

EGU – Session GD6.1

*“The Earth-System-Science-based EGU-session [...] aims to test the hypothesis: That material flow in the **earth’s mantle** affects the **earth’s lithosphere**, including its **surface topography** and the **sedimentary record**, far beyond **plate boundaries**”*

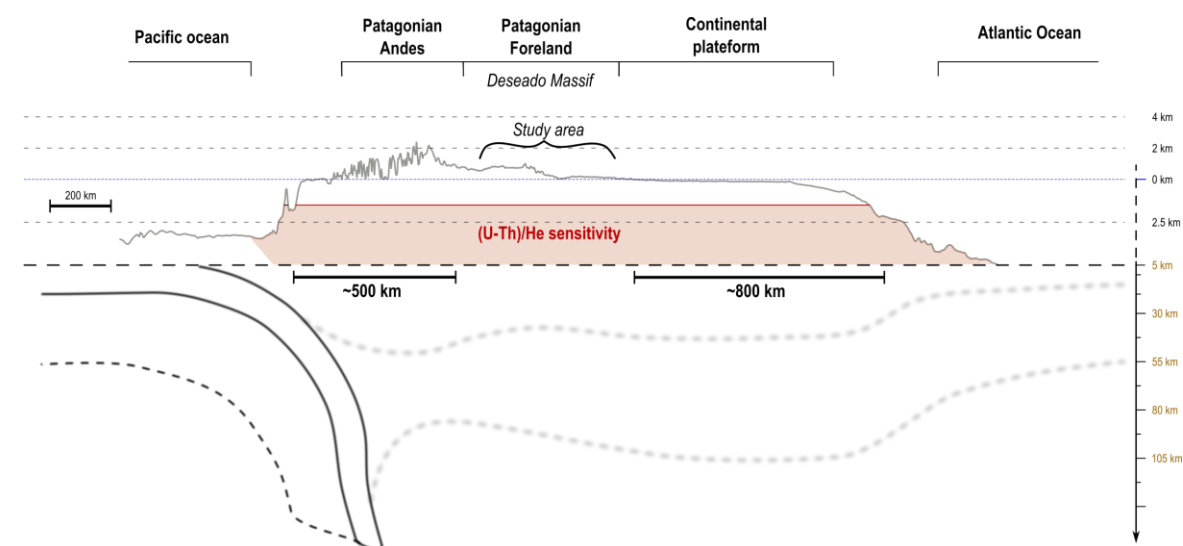
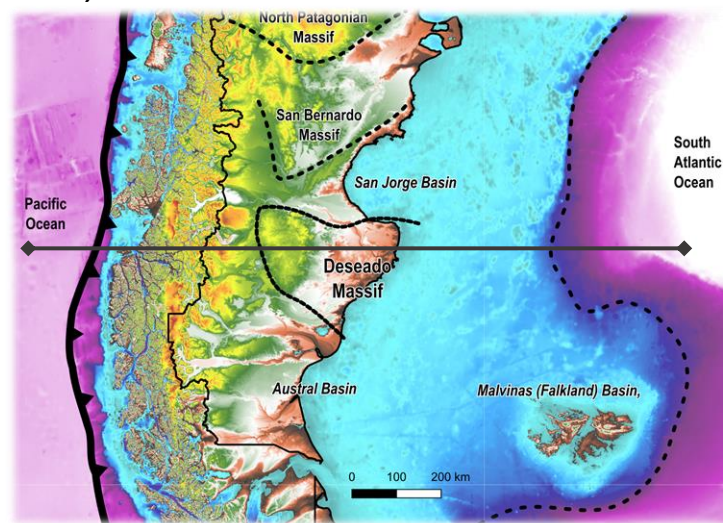
This study is a part of a project that aims to constrain the formation and evolution of the south Patagonian foreland, using a multi-methods approach. One part of this project applied low temperature thermochronology on the intraplate Deseado Massif to constrain its surface evolution,

Here are presented a preview of our results, revealing a heating-cooling phase during the Mesozoic, which we attributed to a burial/exhumation event. The timing of this event is put in perspective with the regional geodynamics evolution, in order to constrain the mechanism that could produce it.

Patagonian foreland burial and exhumation during Mesozoic revealed by low temperature thermochronology: a response to mantle processes?

