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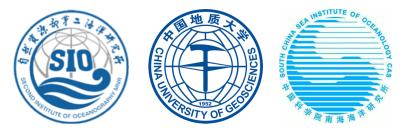
## Extension discrepancy distribution of the hyper-thinned continental crust in the Baiyun Rift, northern margin of the South China Sea

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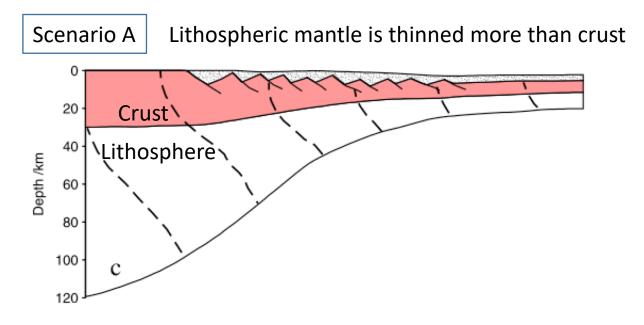
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#### Background



# Scenario B Crust is thinned more than lithospheric mantle

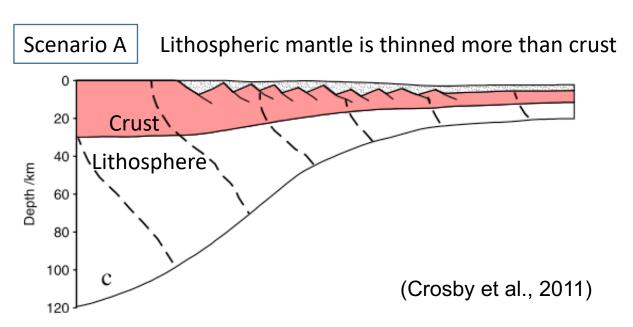
## **Extension discrepancy**, upper crust thinned far less than whole crust or lithosphere.

- North Sea (Ziegler, 1983; White 1990)
- Norwegian rifted margin (Kusznir et al., 2005)
- Brazil-Angolan margins (Moulin et al., 2005)

**Inverse extension discrepancy**, in the hyperextended domain of rifted margins, the upper crustal extension could exceed the whole crust.

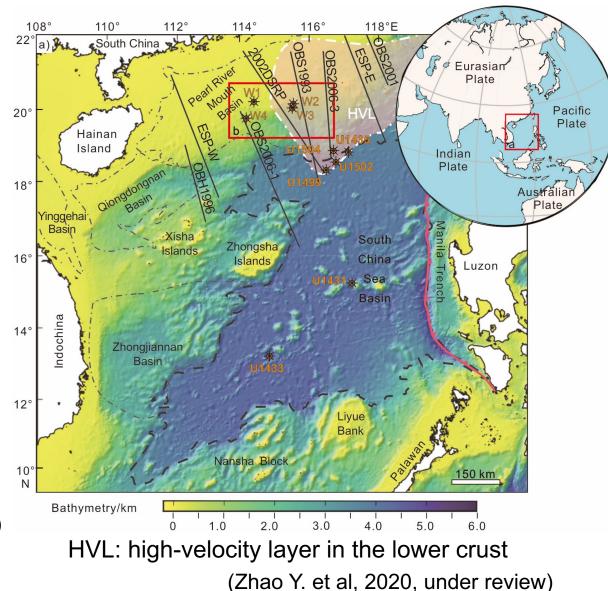
- Brazil-Angolan margin (Crosby et al., 2011)
- Galicia margin (McDermott & Reston, 2015).

#### Background



## In the South China Sea margins, most studies reported extension discrepancy

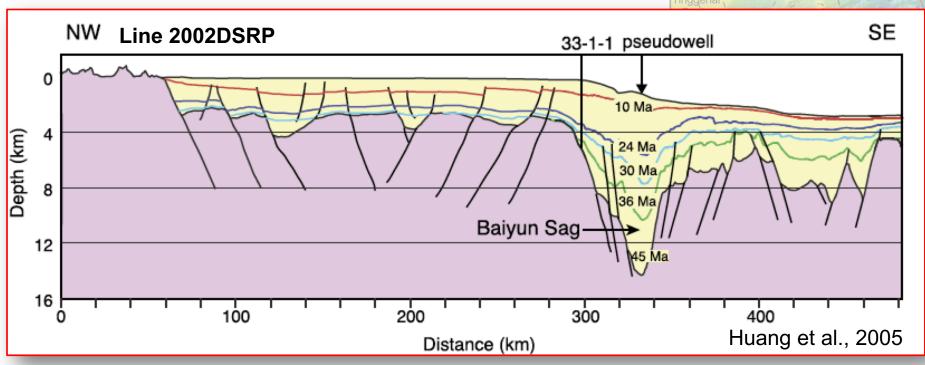
- Pearl River Mouth basin (Bai et al., 2019; Clift et al., 2002; Hsu et al., 2004)
- Qiongdongnan basin (Tong et al., 2009; Zhao Z. et al., 2018)
- Northwest Palawan margin (Franke et al., 2011)
- Offshore Vietnam (Savva et al., 2014)

