



**EGU2020 SSP2.13**  
**Asian Climate and Tectonics**  
**D840 | EGU2020-4042**

**A new negative feedback mechanism for balancing  
Tibet uplift-driven CO<sub>2</sub> drop:  
Evidence from Paleogene chemical weathering records  
in the northern Tibetan Plateau**

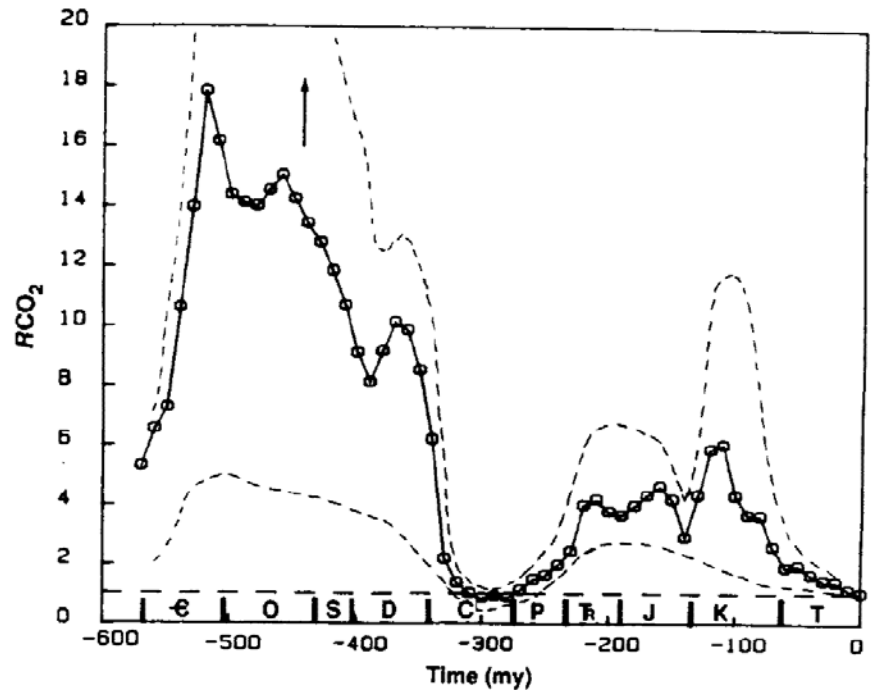
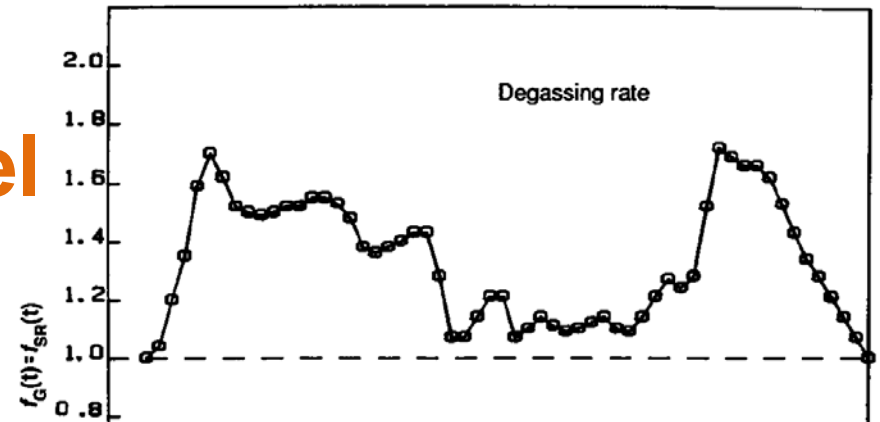
**Xiaomin Fang<sup>1</sup>, Albert Galy<sup>2</sup>, Yibo Yang<sup>1</sup>, Weilin Zhang<sup>1</sup>,  
Chengcheng Ye<sup>1</sup>, Chunhui Song<sup>3</sup>**

- 1. Institute of Tibetan Plateau Research, CAS**
- 2. CRPG-CNRS-University of Lorraine**
- 3. Lanzhou University**

