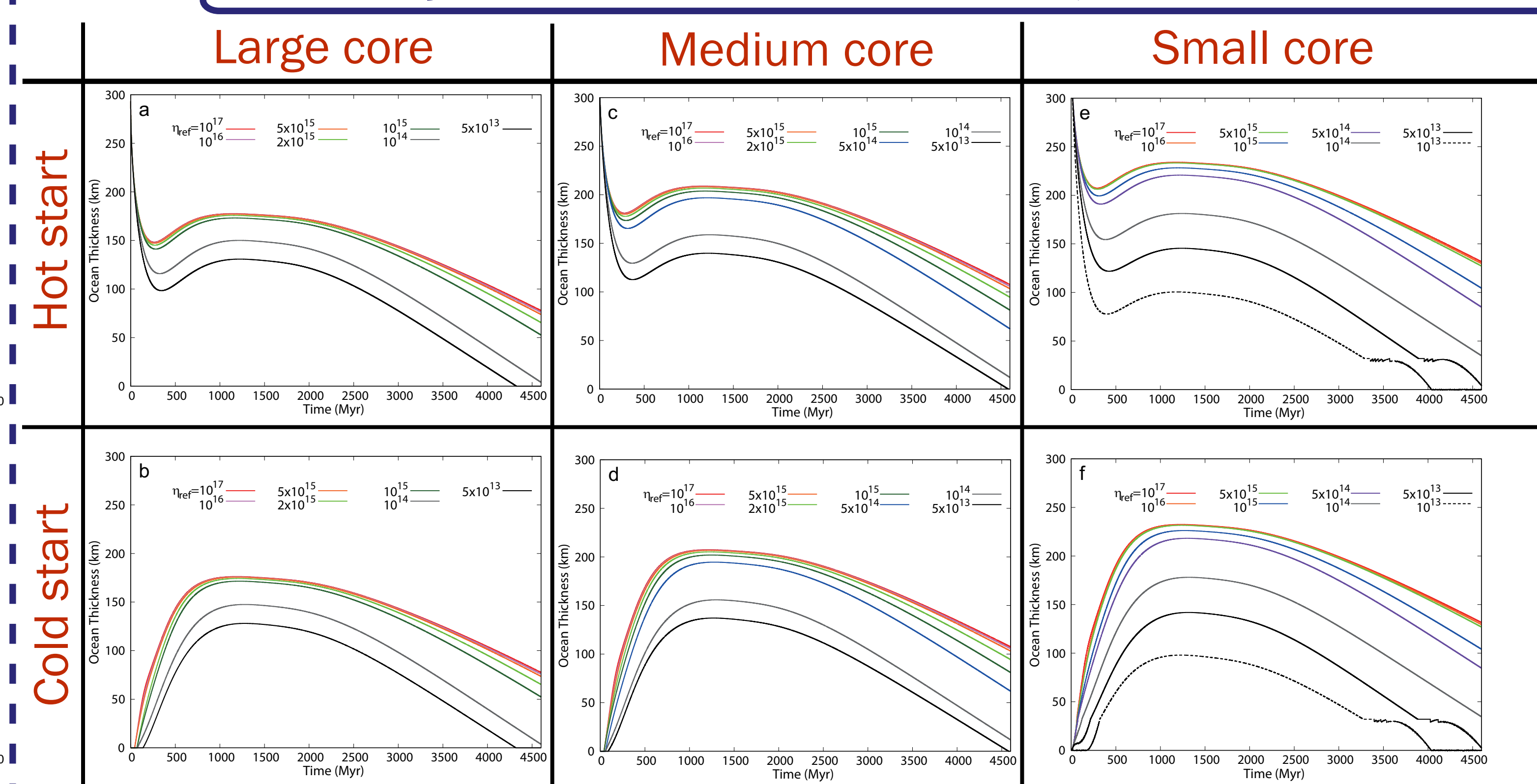


Results: Evolution of the Pluto's ocean thickness

CI chondritic abundance of radioactive isotopes



Ordinary chondritic abundance of radioactive isotopes