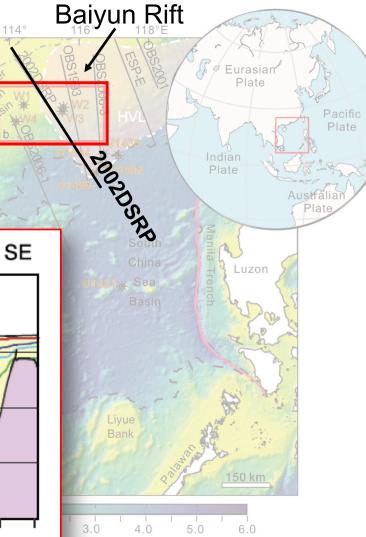


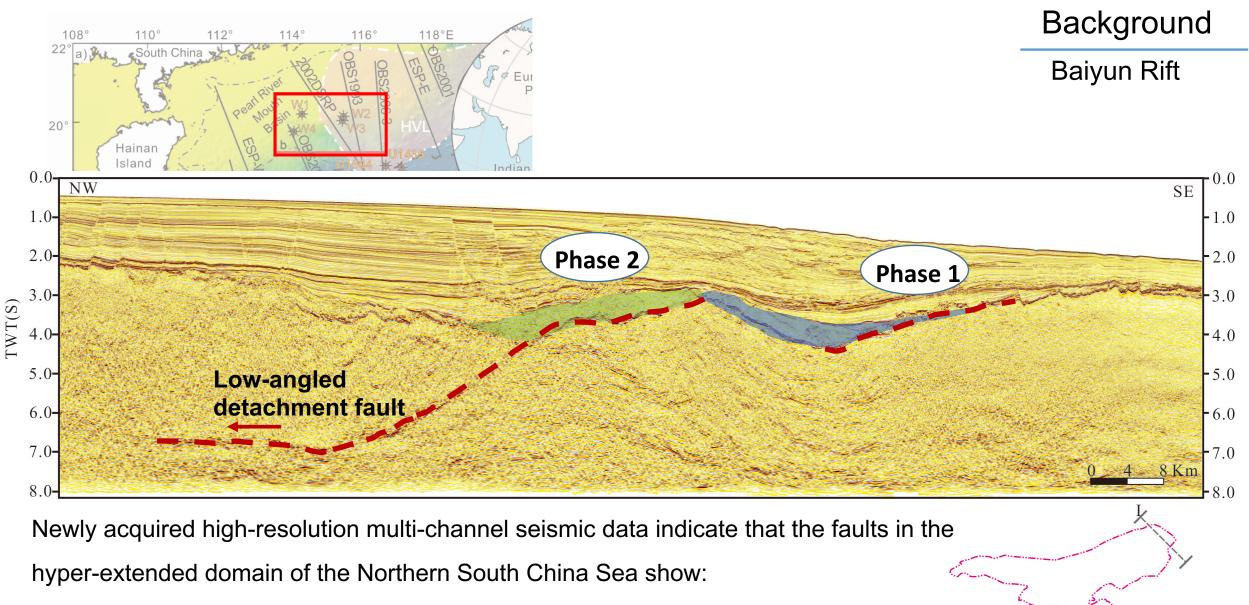
#### Background

However, previous studies may have underestimated upper crustal thinning:

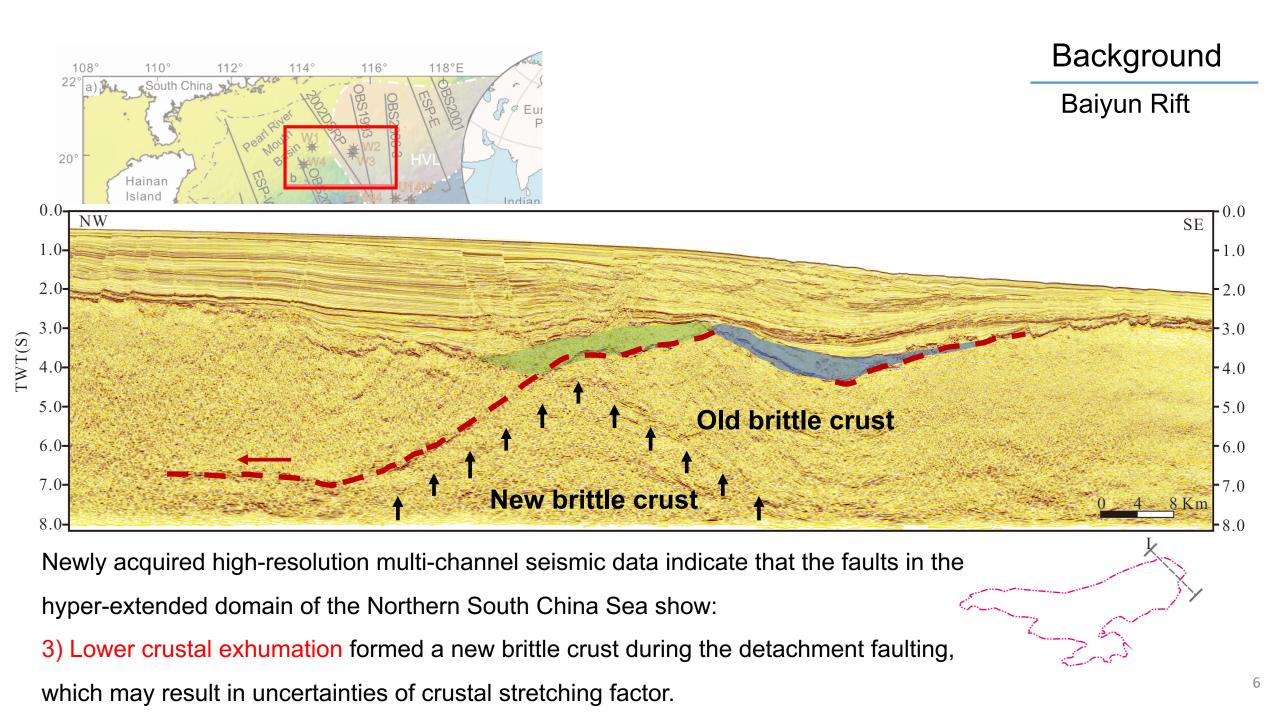
- 1) based on a classic rifted basin model that the rifting was controlled continuously by high-angle normal faulting,
- 2) under an assumption that the observed faulting is close to the total amount of brittle extension.