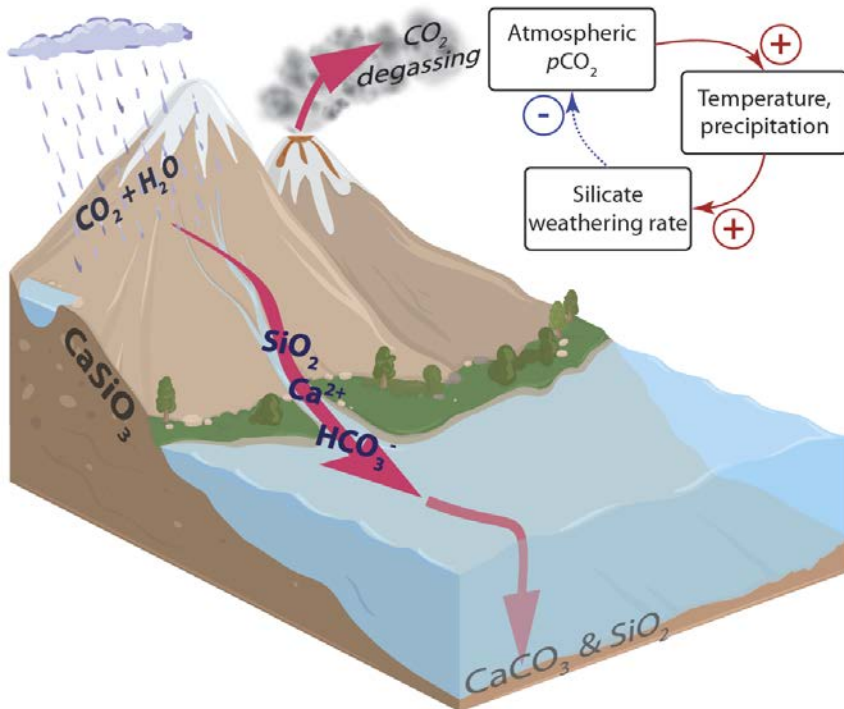


# BLAG model

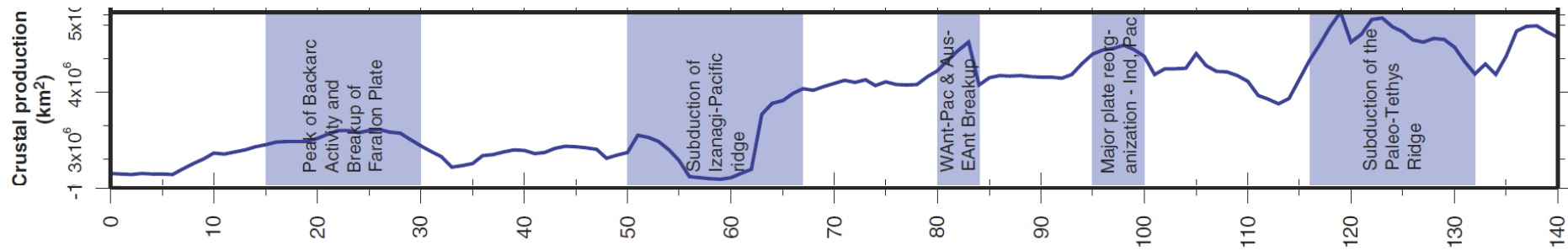
Robert A. Berner



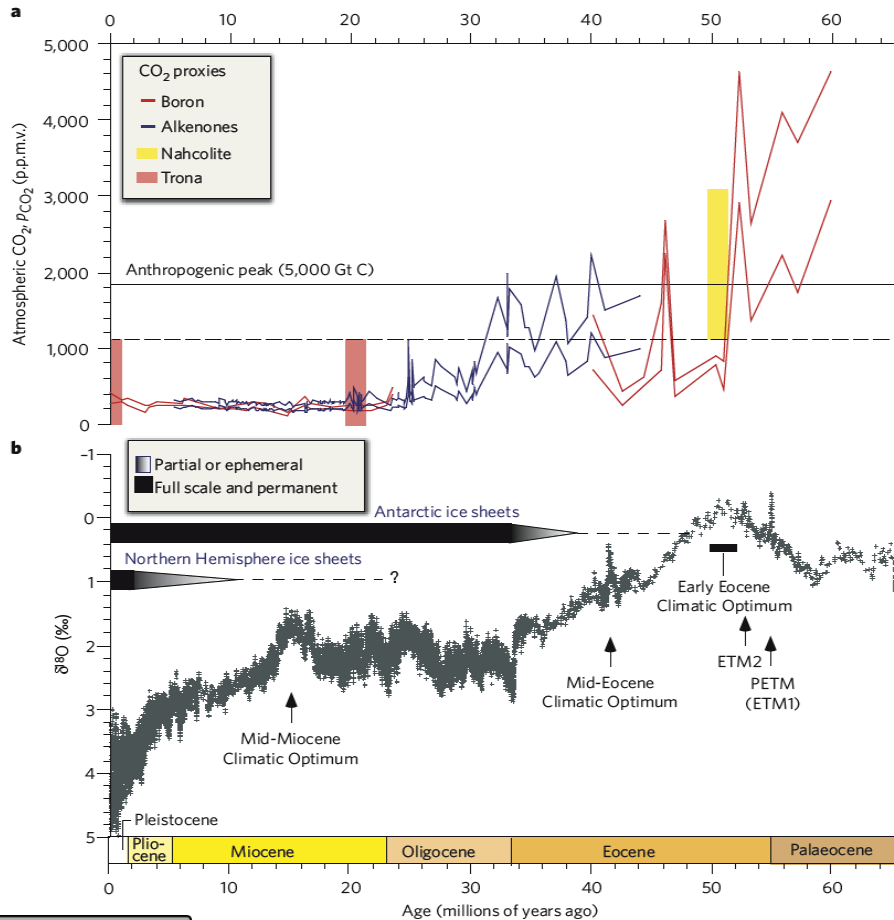
Berner, 1990



Frings, 2019



Muller et al., 2008

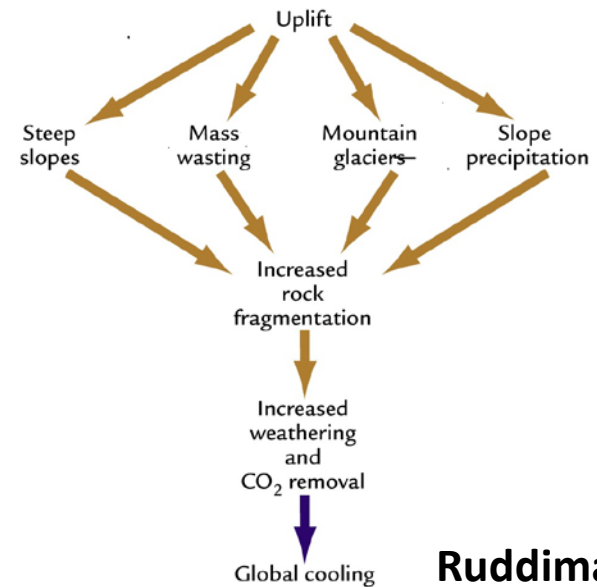


Zachos et al., 2008



## Uplift-Weathering Hypothesis

Maureen Raymo

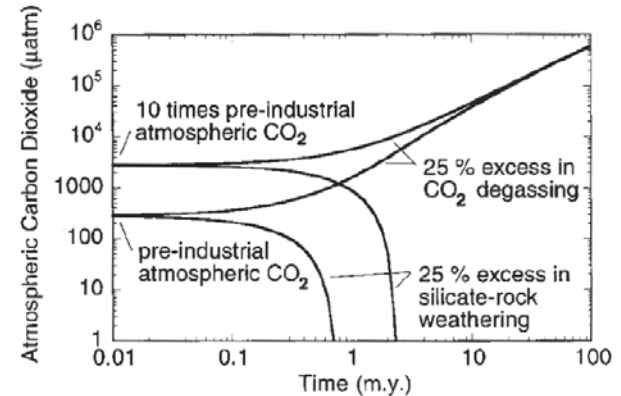


Ruddiman, 2008

# Negative Feedbacks to maintain geological carbon cycle

The imbalance resulting from accelerated  $\text{CO}_2$  consumption and a relatively stable  $\text{CO}_2$  input from volcanic degassing during the Cenozoic should have depleted atmospheric  $\text{CO}_2$  within a few million years.

Therefore, a negative feedback mechanism must have stabilised the carbon cycle.



**Bener and Caldeira, 1997**

- 1 Reduce organic carbon burial (Raymo and Ruddiman, 1992)
- 2 Enhance reversal weathering and  $\text{CO}_2$  release (Raymo and Ruddiman, 1992)
- 3 Enhance metamorphism degassing (Bickle, 1996)
- 4 Enhance sulfide weathering and  $\text{CO}_2$  release (Torres et al., 2014)
- 5 Reduce Temperature-regulated weathering (Kump and Arthur, 1997)

5.1 Ocean crust basalt (Coogan and Dosso, 2015)

5.2 Ocean island basalt (Li et al., 2013)

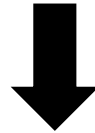
5.3 Continental arc (Lee et al., 2015)

