

Where is the Toba eruption in the Vostok ice core? Clues from the O and S isotopes

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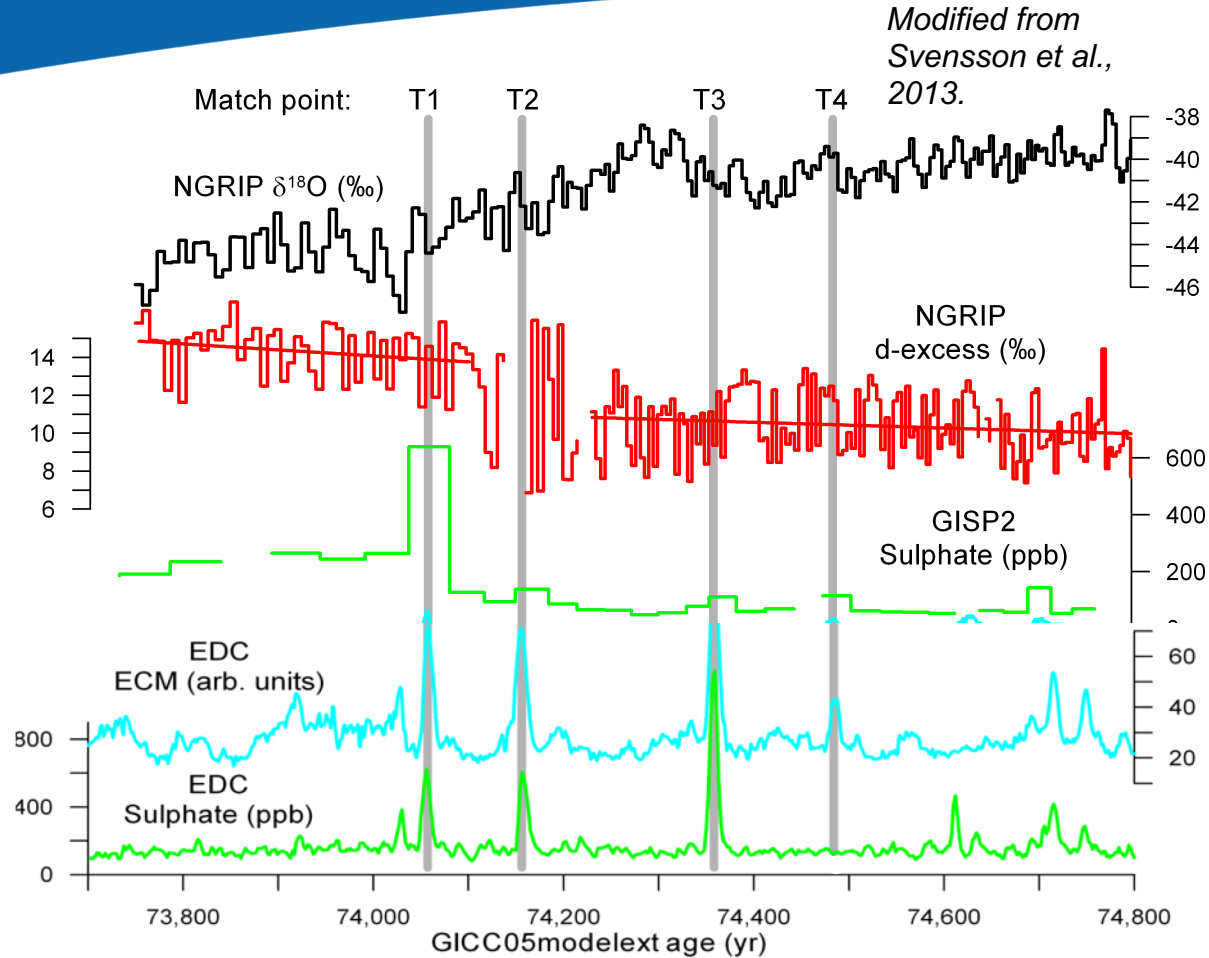
Caldera size Tambora vs Toba