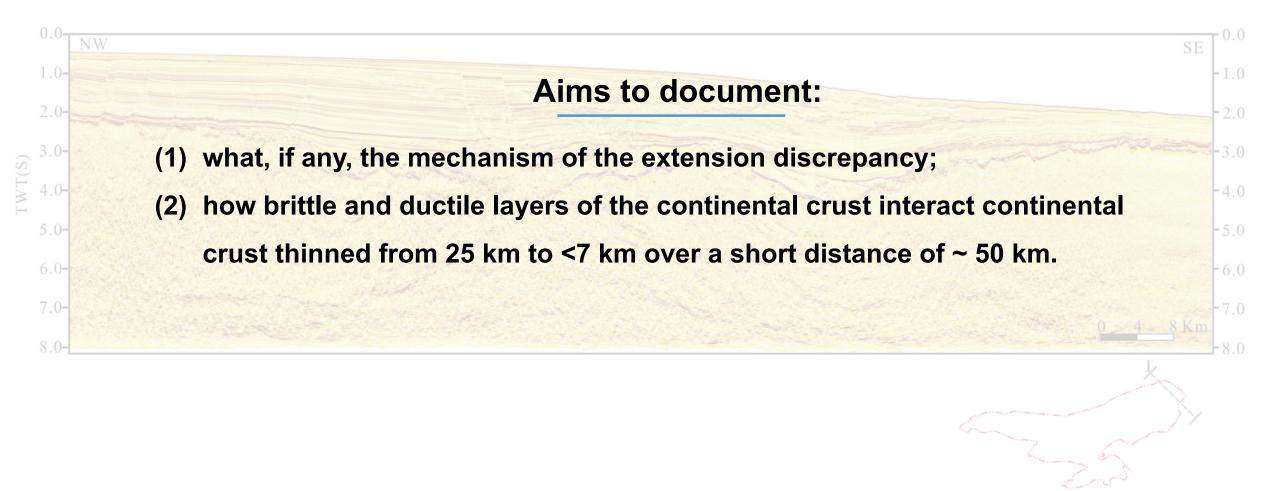






- 1) Low-angled detachment fault dominated the rifted basin evolution (Zhao Y. et al., 2018);
- 2) Polyphase faulting was failed to restored in previous studies;





#### Methods

## 1. Seismic interpretation

- Using basincovered 3D seismic data
- Sedimentary basin structure
- Crustal structure

## 2. Time-depth conversion

Stratigraphic velocity
after 41 drilling wells
Crustal velocity after
seismic modeling of
OBS profiles

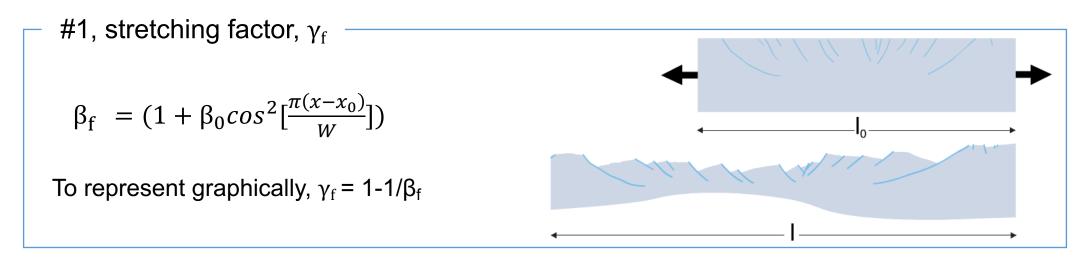
#### 3. Rift structure restoration

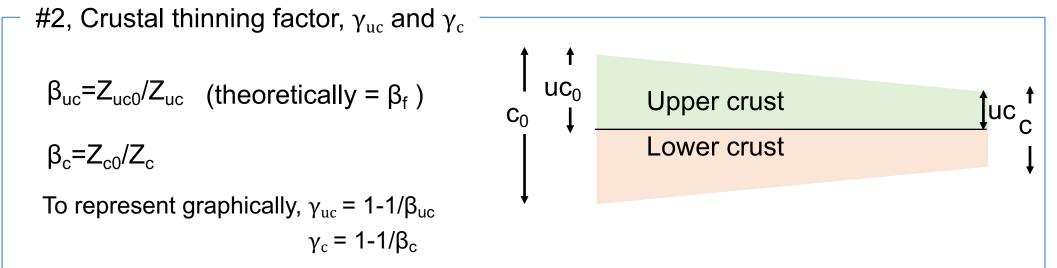
Software Move(Midland Valley)Following arealbalancing principle