5.4 Weatherability (Caves et al., 2016)

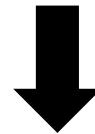
5.5 Weathering in other regions (Kump and Arthur, 1997)

# Working hypothesis

CO<sub>2</sub> consumption increase **in tectonically active region** (e.g., Himalayas, where erosion rate increases)



CO<sub>2</sub> decline reduced the degree of silicate alternation **in tectonically less active region** (where erosion rate is stable)

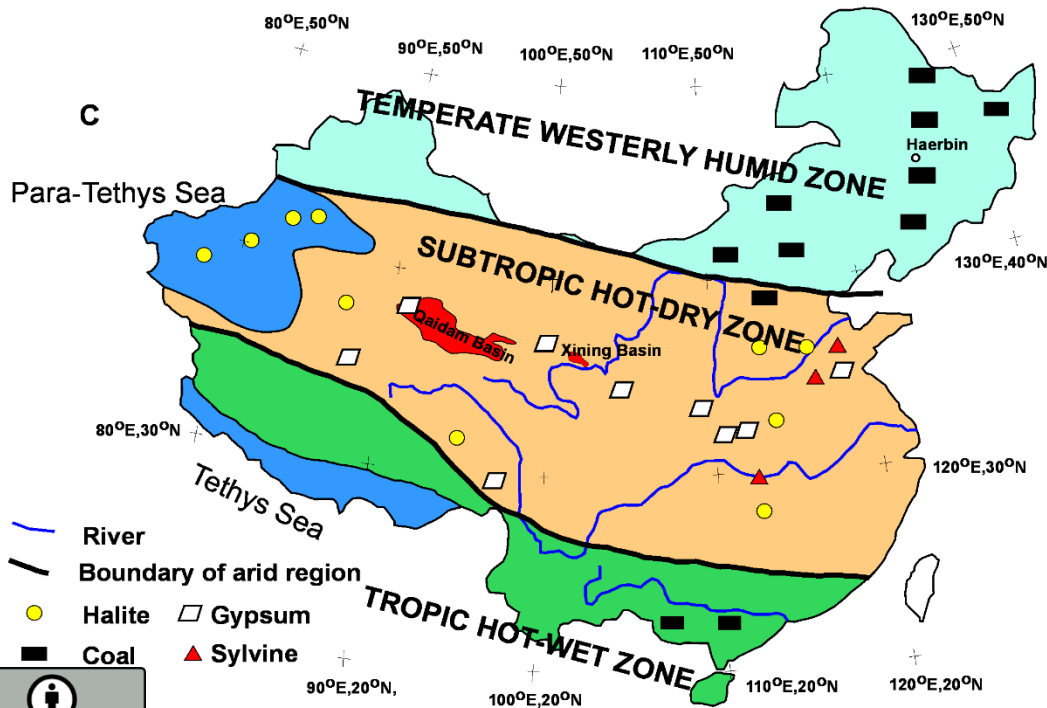
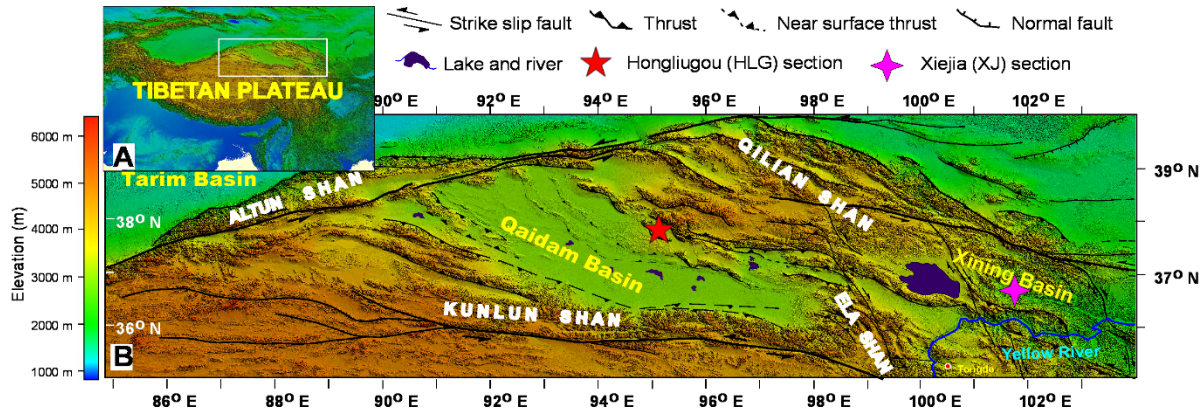


