

Modeling morphological changes by tsunami Induced currents

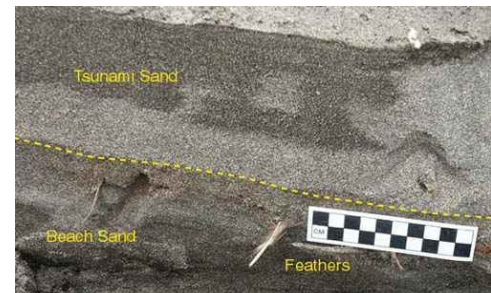
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Motivation

- Tsunamis have wide-spread and profound effects across the ocean.
- Among those, morphodynamic changes were found in harbor areas where significant strong currents occurred.(e.g., Tsunamiites)
- Tsunami physics in the nearshore area tend to be much complex due to the existence of **turbulence sources**(e.g., bottom boundary layer, wave-breaking).
- **Prediction of morphological processes** associated with **tsunami-induced currents** are highly challenging.



Outlines

- **Objectives**

- Develop a numerical solution to predict complex morphodynamic processes in the nearshore.

- **Strategy**

- *Set-up modeling tools by integrating sub-module such as*
 - *Hydrodynamic model*
 - *Sediment transport model*
 - *Morphodynamic model*
- Finite volume method coupled w/ HLL Riemann solver
- Sensitivity tests on parameters included in closure models

- **Scopes**

- 1. Model development
- 2. Validations : Dam-break flows, Breaking solitary waves(1D/2D)
- 3. Application

1. Model Development

Model Equations

- Hydrodynamic model : Boussinesq-type equations with rotational terms

$$\frac{\partial H}{\partial t} + \frac{\partial H \hat{u}_i}{\partial x_i} + (\mathcal{M} + \mathcal{M}^\nu) = \boxed{\frac{e - d}{1 - p}} \quad \leftarrow \text{Erosion/deposition}$$

$$\begin{aligned} \frac{\partial H \hat{u}_i}{\partial t} + \frac{\partial H \hat{u}_i \hat{u}_j}{\partial x_j} + gH \frac{\partial \zeta}{\partial x_i} + H (\mathcal{D}_i + \xi_i + \mathcal{D}_i^\nu + \xi_i^\nu) \\ + \hat{u}_i (\mathcal{M} + \mathcal{M}^\nu) - \boxed{H \frac{\partial}{\partial x_i} \left(\nu_t^h \frac{\partial \hat{u}_i}{\partial x_i} \right)} + \boxed{H \frac{\partial}{\partial x_i} \left(\nu_t^v \frac{\partial \hat{u}_j}{\partial x_i} \right)} + \frac{\tau_i^b}{\rho} \boxed{HF_i - HR_i} = \boxed{-\frac{(e - d) \hat{u}_i}{1 - p}} \end{aligned}$$

Turbulent dissipation
Wave-breaking/BSM

- Sediment Transport Model

$$\frac{\partial H \bar{c}}{\partial t} + \frac{\partial H \bar{c} \hat{u}_i}{\partial x_i} = \frac{\partial}{\partial x_i} \left(K_h H \frac{\partial \bar{c}}{\partial x_i} \right) + e - d$$

- Morphodynamic Model

$$\frac{\partial h}{\partial t} = \frac{e - d}{1 - p}$$

Closures

- Erosion and Deposition Fluxes

$$e = \varphi (\theta - \theta_c) (\bar{u}_i \bar{u}_i)^{0.5} H^{-1} (D_{50})^{-0.2} \quad d = \alpha \bar{c} w_0$$

- Bed-induced Turbulent Eddy Viscosity(Smagorinsky's)

$$\nu_t^h = (C_s \Delta x_i)^2 \sqrt{2 \mathcal{S}_{ij} \mathcal{S}_{ij}}$$

- Wave breaking (Kennedy et al.(2000))

$$R_i = \frac{1}{H} \left[\frac{(1 + \delta_{ij})}{2} \frac{\partial}{\partial x_j} \left\{ \nu_b \frac{\partial (H \hat{u}_i)}{\partial x_j} \right\} + \frac{|\epsilon_{ij}|}{2} \frac{\partial}{\partial x_j} \left\{ \nu_b \frac{\partial (H \hat{u}_j)}{\partial x_i} \right\} \right]$$

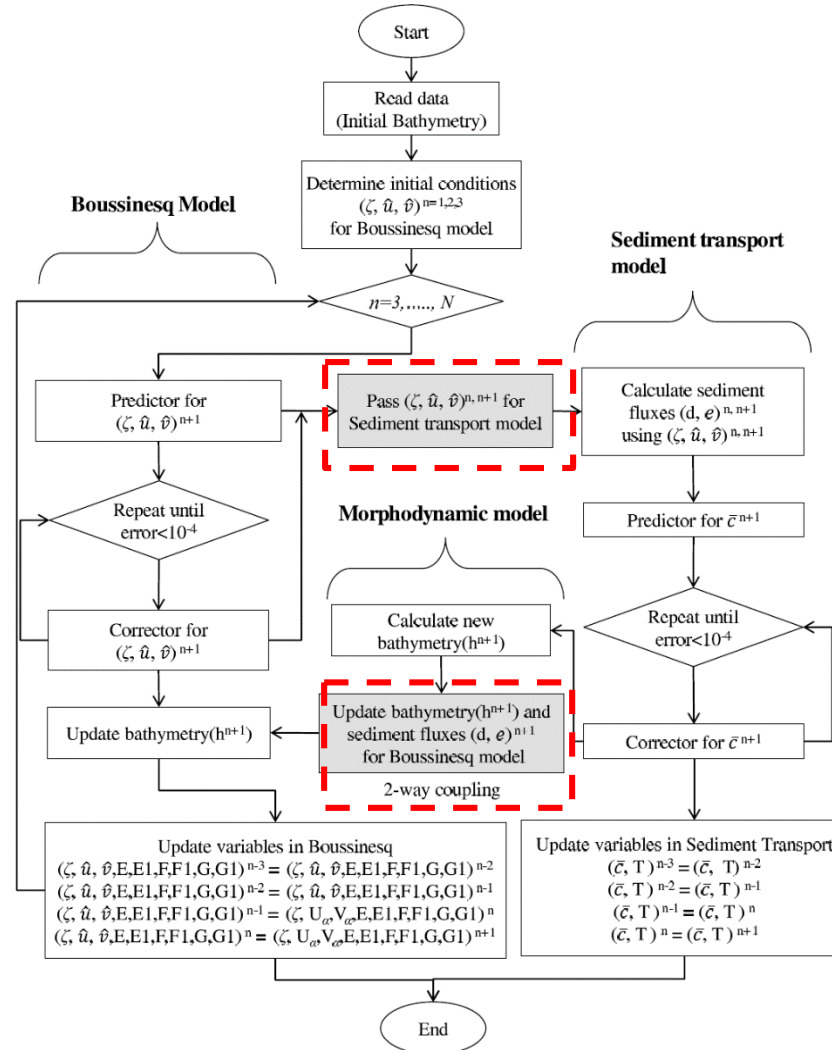
- Turbulence Backscattering Model(Hinterberger et al.(2007))

$$F_i = C_b \frac{\sqrt{\hat{u}_j \hat{u}_j}}{H} \sqrt{\frac{\nu \sqrt{C_f}}{\Delta t}} r_i$$