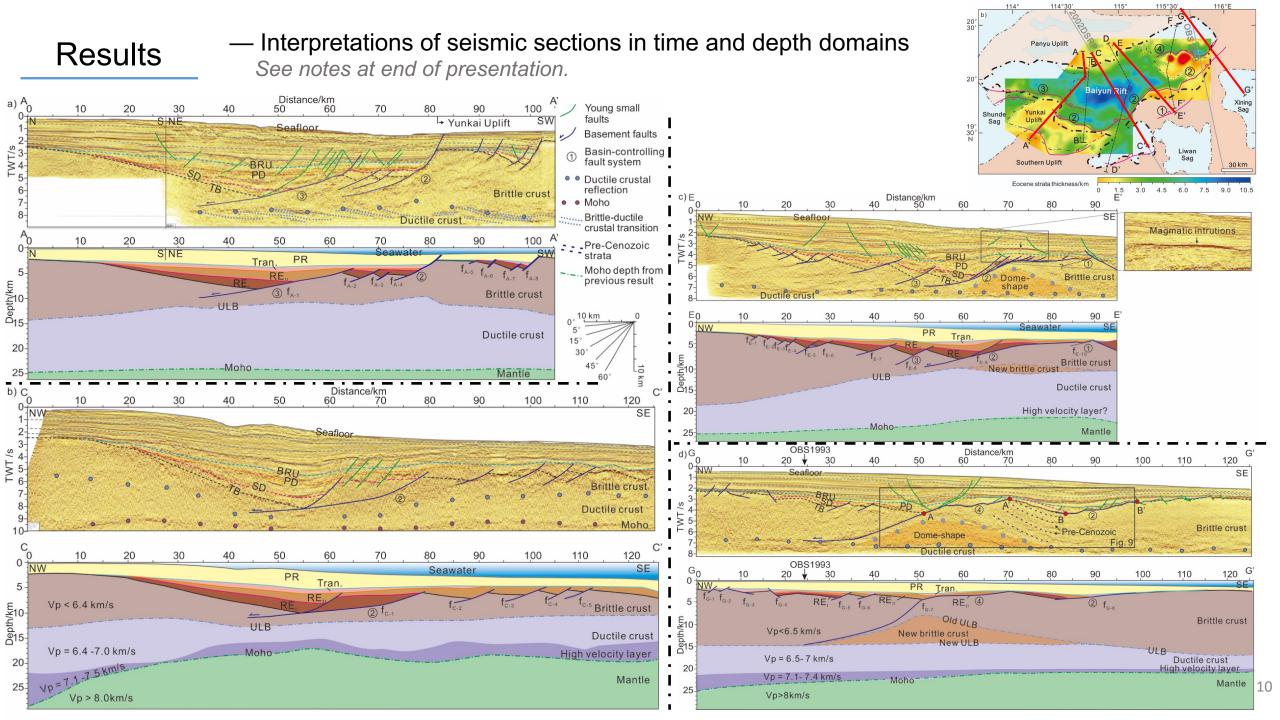
## 4. Crustal extension quantification

- Crustal thinning factor
- Fault-derived stretching factor
- Sub-resolution faulting
- examination

#### Two methods to quantify crustal stretching/thinning amount:

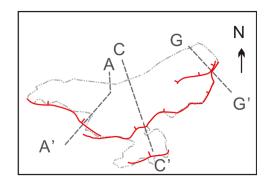


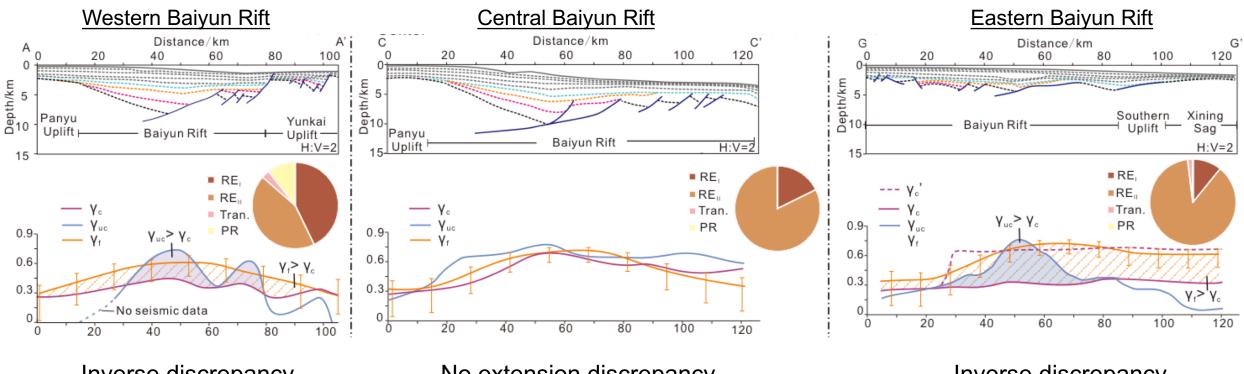




#### Results

— Selected depth sections and the comparisons between whole crustal thinning factor ( $\gamma_c$ ), upper crustal thinning factor ( $\gamma_{uc}$ ), and stretching factor ( $\gamma_f$ ). See notes at end of presentation.