Decrease in silicate weathering flux and CO<sub>2</sub> consumption **in tectonically less active region**



**Carbon cycle balanced!!**

# Study region

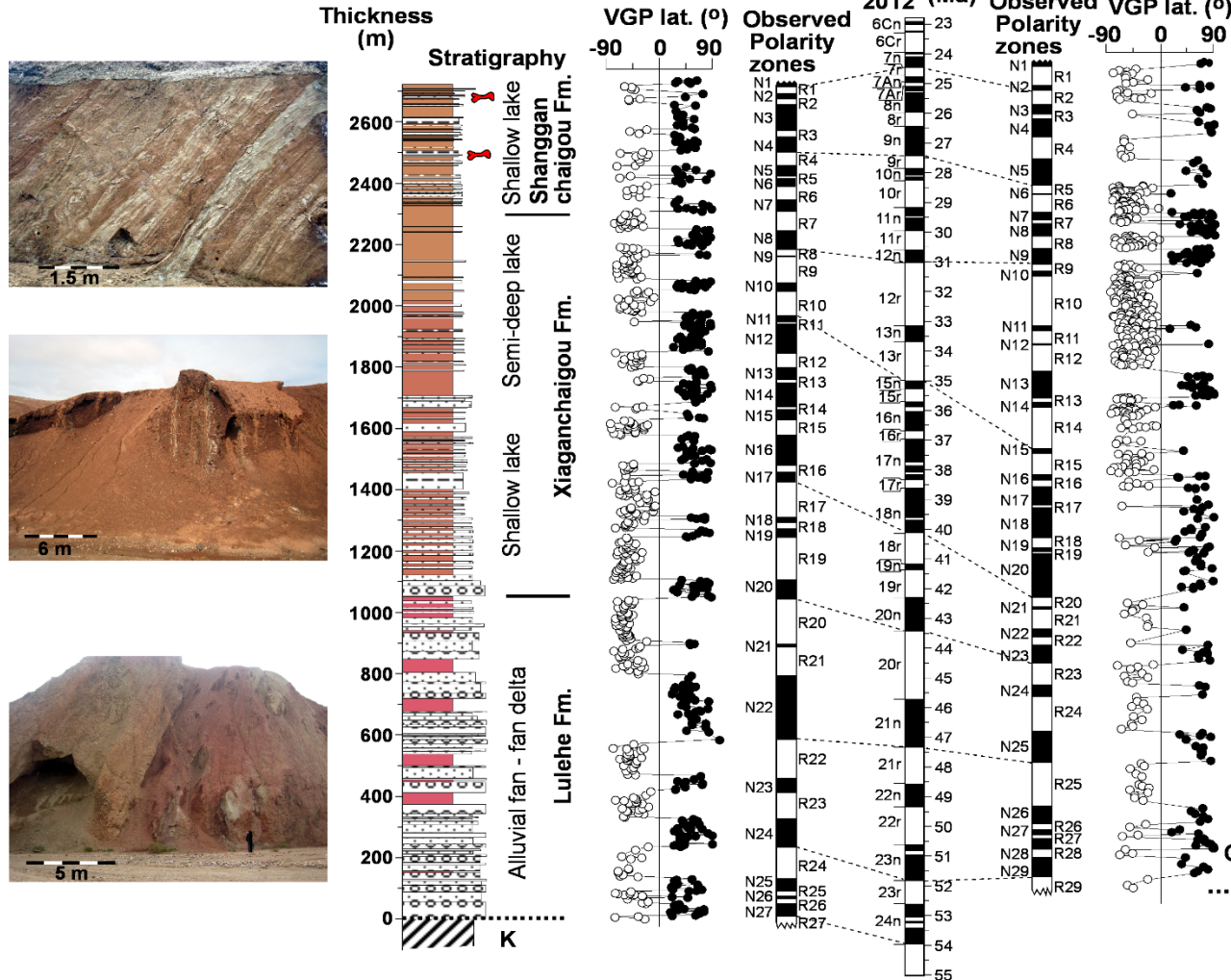


- 1) During the Paleogene, the northern Tibetan Plateau was a tectonically less active area;
- 2) The Asian summer monsoon did not reach this region until the beginning of the Neogene

