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# The role of water vapor and cloud feedback on the evolution of the Indian summer monsoon over the last 22,000 years

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Chetankumar Jalihal

**Guided by:** Prof. J. Srinivasan &  
Prof. Arindam Chakraborty

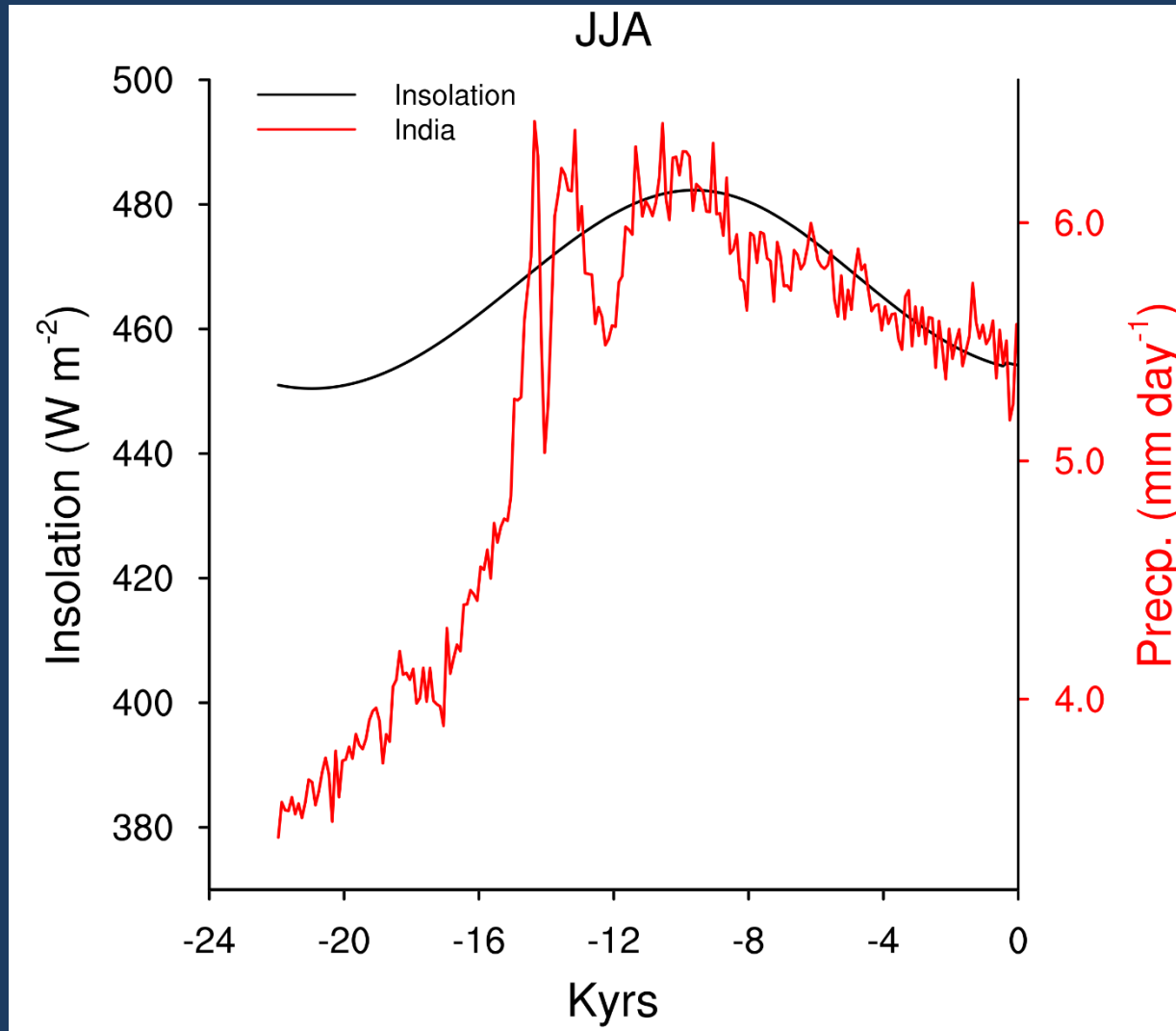


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# Short version

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# DIFFERENT SENSITIVITY OF MONSOON TO INSOLATION



**TraCE21K** simulation

**Model:** CCSM3,  
fully coupled ESM,  
with transient  
boundary conditions,  
 $3.75^\circ \times 3.75^\circ$

**Duration:** 22 Kyrs

# OBJECTIVE

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**Identify** the feedbacks.

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**Quantify** the role of forcings and feedbacks.

# ENERGETICS FRAMEWORK

(Neelin & Held 1987; Raymond 2009)

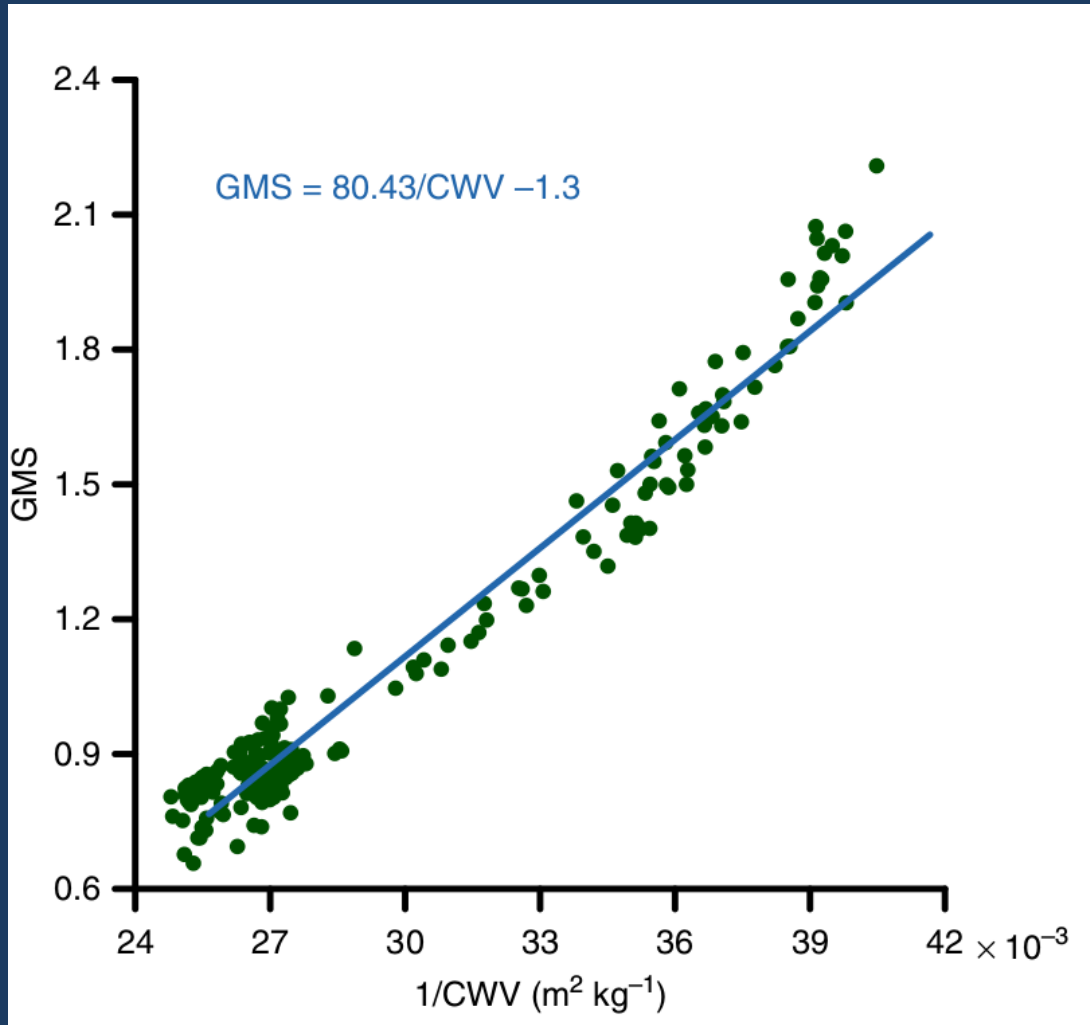
$$P - E = \frac{Q_{div}}{GMS}$$

Net energy flux into atmosphere

Gross Moist Stability

$$GMS = - \frac{\int_{P_B}^{P_T} \vec{U} \cdot \nabla m + \omega \frac{\partial m}{\partial p} dp}{L_v \int_{P_B}^{P_T} \vec{U} \cdot \nabla q + \omega \frac{\partial q}{\partial p} dp}$$