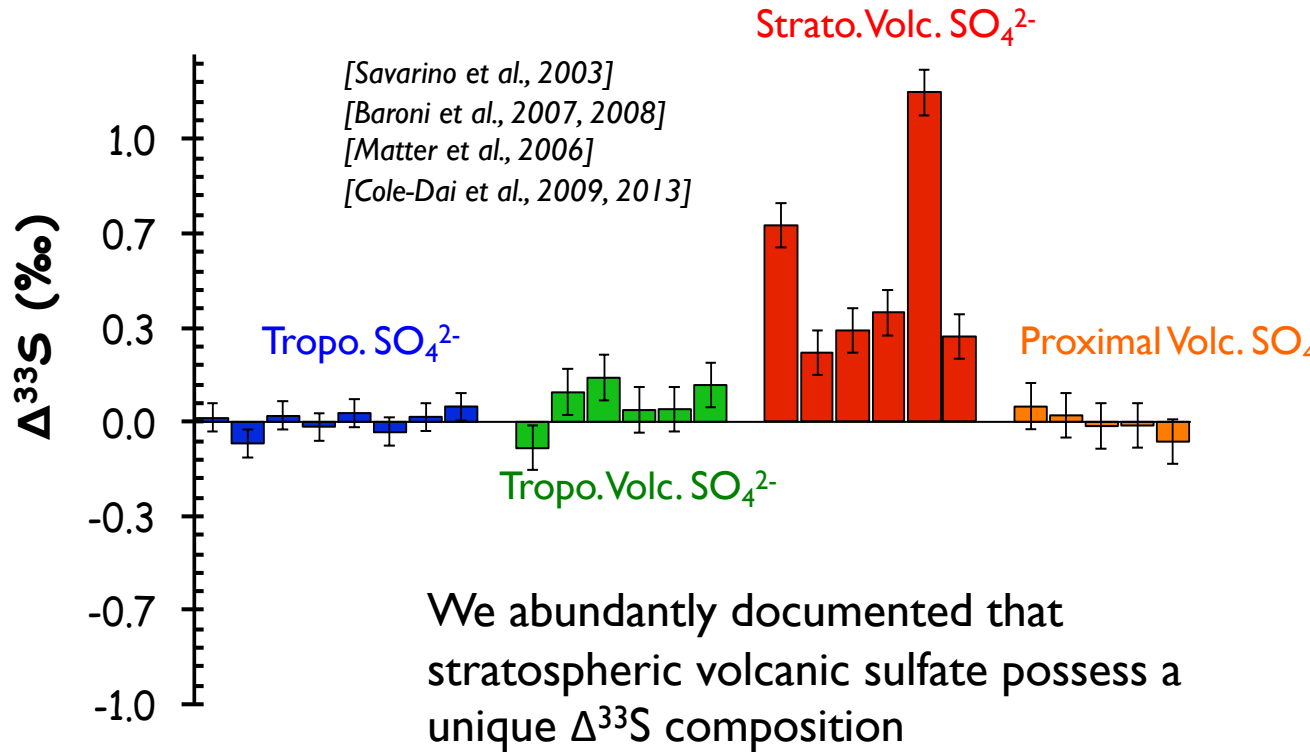
	Tambora (Stothers, 1984, Self et al. 2004, Sigl et al., 2015)	Toba (Robock et al. 2009, Timmreck et al. 2010, William, 2012, Smith et al. 2018)
Year of eruption	1816	≈ 74 ka
Caldera size	100 km ²	3 000 km ²
Dense Rock Equi.	33 km ³	2 500 – 3 000 km ³
Ash distance	550 km	9 000 km
Sulfur emitted	27-29 Tg	35 – 850 Tg
Cooling forcing	17 W m ⁻²	18 – 100 W m ⁻²
Global temp. drop	1.0 to 1.5 °C	3 to 17 °C
Highlight	The year without summer	MIS 4-5, Population bottleneck,

No doubt Toba was a massive eruption, still dynamics and characteristics of the eruption remain largely unknown

Based on bipolar comparison, four possible sulfate peaks in ice cores for the Toba has been reported



We used the $\Delta^{33}\text{S}$ proxy to verify that these sulfate peaks are at least of stratospheric origin as it should be for mega-eruptions