Numerical Methods

- Leading order terms :
 - 4th order MUSCL-TVD scheme by Yamamoto & Daiguji(1993) incorporated with approximated HLL Riemann Solver
- Higher order terms :
 - Cell-centered Finite Volume Method by Lacor(2004)
- Time marching :
 - 3rd order Adams-Bashforth predictor and
 - 4th order Adams-Moulton corrector

Flow Chart

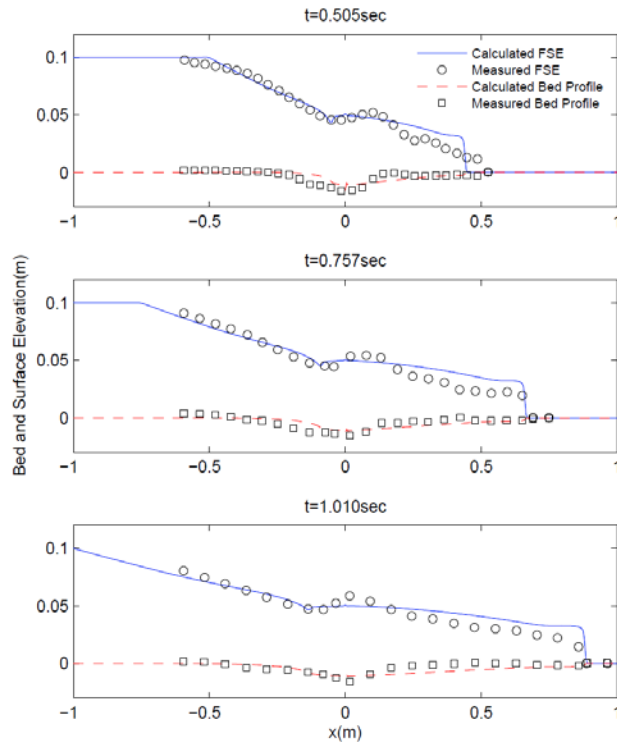


Two way coupling

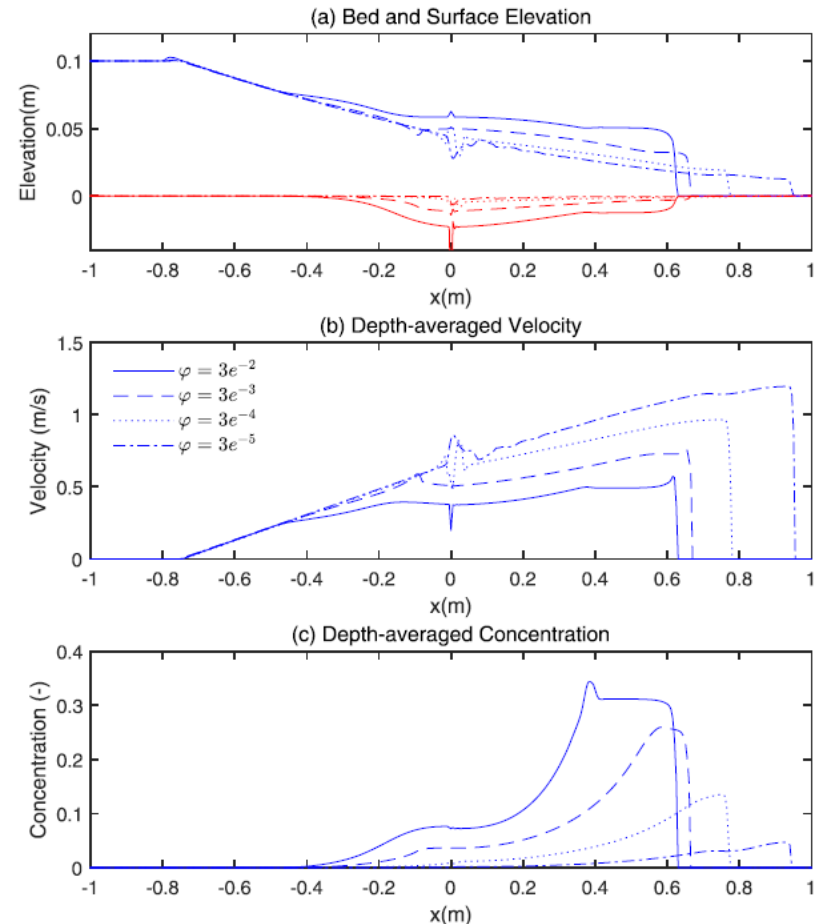
2. Model Validations

Validations : Case 1

- CASE1 : 1D Dam Breaking
over a Movable Bed
(Fraccarollo and Capart [2002])*

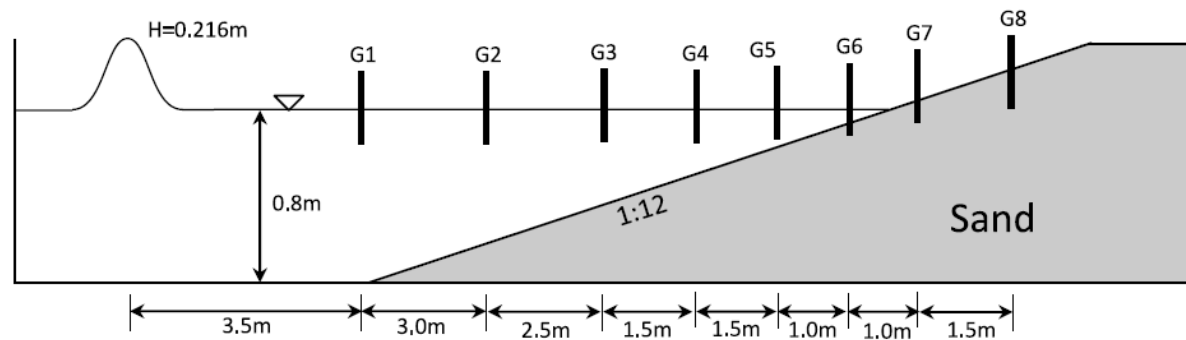


Sensitivity tests on φ

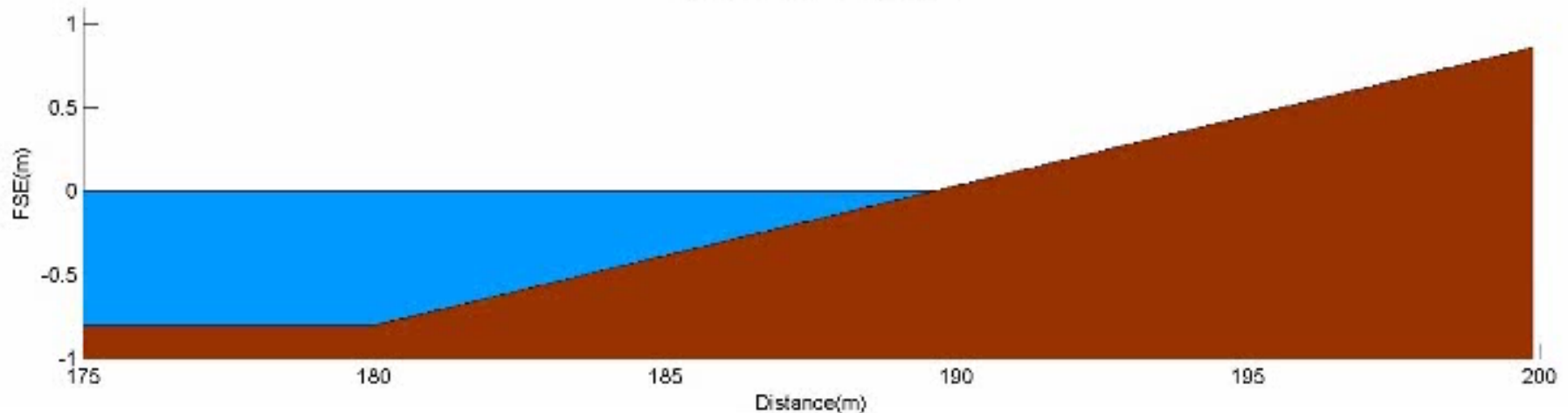


Validations : Case 2

- CASE2 : Breaking Solitary Waves on Sloping Beach
(Kobayashi and Lawrence [2004])*