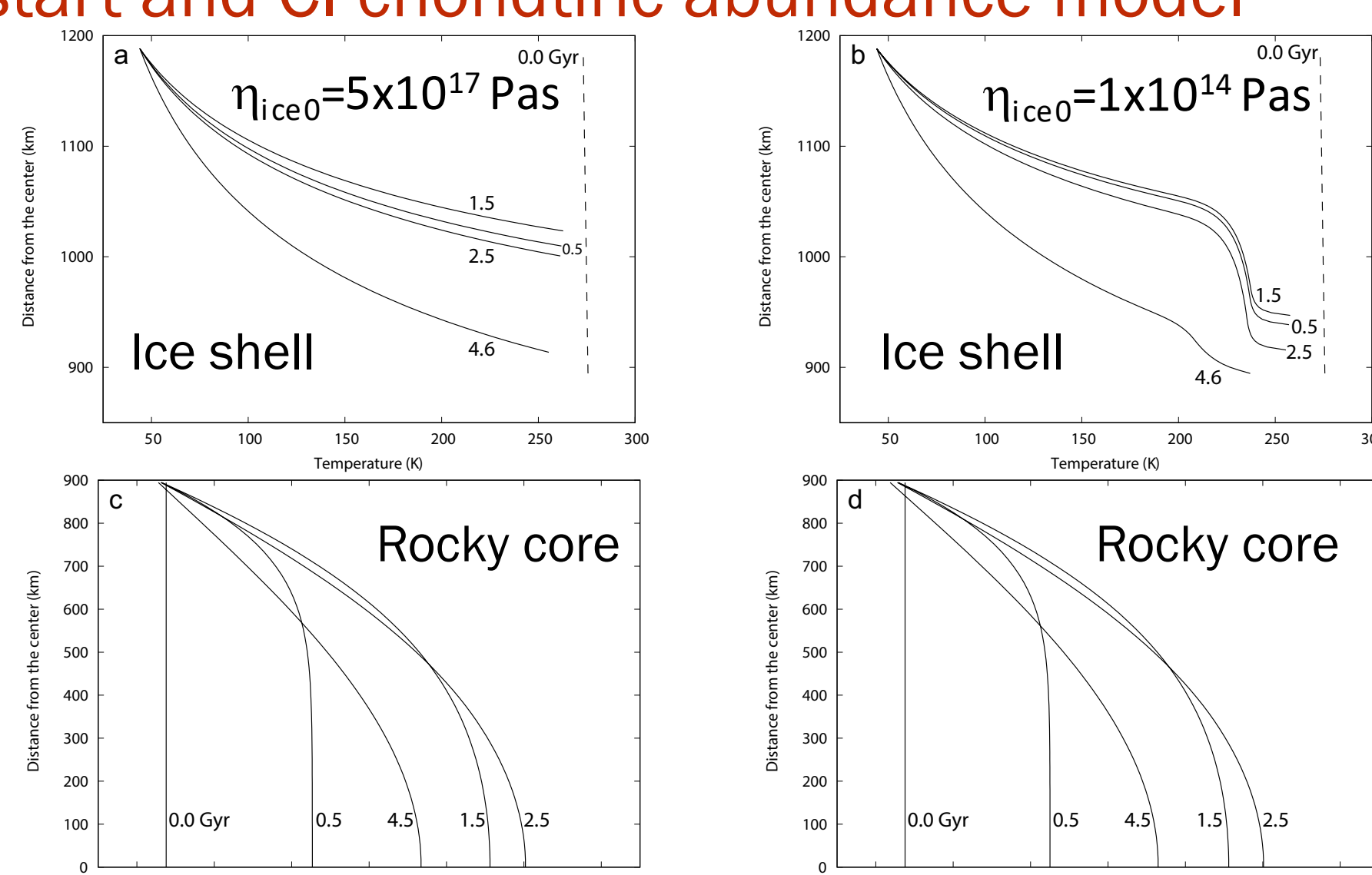


Time evolution of temperature profiles

ex.) Large core, hot start and CI chondritic abundance model

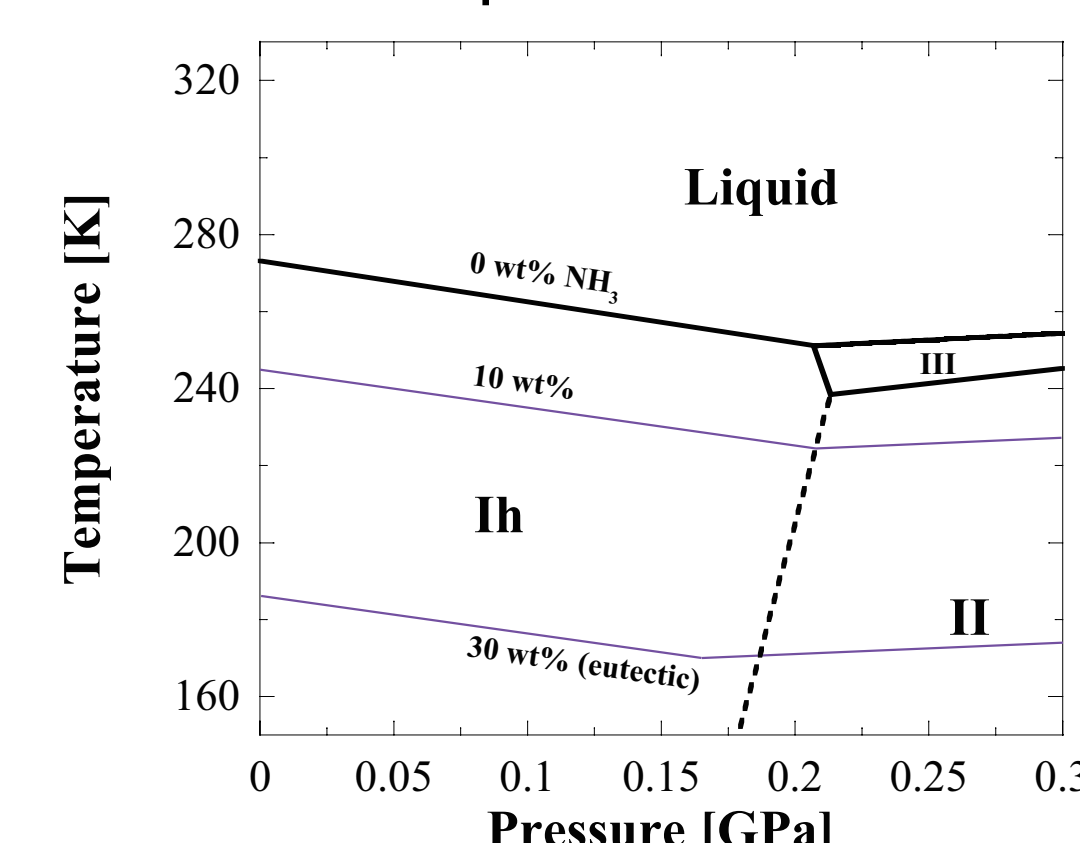
- In case of smaller ice viscosity, the ice shell is convective and effectively transports heat, resulting in a short duration of the ocean.

- Temperature in the rocky core cannot be affected by a thermal state of the ocean and the ice shell because of its large viscosity.

