CLIMATE SHIFT OF THE ATLANTIC MERIDIONAL OVERTURNING CIRCULATION (AMOC) IN REANALYSES (ORAS5): POSSIBLE CAUSES, AND SOURCES OF UNCERTAINTY Vincenzo de Toma¹, Chunxue Yang², Vincenzo Artale³

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Introduction & Motivation

The ECMWF Reanalysis ORAS5 (Zuo et al. 2018) has been used to study the long term variability of the AMOC: a comparison has been carried out both with the in-situ data of the RAPID program, which monitor the Volume Transport at 26.5°N from 2004, and the members of CMEMS_GLO_RAN project, which gather other reanalyses from 1993 onward. We observed a declining signal of the AMOC strength signal in the mid-1990s, of about 5Sv, which we tried to characterize from a dynamical point of view. The AMOC strength is defined for each latitude y, depth z and time t as:

$$\psi(y, x, t) = \max_{z} \int_{\eta}^{z} dz' \int_{x_{w}}^{x_{e}} dx \ v(x, y, z', t). \tag{1}$$

where η is the free surface of the ocean, v is the meridional velocity and x_w, x_e the Atlantic Boundaries for each latitude.





Fig. 1: The shift at the RAPID latitude: Comparison with CMEMS and in-situ data. All the data show a good agreement on the common period. We show also that the shift is present at other latitudes, in particular involving depths between 1000m and 3000m, and latitudes between $25^{\circ}N$ and $45^{\circ}N$. Searching for reasons of the shift in the Labrador Sea, we found a transition pattern in the same years, regarding temperature, salinity, potential density and also heat fluxes. The use of a dynamical metric (Liu et al. 2017) helped us to understand that this declining signal can be thought as a stability regime transition of the AMOC.

Fig. 2: The shift in the entire Stream Function, and variablity in the Labrador Sea

Analysis & Conclusions

The dynamical metrics Σ characterizing the AMOC-induced Freshwater transport follows Liu et al. 2017 definition:

$$\Sigma = M_{ov,S} - M_{ov,N} \quad \text{where} \quad M_{ov}(\phi) = -\frac{1}{S_0} \int_{-D}^0 \bar{v^*}(\phi, z) \langle S(\phi, z) \rangle dz.$$

We found that passing the transition there's an activation of the Eastern Atlantic expressed in a



redistribution of surface Kinetic Energy Extreme Events, which we have defined as

$$\mathcal{N} = \sum_{n=1}^{T} \Theta \left(|\mathcal{K}_n| - 3\sigma_{\mathcal{K}} \right) \text{ where } \Theta \left(x \right) = \begin{cases} 1 \text{ if } x > 0\\ 0 \text{ if } x < 0 \end{cases}$$

this redistribution can be interpreted as a balance mechanism for the loss of Gravitational Potential Energy suffered by the AMOC (Kuhlbrodt et al. 2007).

Conclusion: We have shown that this transition is coherent with physical processes descending from a decreased mass transport. Through dynamical metrics we found a different connectivity ($\Sigma > 0 \Rightarrow \Sigma < 0$) among the North and South Atlantic domain, linked to the West-East redistribution of activity in the two periods (\mathcal{N}). Mechanisms at the origins of this link need further investigations.

Fig. 3: Regime Stability Diagram and Kinetic Energy Extreme Events redistribution

Short List of References:

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