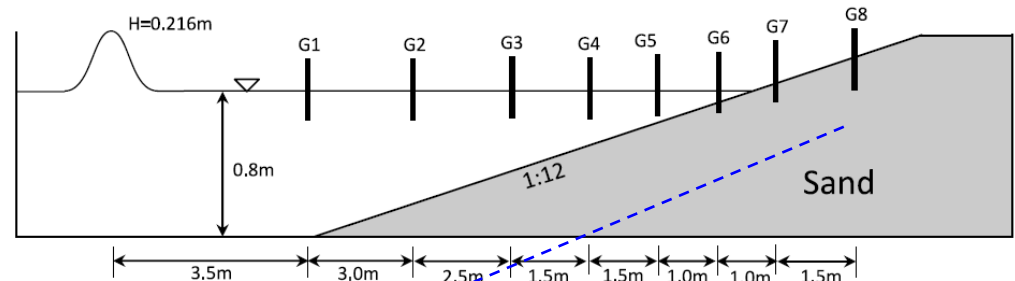
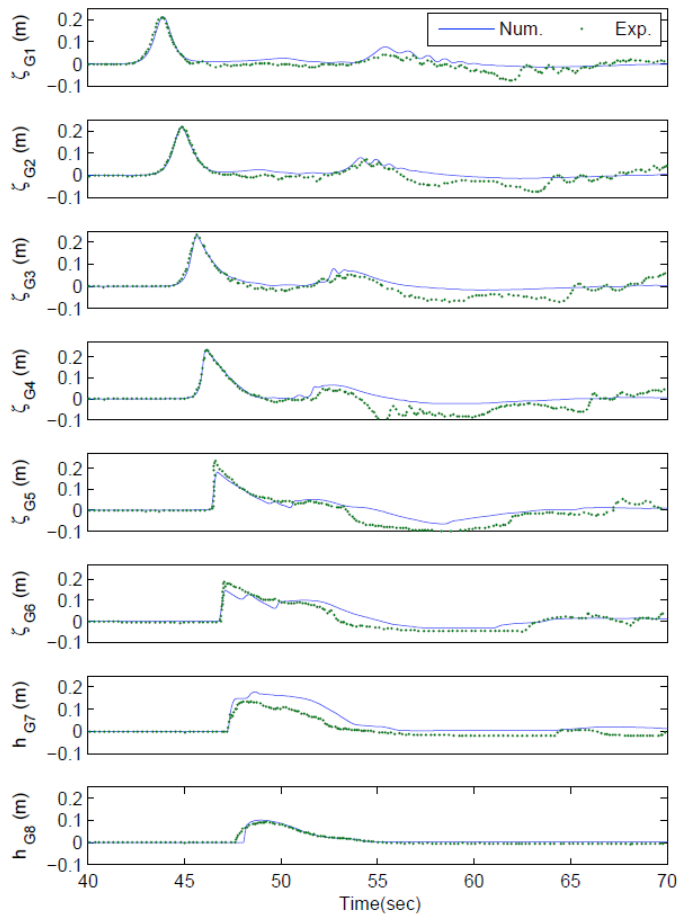