A. Derycke ^(1,2), C. Gautheron ⁽²⁾, M.C. Genge ^(2,4), M. Zattin ⁽²⁾, S. Mazzoli ⁽³⁾, C. Witt ⁽⁴⁾, M. Marquez ⁽⁵⁾.

Study area:



Background

Deseado Massif:

- Topographic high (400 to 1000 m)
- > **400 km** from plate boundary
- Consider stable since the Middle Jurassic

Main geodynamic events:

- **Andean subduction** since ~250 Ma, with a **flat slab** morphology **between ~220 Ma and ~200 Ma**
- Gondwana breakup starting at ~180 Ma with the **Karoo-Ferrar Plume** and finish at 124 Ma by the oceanisation
- **Chon Aike LIP event** between ~190 to ~150 Ma

The Deseado Massif is the southern massif of the Patagonian Foreland and it is mainly covered by a massif volcanic deposit associated with the Chon Aike LIP.

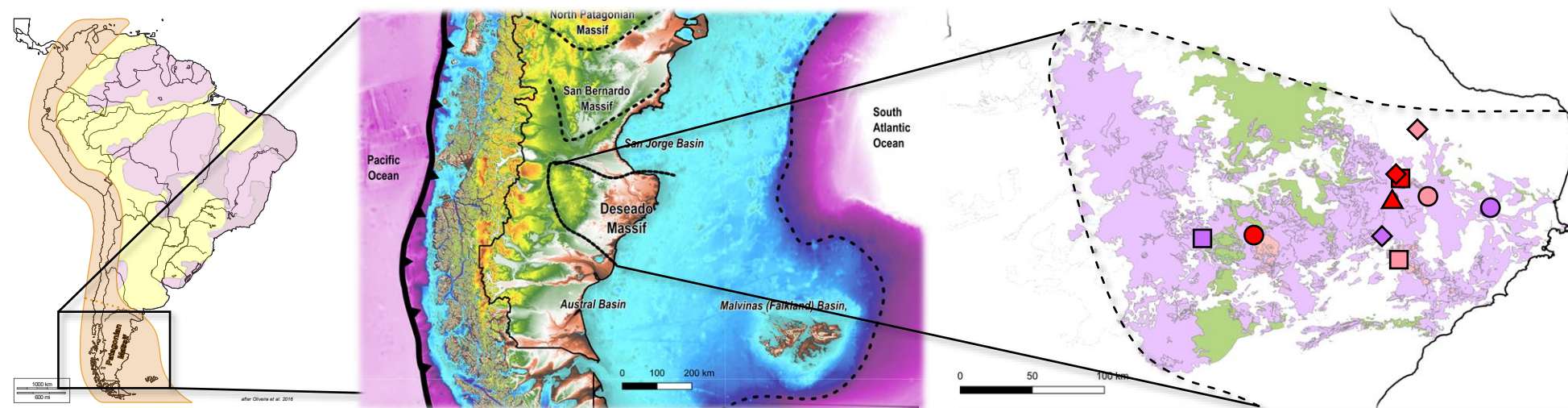
Local geology:

Cretaceous cover: restrain continental deposit (~100 Ma)

Chon Aike deposit: major volcanic Jurassic cover (~180 to ~150 Ma)

Basement: Paleozoic to Early Mesozoic rock, in surface at ~190 Ma

Deseado Massif position and morphology:



Geological background:

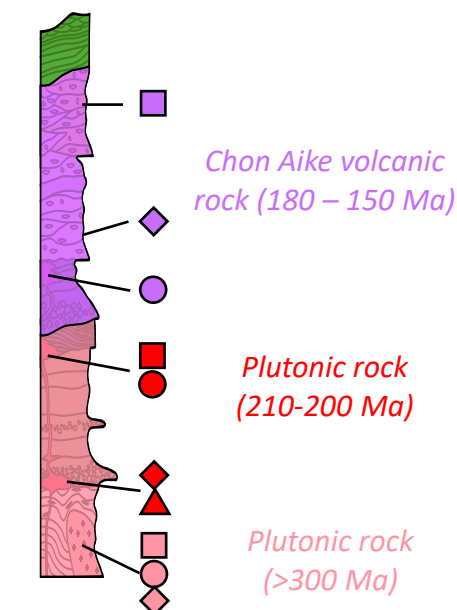
- Patagonia is divided in the **Andean range** itself, and the **Foreland**
- The Foreland is constituted by multiple **basins and massifs**, formed during the **Meso-Cenozoic periods**
- The lack of sedimentary deposits on those massifs conducted to suppose there were “**stable topographic high**” since the **Mesozoic**.

Navarrete et al. (2019) demonstrated that, between ~220 and ~150 Ma, a **complex subduction setting** occurred under the **Deseado Massif**, highlighting its potential **impact and control** the Deseado Massif surface evolution.