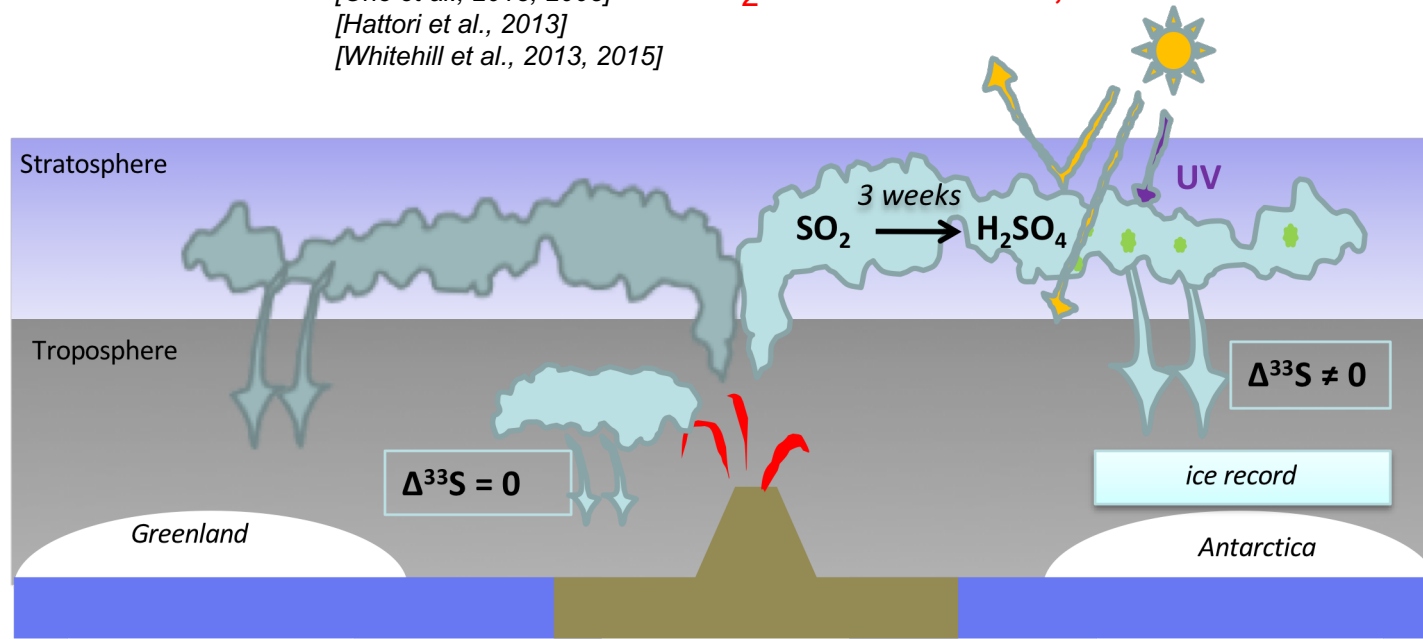
Mechanism to generate a non-zero $\Delta^{33}\text{S}$

[Farquhar et al., 2001]

[Ono et al., 2013, 2008]

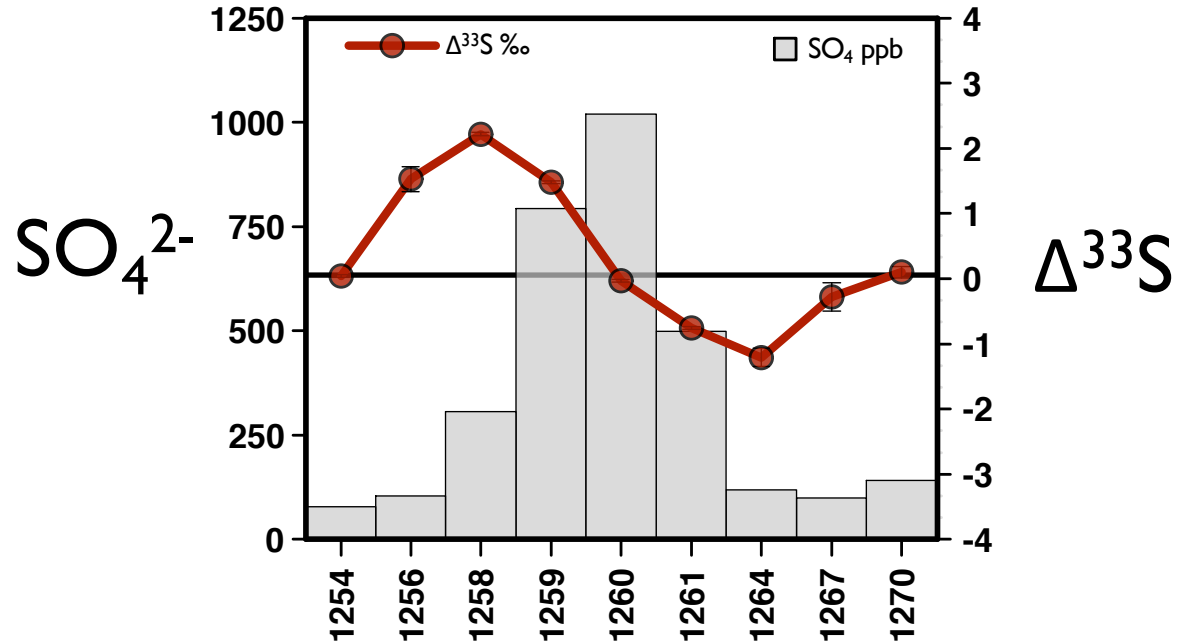
[Hattori et al., 2013]