WAVE #1 Time = 36.944 sec



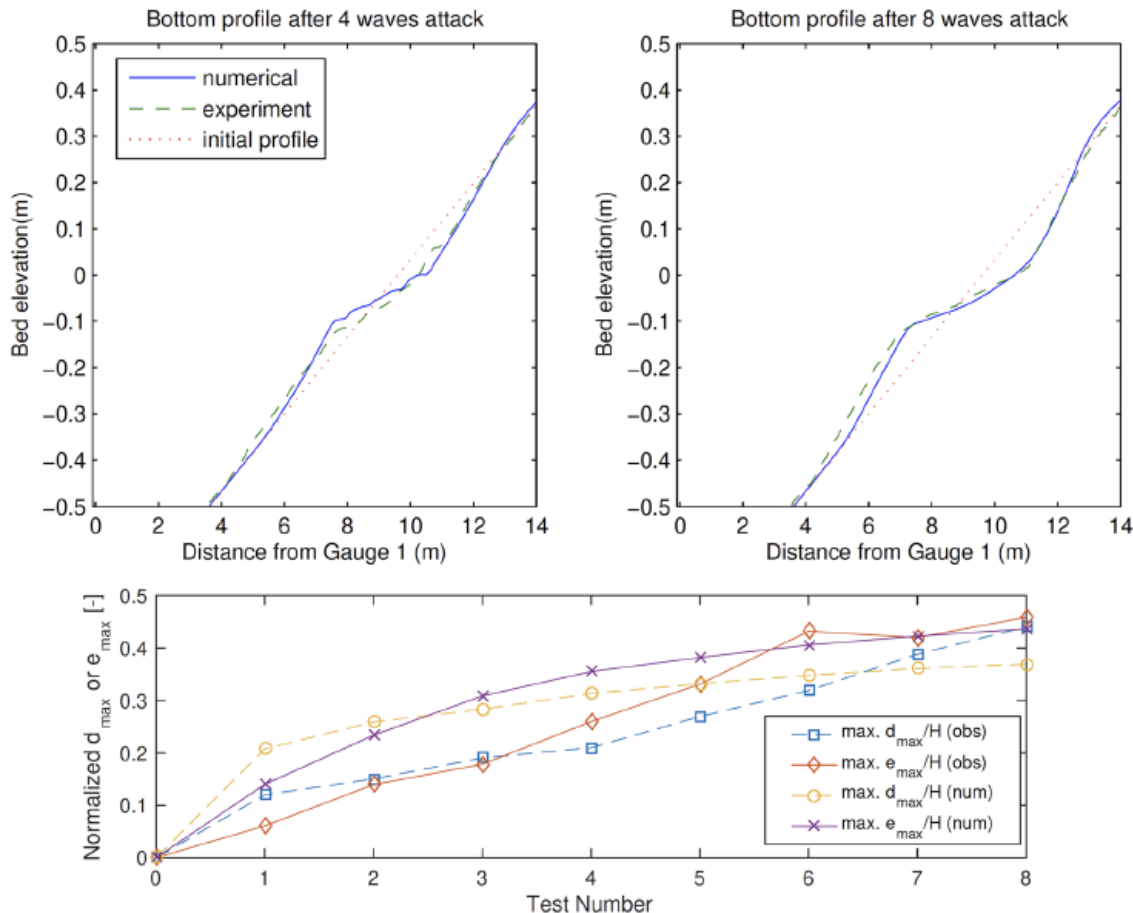
Validations : Case 2

- CASE2 : Breaking Solitary Waves on Sloping Beach*



Validations : Case 2

- CASE2 : Breaking Solitary Waves on Sloping Beach*

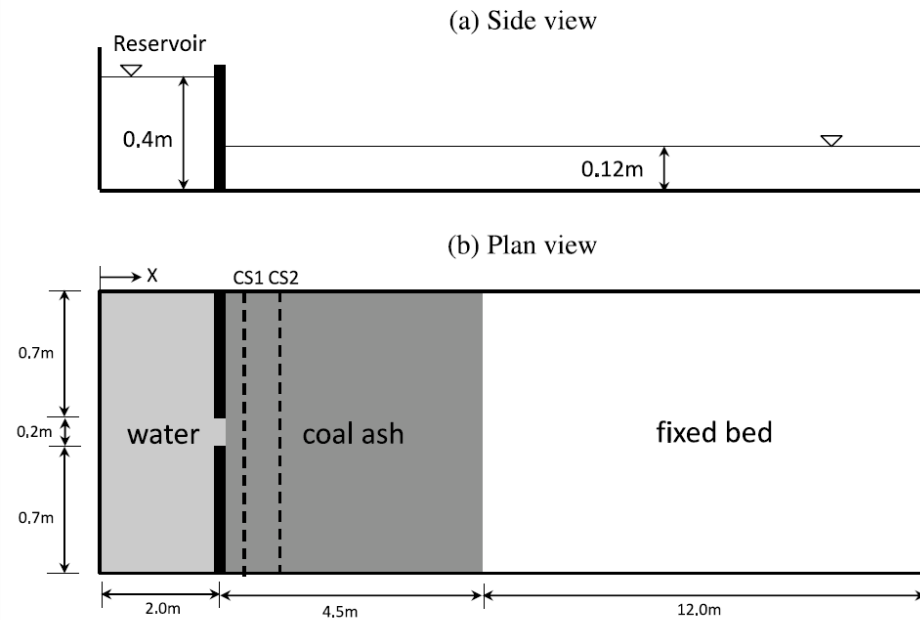
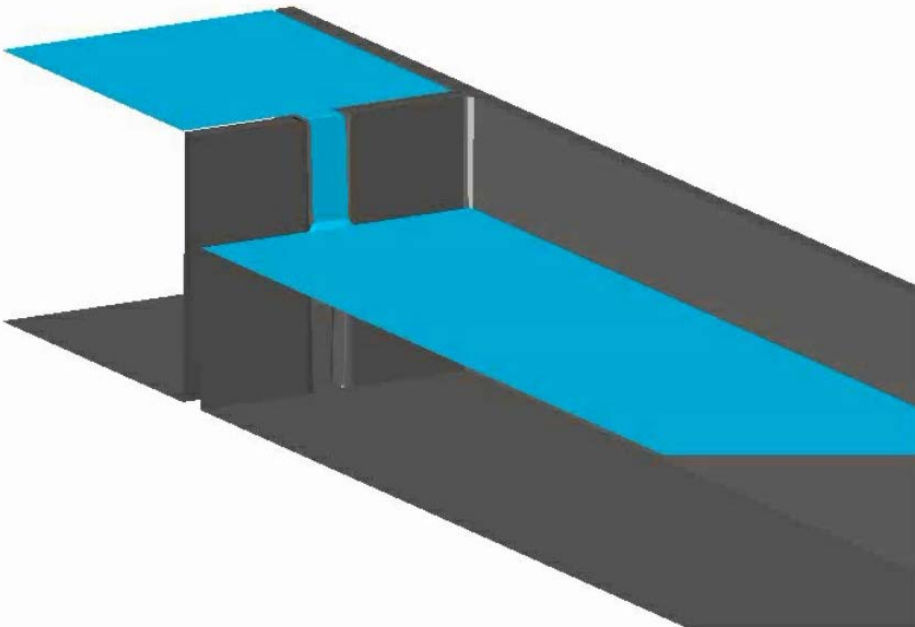


A significant erosion process at the foreshore may be explained by the strong backwash current initiated when a solitary wave retreats.

Validations : Case 3

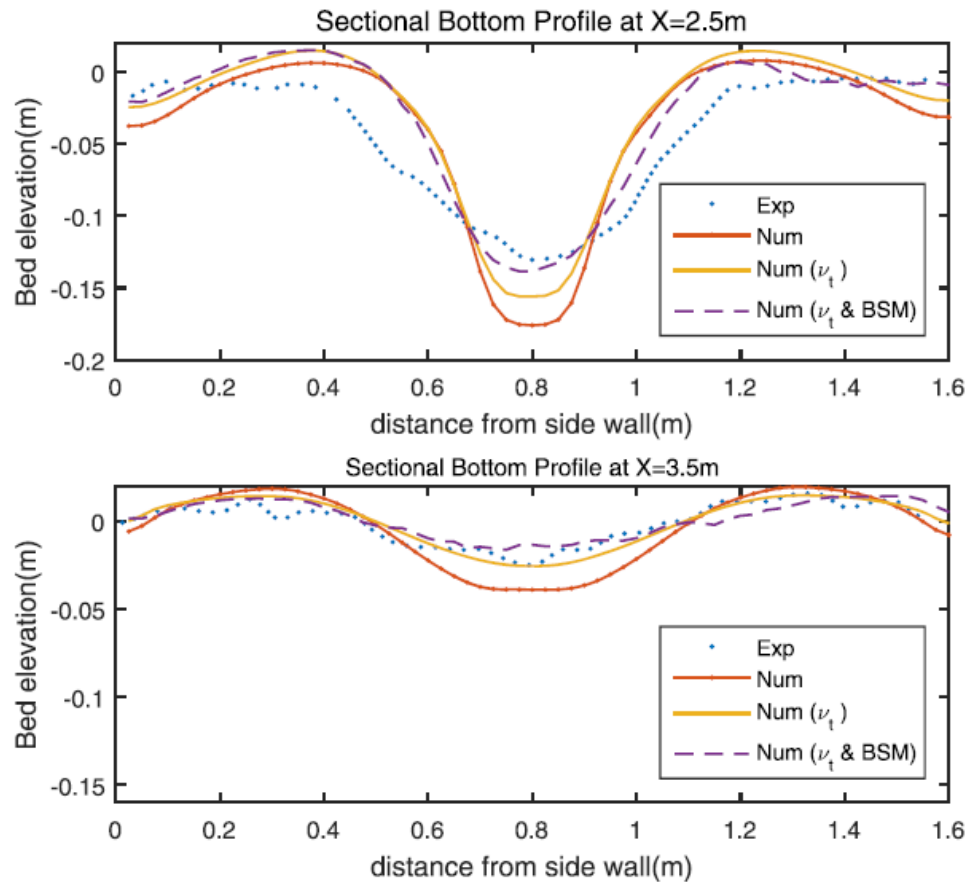
- CASE3 : Dam-Break Flow through a Partial Breach*