Objective of this study:

Constrain the **Deseado Massif thermal history** in order to **reconstruct its surface evolution** and questioned its linked to the complex mantle/subduction setting.

Simplified log:



Geodynamics statement & mantle setting

Geodynamics events:

The pre-Atlantic opening history of Patagonia corresponds to a **flat slab** emplacement, follow by the arrival of the **Karoo-Ferrar mantle plume** and **Chon Aike SLIP**.

This complex evolution **ended with the Atlantic opening**.

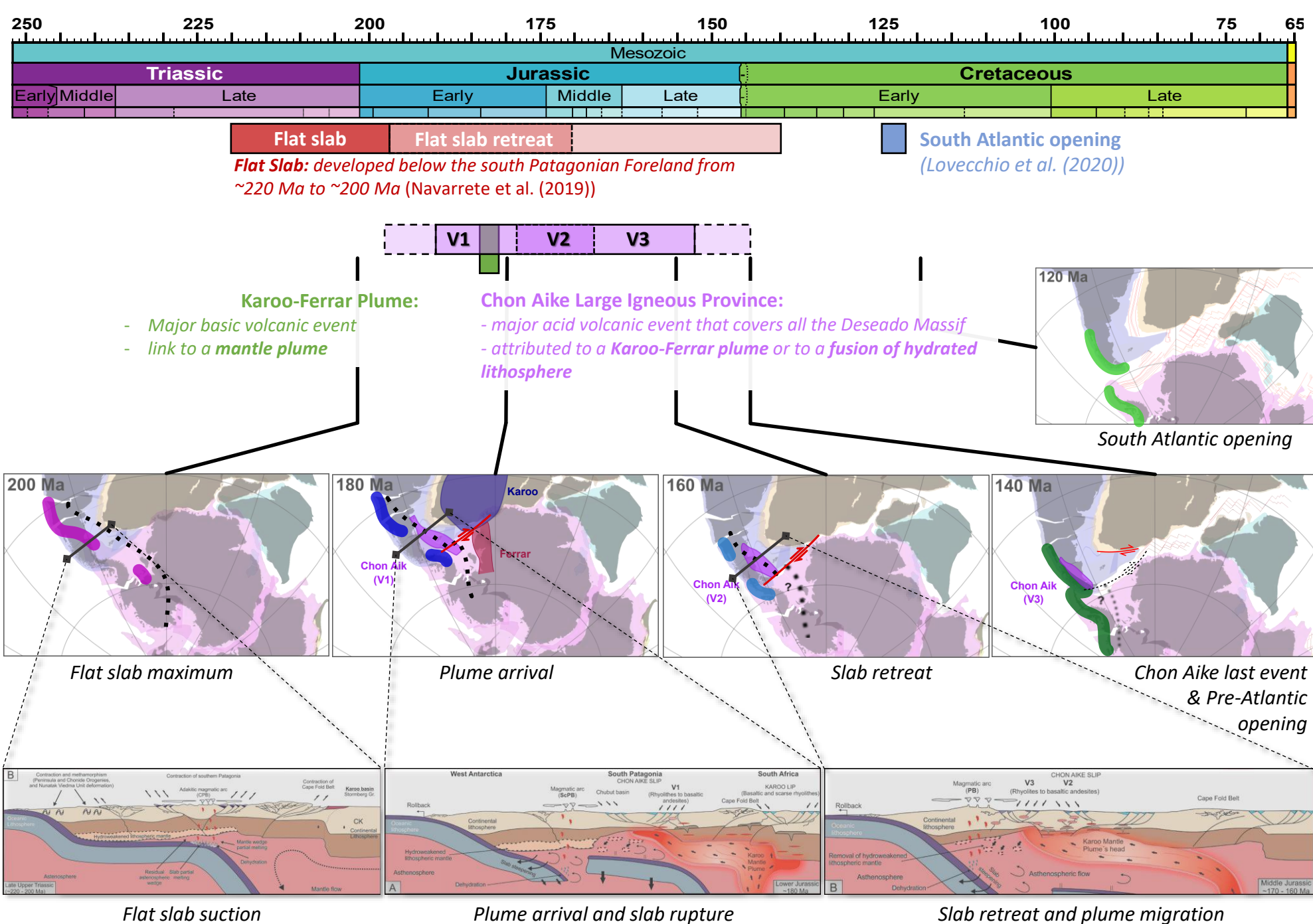
Paleogeographic evolution: (base on Müller et al. (2016))

Legend:

- position of subduction related magmatism
- supposed position of flat slab
- position of Large magmatic events