[Whitehill et al., 2013, 2015]



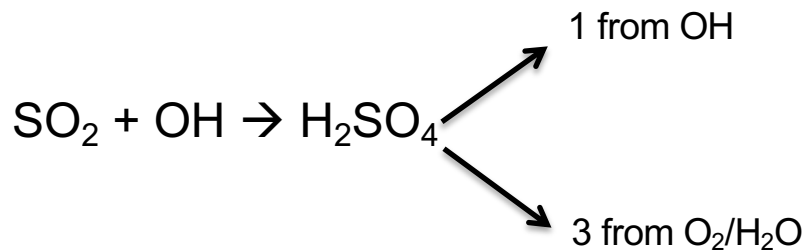
$\Delta^{33}\text{S}$ tracer: → independent any chronology
→ independent any location
→ independent of deposition flux

Typical $\Delta^{33}\text{S}$
imprint in ice for
stratospheric
eruptions

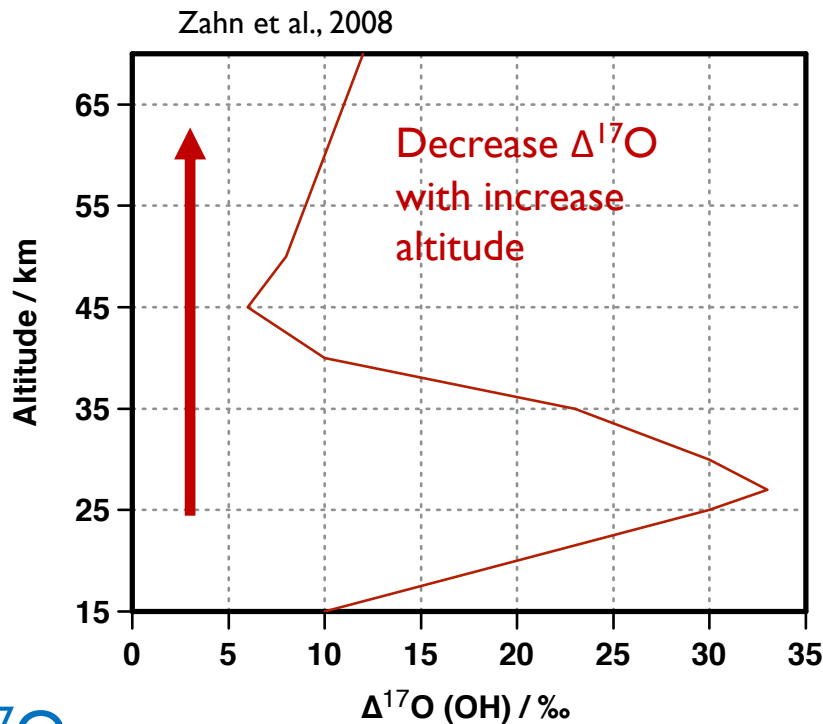


We also used the $\Delta^{17}\text{O}$ of sulfate as a qualitative proxy of altitude injection