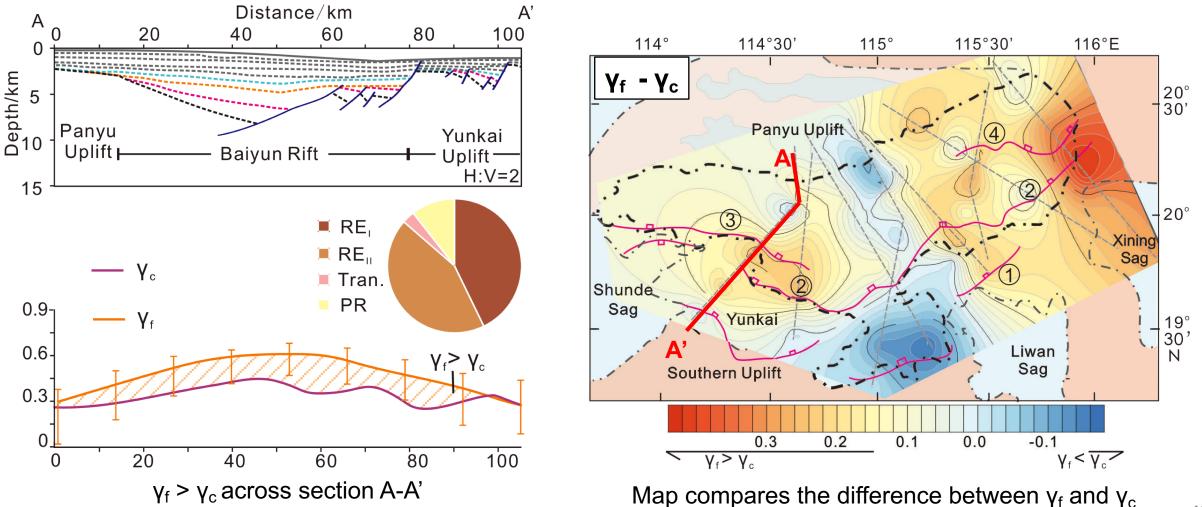
Inverse discrepancy

No extension discrepancy



#### Western Baiyun Rift

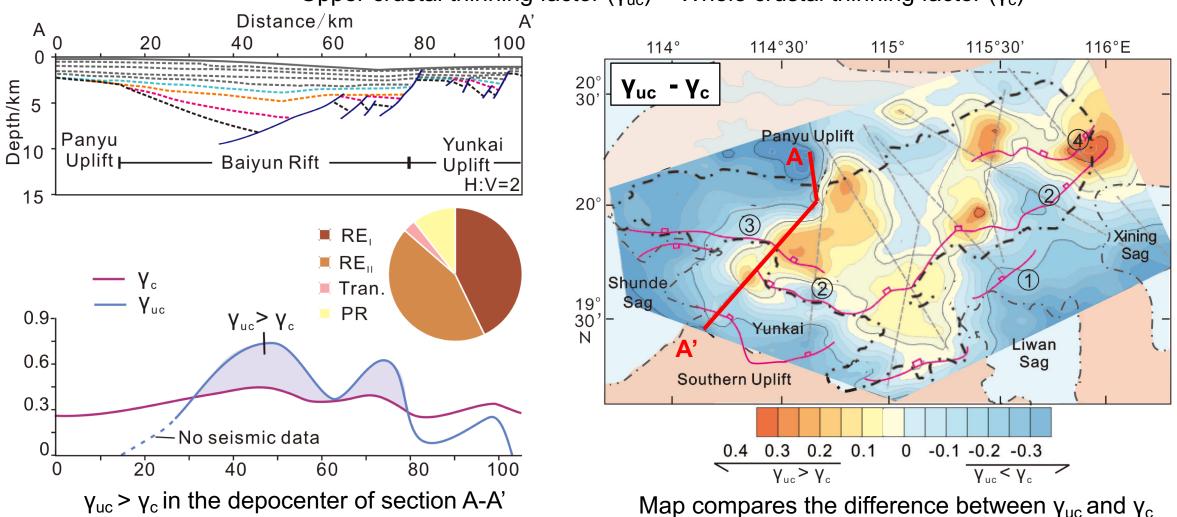
#### Inverse discrepancy



Stretching factor ( $\gamma_f$ ) > Whole crustal thinning factor ( $\gamma_c$ )