Time = 0.0018931 sec



Validations : Case 3

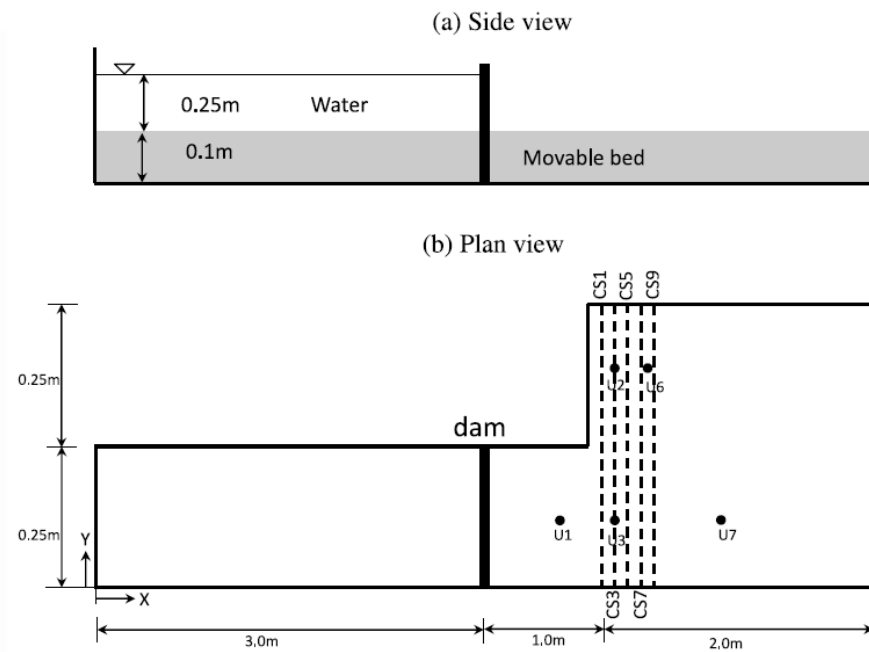
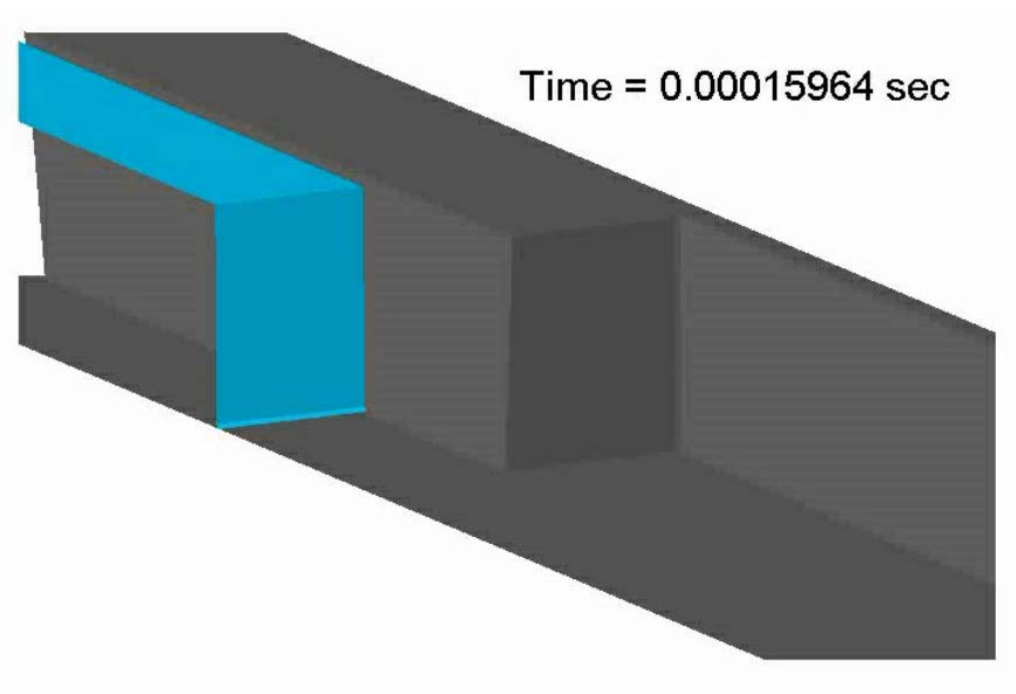
- CASE3 : Dam-Break Flow through a Partial Breach*



Reasonable agreement found when including turbulence effects, while the calculations overestimated the peak erosion depth at CS1.

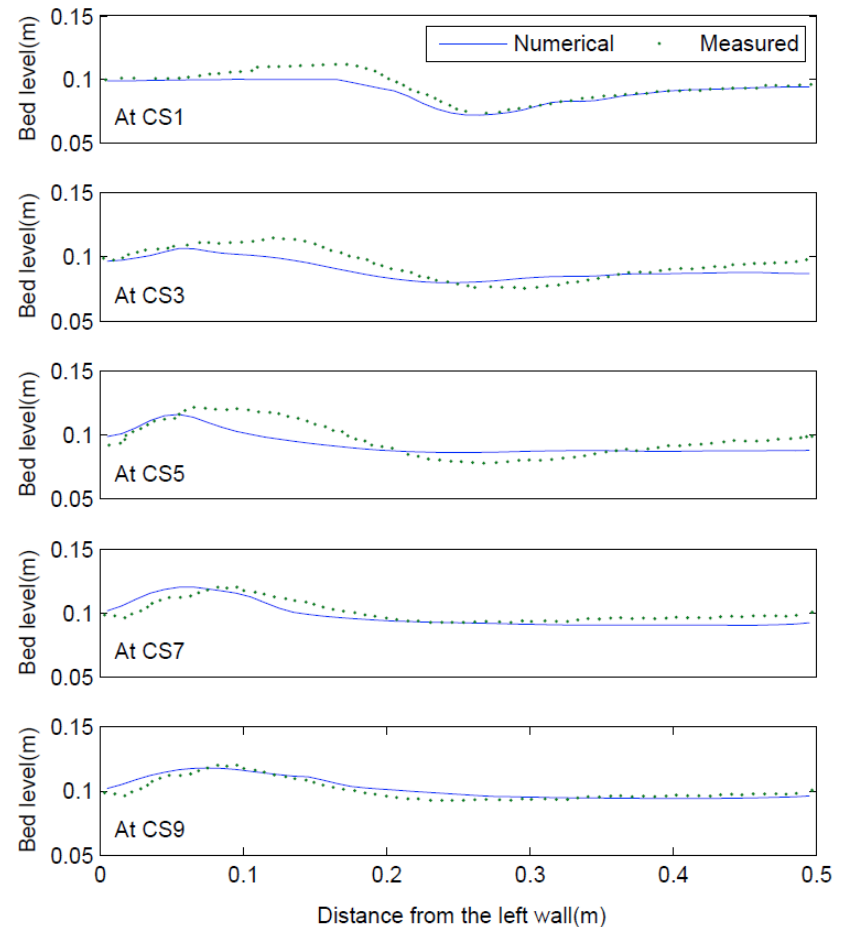
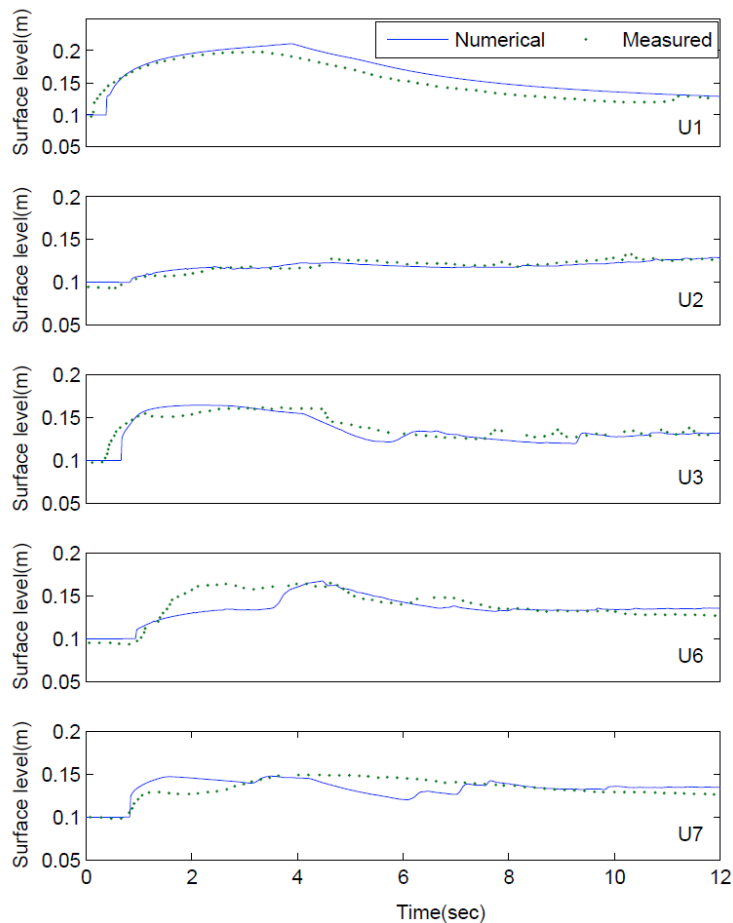
Validations : Case 4