Main oxidation pathway in the stratosphere

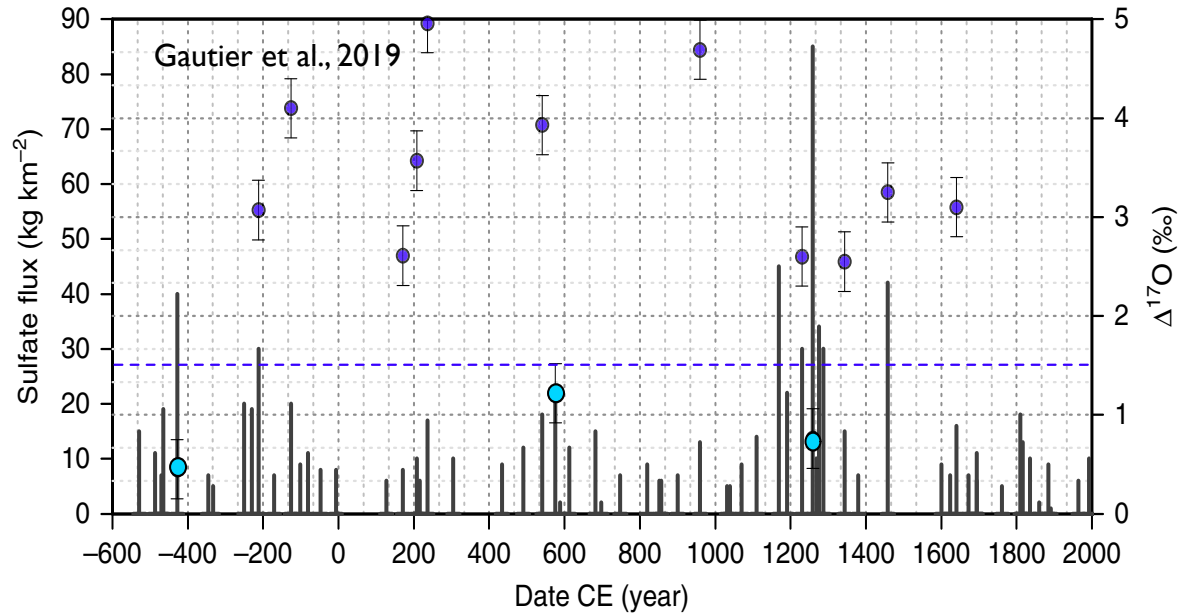


High altitude eruptions have low $\Delta^{17}\text{O}$



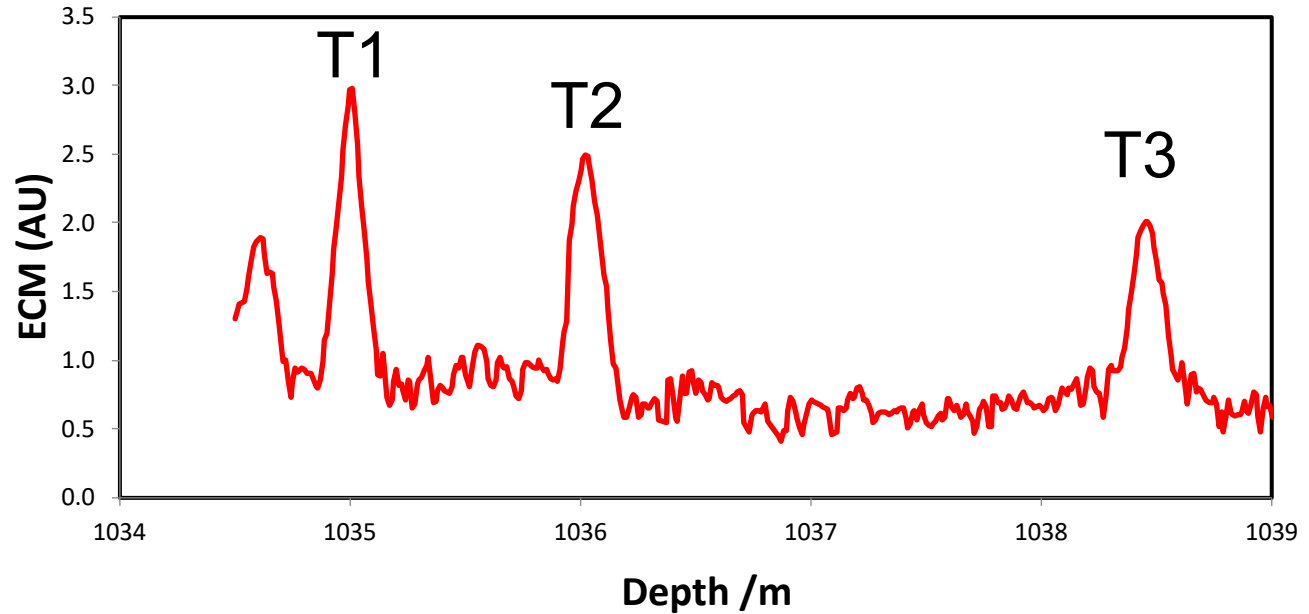
Large eruptions: 1259
CE, 575 CE, 426 BCE
have low $\Delta^{17}\text{O}$