#### Western Baiyun Rift

#### Inverse discrepancy



Upper crustal thinning factor ( $\gamma_{uc}$ ) > Whole crustal thinning factor ( $\gamma_c$ )

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#### Western Baiyun Rift

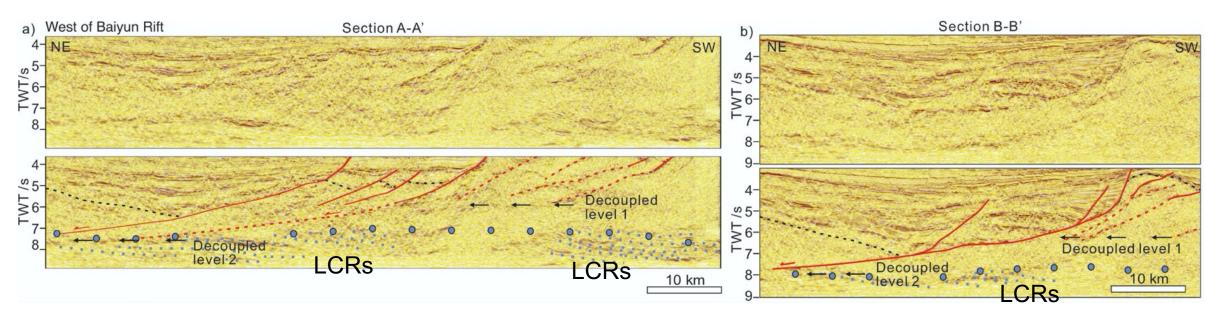
Possible cause: ductile shearing deformation

Evidence In the western Baiyun Rift, lower crustal reflections (LCRs) are identified near the brittle-ductile transition. These LCRs are symmetric to the continentward-dipping faults and notably abundant beneath the low-angled normal faults (LANFs)

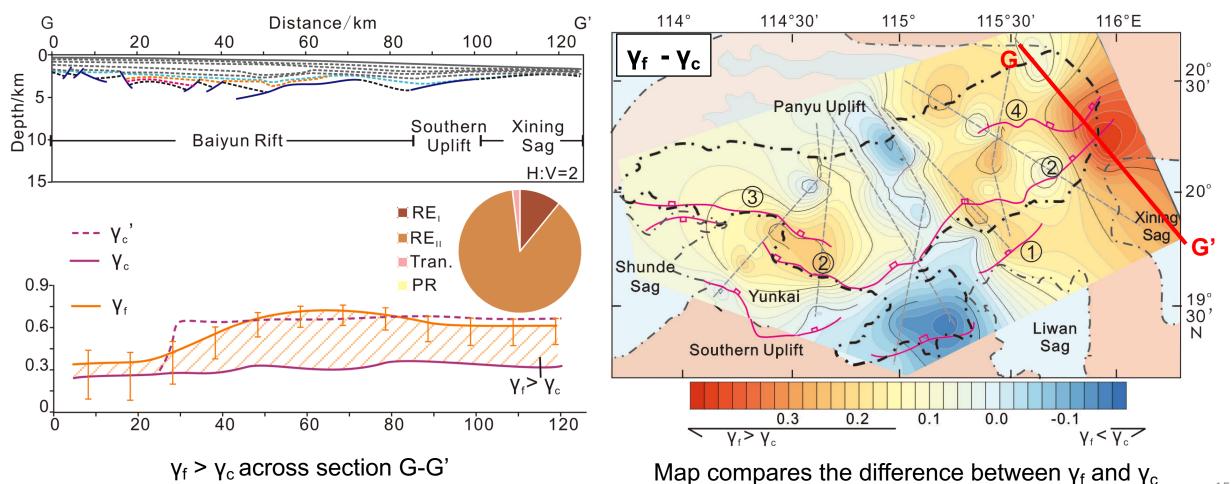
114° 114°30' 115° 115°30' 116°E 20° Panyu Uplift Panyu Uplift Shunde Shunde Southern Uplift Eccene strata thickness/km

reflections, which indicates a deformation dominated by a simple shear towards the continent.

Hence, inverse discrepancy in the western Baiyun Rift is achieved by an intense tectonic faulting in the upper crust and a relatively weak ductile shearing in the lower crust.