# GMS IS A UNIQUE FUNCTION OF WATER VAPOR

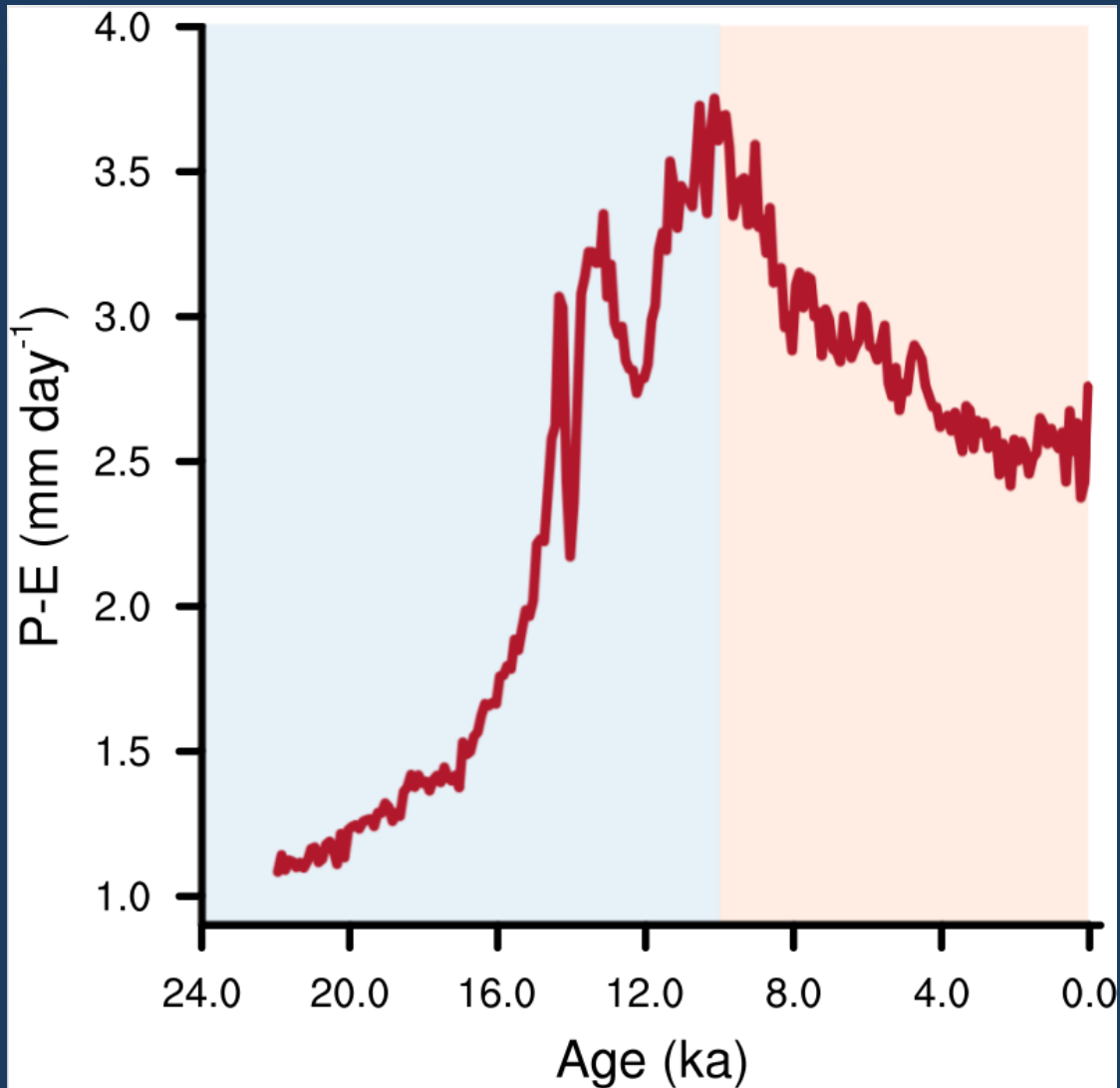


*Jalihal et al. (2019) Nat. Comm.*

$$P - E = \frac{Q_{div}}{GMS}$$

$$P - E = \frac{Q_{div}}{80.4/CWV - 1.3}$$

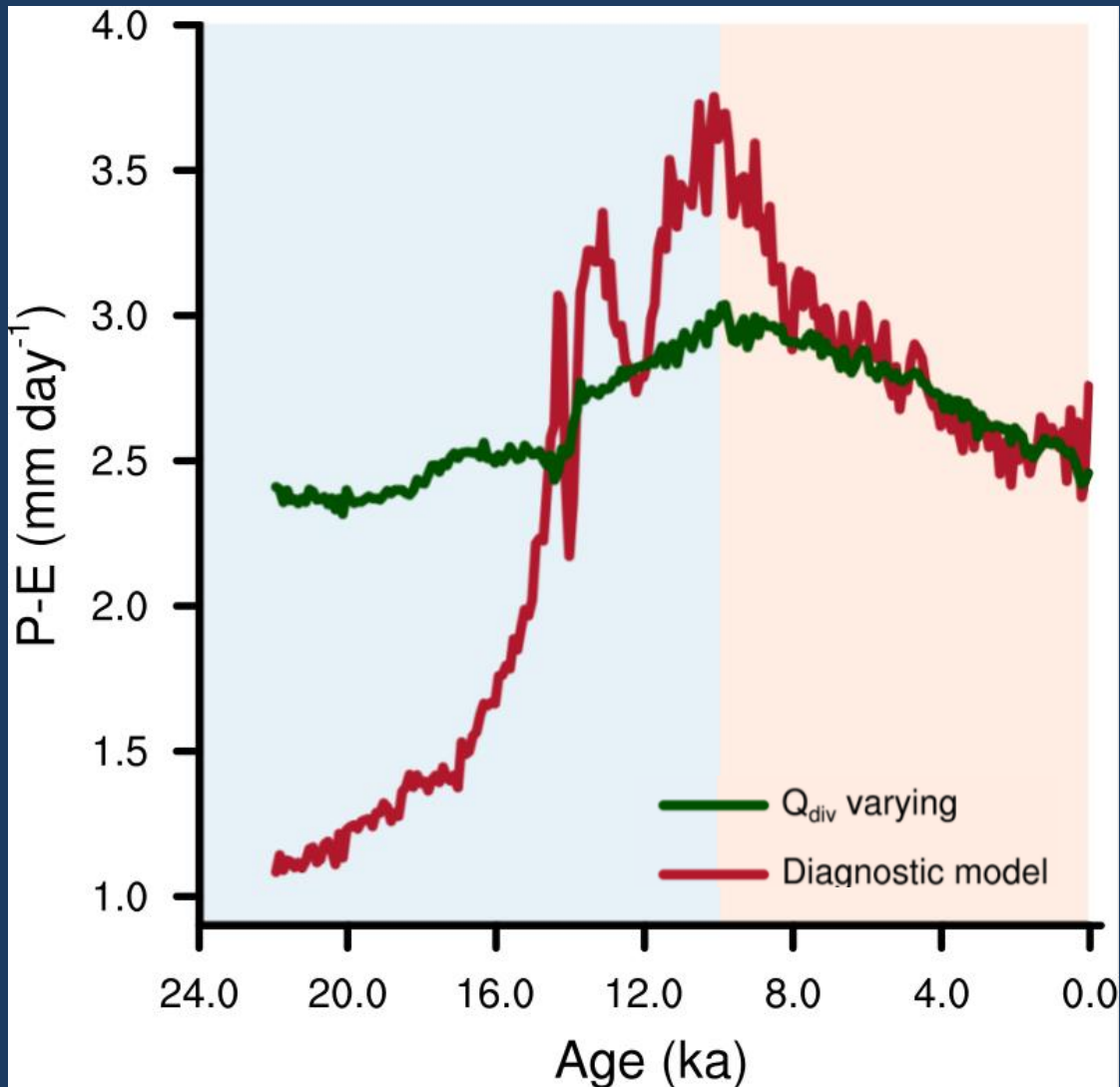
# (P-E) OVER THE LAST 22,000 YRS



$$P - E = \frac{Q_{div}}{80.4 / CWV - 1.3}$$

*Jaliha et al. (2019) Nat. Comm.*

# $Q_{DIV}$ DRIVES (P-E) DURING HOLOCENE



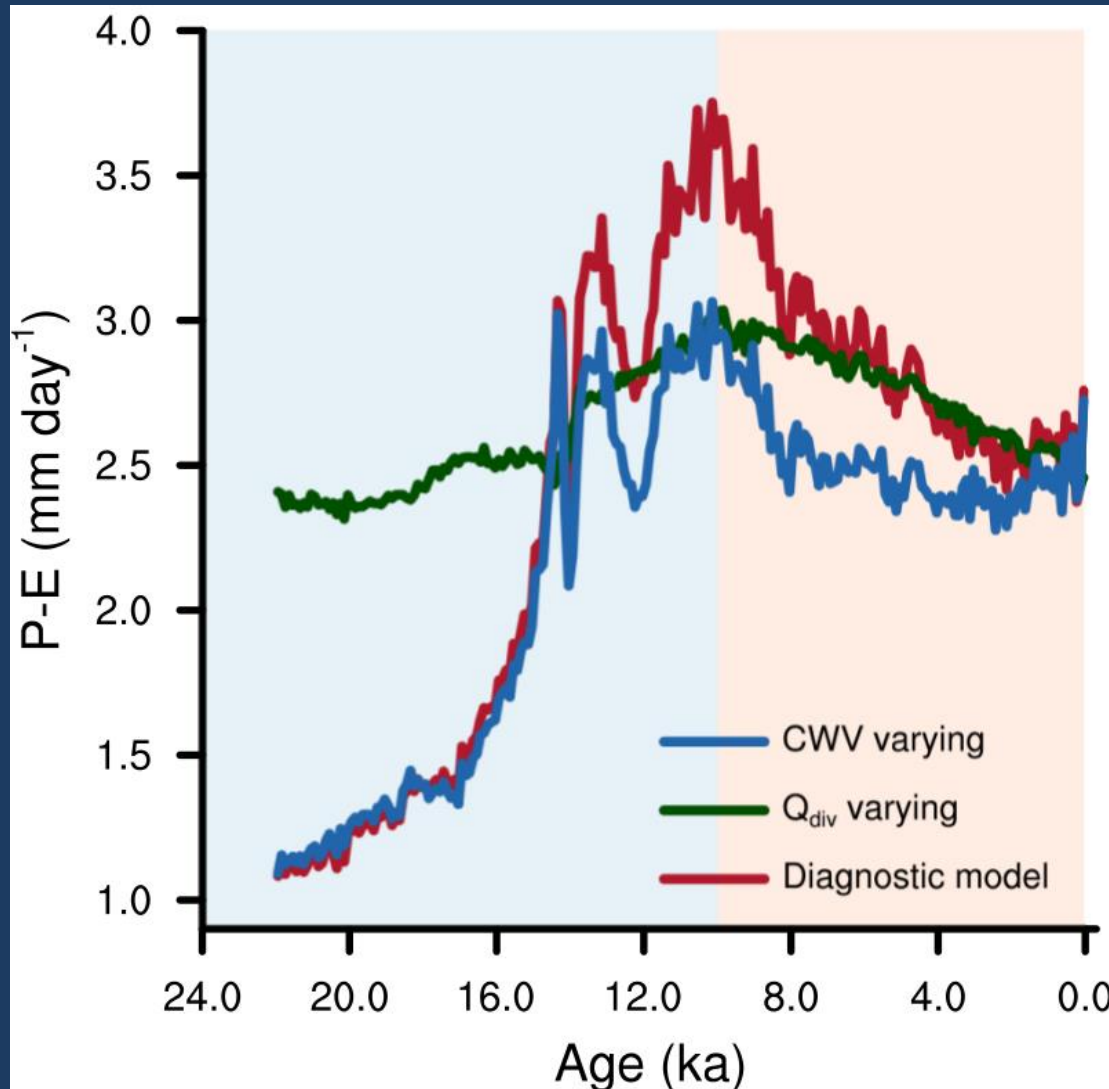
$$P - E = \frac{Q_{div}}{80.4 / \text{CWV} - 1.3}$$

**CWV fixed at pre-industrial values.**

*Jaliha et al. (2019) Nat. Comm.*



# WATER VAPOR PLAYS CRUCIAL ROLE DURING DEGLACIAL

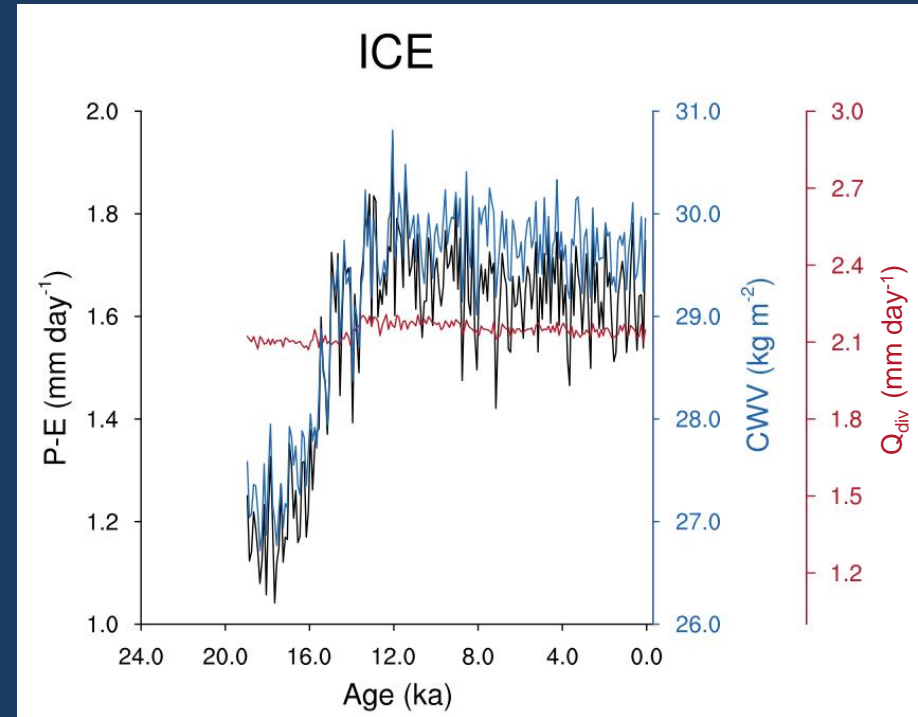
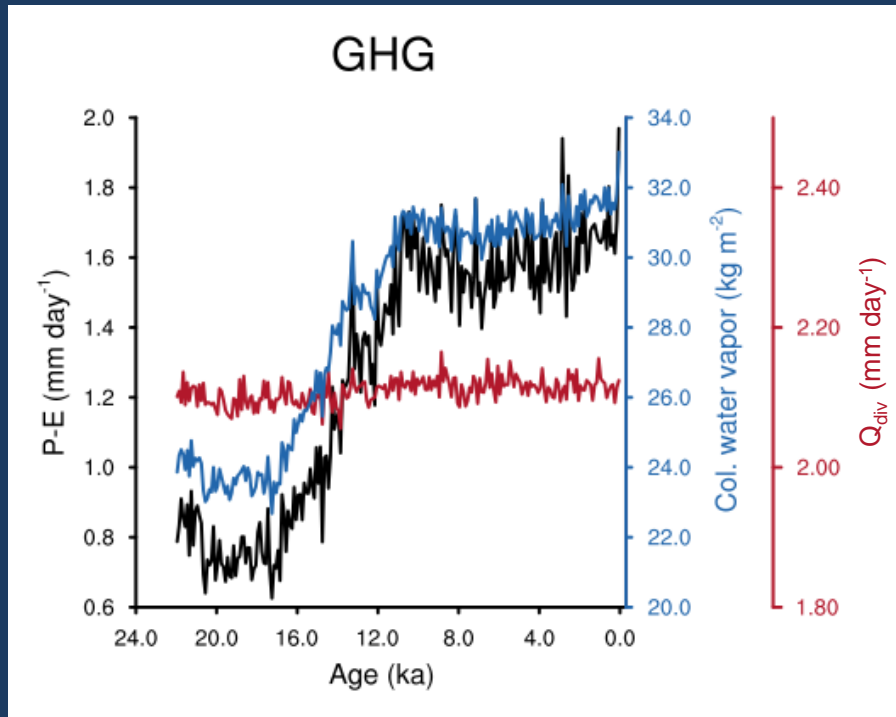


$$P - E = \frac{Q_{div}}{80.4 / CWV - 1.3}$$

**Q<sub>div</sub> fixed at  
pre-industrial values.**

*Jaliha et al. (2019) Nat. Comm.*

# GREENHOUSE GASES AND ICE SHEETS AFFECT GMS ALONE

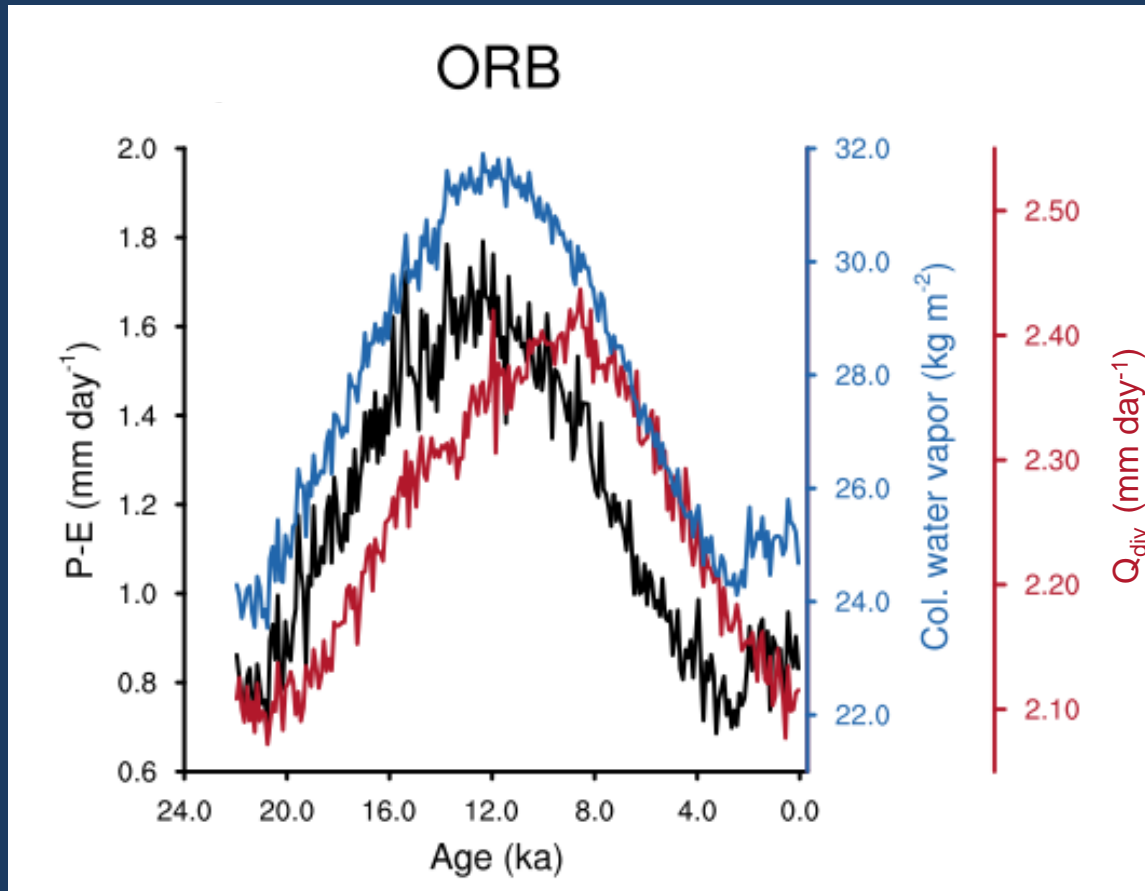


$$P - E = \frac{Q_{div}}{80.4 / CWV - 1.3}$$

**GHG and ice sheets  
do not influence  $Q_{div}$**

*Jaliha et al. (2019) Nat. Comm.*

# ORBITAL FORCING AFFECTS BOTH GMS AND $Q_{DIV}$



$$P - E = \frac{Q_{div}}{80.4 / CWV - 1.3}$$

Even though both  $Q_{div}$  and GMS are influenced, GMS is dominant, when boundary conditions are that of LGM.

Jaliha et al. (2019) Nat. Comm.