Volcanic $\Delta^{17}\text{O}$



We have identified
in the Vostok ice
core, three of the
potential Toba
sulfate peak

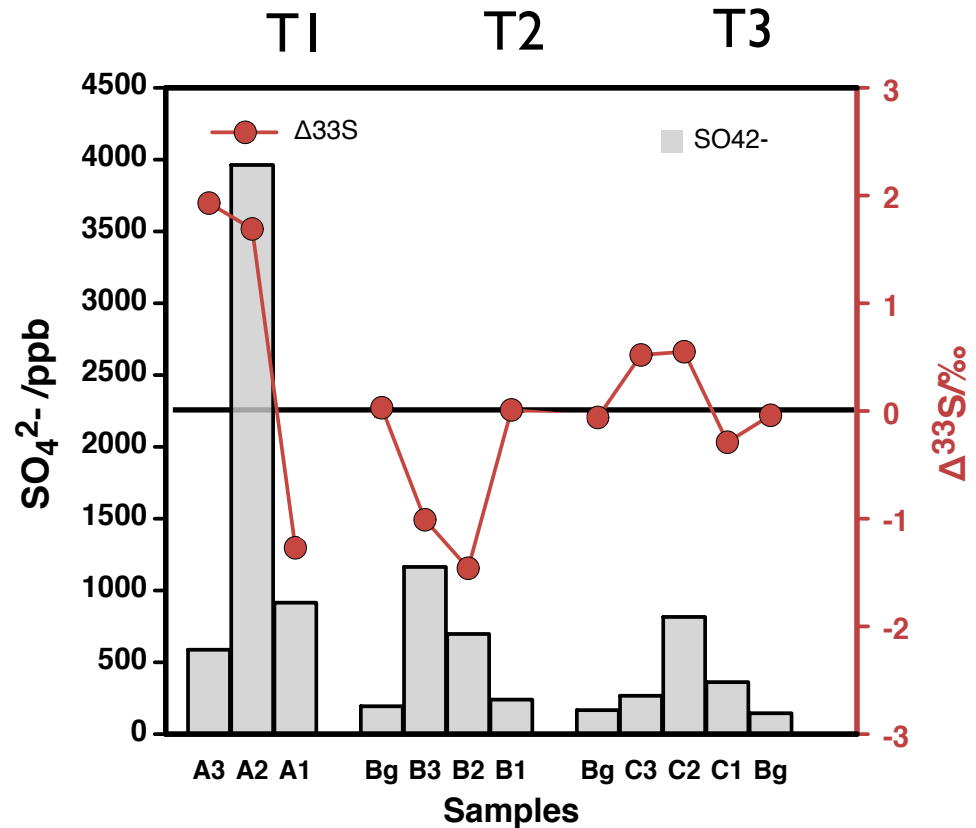
ECM profil Vostok Ice core
with the potentiel Toba eruption(s)