Est – West cross section evolution:

Theoretical section from **Navarrete et al. (2019)**, supposing a westward migration of the plume, controlled by the flat slab morphology, and that triggered the Chon Aike event evolution



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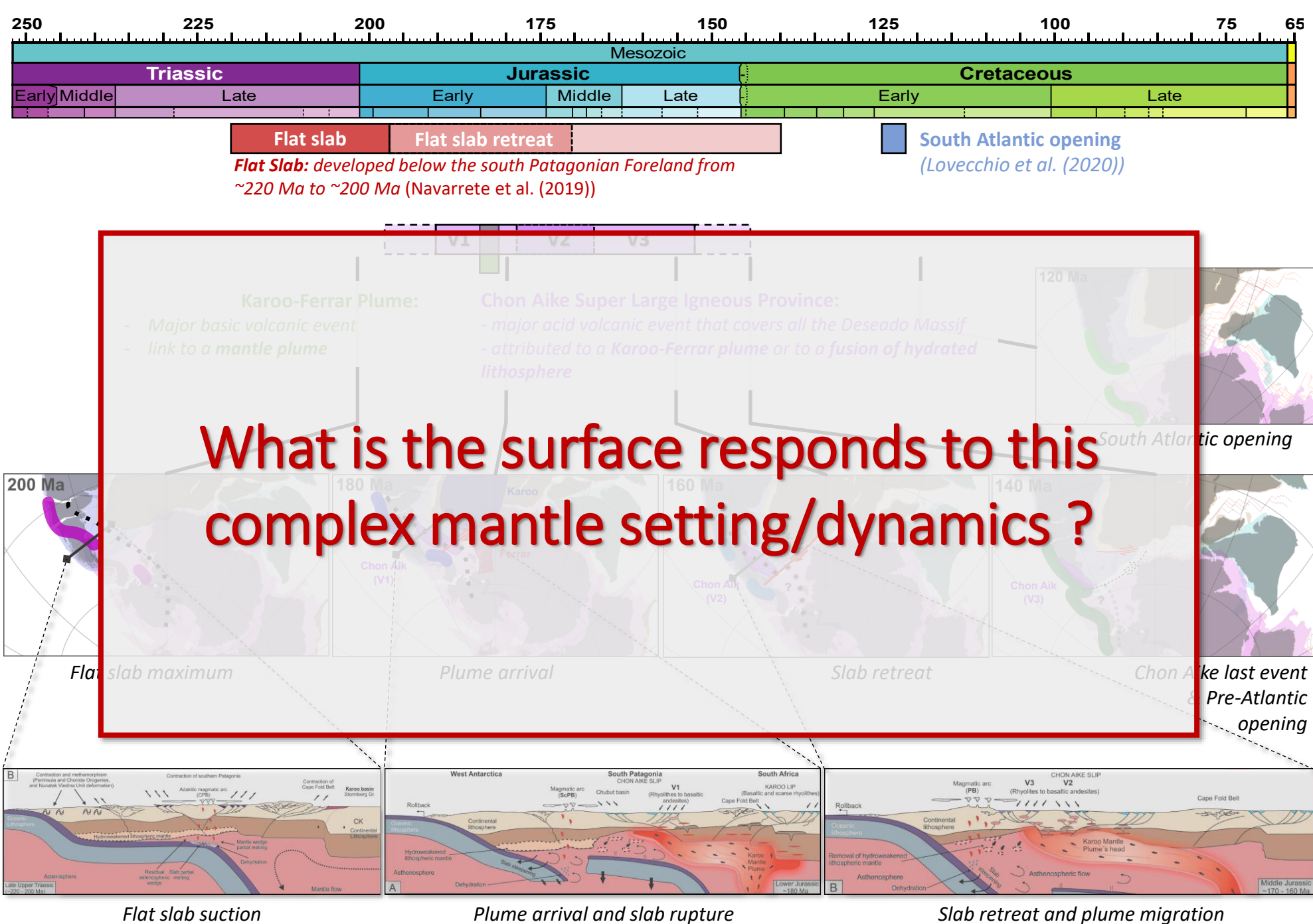
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Methods & Results

Methods: (U-Th)/He on apatite (AHe)

- Thermal sensitivity: 40° to 120°C
- 2 – 5 km depth

Results:

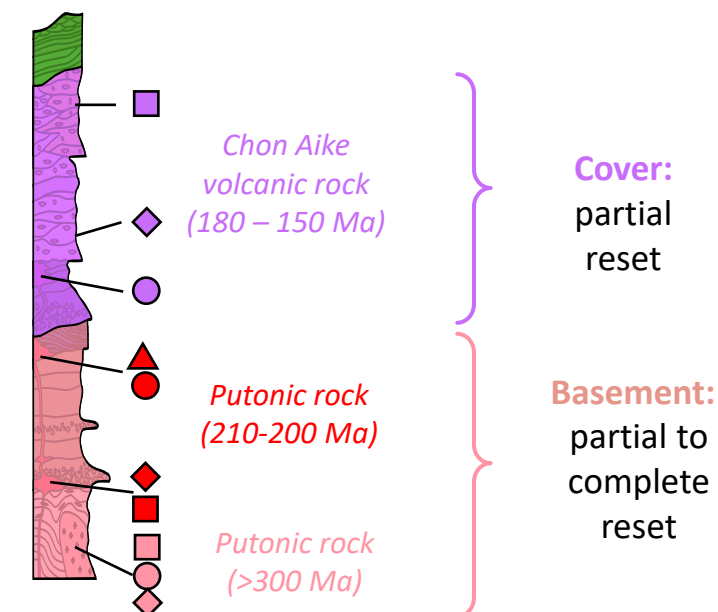
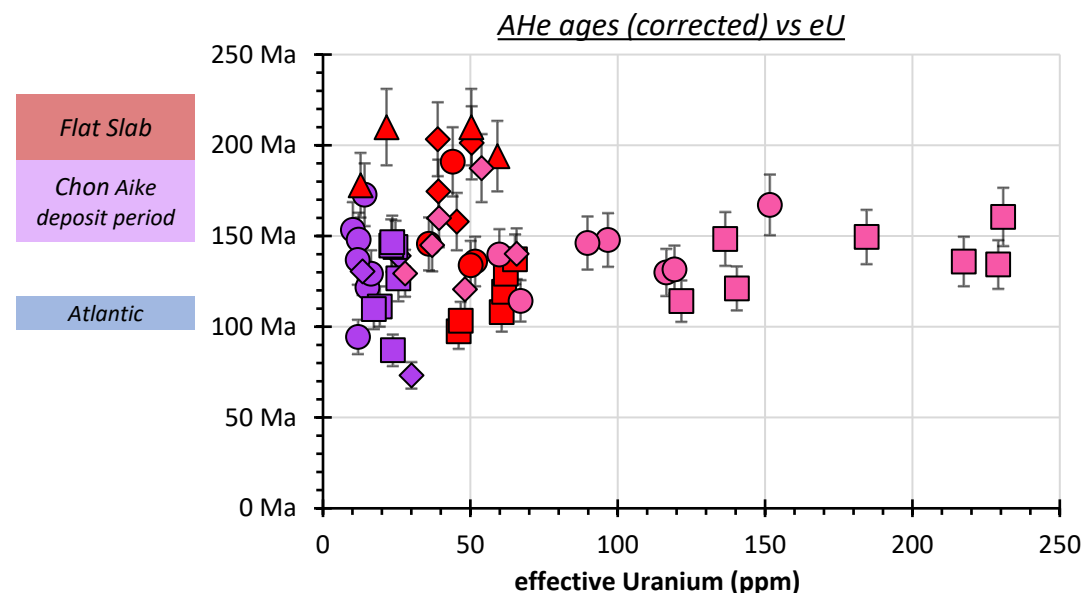
- 11 samples dated in AHe
- 4 samples were dated in AFT by Fernandez et al. (2019) and were used during the modeling

Ages spread from a ~200 Ma to ~100 Ma, with a large range of eU

Thermal modeling:

- Thermal inversion using QTQt software (Gallagher, 2012)
- Apatite Helium diffusion using Gerin et al. (2017) model

(U-Th)/He on apatite results:

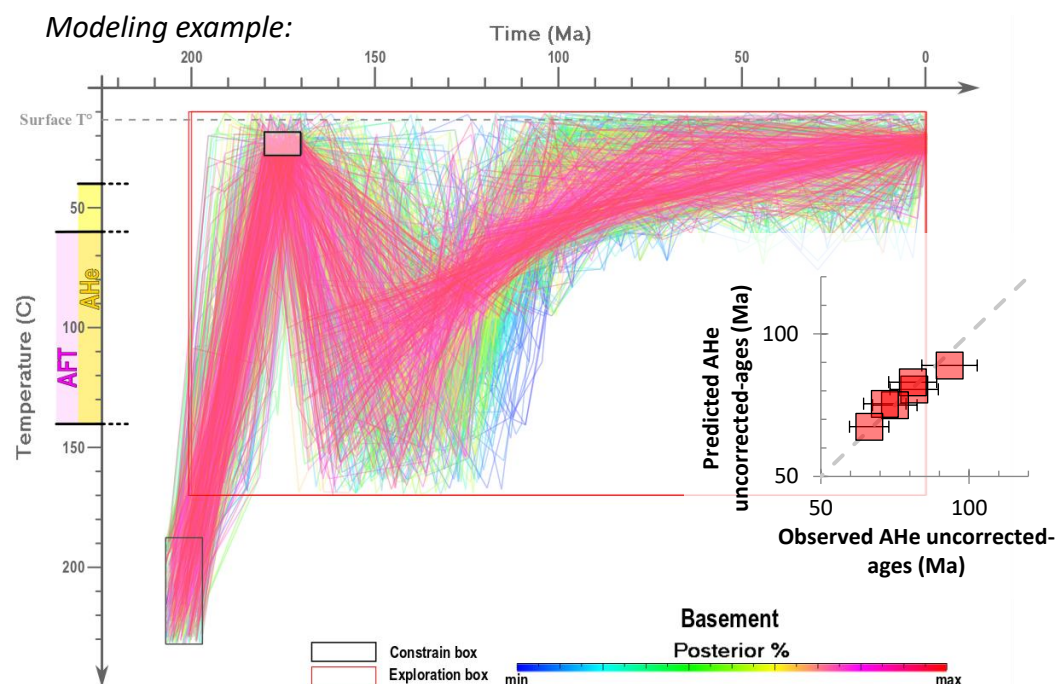


Modeling results:

- All samples show a clear **heating event**, up to ~100°-120°C, **before ~150 Ma**
- Followed by a **cooling up to 100 Ma**

This event could be interpreted as:

- A burial and exhumation event
 - A **thermic event**, linked to the geodynamic setting/evolution
- ➡ Necessity to characterize the impact of the thermic event



Thermal Event ?

The Chon Aike LIP impact:

- LIP link to a mantle plume or a hydrated lithosphere fusion.

➡ No significant gradient modification

- surface hot deposit (rhyolite/ignimbrite)

➡ Only a local effect

Interpretation:

The impact of the Chon Aike thermal event is negligible regarding to the heating recorded.

Chon Aike SLIP could produced two potential thermal impacts:

- **Geothermal gradient increased**, linked to the plume arrival/life or lithosphere fusion.

Tested by 1D lithospheric model:

No significant impact on the geothermal gradient
($< 20^{\circ}\text{C}$ increase in shallow surface)