- CASE4 : Dam-Break Flow over a Movable Bed Channel with Sudden Enlargement*



Validations : Case 4

- CASE4 : Dam-Break Flow over a Movable Bed Channel*



3. Model Applications

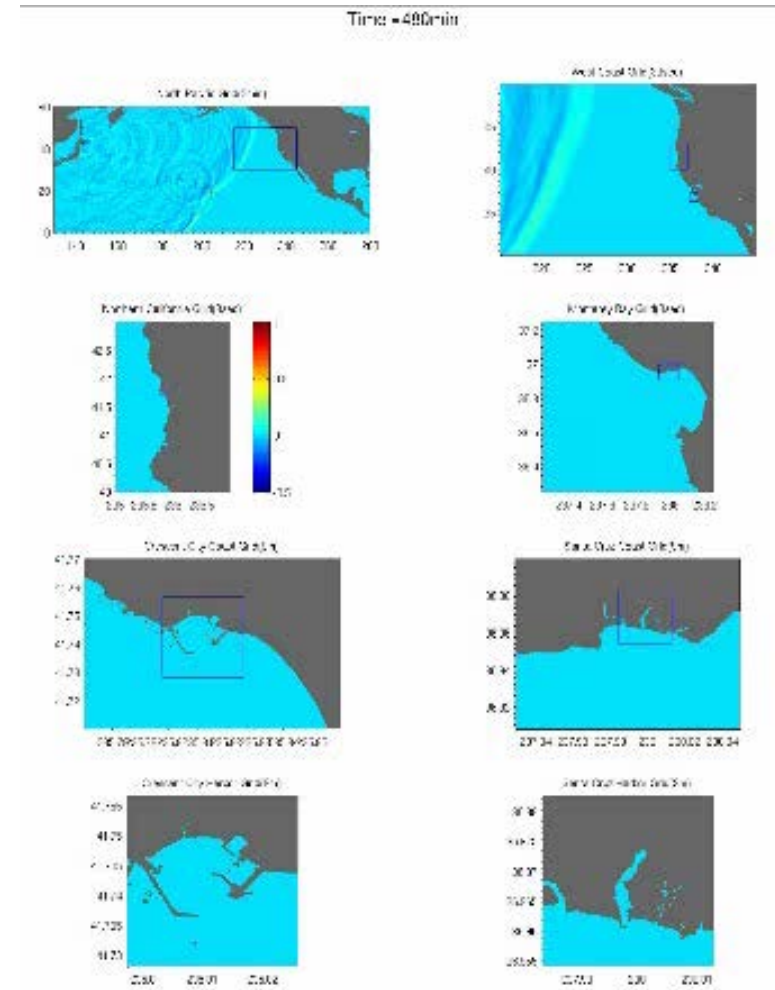
Applications

- *2011 Japan Tohoku Tsunami*

- **Simulation Result**
Layer level 1-6 : SWE,
Layer level 7-8 : Boussinesq)
- Successful recreation of complex hydrodynamic process by tsunami waves at the nearshore harbor area.
- Local focus : Crescent City Harbor, Santa Cruz Harbor