GHG & Ice sheets

Insolation

①

②

Surface temperature

Total column water vapor

Energy available

Precipitation

① - Dominant during the Deglacial

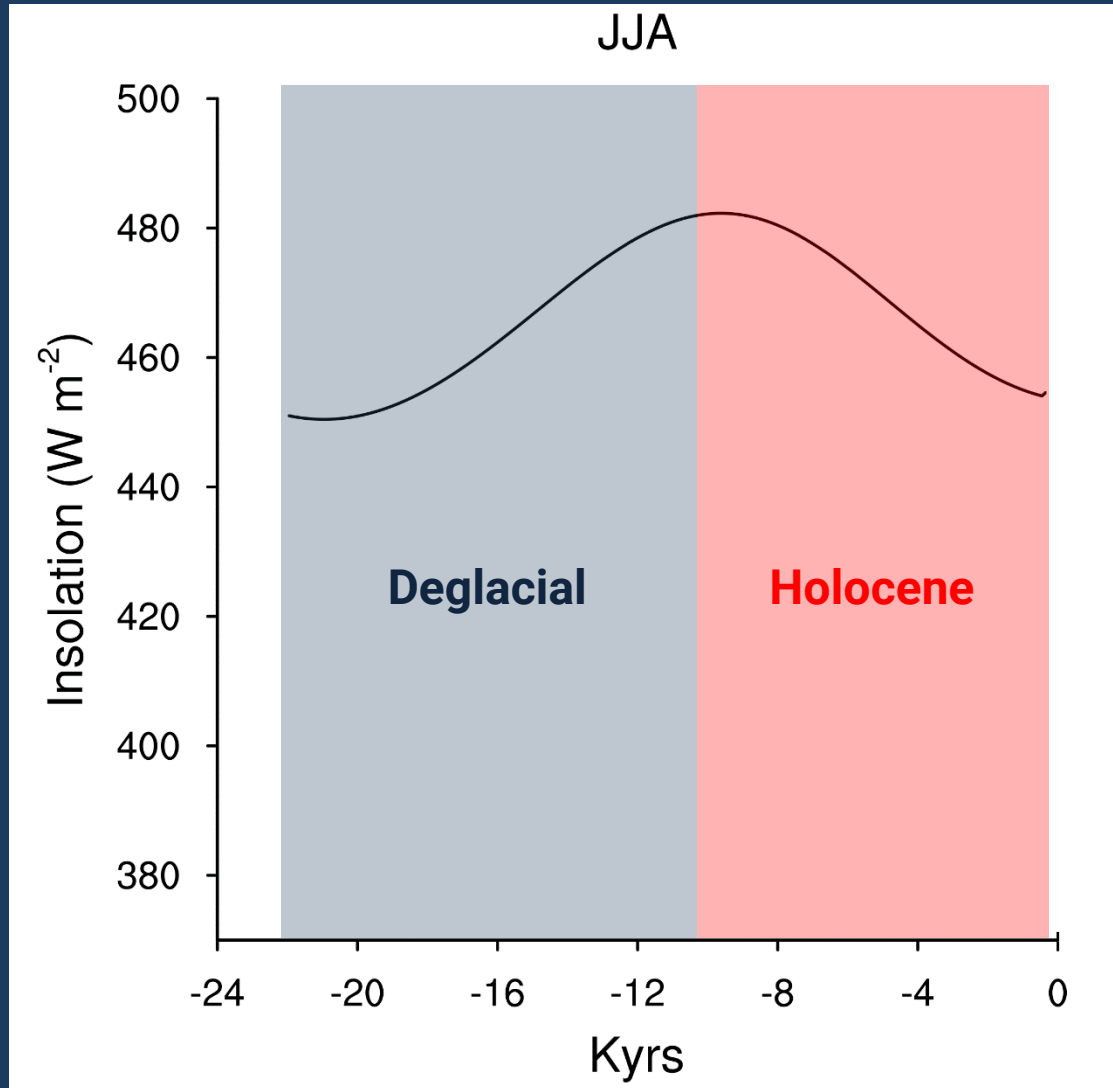
② - Dominant during the Holocene

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# Long version

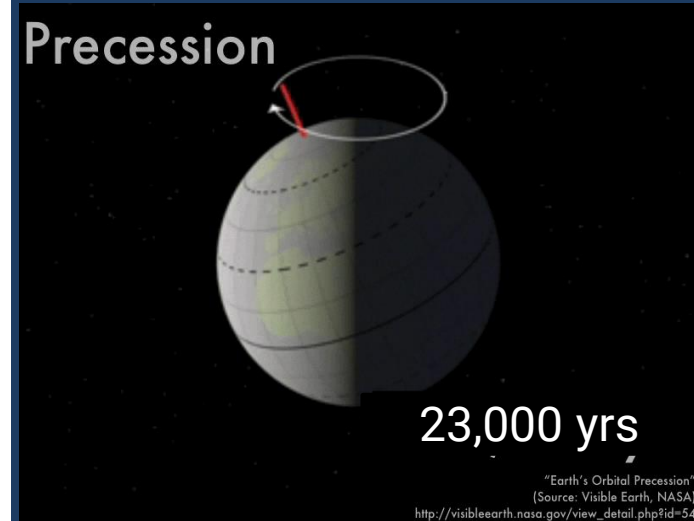
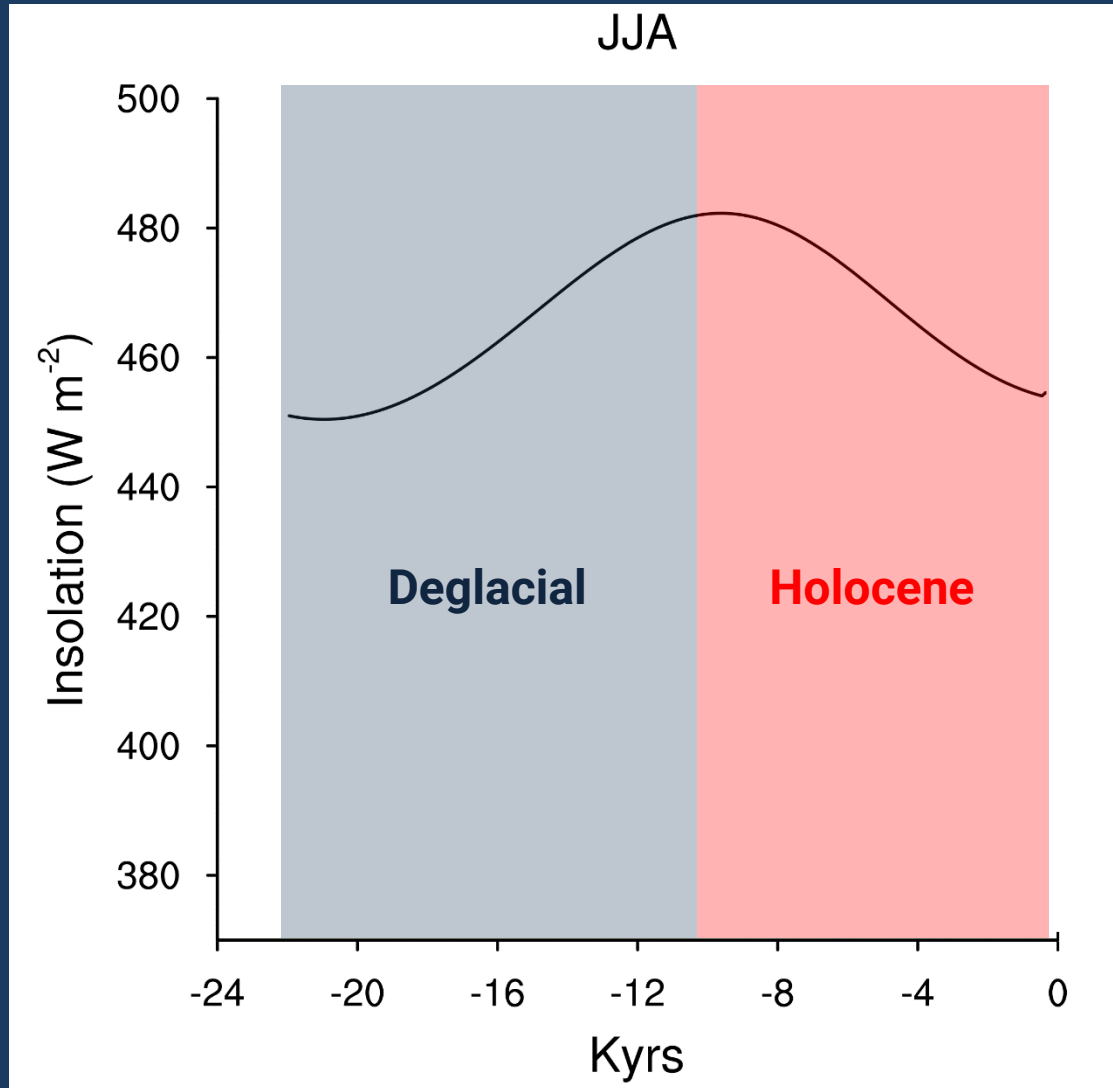
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# INSOLATION OVER LAST 22,000 YRS

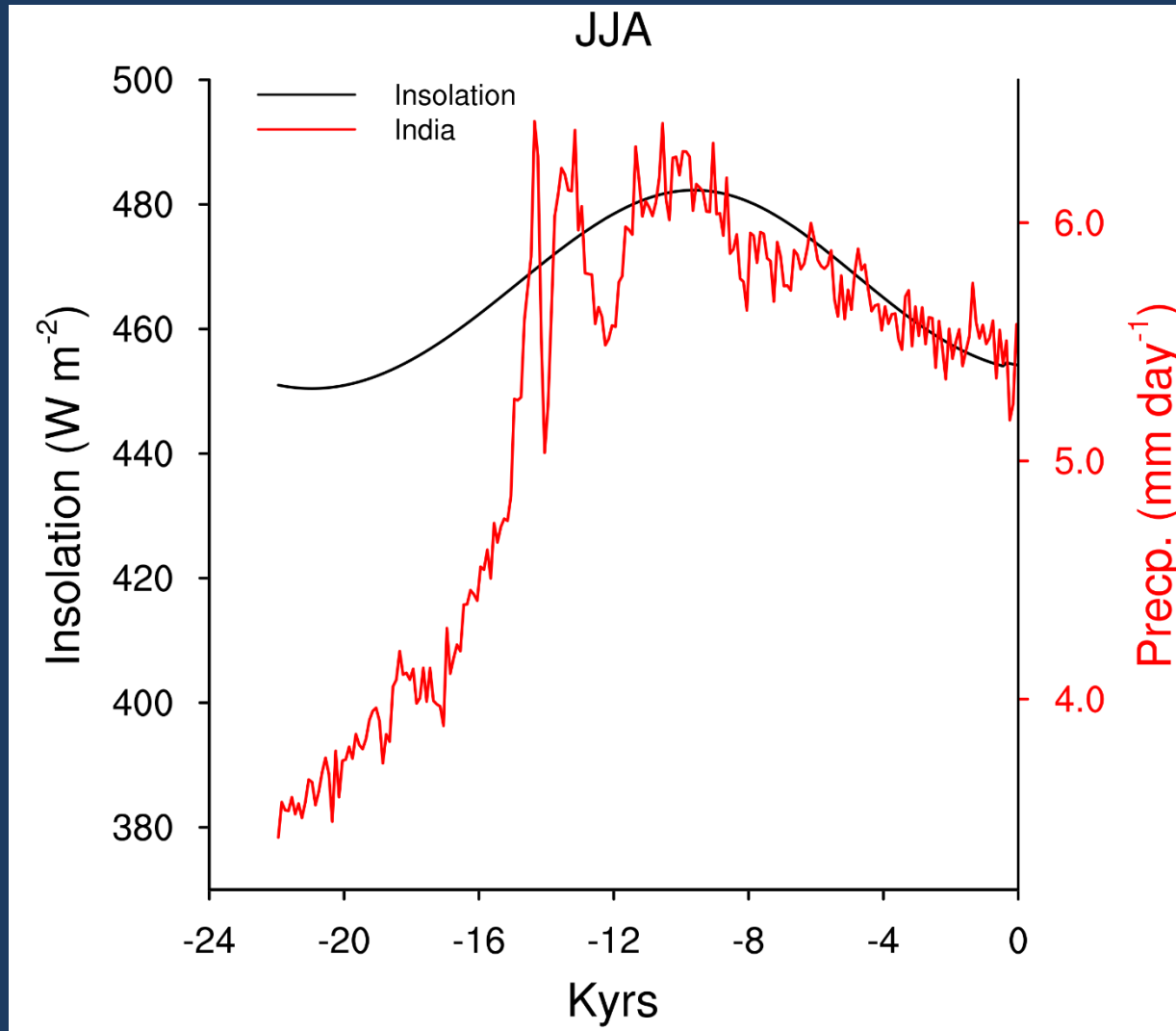


Summer insolation over India

# THIS VARIATION IS DUE TO PRECESSION OF EARTH



# DIFFERENT SENSITIVITY OF MONSOON TO INSOLATION



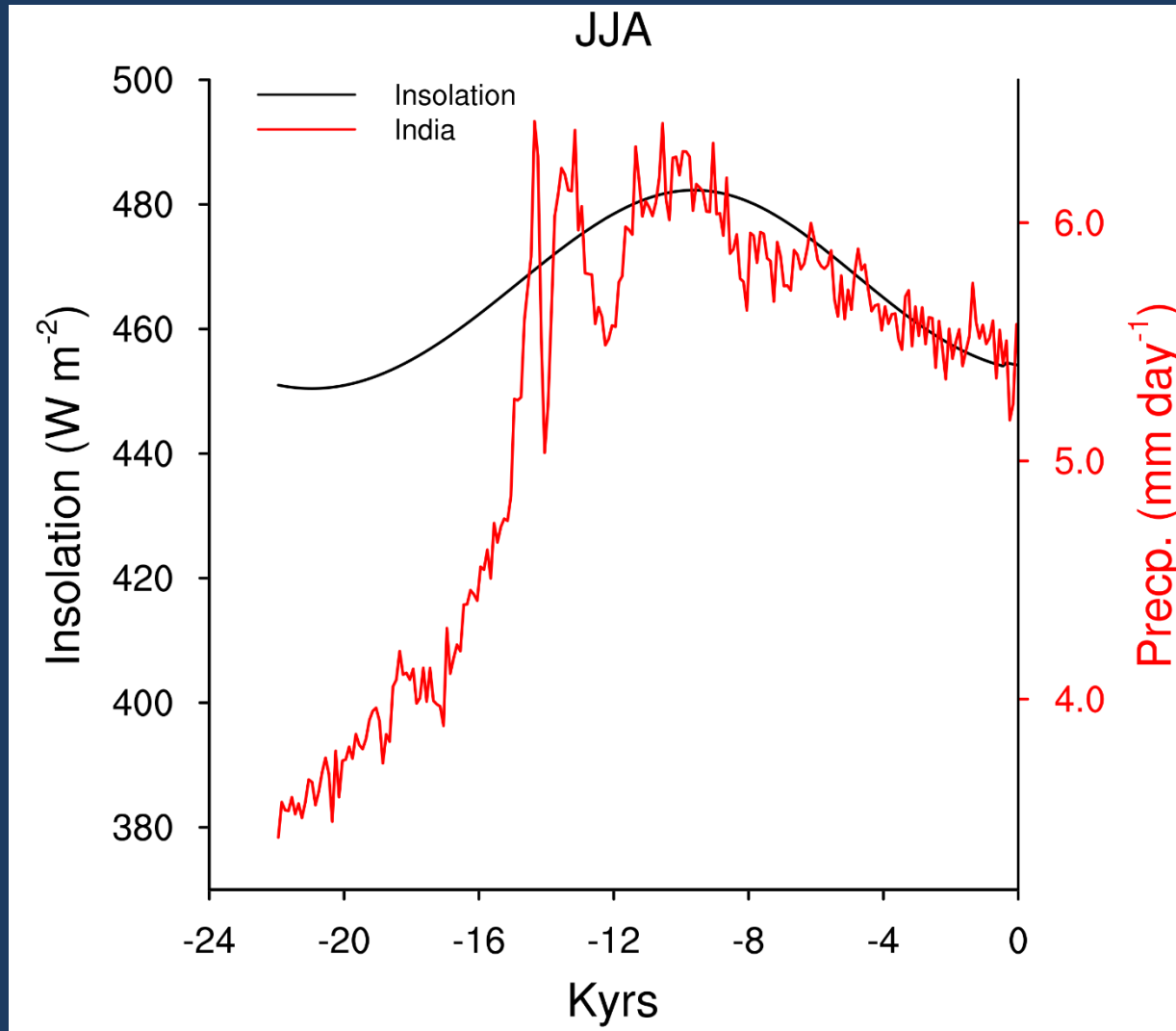
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**Model:** CCSM3,  
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 $3.75^\circ \times 3.75^\circ$

**Duration:** 22 Kyrs



# DIFFERENT SENSITIVITY OF MONSOON TO INSOLATION



**Deglacial forcings:**  
Insolation  
greenhouse gases  
ice sheets

**Holocene forcings:**  
Mainly insolation

# OBJECTIVE

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**Identify** the feedbacks.

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**Quantify** the role of forcings and feedbacks.

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# Method

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# ENERGETICS FRAMEWORK

Based on the equations for **conservation of moisture & Moist Static Energy**.

$$\int_{P_t}^{P_b} \left( \vec{U} \cdot \nabla m + \omega \frac{\partial m}{\partial p} \right) dp = g[Fb - Ft]$$

$$\int_{P_t}^{P_b} \left( \vec{U} \cdot \nabla q + \omega \frac{\partial q}{\partial p} \right) dp = g[E - P]$$

**m** - Moist Static Energy;  
( $m = C_p * T + g * Z + L_v * q$ )

**Fb** - Bottom Fluxes

**Ft** - Fluxes at Top of Atm.

**u** - x component of velocity

**q** - Specific Humidity (kg/Kg)

**ω** - Vertical velocity (Pa/s)

**Pb** - Surface Pressure

**Pt** - Top Pressure

**v** - y component of velocity

# ENERGETICS FRAMEWORK

*(Neelin & Held 1987; Raymond 2009)*

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$$P - E = \frac{Q_{div}}{GMS}$$

$$GMS = - \frac{\int_{P_B}^{P_T} \vec{U} \cdot \nabla m + \omega \frac{\partial m}{\partial p} dp}{L_v \int_{P_B}^{P_T} \vec{U} \cdot \nabla q + \omega \frac{\partial q}{\partial p} dp}$$

# ENERGETICS FRAMEWORK

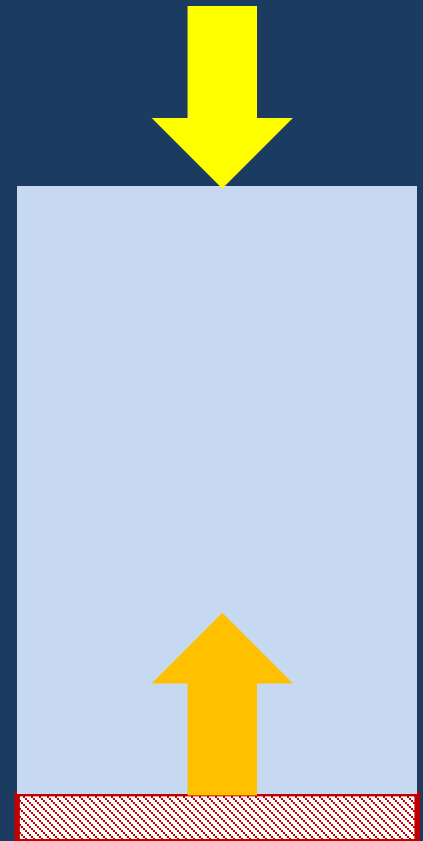
(Neelin & Held 1987; Raymond 2009)

$$P - E = \frac{Q_{div}}{GMS}$$

Net energy flux into atmosphere (top + bottom)

Over land net surface energy fluxes are small.

$$GMS = - \frac{\int_{P_B}^{P_T} \vec{U} \cdot \nabla m + \omega \frac{\partial m}{\partial p} dp}{L_v \int_{P_B}^{P_T} \vec{U} \cdot \nabla q + \omega \frac{\partial q}{\partial p} dp}$$



# ENERGETICS FRAMEWORK

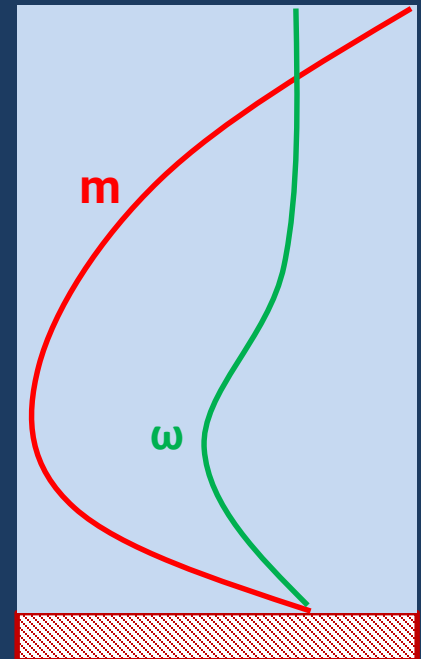
(Neelin & Held 1987; Raymond 2009)

$$P - E = \frac{Q_{div}}{GMS}$$

Gross Moist Stability.