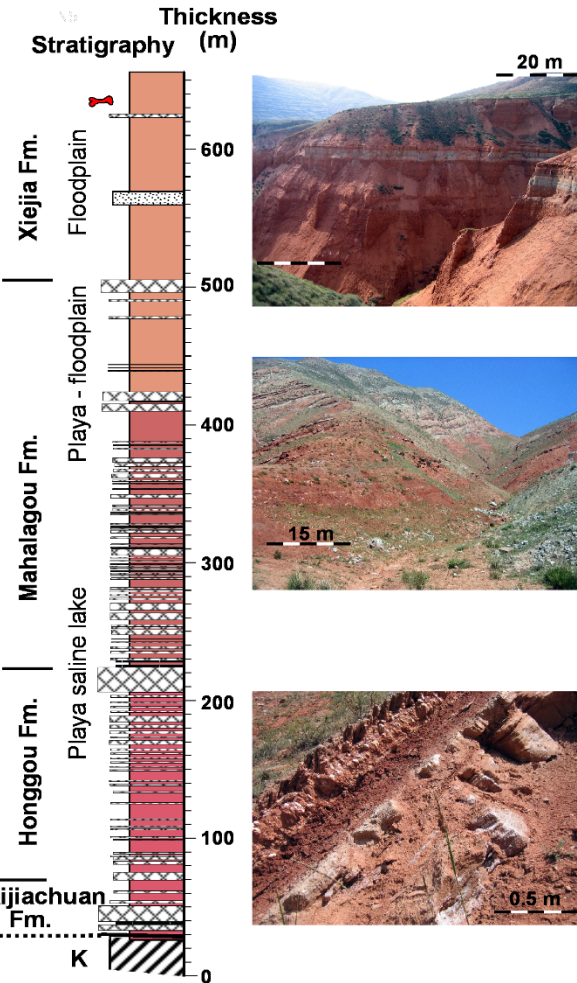


# Lithology and Age control