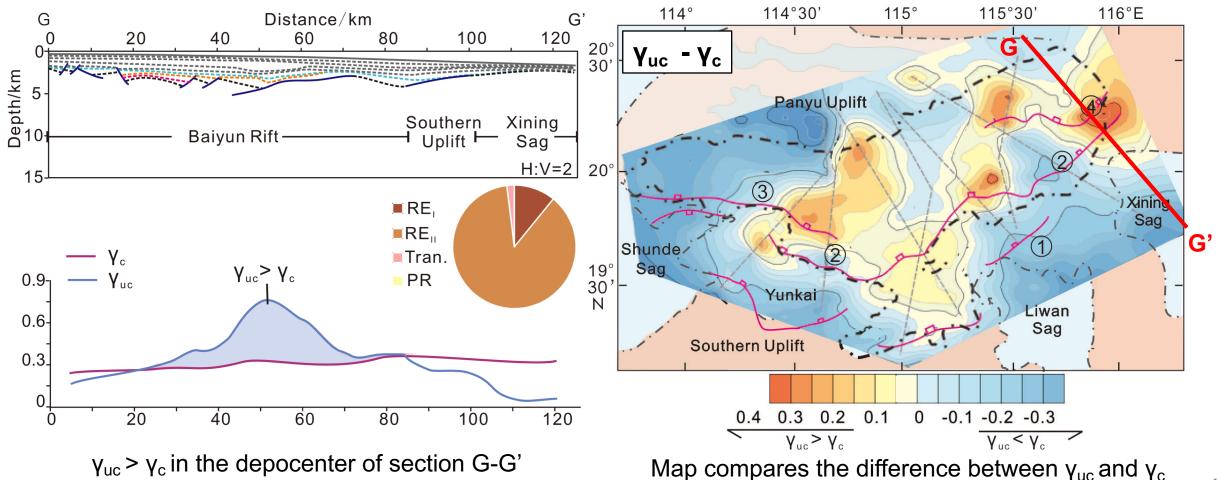
#### Inverse discrepancy



Stretching factor ( $\gamma_f$ ) > Whole crustal thinning factor ( $\gamma_c$ )

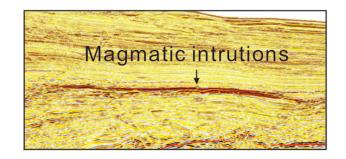
#### Inverse discrepancy

Upper crustal thinning factor ( $\gamma_{uc}$ ) > Whole crustal thinning factor ( $\gamma_{c}$ )



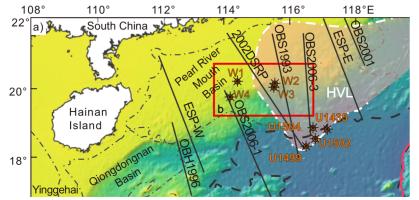
Possible cause: Lower crustal exhumation

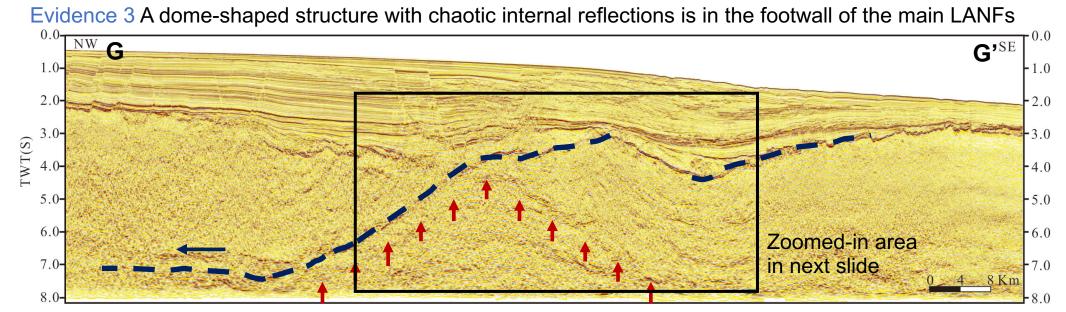
Evidence 1 More intense magmatism in the eastern Baiyun Rift (Zhao F. et al., 2016)



Evidence 2 High-velocity layer in the lower crust extending to

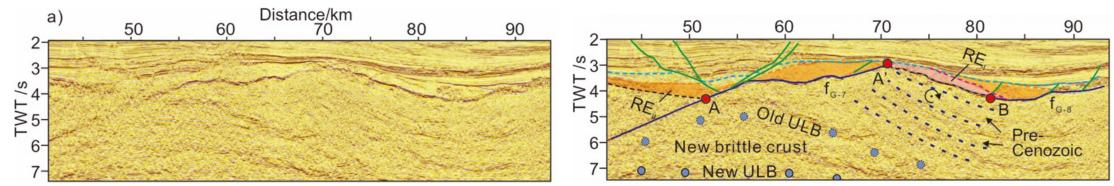
the northeast along the NSCS (Yan et al., 2001; Wei et al., 2011)



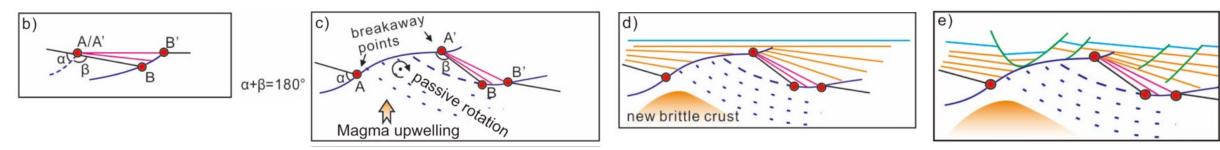


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Polyphase detachment faulting caused the exhumation of the lower crust. Due to isostasy, the magma passively upwelled and thickened the crust and thus results in an underestimation of the whole crustal extension.



a) At least two phases of faulting with two couples of breakaway points, A/A' and B/B', are identified.



b) Sequential development of  $f_{G-7}$  and  $f_{G-8}$  hyper-stretched brittle crust.