Primarily depends on vertical profile of  $m$  and  $\omega$

$$GMS = - \frac{\int_{P_B}^{P_T} \vec{U} \cdot \nabla m + \omega \frac{\partial m}{\partial p} dp}{L_v \int_{P_B}^{P_T} \vec{U} \cdot \nabla q + \omega \frac{\partial q}{\partial p} dp}$$

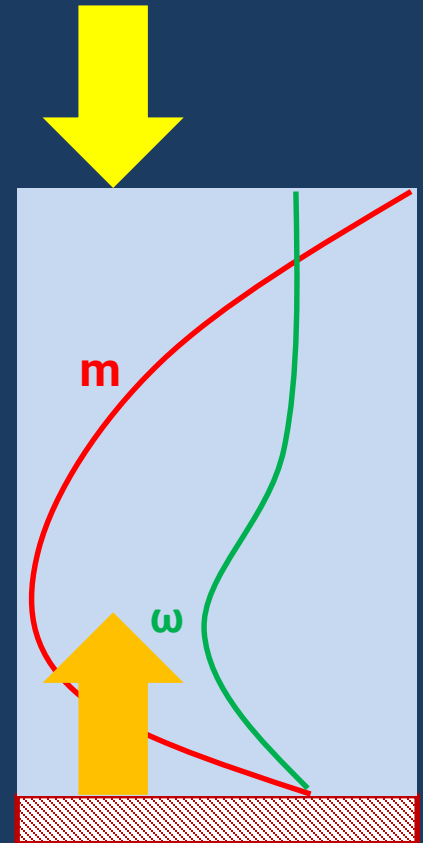


# ENERGETICS FRAMEWORK

(Neelin & Held 1987; Raymond 2009)

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# A CLOSER LOOK AT $Q_{DIV}$

$$Q_{div} = TOA\_Net\_SW - TOA\_OLR + SFC\_RAD + SHF + LHF$$

$Q_{net}$

**TOA\_Net\_SW-** Net Short Wave at Top of Atm

**TOA\_OLR** - Net Outgoing Longwave Radiation

**$Q_{div}$**  over land is influenced by insolation and cloud radiative feedback.

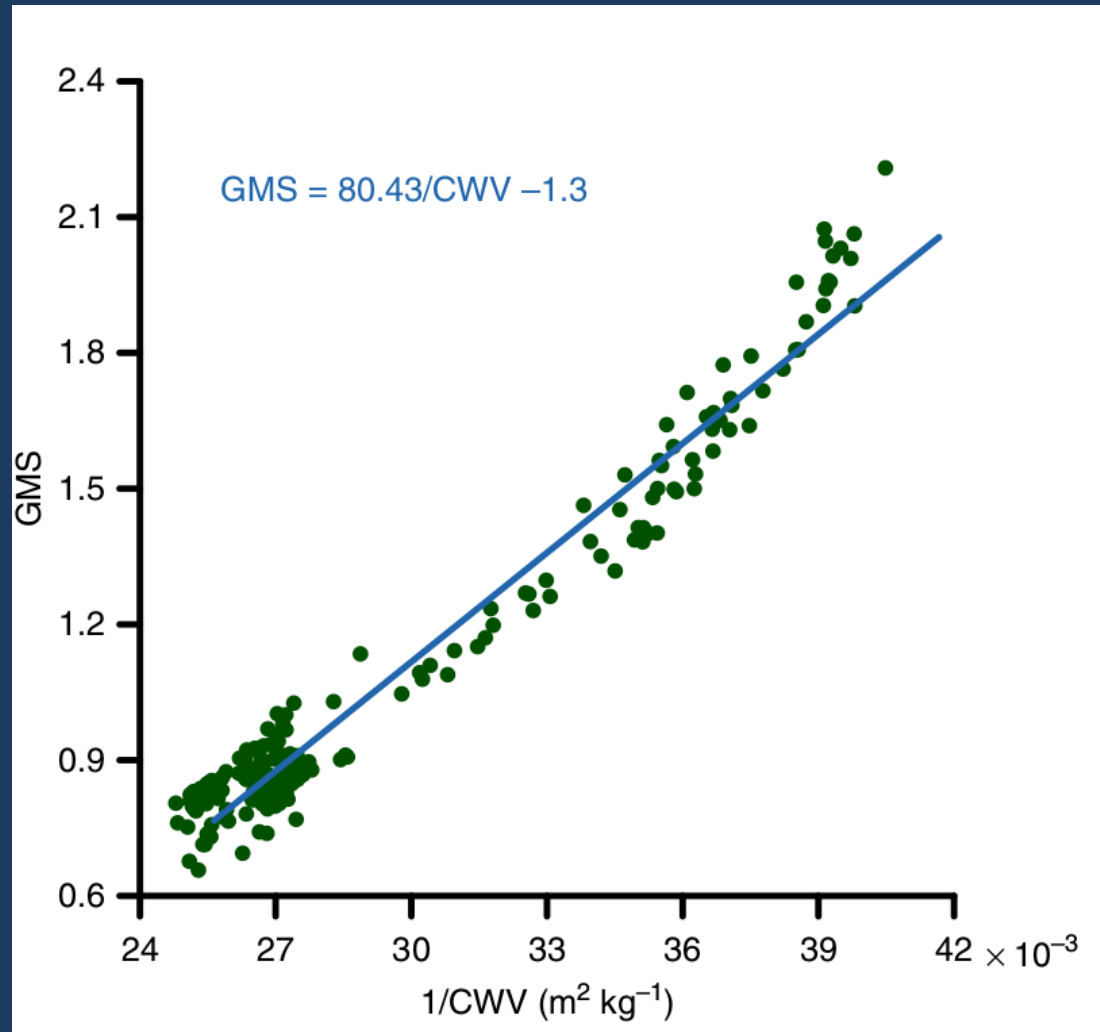


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# Results

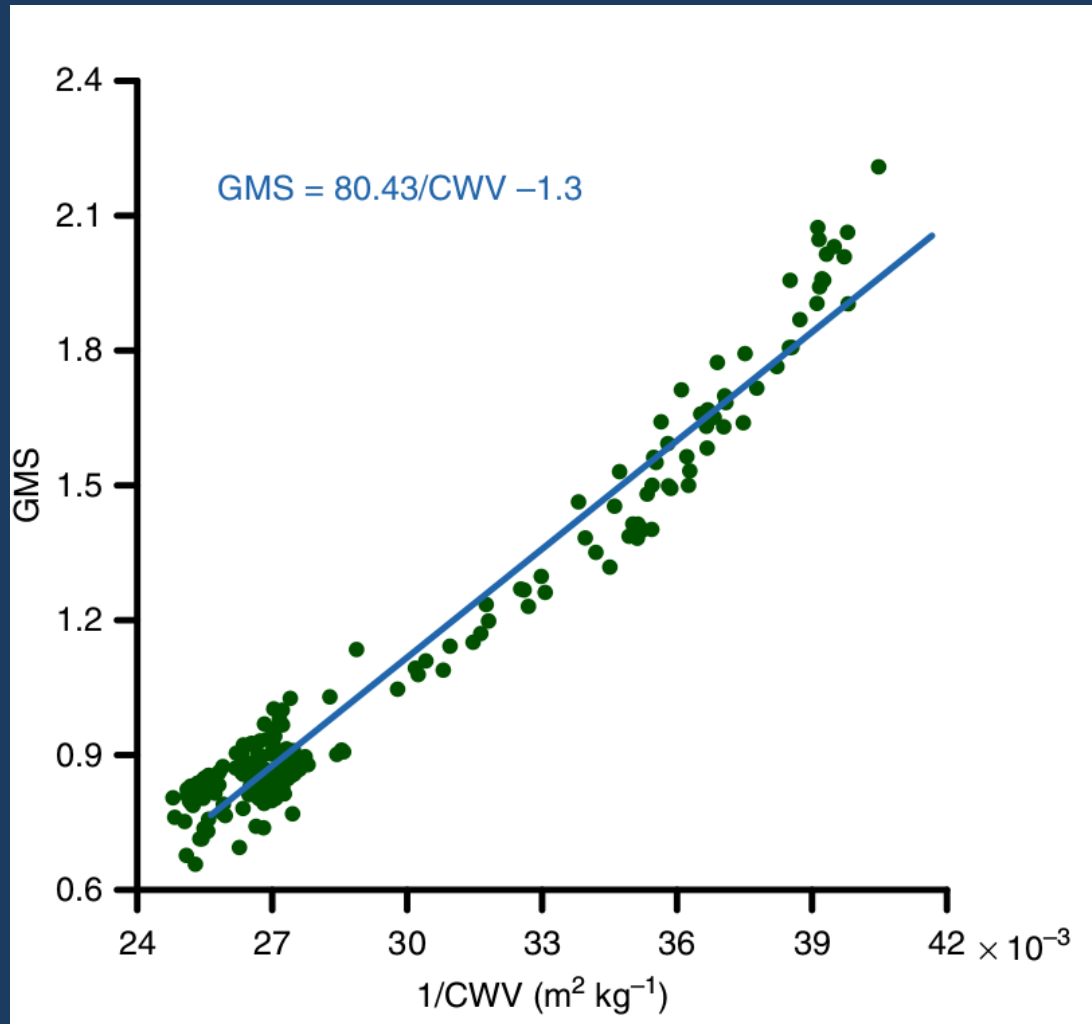
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# GMS IS A UNIQUE FUNCTION OF WATER VAPOR



*Jalihal et al. (2019) Nat. Comm.*

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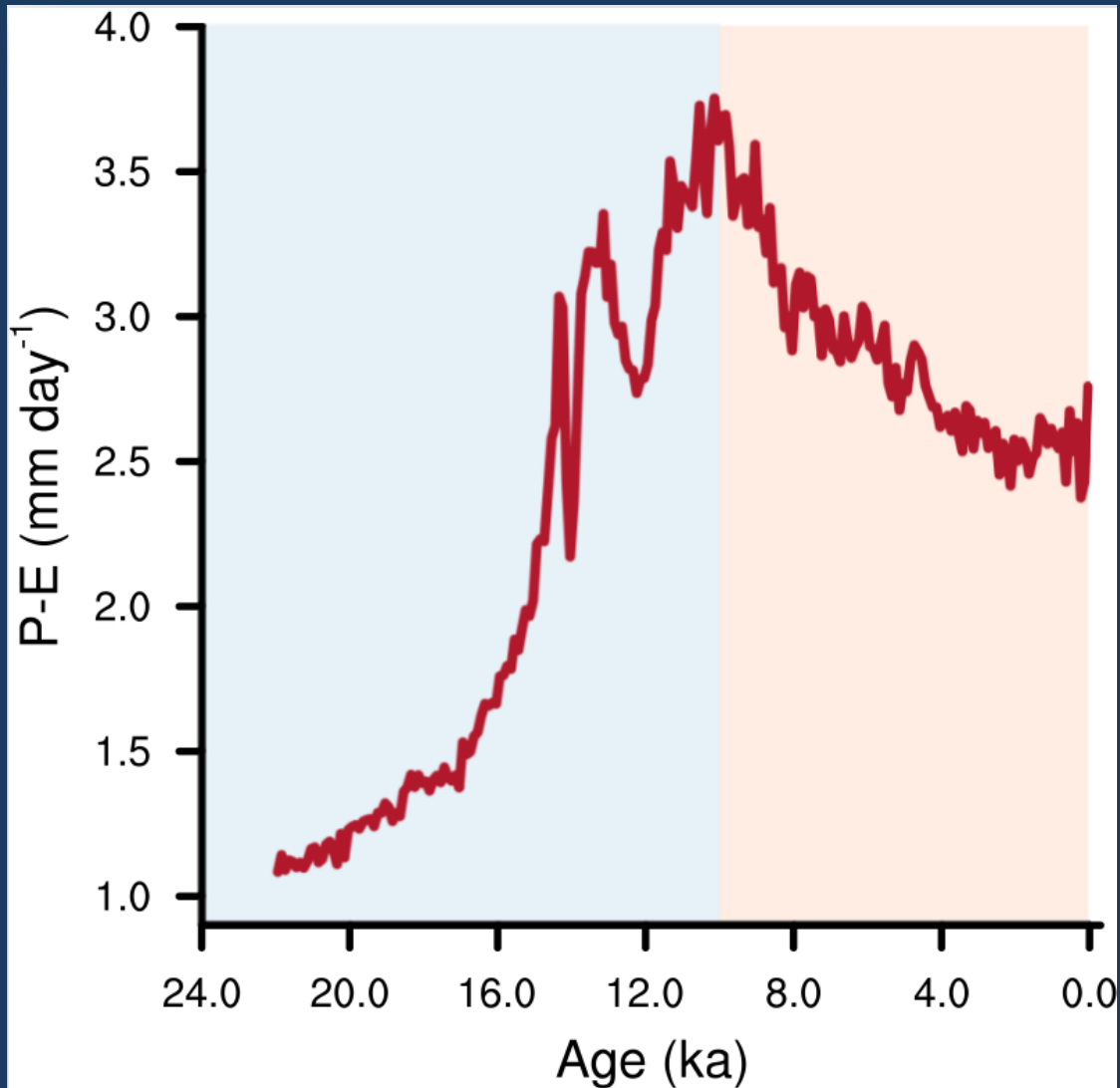


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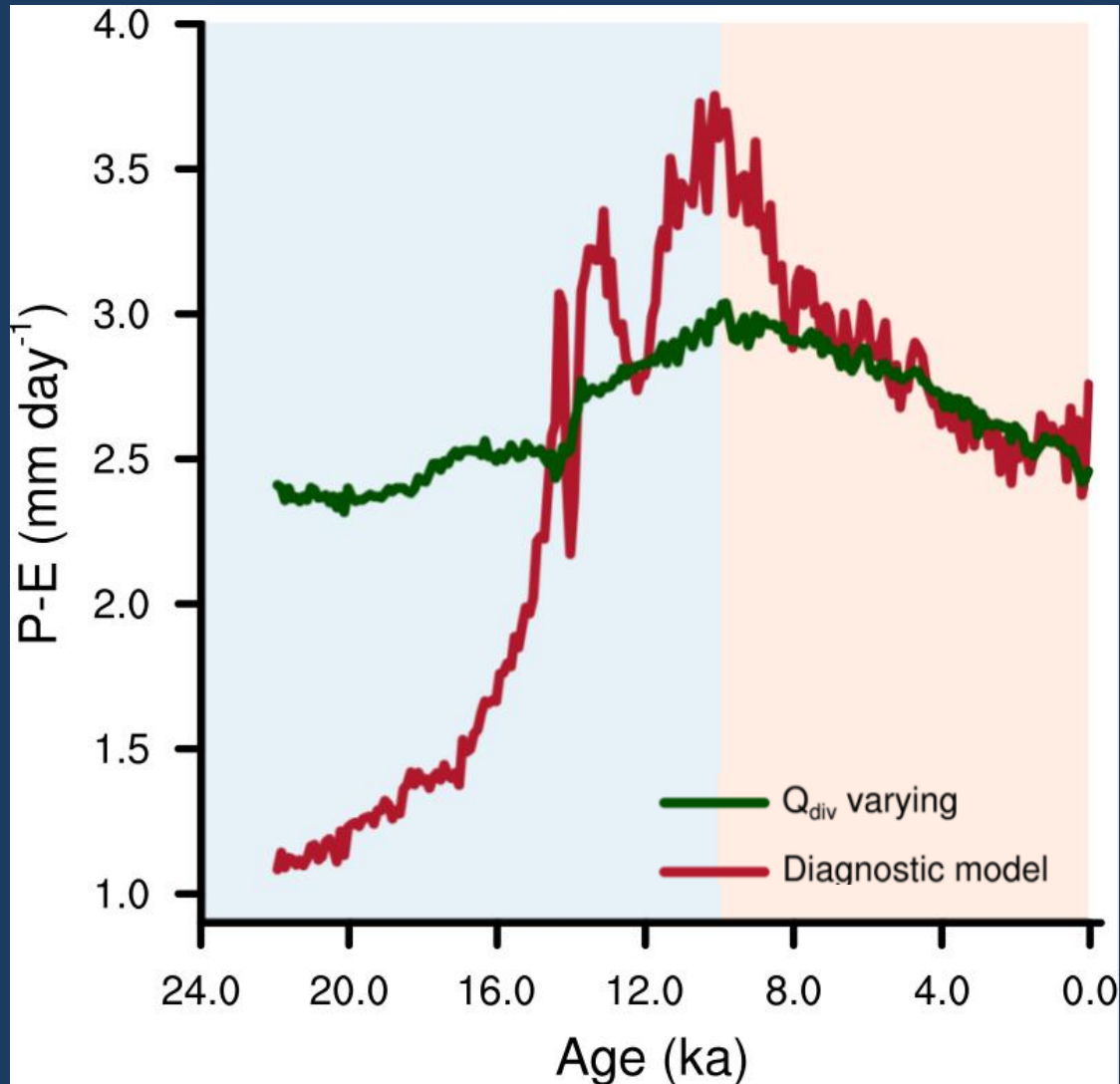
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$$P - E = \frac{Q_{div}}{80.4 / CWV - 1.3}$$