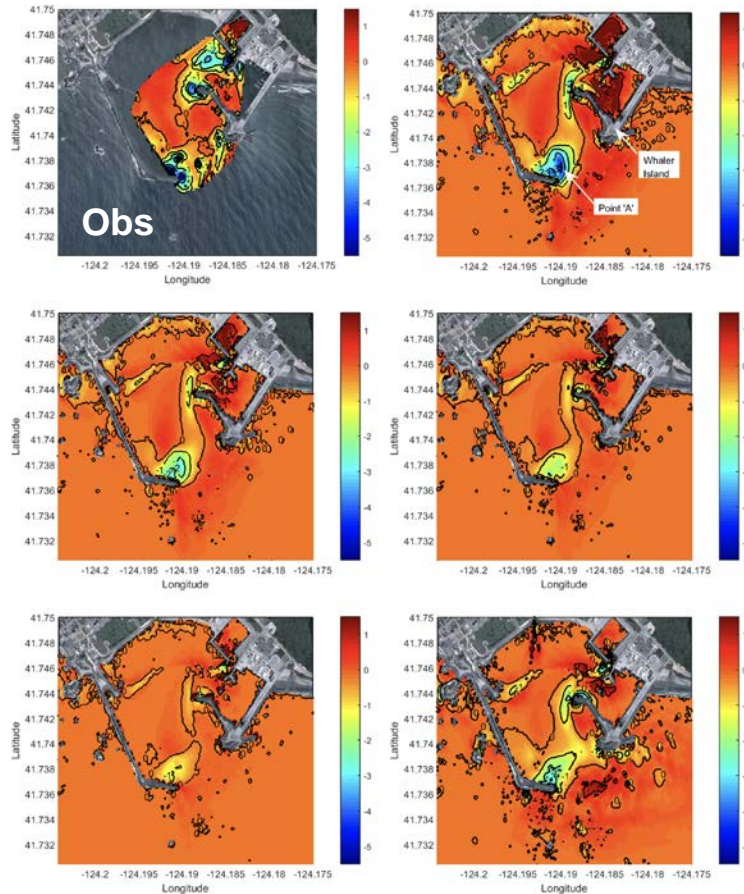
SWE
Layers

Boussiensq
Layers

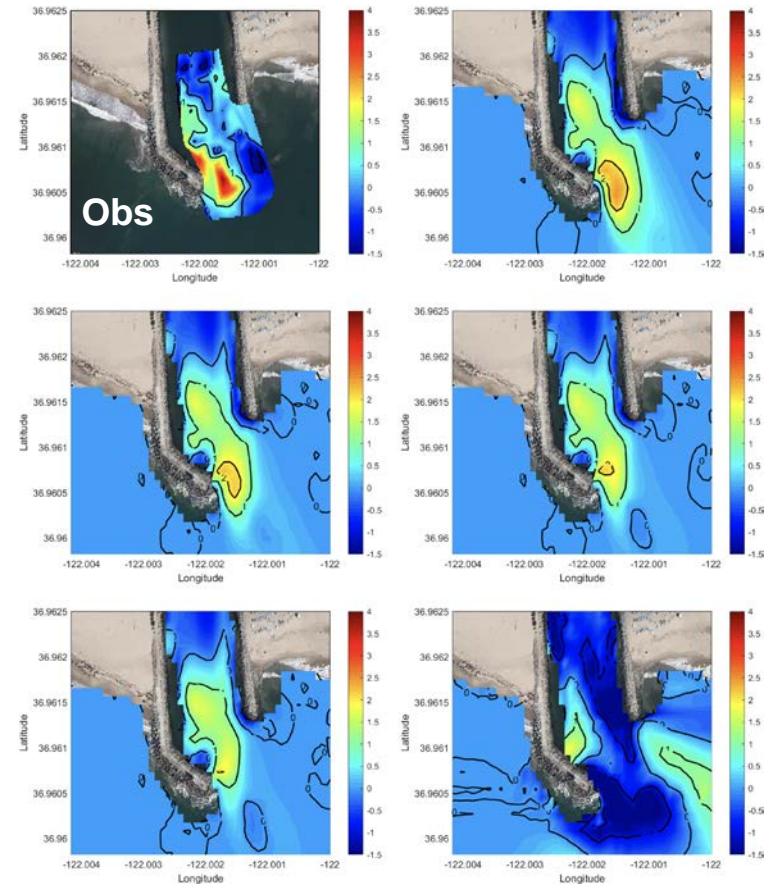


Applications

- Crescent City Harbor

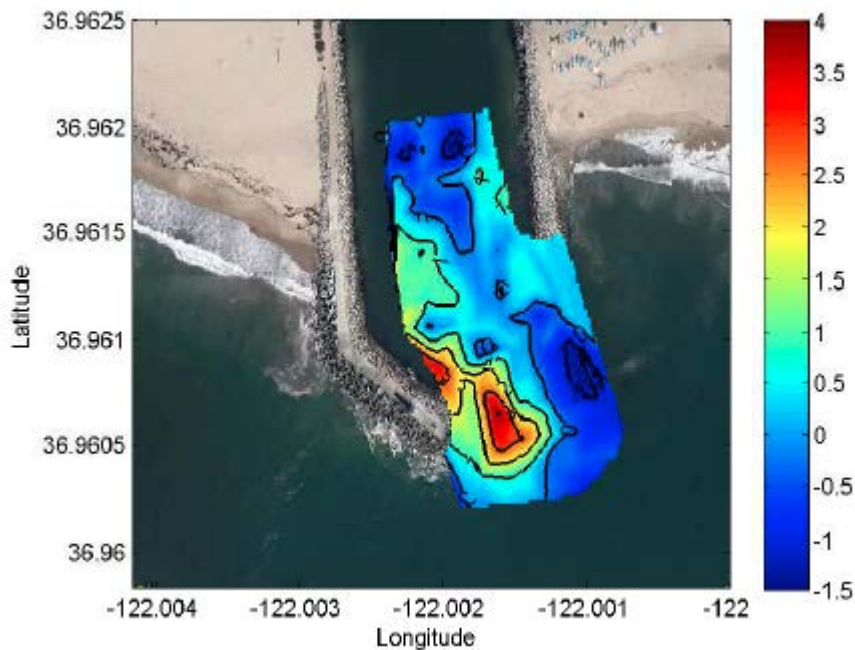


- Santa Cruz Harbor

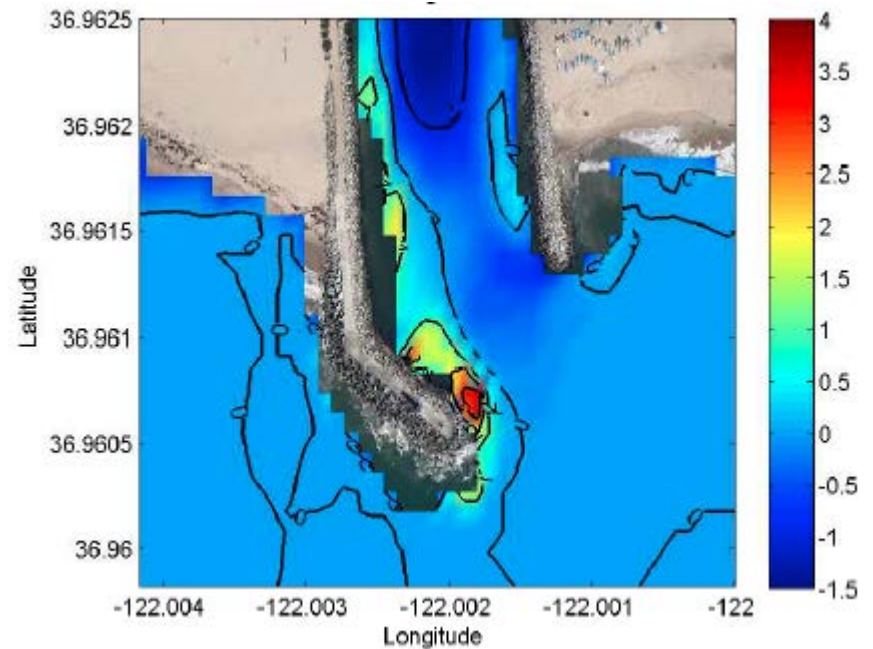


Applications

- Santa Cruz Harbor



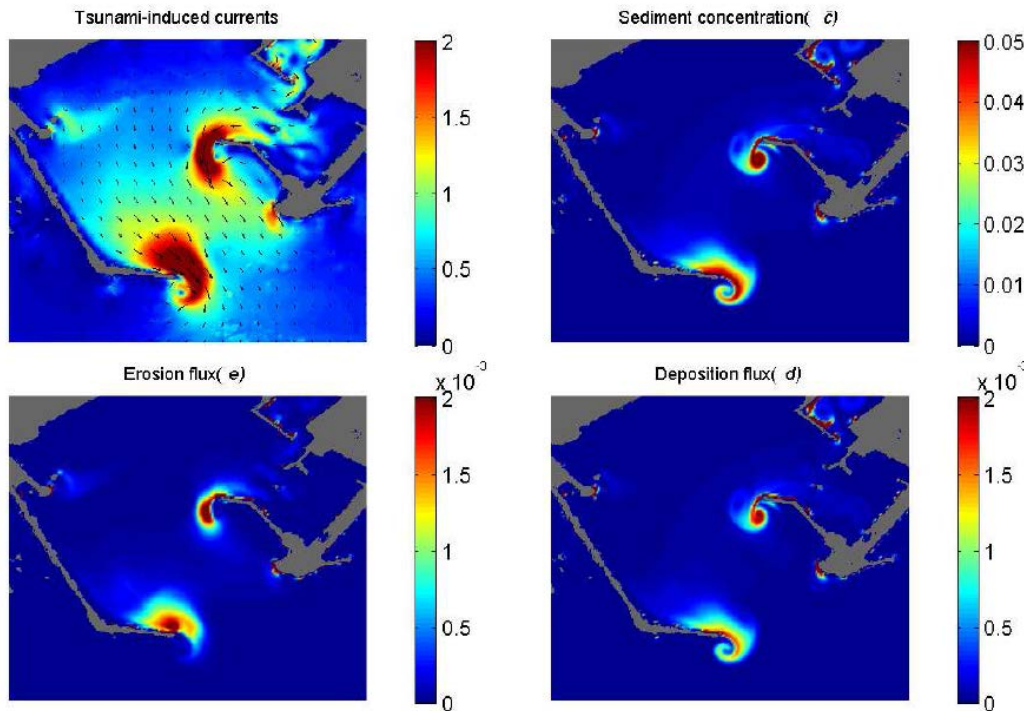
Observed data (Wilson et al.(2012))



Modeled results

Applications

- Numerical results of bathymetric changes in Crescent City Harbor at time $t = 637$ min since EQ.