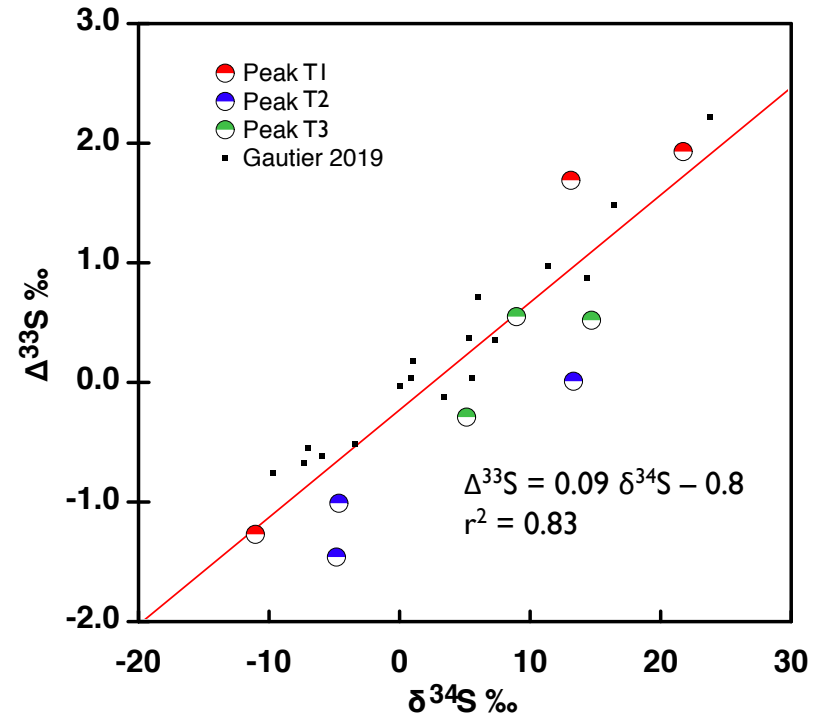
$\Delta^{33}\text{S}$ of the proposed
Toba's layers

All three eruptions are
clearly stratospheric

Peak B preserved only
 \ominus phase



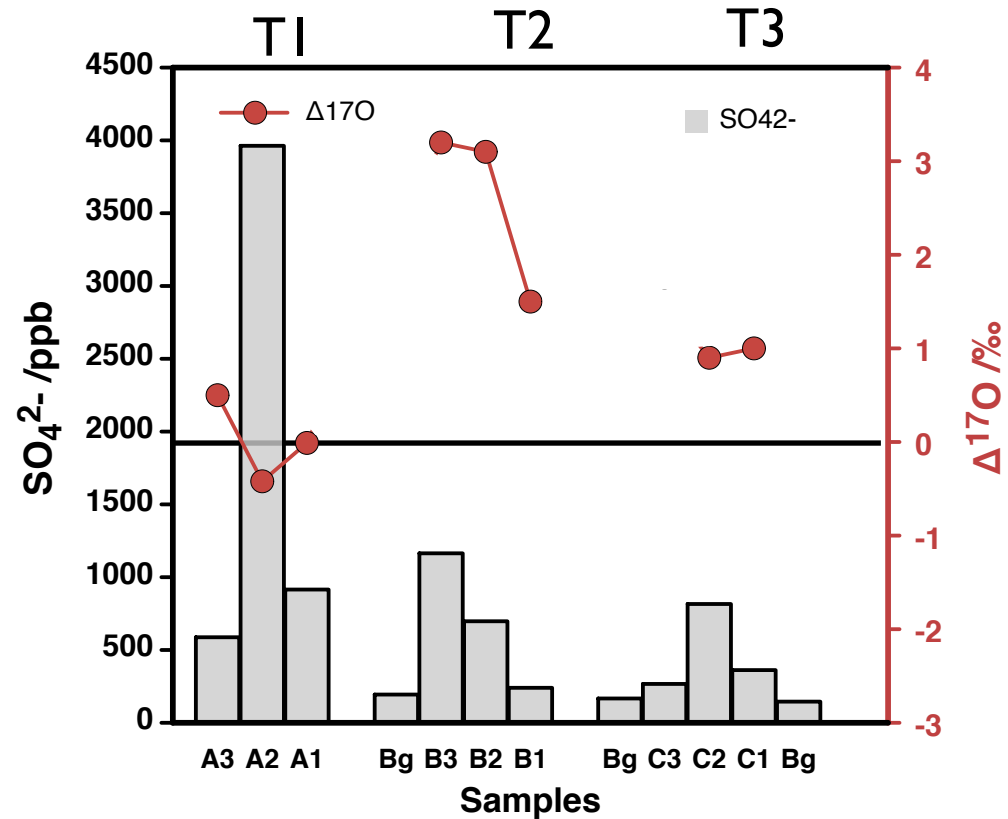
The sulfur isotopic composition of the T1,T2,T3 eruptions is indistinguishable from previous stratospheric eruptions