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# $Q_{DIV}$ DRIVES (P-E) DURING HOLOCENE

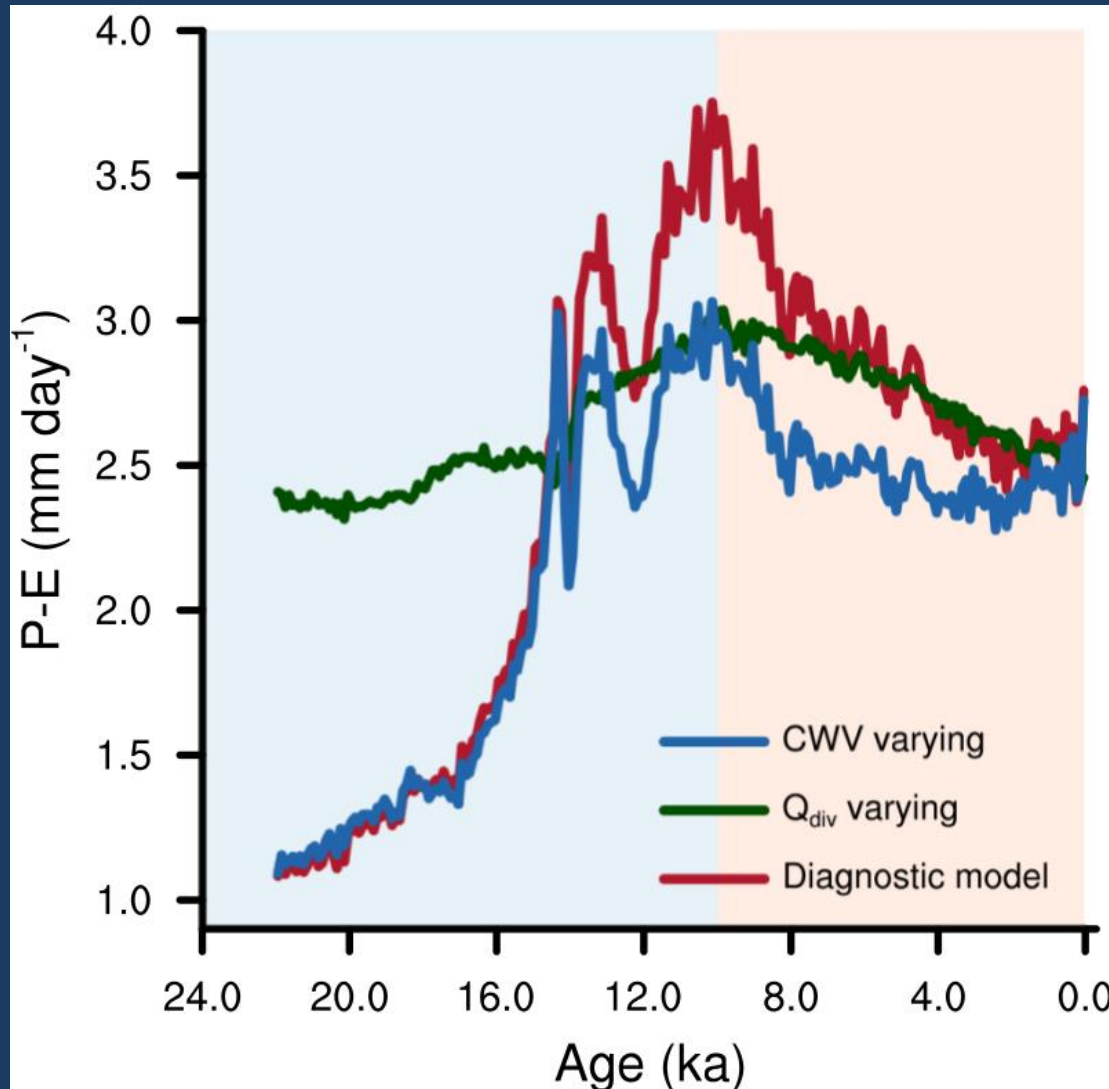


$$P - E = \frac{Q_{div}}{80.4 / \text{CWV} - 1.3}$$

**CWV fixed at pre-industrial values.**

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# WATER VAPOR PLAYS CRUCIAL ROLE DURING DEGLACIAL



$$P - E = \frac{Q_{div}}{80.4 / CWV - 1.3}$$

**Q<sub>div</sub> fixed at  
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*Jaliha et al. (2019) Nat. Comm.*

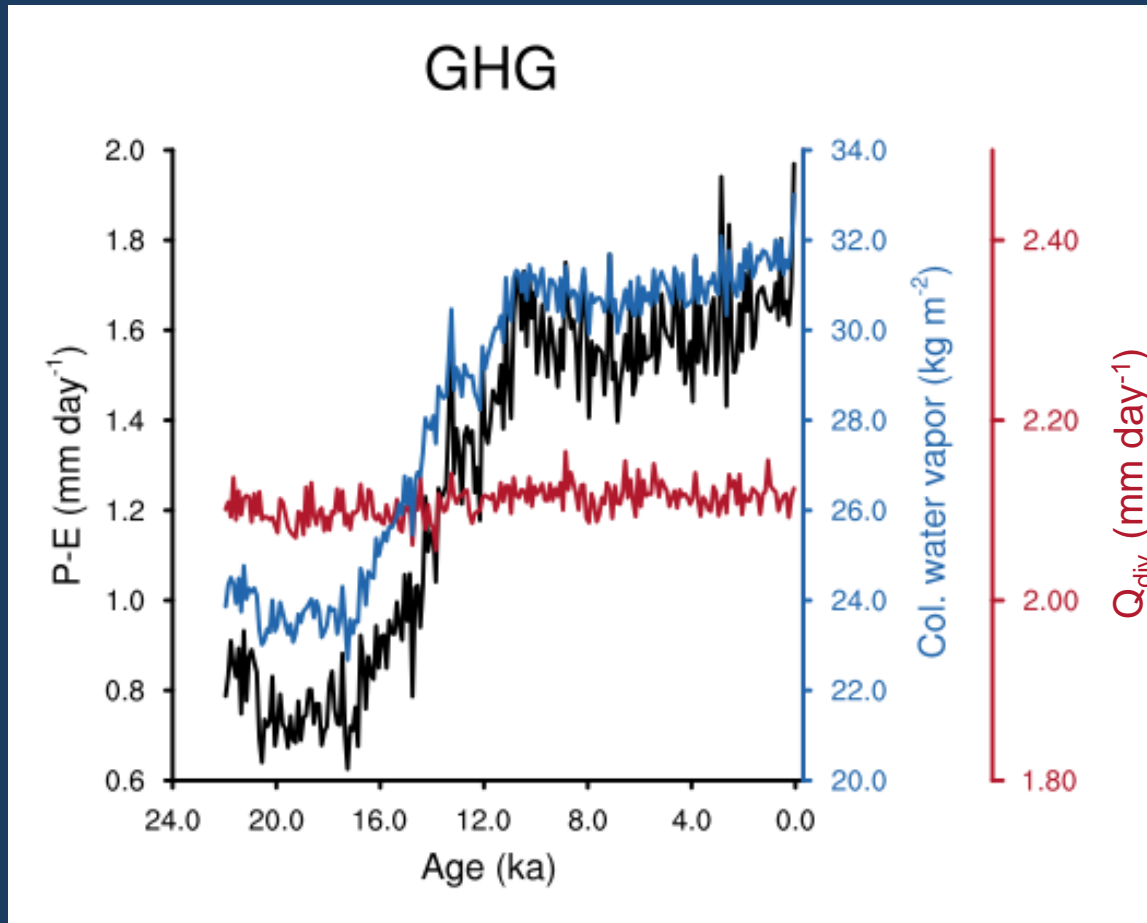
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# Role of different forcings

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# GREENHOUSE GASES AFFECT GMS ALONE

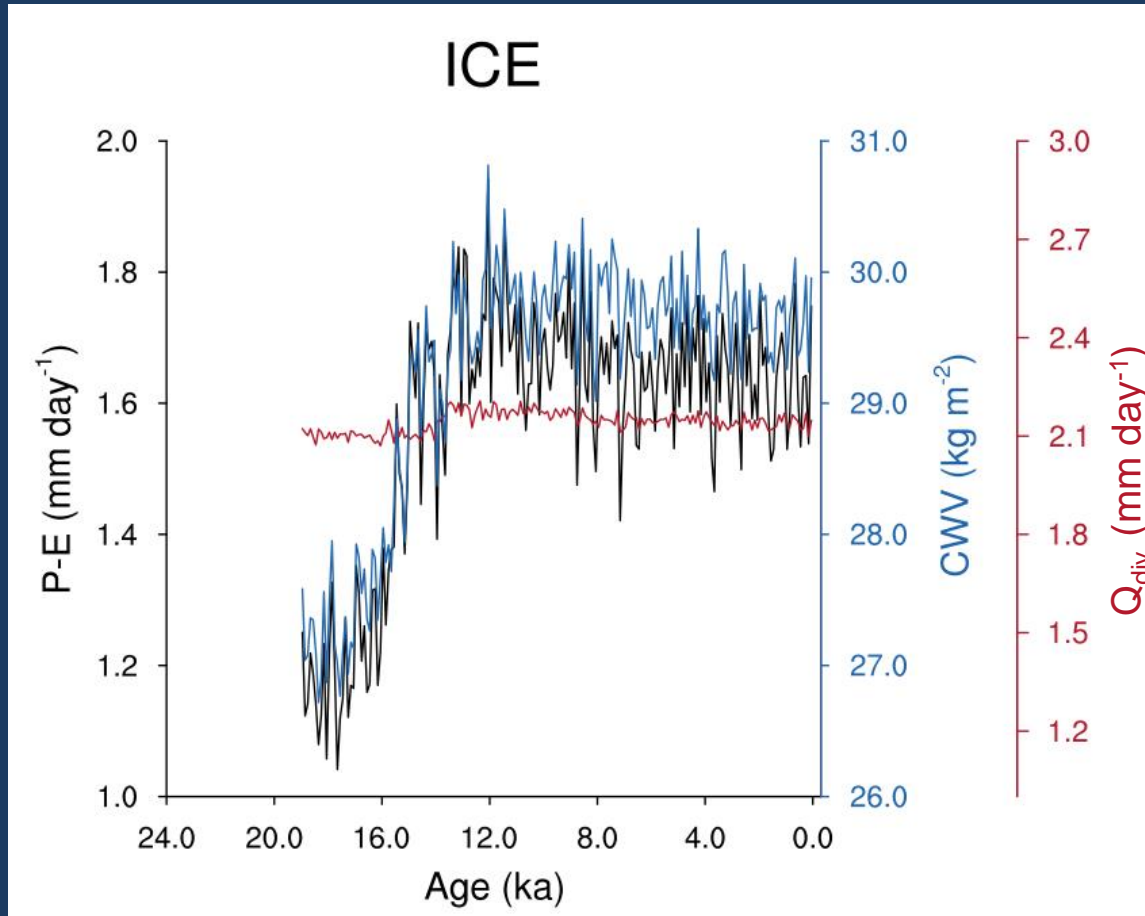


$$P - E = \frac{Q_{div}}{80.4 / CWV - 1.3}$$

**GHG does not influence  $Q_{div}$**

*Jalihal et al. (2019) Nat. Comm.*

# ICE SHEETS AFFECT GMS ALONE

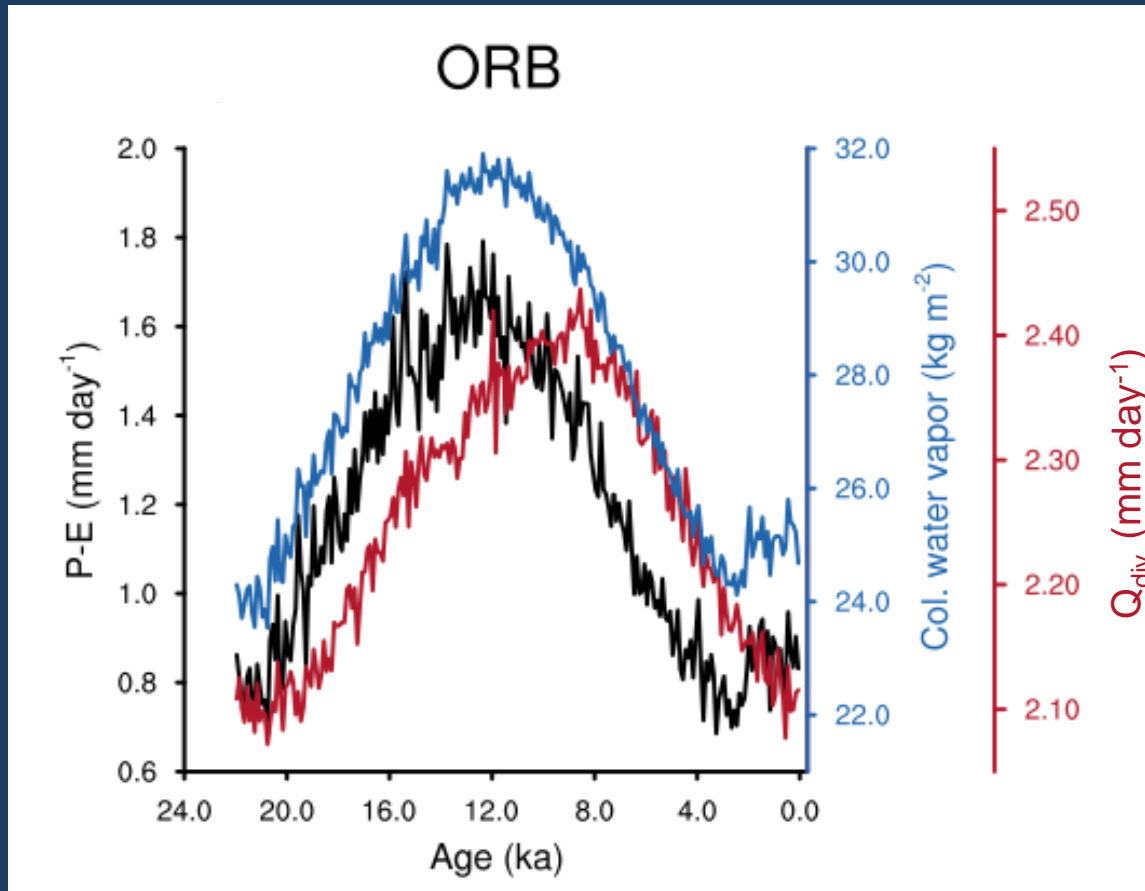


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Jaliha et al. (2019) Nat. Comm.