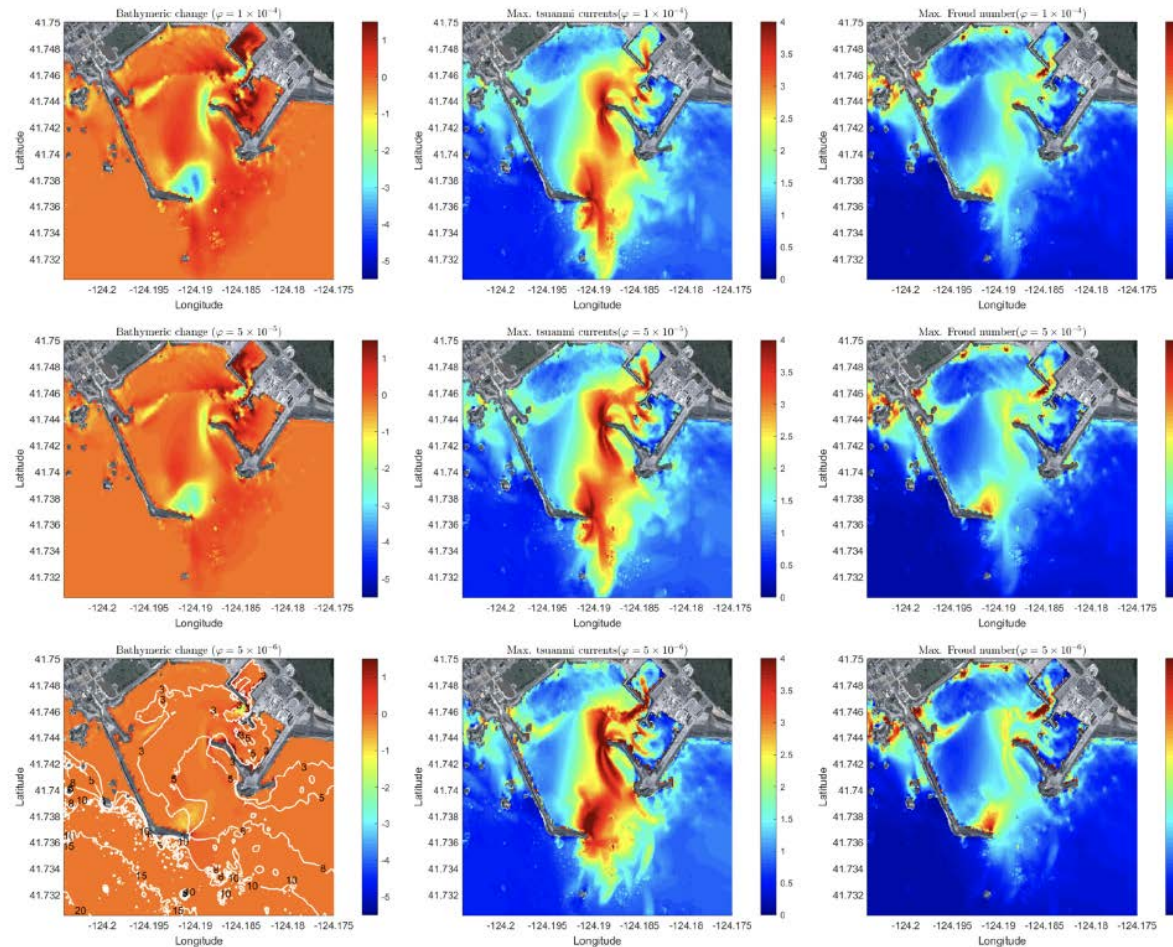
The strong current field formed by geometric contraction



Very active sedimentation processes in the vicinity of the outer and inner harbour entrances

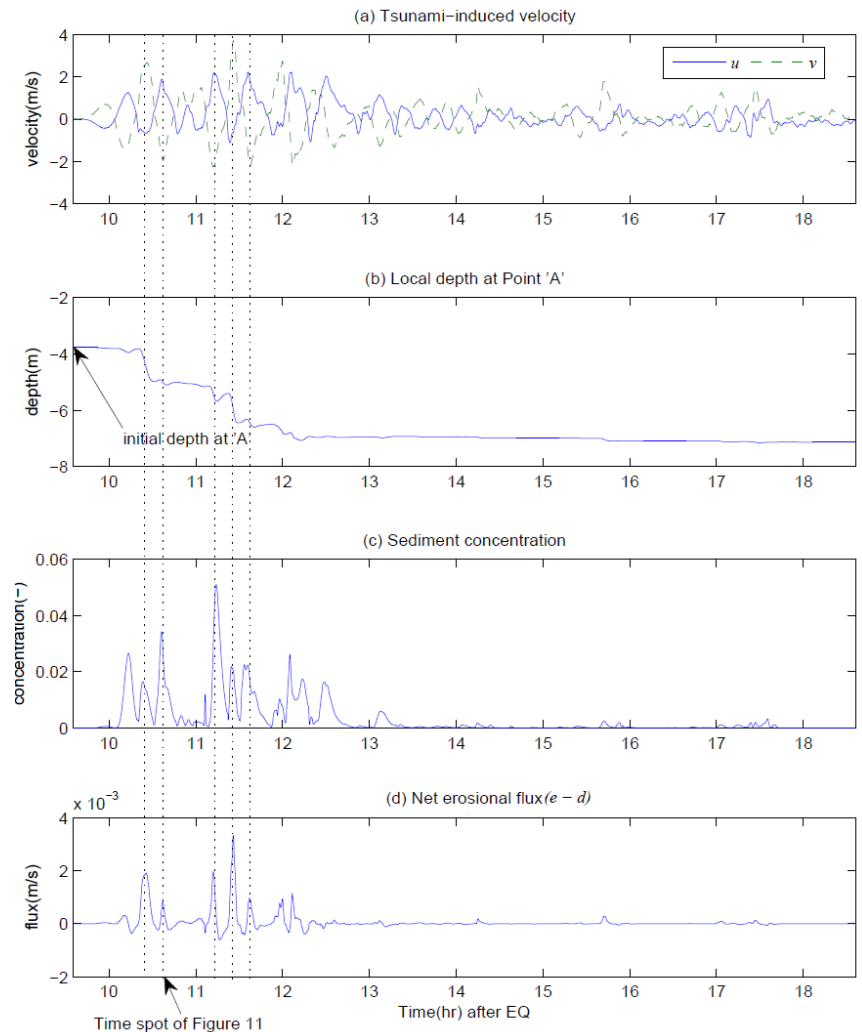
Applications

- Sensitivity tests on φ show that smaller φ produced less erosive sedimentation process
- Tsunami momentum is damped due to the strong interaction between surface processes and landforms, so the propagation gets retarded.



Applications

- Calculated time history at local point 'A'
 - (a) velocity
 - (b) local depth
 - (c) sediment concentration
 - (d) net erosion flux
- Coastal morphologic evolution is significant during the initial 3–4 hrs.
- Physics of tsunami-induced sedimentation is confined to a time scale of $O(\text{hours})$.



Conclusions

1. *A numerical model to predict morphologic changes by tsunami-induced current is developed by coupling 'hydrodynamic-morphologic models.*
2. *The model is validated through 1D and 2D sediment problems, with a comprehensive test on parametric sensitivity.*
3. *Application to the real tsunami event(2011 Japan Tohoku Tsunami) revealed that*
“the simulated bathymetric changes provide good approximations to the observation as long as the parameters is properly chosen”