- Basement heating by **hot deposit emplacements**

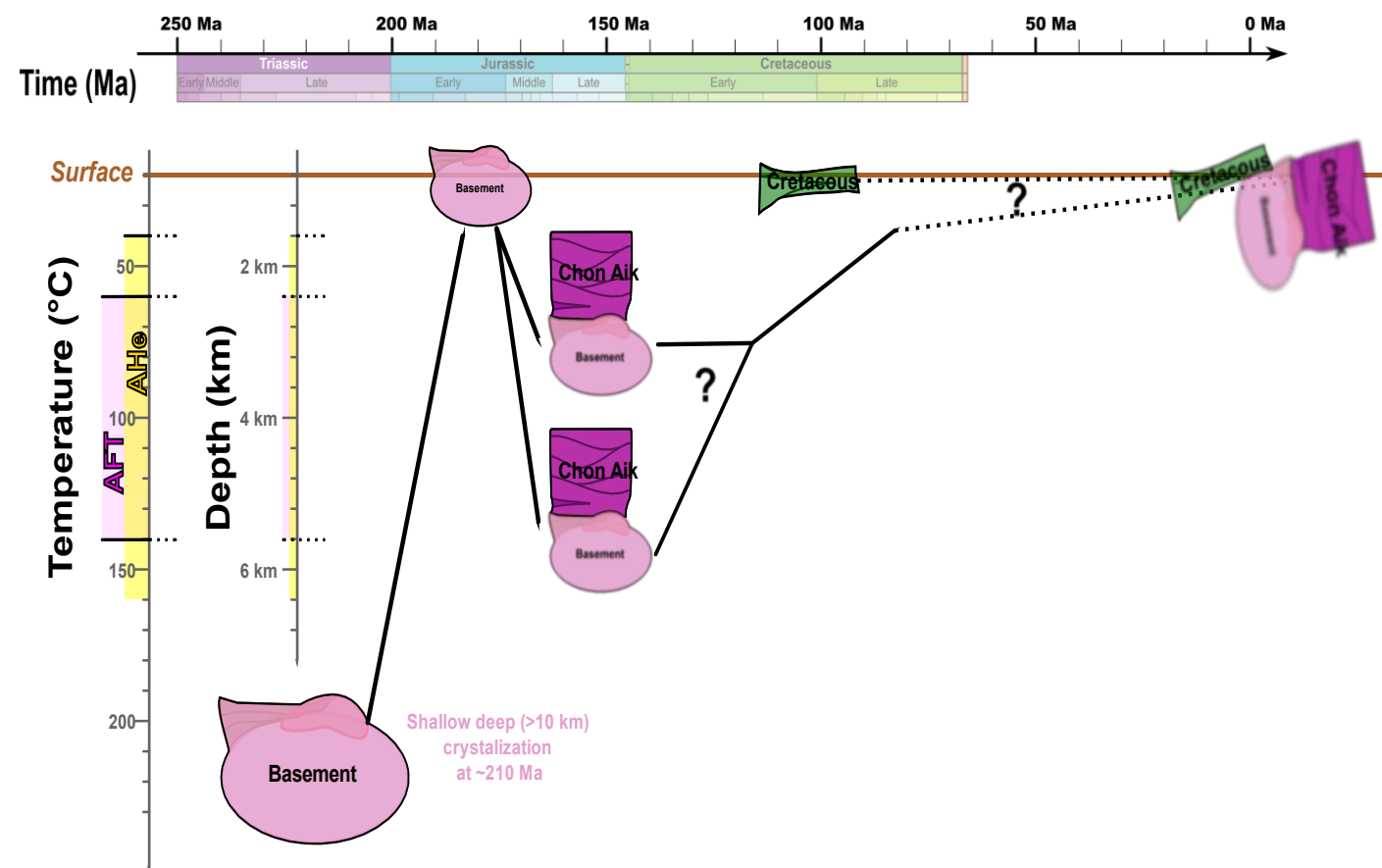
Tested by 1D surface model:

Short-term event ($< 2\text{ Ma}$),
with a local impact ($\sim 200\text{m}$ deep)

Surface evolution:

- ~ 205 to $\sim 195\text{ Ma}$: **first exhumation**, of $\sim 10\text{ km}$, put the basement in surface.
- ~ 170 to $\sim 150\text{ Ma}$: the basement was buried **under 3 to 5 km** of deposits, probably the Chon Aike formation
- ~ 150 to 100 Ma : the basement last exhumation

What mechanism triggered this surface evolution?



Burial/exhumation origin ?

The early Mesozoic exhumation:

This exhumation was **produced by the flat slab emplacement** between 220 Ma and 200 Ma.

The Jurassic burial – Chon Aike event:

- ~50 Ma duration,
- **deposits and conservation of 3 to 5 km of material**
- the **Deseado Massif/Chon Aike event position** through time

The Cretaceous exhumation:

- ~50 Ma duration
- structurally asymmetric

So, this J-K surface evolution, far from any plate border, was controlled by two large scale phenomenon:

- **the plate strain/stress propagation**, triggered by the flat slab retreat and Atlantic opening
- **the crustal equilibration** and evolution, triggered by the mantle dynamics and lithosphere fusion

The burial and exhumation can be linked to:

1) Plate strain/stress propagation (Plate model)