# ORBITAL FORCING AFFECTS BOTH GMS AND $Q_{DIV}$

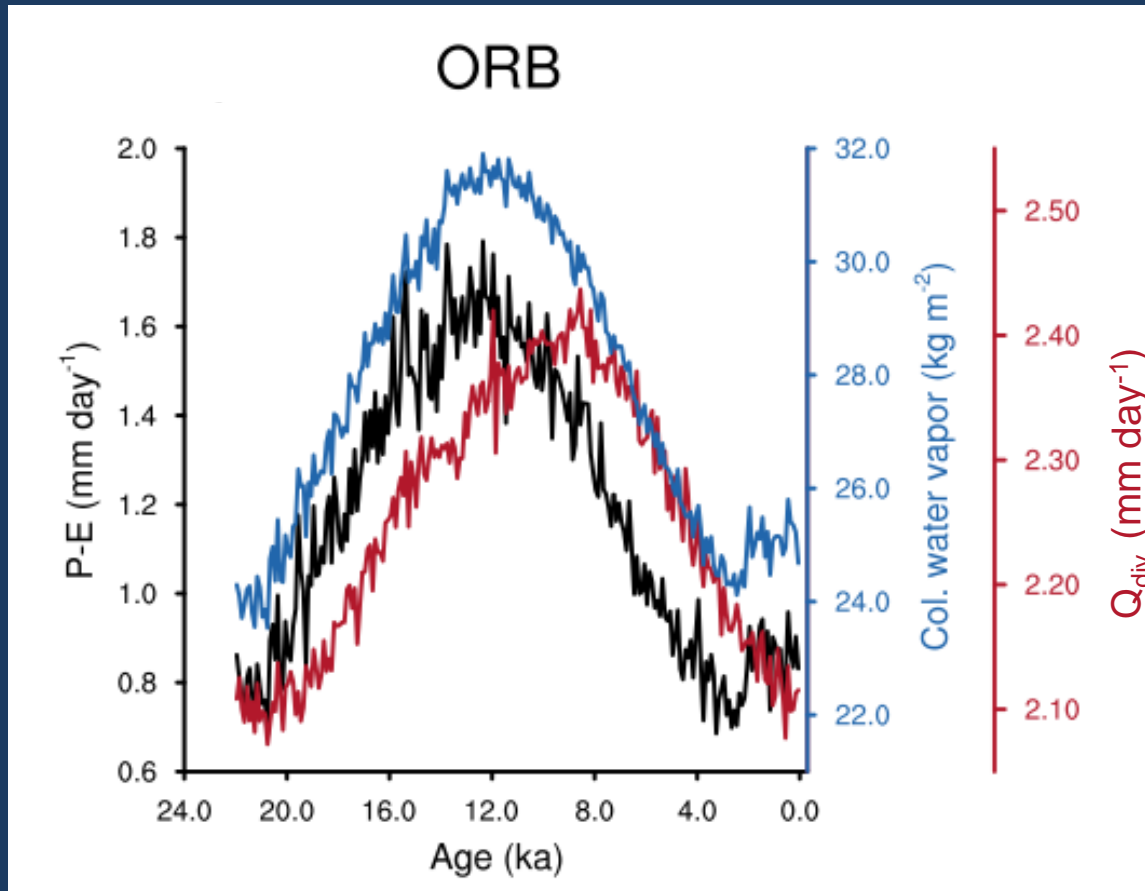


$$P - E = \frac{Q_{div}}{80.4 / CWV - 1.3}$$

Even though both  $Q_{div}$  and GMS are influenced, GMS is dominant.

Jalihal et al. (2019) Nat. Comm.

# ORBITAL FORCING WITH GLACIAL BOUNDARY CONDITIONS

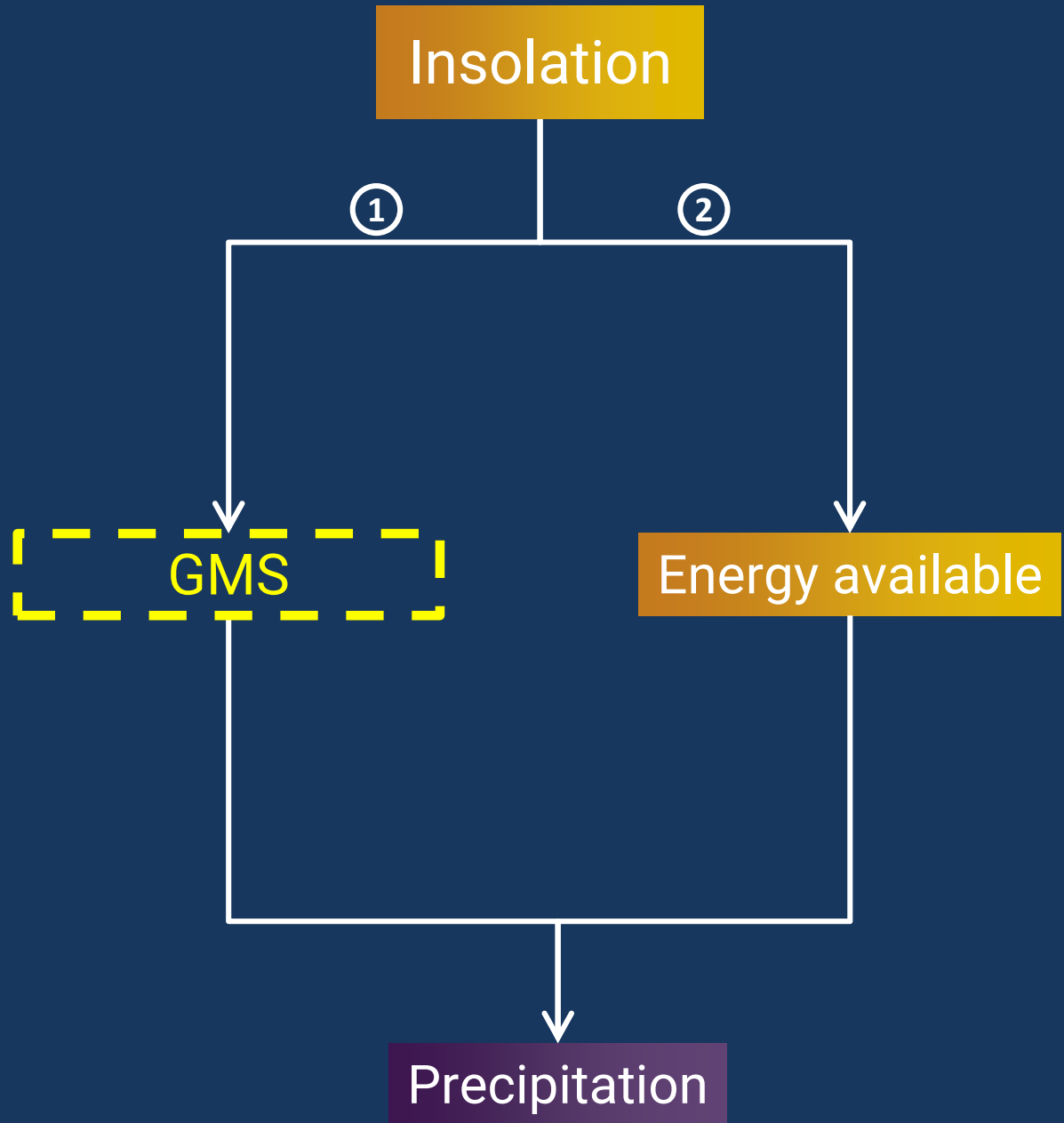


$$P - E = \frac{Q_{div}}{80.4 / CWV - 1.3}$$

**ORB:**  
orbital parameters  
only. Constant LGM  
values for all other  
boundary conditions.

**During cold periods,  
orbital forcing drives  
monsoon through  
GMS.**

*Jaliha et al. (2019) Nat. Comm.*



① - Dominant during the Deglacial

② - Dominant during the Holocene

GHG & Ice sheets

Insolation

①

②

GMS

Energy available

Precipitation

① - Dominant during  
the Deglacial

② - Dominant during  
the Holocene

GHG & Ice sheets

Insolation

①

②

Surface  
temperature

Total column  
water vapor

Energy available

Precipitation

① - Dominant during  
the Deglacial

② - Dominant during  
the Holocene

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# Conclusions

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# CONCLUSIONS

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- Changes in insolation is an initial trigger, final response involves feedbacks as well.
- 
- Water vapor feedback amplifies solar forcing during deglacial.
- 
- Cloud radiative feedback influences monsoon during Holocene.

# CONCLUSIONS

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- Insolation can drive monsoon through two different pathways.
- 
- Through energy available during warm periods, and through GMS during cold periods
- 
- Greenhouse gases and ice sheets only affect GMS.

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Thank you

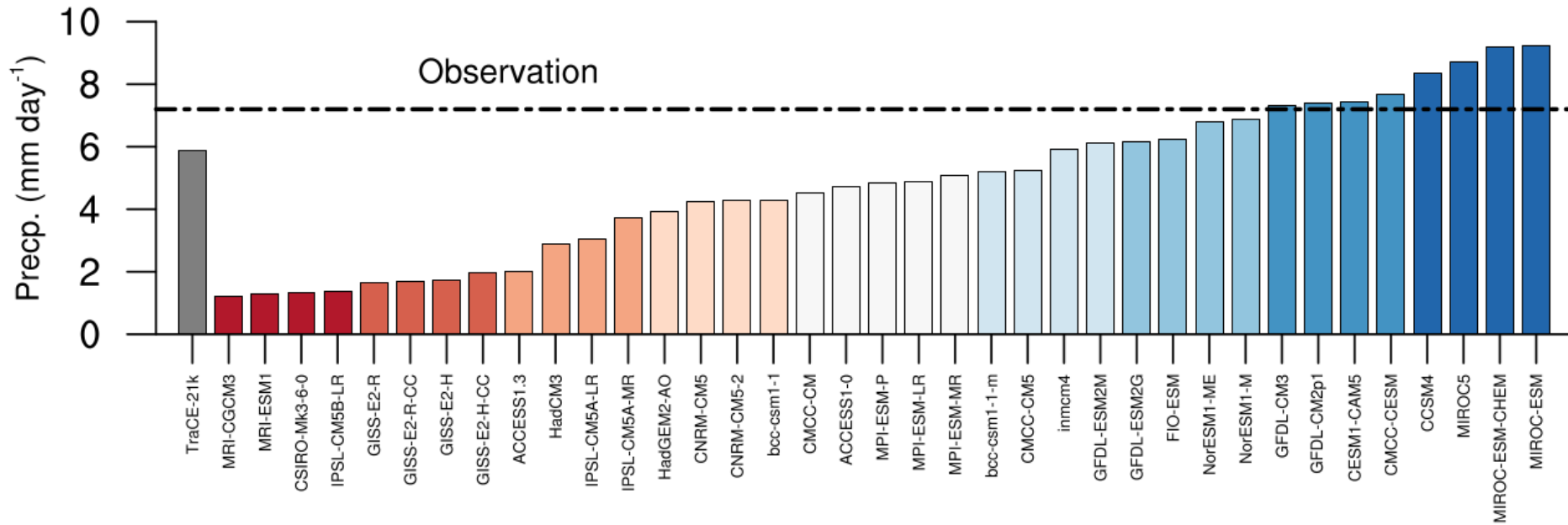
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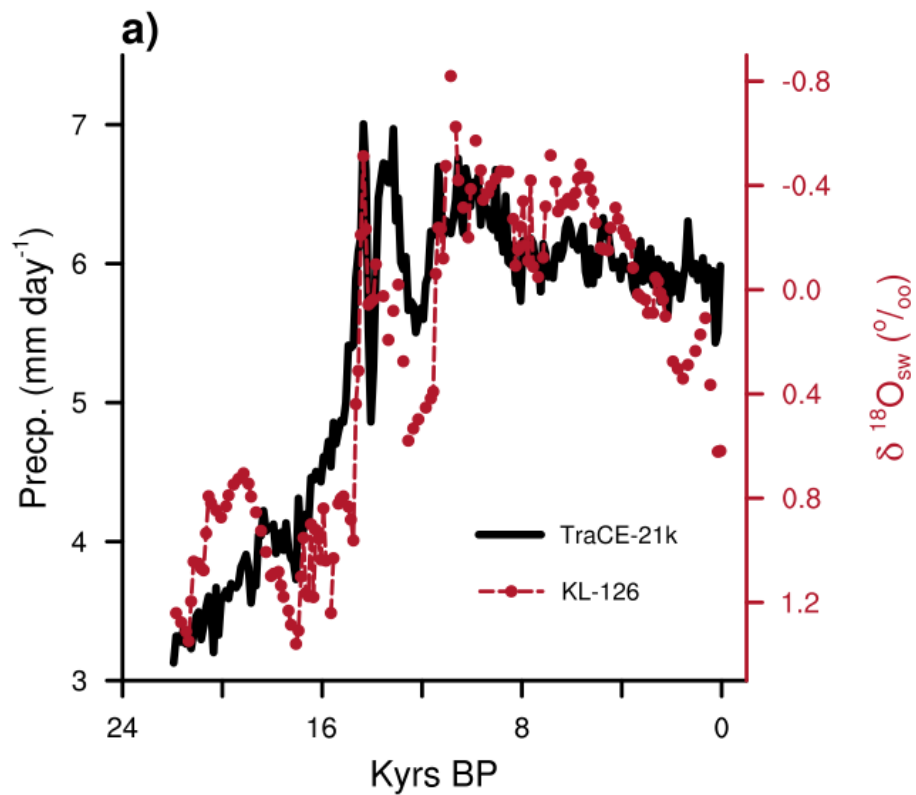
# Extra Slides

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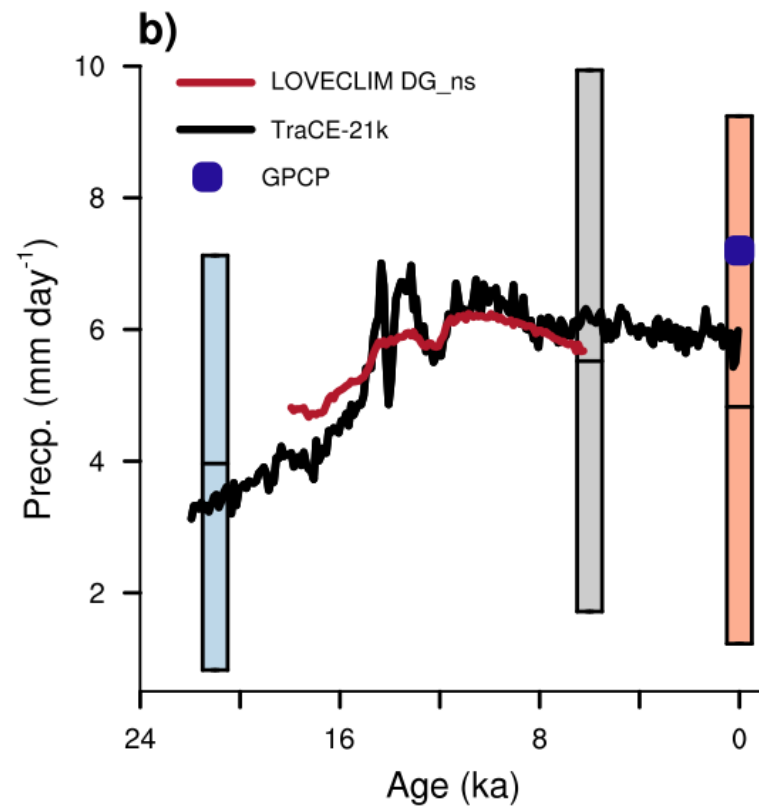
## Historical