Peak T1 shows extremely low $\Delta^{17}\text{O}$, never observed before

Peak T2 as expected for standard stratospheric eruptions

Peak T3 shows low $\Delta^{17}\text{O}$, in range of previous observation



Summary & Conclusions

Peak T1: High SO_4^{2-} , +/- $\Delta^{33}\text{S}$, Lowest $\Delta^{17}\text{O}$ observed

➡ Possible candidate for Toba

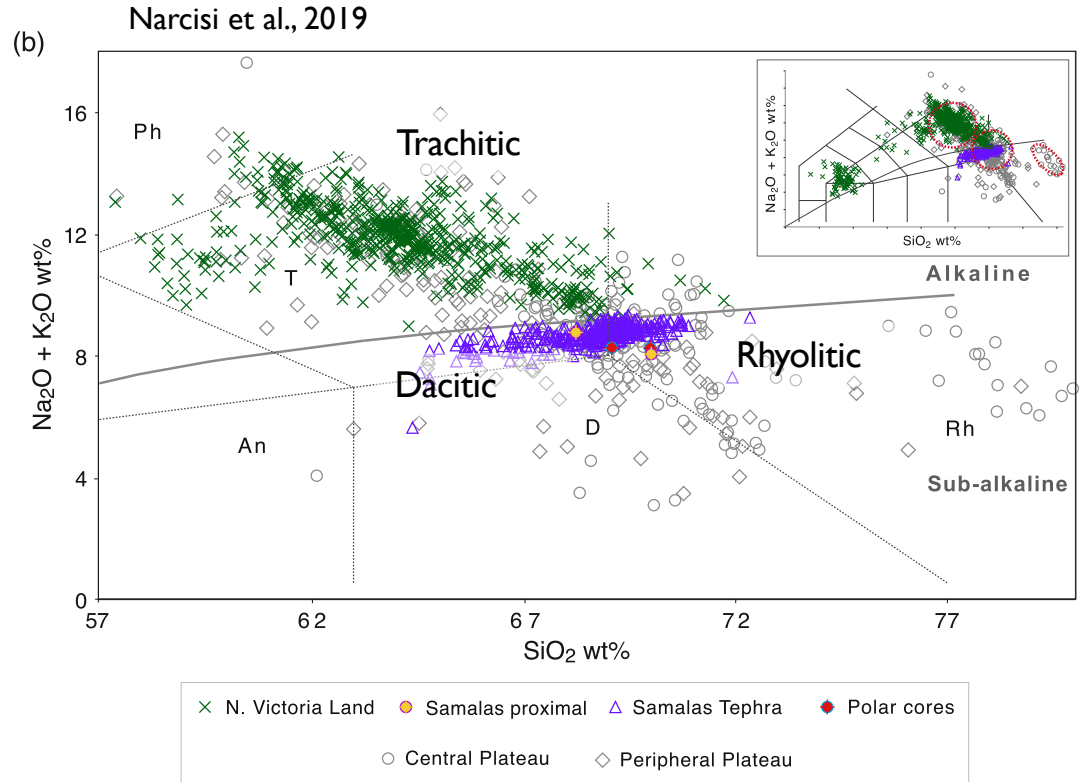
Peak T2: Medium SO_4^{2-} , - $\Delta^{33}\text{S}$, high $\Delta^{17}\text{O}$

➡ Missing positive component that precludes any conclusion

Peak T3: Medium SO_4^{2-} , +/- $\Delta^{33}\text{S}$, Low $\Delta^{17}\text{O}$

➡ Possible candidate for Toba

In progress ...
Geochemical analysis if
shards are present



Svensson et al., Clim Past, 2013

*« Based on ice core data there are thus several possible candidates for Toba eruption(s), and it cannot **be excluded that several of them are related to Toba.** ../..*

*Several closely spaced large Toba eruptions could, however, help to explain why **none of the identified ice core events are as strong as would be expected from current geological evidence** »,*

The End