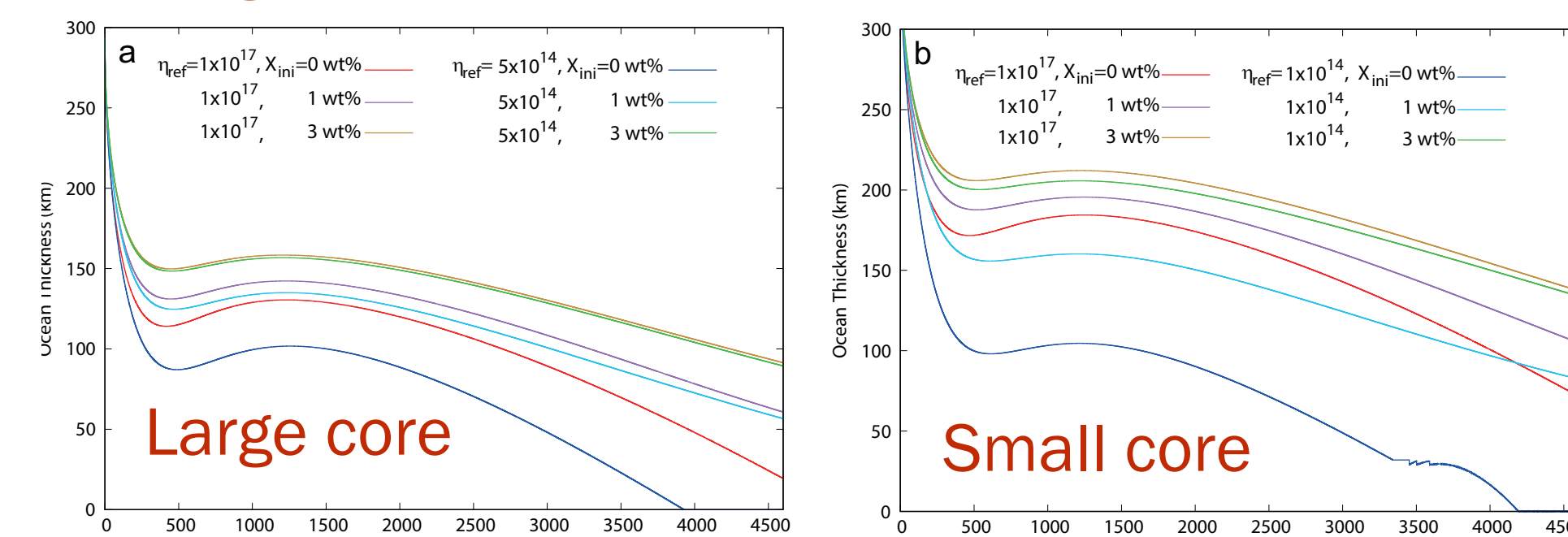


Effect of antifreeze

Presence of NH₃ decreases the melting point of the ocean up to 175 K.



ex.) Large core, hot start & CI chondritic abundance model



An inclusion of ammonia in the ocean increases the possibility of the current presence of a subsurface ocean even in the case of 1 wt% NH₃ at the initial.