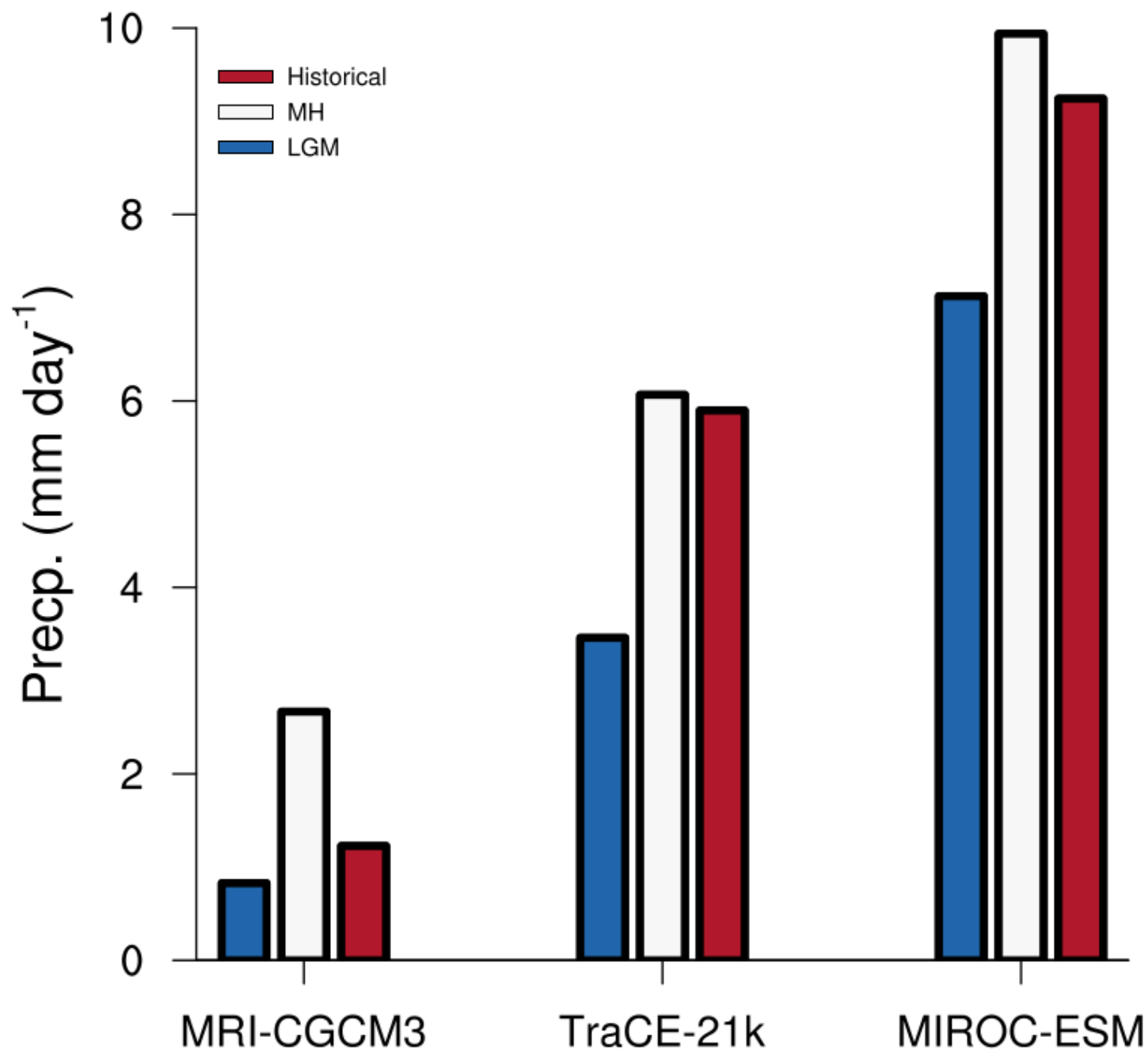


TraCE-21k vs Proxy

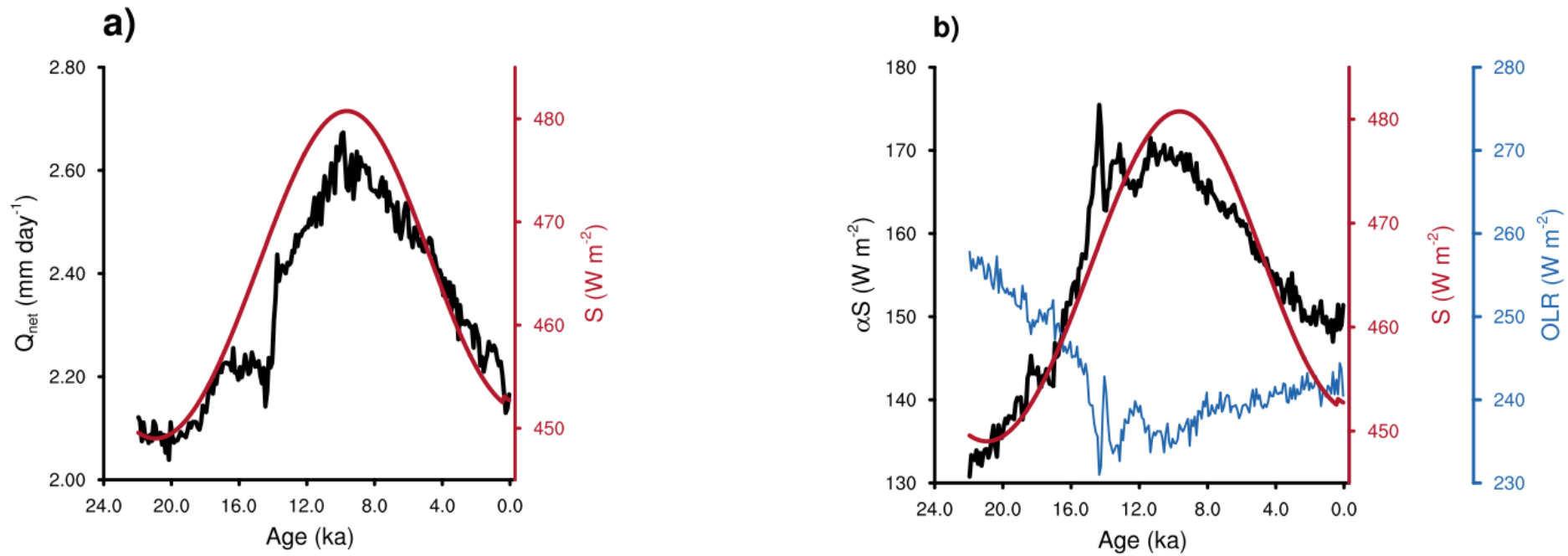


TraCE-21k vs PMIP3



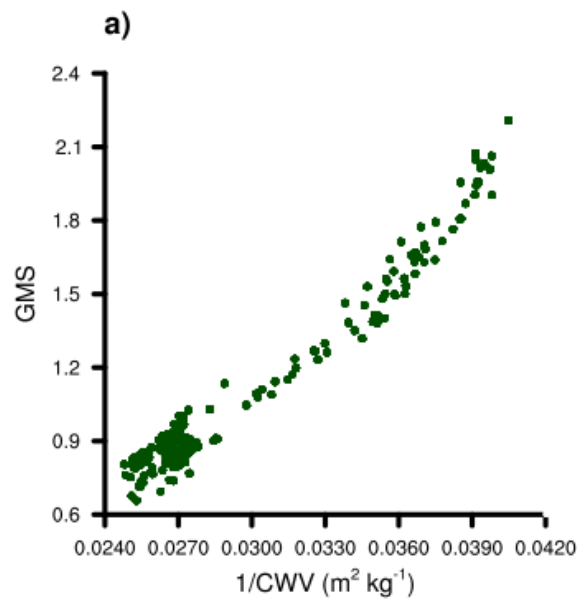


## TraCE-21k (JJA)

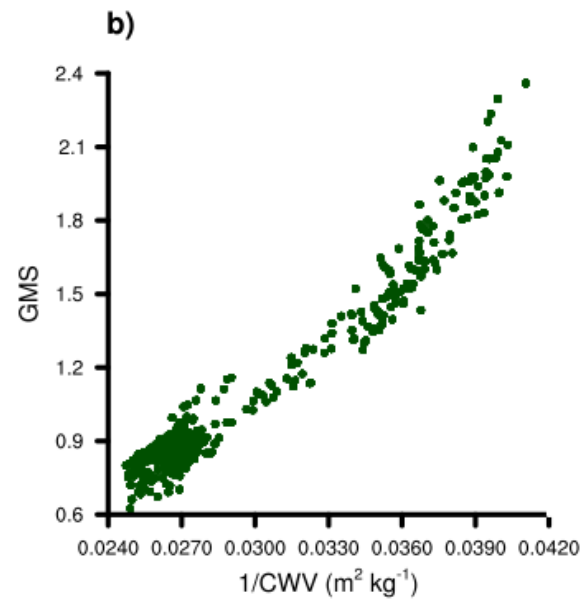




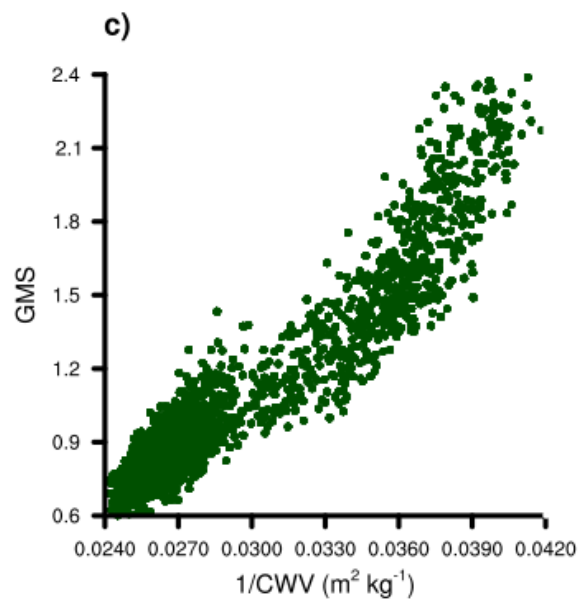
100-yr avg



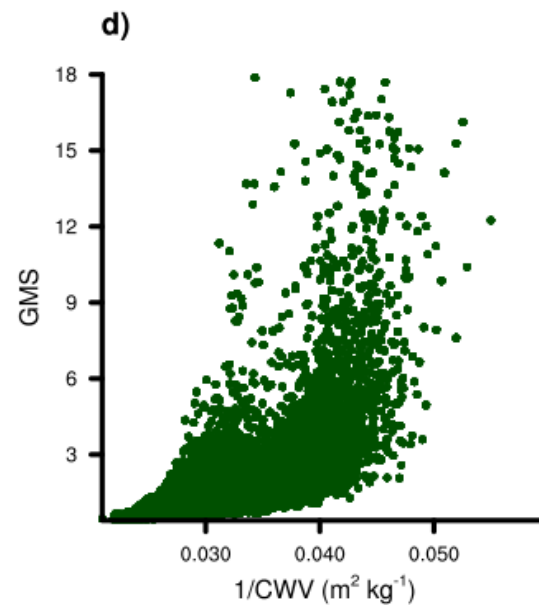
50-yr avg



10-yr avg



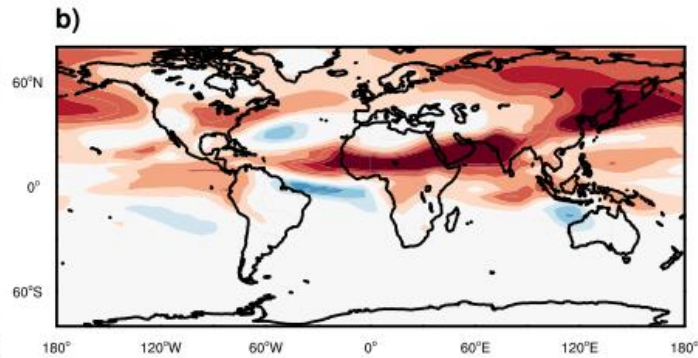
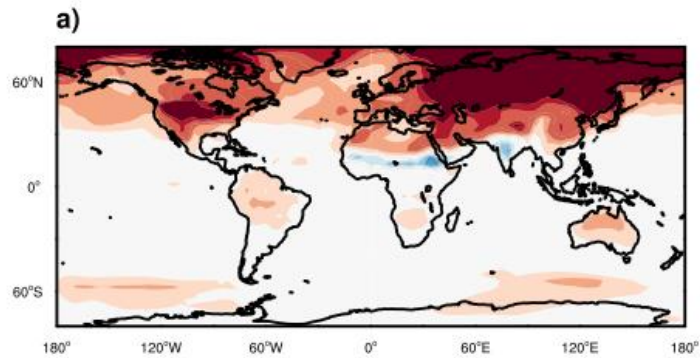
Inter-annual



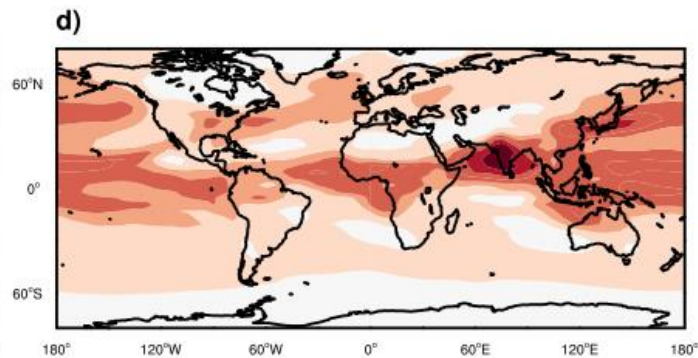
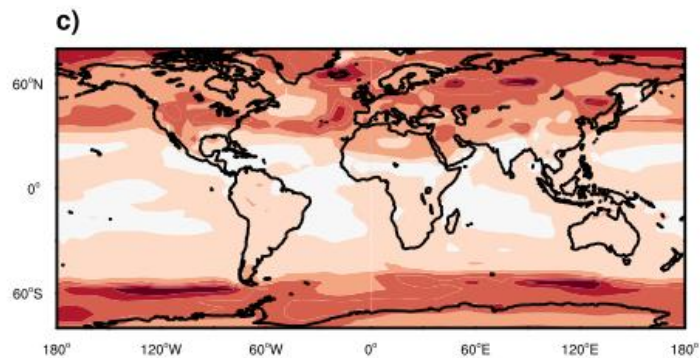
ORB (JJA; 12 ka - 22 ka)

Sfc. temp (K)

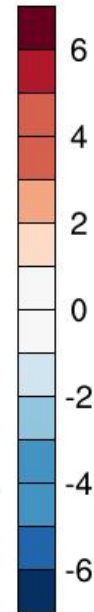
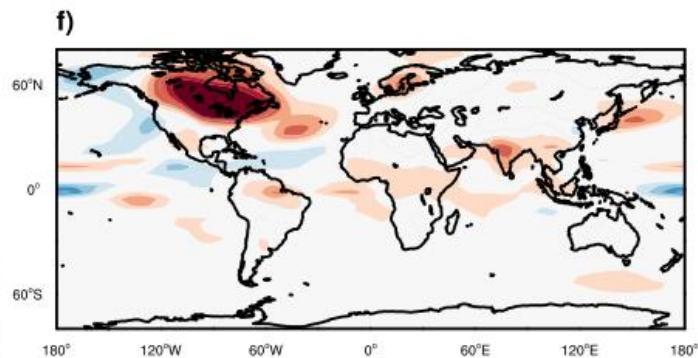
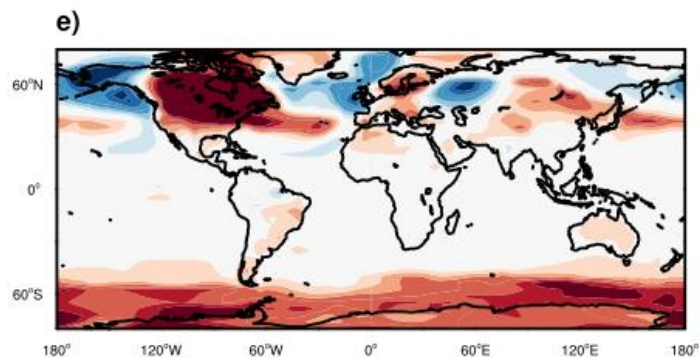
Col. Water Vapor (kg m<sup>-2</sup>)