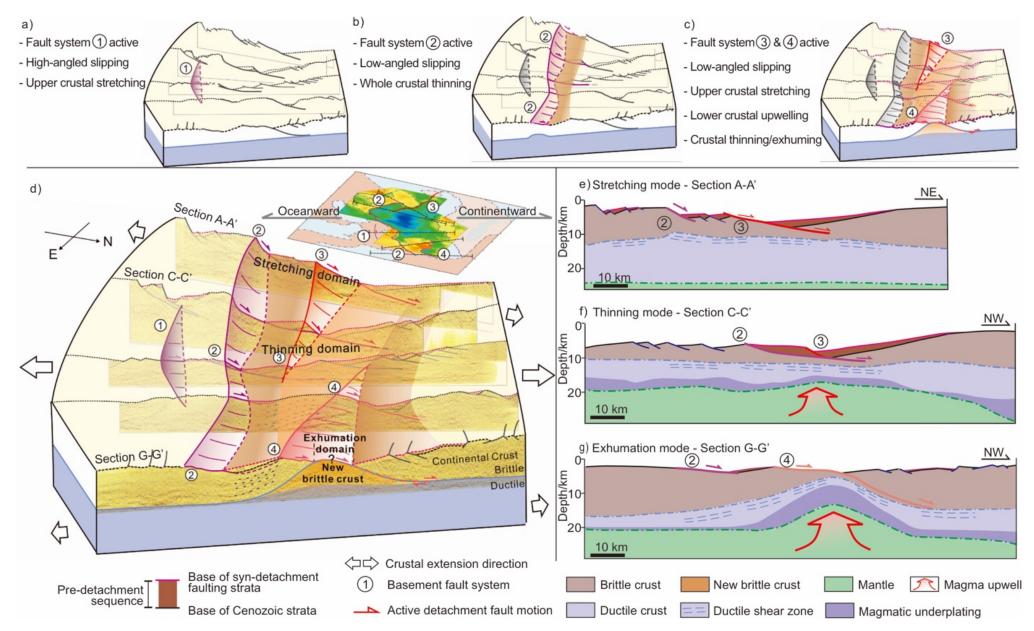
c) Due to isostasy, magmas migrated upward from the mantle, forming the dome-shaped structure at the footwall of the f<sub>G-7</sub>.

d) With the cooling of magmas, the ductile materials become brittle. New brittle crust and new ULB formed.

e) During post-rifting stage, faults re-activated while have not changed rift geometry.

#### Summary

- Models of polyphase faulting leading to a later detachment faulting associated with a lower crustal exhumation in the eastern Baiyun Rift. See notes at end of presentation.



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#### Notes

#### Slide # 10

Interpretations of sections A-A', C-C', E-E', and G-G' in time and depth domains. The pattern of the basement faults changes from a stairstep combination in the west, to a deformed singular fault with large displacements in the east. Shallow dipping basement faults sole out at a depth of 7s-8s TWT, where high-amplitude and continuous reflections developed. We interpret a brittle-ductile transition is at this depth. Moho depth is extracted from the previous result (Li et al., 2019). P-wave velocity is after Yan et al., (2001). RE<sub>I</sub>, rift episode I; RE<sub>II</sub>, rift episode II; Tran., transition stage; PR, post-rift stage.

In the west, basement faults are rooting at different depths, suggesting multiple decollement zones. ULB truncates SW-dipping reflections in the lower crust. In the east, a dome-shaped structure with chaotic internal reflections is in the footwall of the main LANFs. Parallel contacting relationship of the dome flank, pre-Cenozoic reflection and the hanging wall of the LANF suggests a synchronic deformation of these structures.

#### Slide # 11

Error bars of  $\gamma_f$  indicate a range of 0-60% of extension which could be underestimated due to a sub-resolution faulting (Clift & Sun, 2006). Blue and orange shaded regions represent  $\gamma_{uc} > \gamma_c$  and  $\gamma_f > \gamma_c$ , indicating inverse discrepancy. Pie graphs demonstrate the proportion of the extension in each structural evolution episode. The light brown category in the pie graph has the largest proportion shows the most intense extension happened in the RE<sub>II</sub> when low-angled faults were active.  $\gamma_c'$  is the thinning factor of the whole crust under an assumption that 50% of the basement is pre-Cenozoic remnants. The actual thinning factor of the crust ranges from  $\gamma_c$  to  $\gamma_c'$ .

#### Slide #19

(a-d) Block diagrams display a perspective view from the NE of the 3D volume, without sediments to expose the top of the acoustic basement. Four sets of the basement fault systems in the Baiyun Rift have been identified. These fault systems were initiated earlier in the central south and migrated to the NE and NW progressively.

(e-g) Models for temporal and spatial evolution of the hyper-thinning process based on observations from seismic sections in the Baiyun Rift. (e) Stretching mode is characterized by listric faulting, a differential subsidence of half-grabens, and a major ductile shear zone exemplified by section A-A', western Baiyun Rift. (f) Thinning mode is characterized by the maximum thinning of the crust and the presence of magmatic underplating in the lowest crust. (g) The exhumation mode is well documented by section G-G', eastern Baiyun Rift. This phase is distinguished by the exhumation and embrittlement of the lower crust from less than 5 km depth along a downward-concave fault (4).

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