1-minute summary

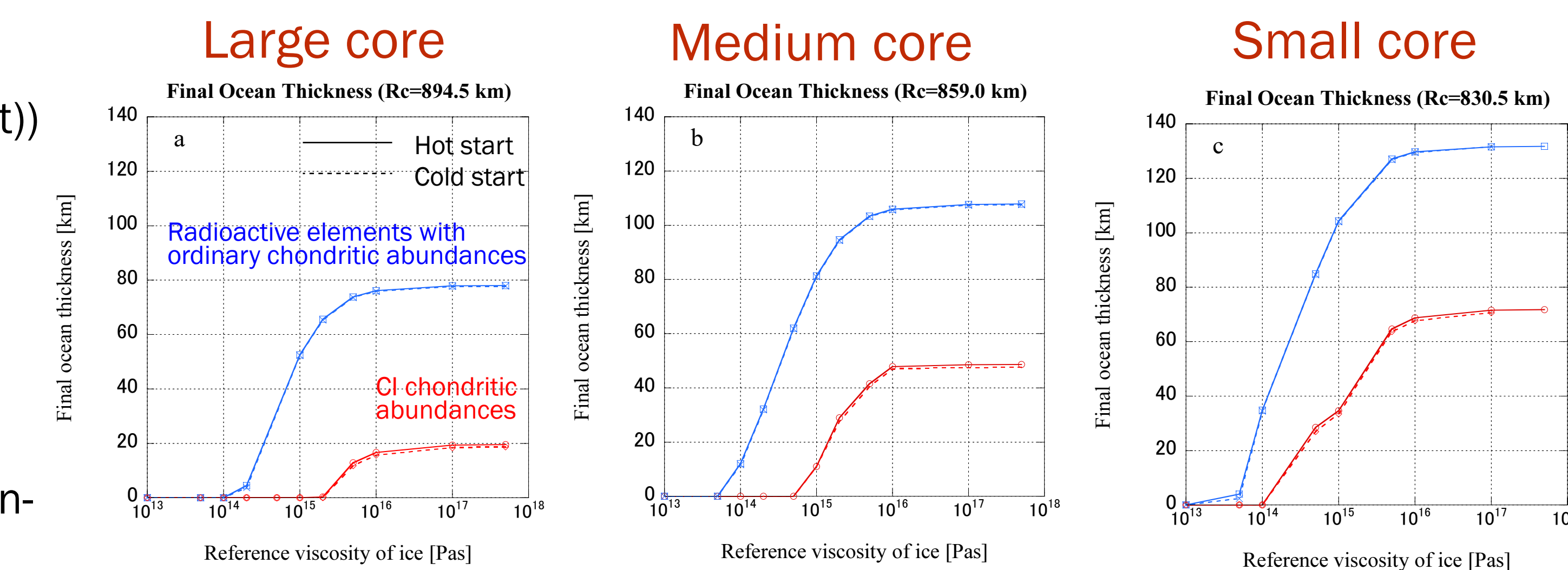
We investigated stability of the Pluto's ocean with variations of:

- Water-rock volume ratio (core size; small/medium/large)
- Radioactive isotope abundances (CI/Ordinary chondritic)
- Initial state of water layer (entirely molten (hot start) / frozen (cold start))
- Ice reference viscosity (1x10¹³ - 5x10¹⁷ Pas)

We've got results as follows:

- Large ice viscosity leads thicker ocean.
- Ordinary chondritic abundances lead thicker ocean.
- Smaller core leads thicker ocean.
- Initial state does not affect the current ocean thickness.
- Present Pluto may have a subsurface ocean if the ice shell is purely conductive or only weakly convective, and its thickness is largely uncertain and could be 20–130 km, depending on the ice reference viscosity.

Final thickness of the subsurface ocean as a function of the ice reference viscosity for three core sizes.



Numerical frameworks:

Equations: 1-D spherically symmetric model for convective/conductive heat transfer incorporating the radial dependence of viscosity and heat source distributions.

- for solid layers (ice shell, HP-ice layer & rocky core), "Mixing Length Formulation" [6]

$$\rho C_p \frac{\partial T}{\partial t} = \nabla \cdot \mathbf{F} + \rho Q$$

$$\mathbf{F} = F_{\text{cond}} + F_{\text{conv}}$$

$$F_{\text{cond}} = k_c \nabla T$$

$$F_{\text{conv}} = k_v (\nabla T - \nabla_{\text{ad}} T)$$

$$k_v = \begin{cases} 0 & \frac{\partial T}{\partial r} < \left(\frac{\partial T}{\partial r} \right)_{\text{ad}} \\ \frac{\rho C_p \alpha g l^4}{18 \nu} \left[\frac{\partial T}{\partial r} - \left(\frac{\partial T}{\partial r} \right)_{\text{ad}} \right] & \frac{\partial T}{\partial r} > \left(\frac{\partial T}{\partial r} \right)_{\text{ad}} \end{cases}$$

t: time
T: temperature
 ρ : density
 C_p : specific heat
 Q : heat generation
 k_c : thermal conductivity
 ν : kinetic viscosity
 r : radius
 α : thermal expansion coef.
 l : mixing length
 g : gravity

- for the liquid layer (subsurface ocean), "Parameterized Convection Theory".