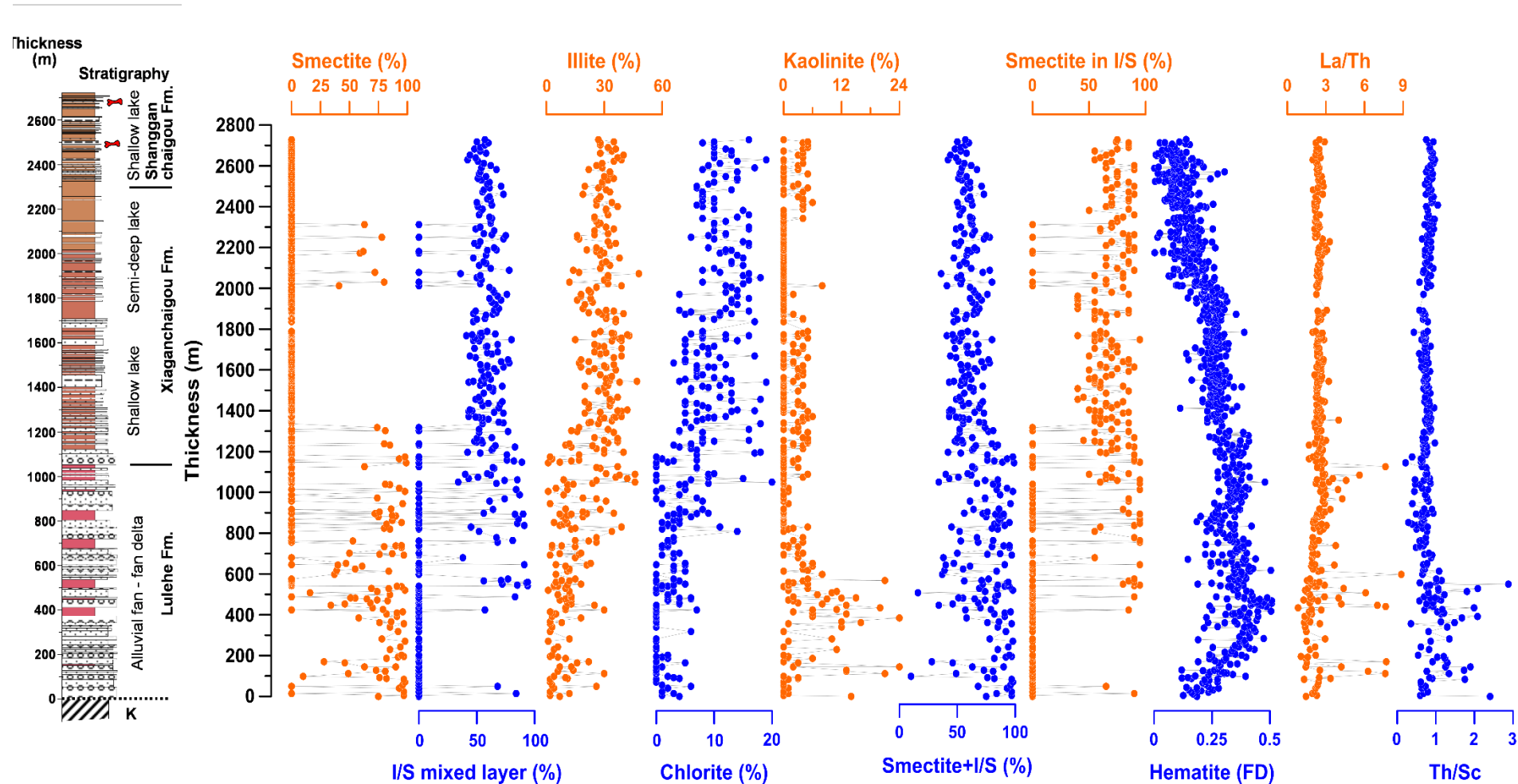
## Qaidam Basin



## Xining Basin

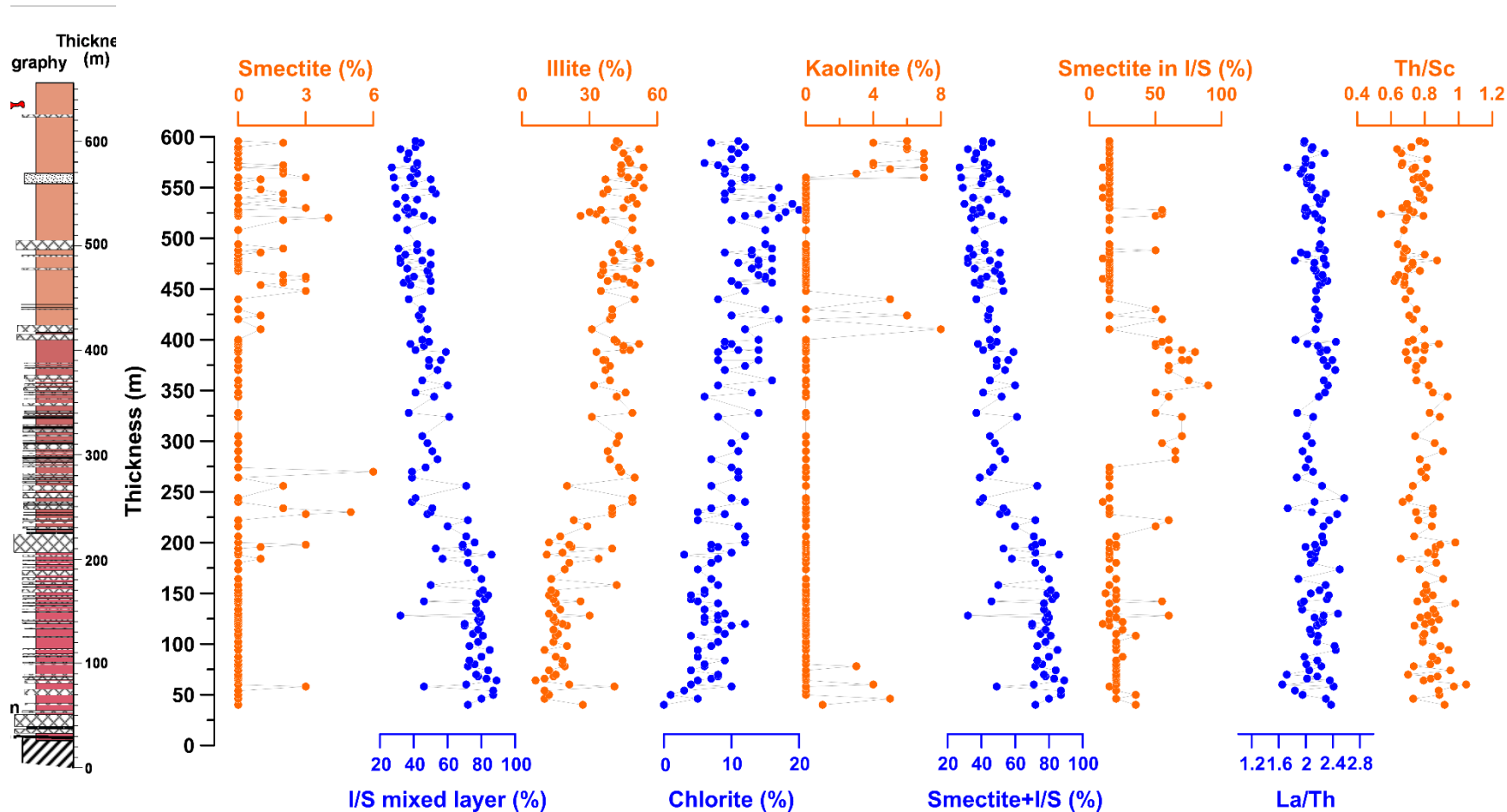