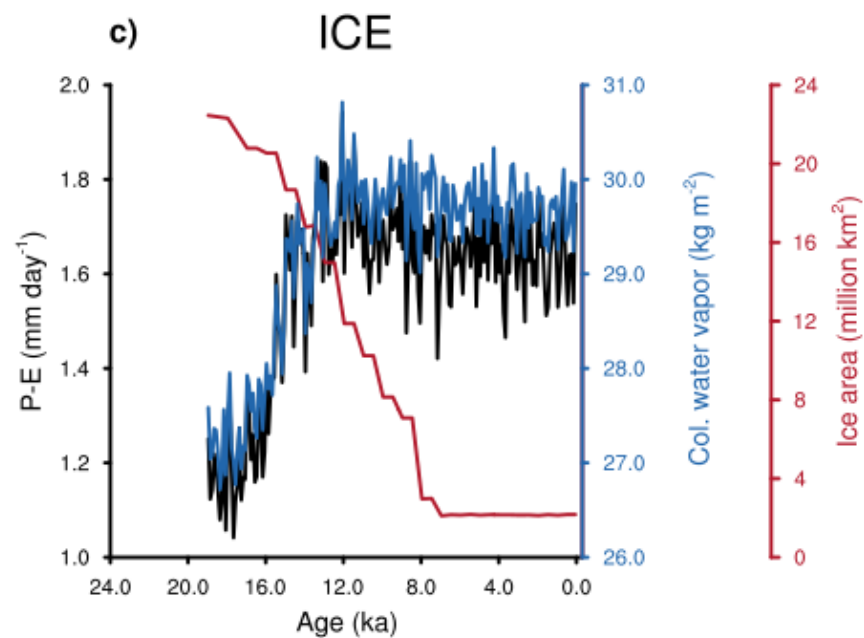
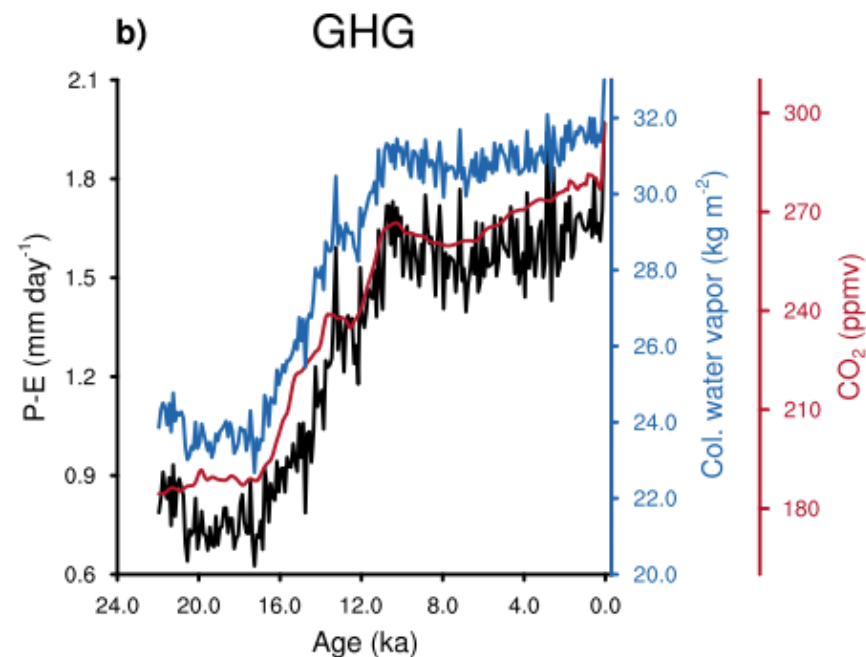
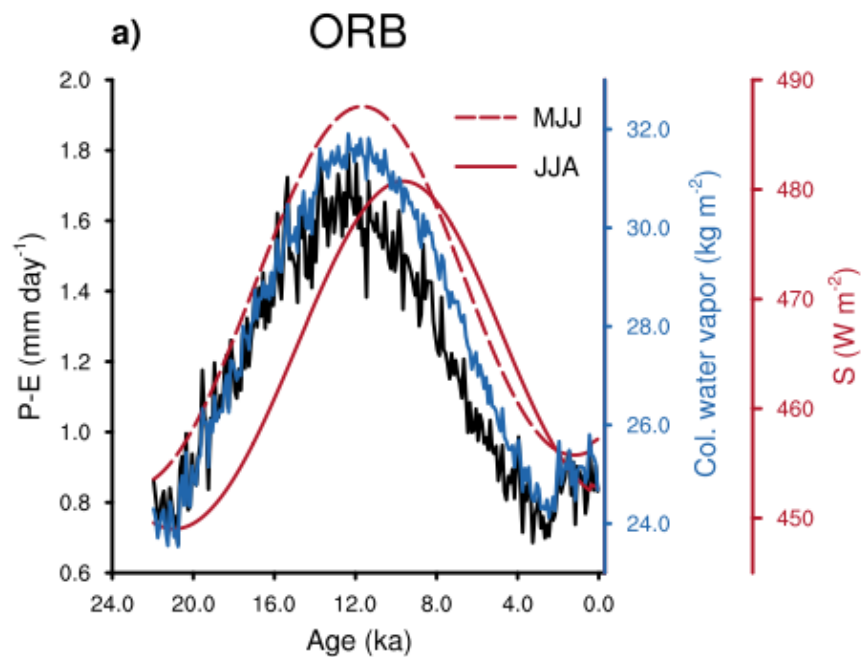


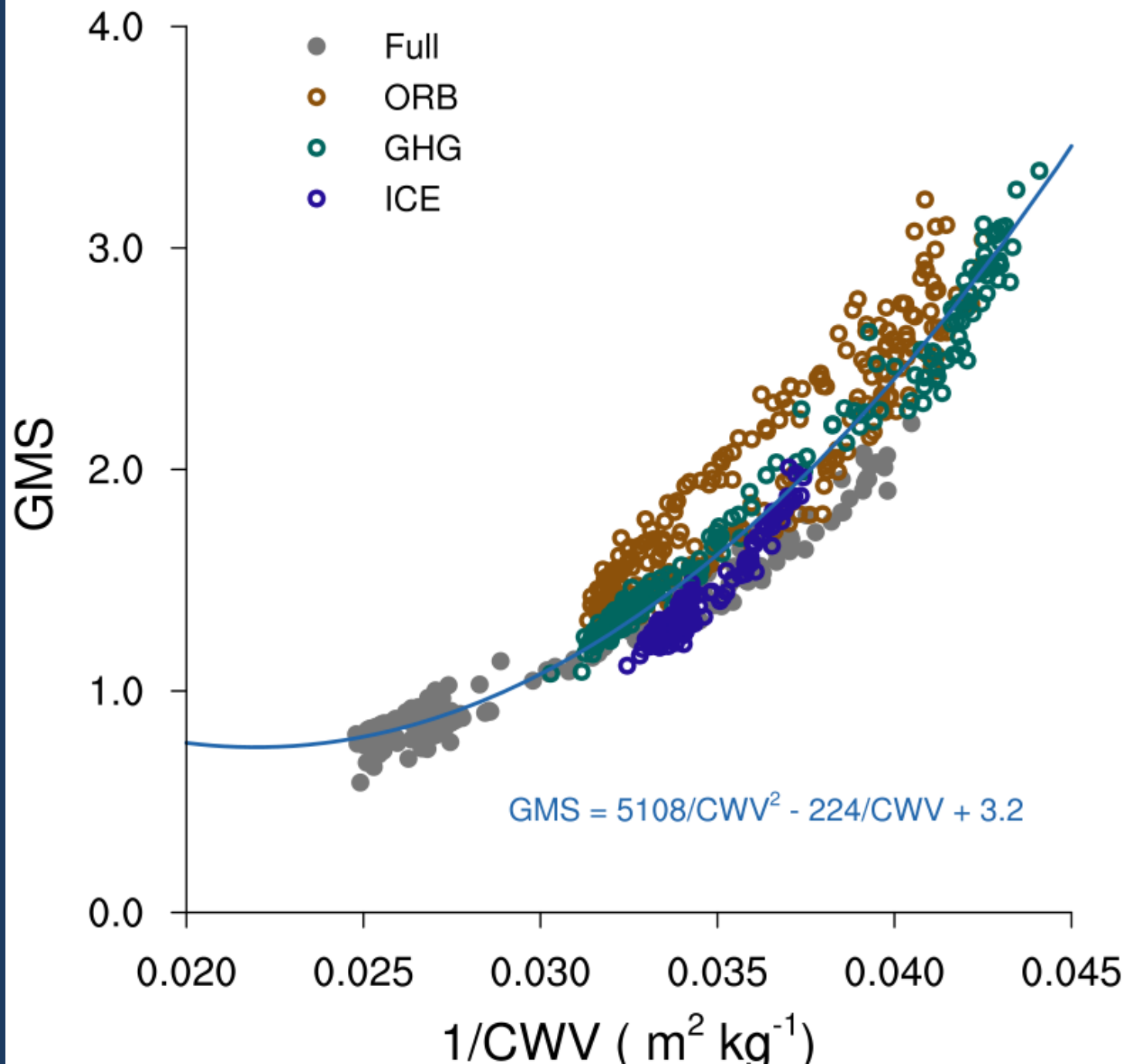
GHG (JJA; 12 ka - 22 ka)



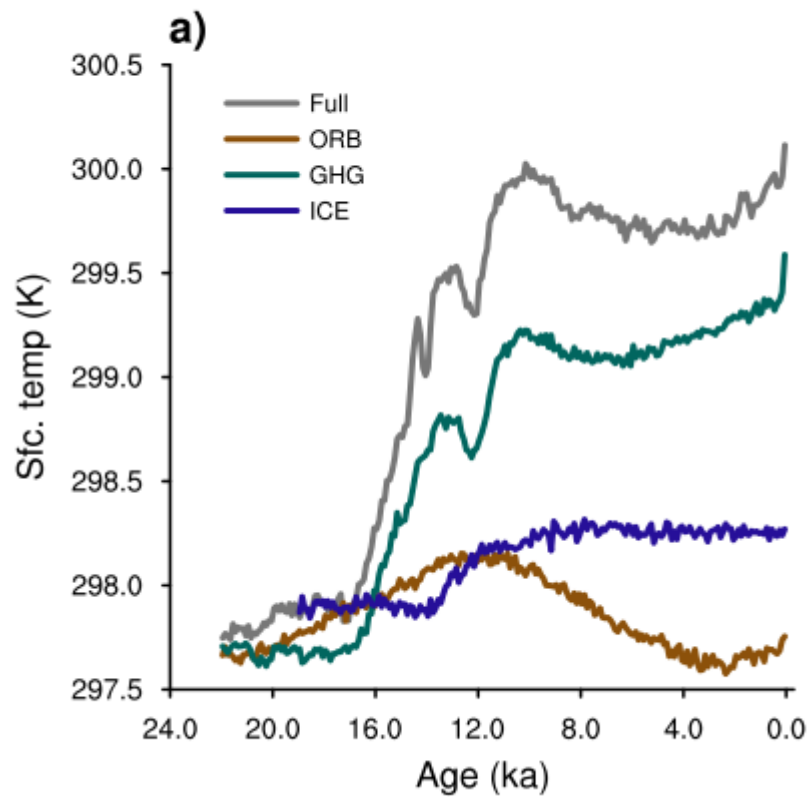
ICE (JJA; 8 ka - 19 ka)



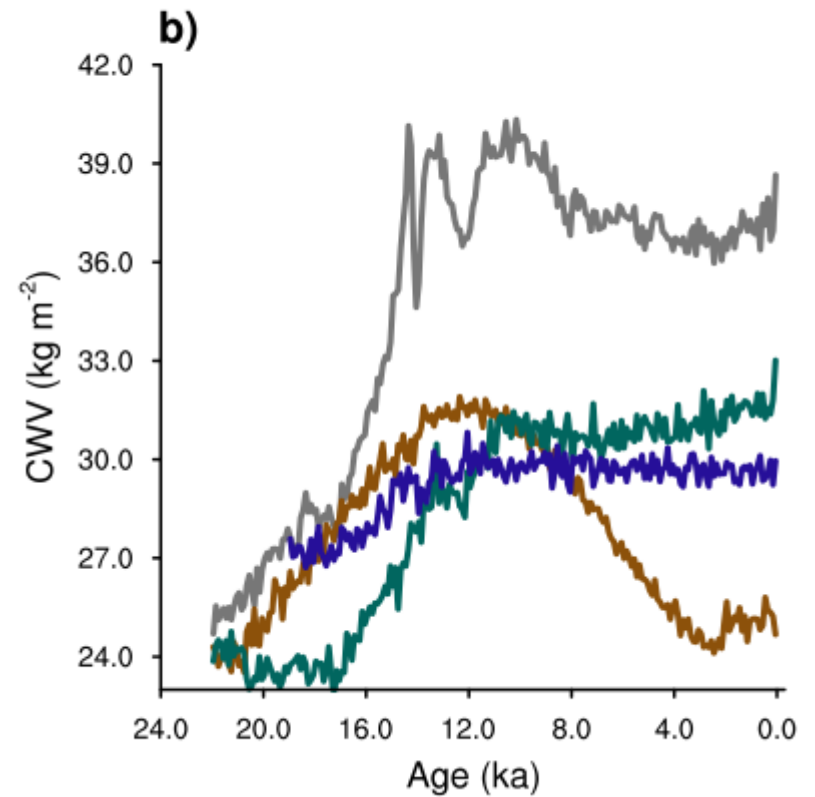




## Arabian Sea

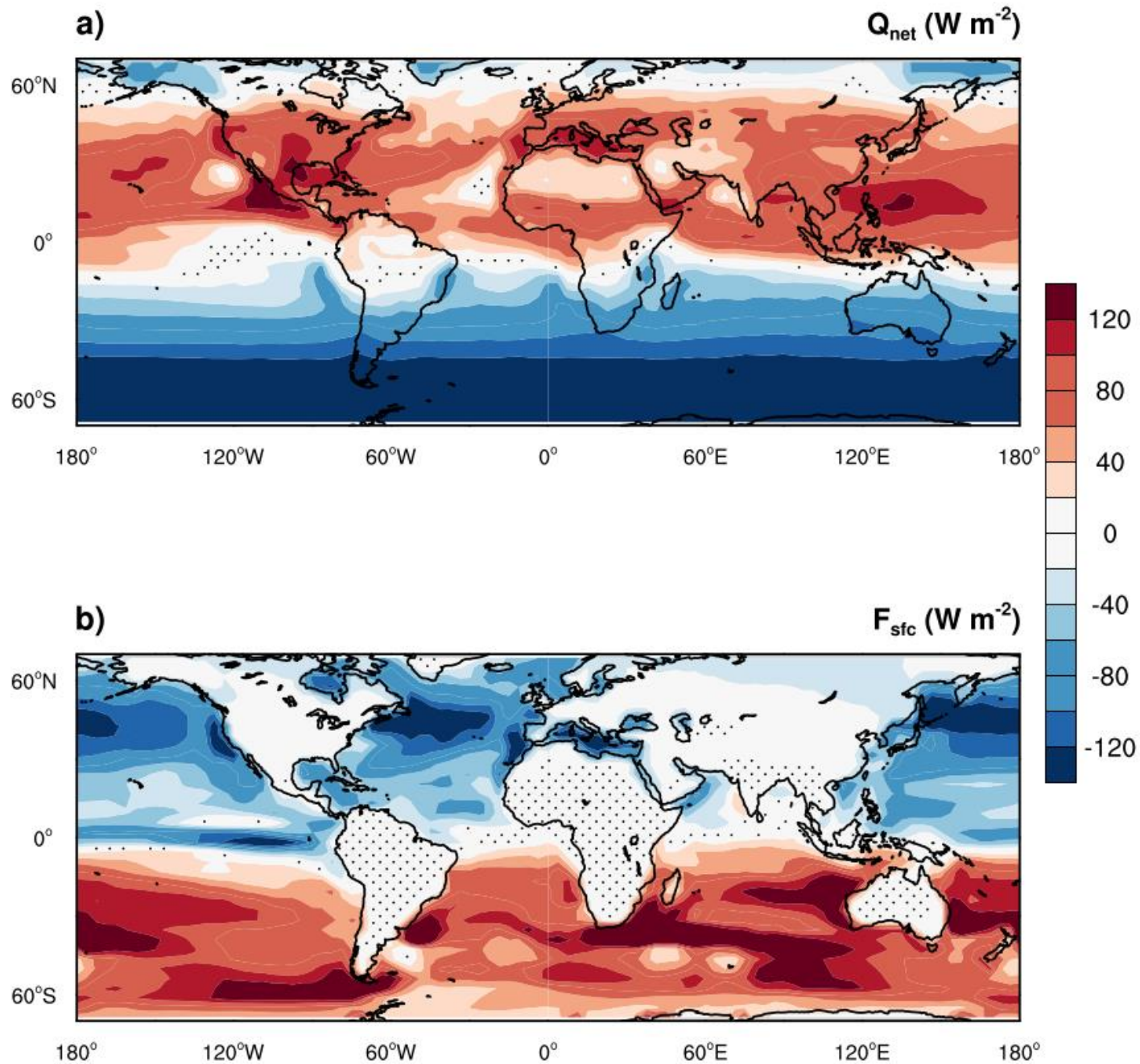


## India





# TraCE-21k (JJA; 1850 - 1950 AD)



# TraCE-21k; (JJA)