Rheologies

- liquid water: $\eta_w = 10^{-3}$ Pas

$$\text{- solid ices: } \eta_{\text{ice}} = \eta_{\text{ref}} \exp \left[\frac{E_a}{R_g T_m} \left(\frac{T_m}{T} - 1 \right) \right] \quad \eta_{\text{ref}} = 1 \times 10^{13} \sim 5 \times 10^{17} \text{ Pas} \quad [7]$$

$$\text{- rock: } \eta_{\text{rock}} = 4.9 \times 10^8 \exp(23.25 T_m / T) \quad [8]$$

Assume that the molten water ascends instantaneously from the core surface to the ocean (or the triple point depth), which corresponds to the extreme case in which the core surface and the ocean are directly connected through a stable upwelling.

Change of the boundary position (M) between solid and liquid (Stefan problem)

- evaluating heat fluxes balance between incoming (F_{in}) and outgoing (F_{out}) at the boundaries. L is the latent heat of ice and M is the position of boundary.

$$\rho L \frac{dM}{dt} = F_{\text{in}} - F_{\text{out}}$$

Heat sources

- Radiogenic heat only in the rocky core.

- No tidal heating

	CI chondritic (ppb)	Ordinary chondritic
²³⁸ U	19.9	26.2
²³⁵ U	5.4	8.2
²³² Th	38.7	53.8
⁴⁰ K	738	1104

Thermal properties

- solid ices:

$$k_c = \frac{632}{T} + 0.38 - 0.00197T$$

$$\alpha = 3.0 * [(2.5 \times 10^{-7} T) - 1.25 \times 10^{-5}]$$

$$C_p = 7.037T + 185.0$$

Ocean composition (salinity)

- pure water
- initially 0, 1, 3 wt% NH₃ (salinity changes with freezing the ocean).

Symbol	Meanings	Dimensional value
ρ_{liq}	Density of the liquid H ₂ O layer	1000 kg/m ³
C_{pliq}	Specific heat of the liquid H ₂ O layer	4.2×10^3 J/K kg
k_{cliq}	Thermal conductivity of the liquid H ₂ O layer	0.566 W/m K
α_{liq}	Thermal expansion coefficient of the liquid H ₂ O layer	2.1×10^{-4} /K
ρ_c	Density of the rocky core	3000–3500 kg/m ³
C_{pc}	Specific heat of the rocky core	920 J/K kg
k_{cc}	Thermal conductivity of the rocky core	3.0 W/m K
α_c	Thermal expansion coefficient of the rocky core	2.4×10^{-5} /K

Concluding Summary

We investigated Pluto's thermal evolution and stability of the subsurface ocean, and their dependency on the interior structure, the ice viscosity, the radioactive isotope abundances, the initial thermal state, and the dissolved component in the ocean.

The minimum reference viscosities of ice required to maintain the ocean at present:

Radioisotope abundances/ Core size	CI chondritic	Ordinary chondritic
Large core (Rc = 894.5 km)	> 5x10 ¹⁵ Pas	> 2x10 ¹⁴ Pas
Medium core (Rc = 859.0 km)	> 1x10 ¹⁵ Pas	> 1x10 ¹⁴ Pas
Small core (Rc= 830.5 km)	> 5x10 ¹⁴ Pas	> 5x10 ¹³ Pas

Initial thermal state does not affect to the current ocean thickness.

Our results find broader conditions to be capable of maintaining the Pluto's ocean at present than previous works depending on the core size. Especially, in case of medium or small core, the ocean can sustain until the present for a typical ice viscosity even without clathrate layer [9].

References

- [1] Nimmo+ Icarus 2017, [2] Nimmo+ Nature 2016, [3] Moore+ Nature 2016, [4] Grundy+ Nature 2016, [5] Robuchon and Nimmo Icarus 2011, [6] Kamata JGR 2018 [7] Goldsby and Kohlstedt JGR 2001, [8] Karato+ JGR 1986. [9] Kamata+ Nature Geoscience 2019.