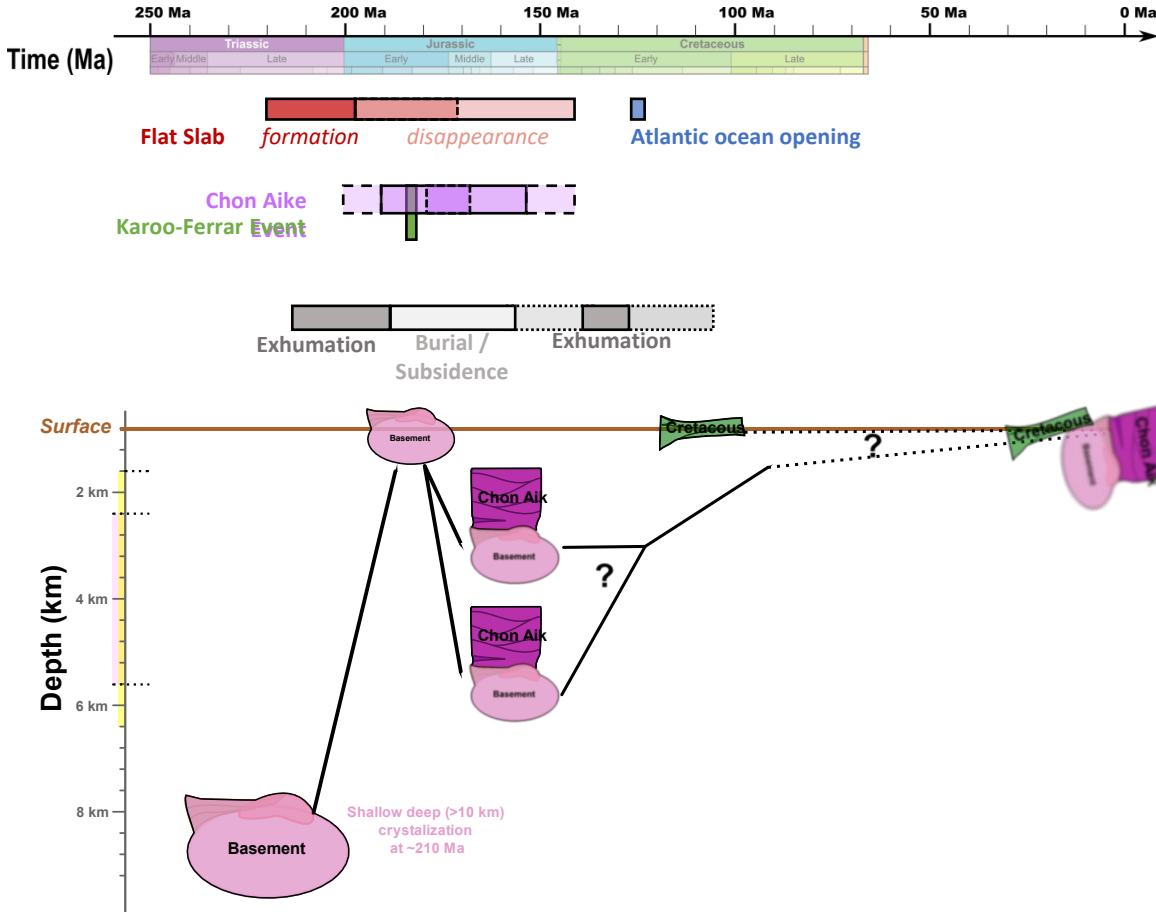
Arguments :

- Consistent with a **Jurassic extensional event**, produced by **the slab retreat**
- Consistent with a **Cretaceous compressional event**, produced by the **Atlantic opening**

2) Crust equilibration (*Plume model*)

Arguments :

- Consistent with the **Jurassic magmatic evolution**
- Could correspond to a **dynamic topography/isostatic equilibration response to the slab evolution and failure ?**



Conclusion:

- the **Patagonian Foreland** was **not completely stable** during the **Mesozoic** (3 to 5 km burial and exhumation).
- Its Mesozoic **surface evolution** was **controlled by large scale phenomenon**.
- Those results could be used, in the future, as **surface evolution constrains to improve the understanding of these phenomenon, here the exact mantle evolution** during and after the Chon Aike event.

References

Thanks for your attention,

*I hope the presentation was clear,
if you have any questions,
feedback and comments,
I'm in the chat room to
answer or clarify!*

Conclusion:

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- Its **Mesozoic surface evolution** was controlled by large scale phenomenon.
- Those results could be used, in the future, as **surface evolution constrains** to improve the understanding of **these phenomenon**, here the **exact mantle evolution** during and after the Chon Aike event.

Thank you for your attention !

Geology:

- Navarrete et al. 2019. *Journal of South American Earth Sciences* 94 (October): 102224.
- Pankhurst et al. 2000. *Journal of Petrology* 41 (5): 605–25.

AFT data:

- Fernández et al. 2020. *Tectonophysics* 775 (January): 228302.

Geodynamics:

- Lovecchio et al. 2020. *Earth-Science Reviews* 203 (April): 103138.
- Navarrete et al. 2019. *Earth-Science Reviews* 194 (July): 125–59.

QTQt modelling (thermochronological data inversion):

- Gallagher. 2012. *Journal of Geophysical Research: Solid Earth* 117 (B2):
- Gerin et al. 2017. *Geochimica et Cosmochimica Acta* 197 (January): 87–103.

Thermal modeling (modified from):

- Zeyen and Fernandez. 1994. *Journal of Geophysical Research: Solid Earth* 99 (B9): 18089–102.
- Fayon and Whitney. 2007. *Tectonophysics* 434 (1–4): 1–13