# Clay mineral records in Qaidam Basin

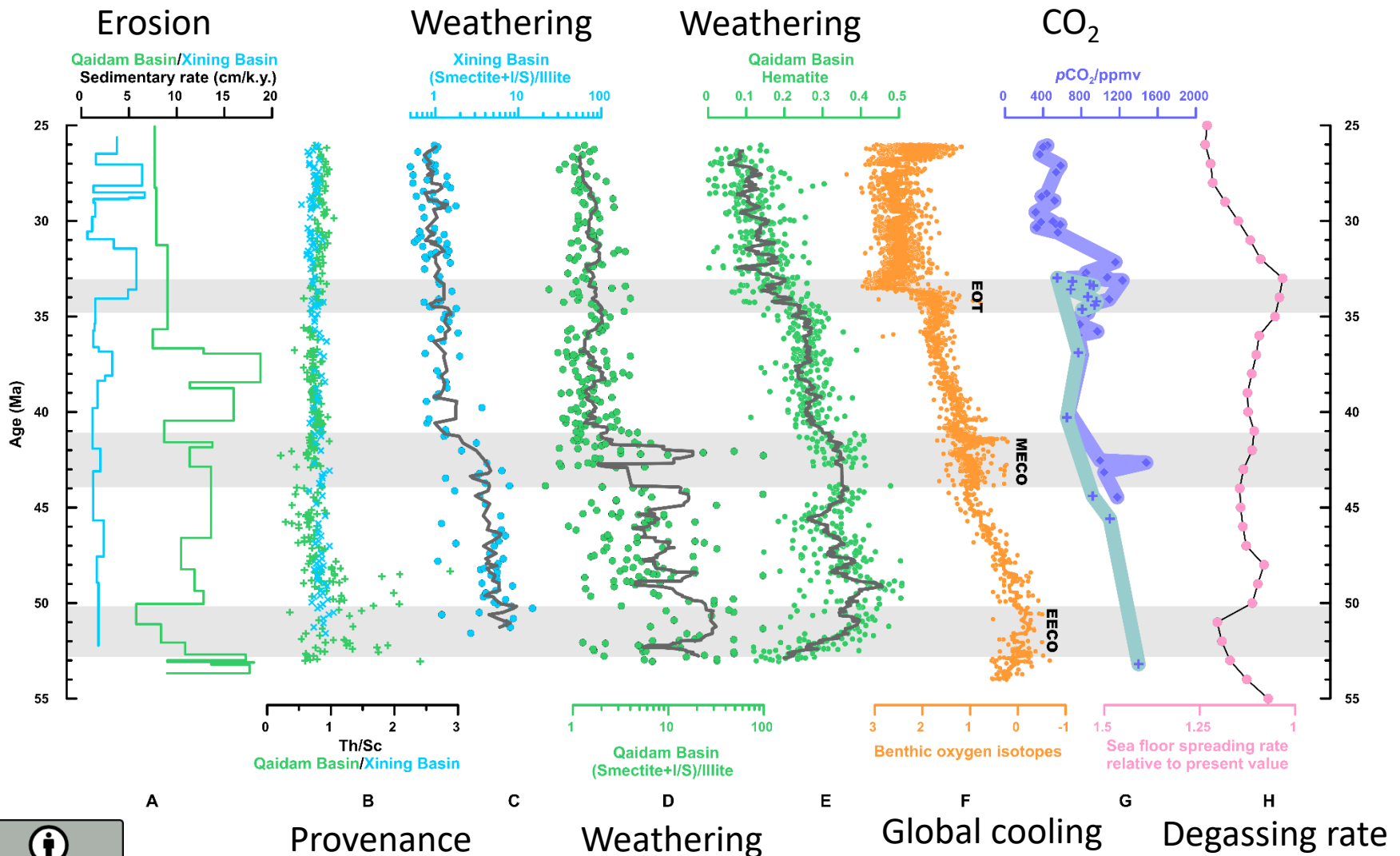




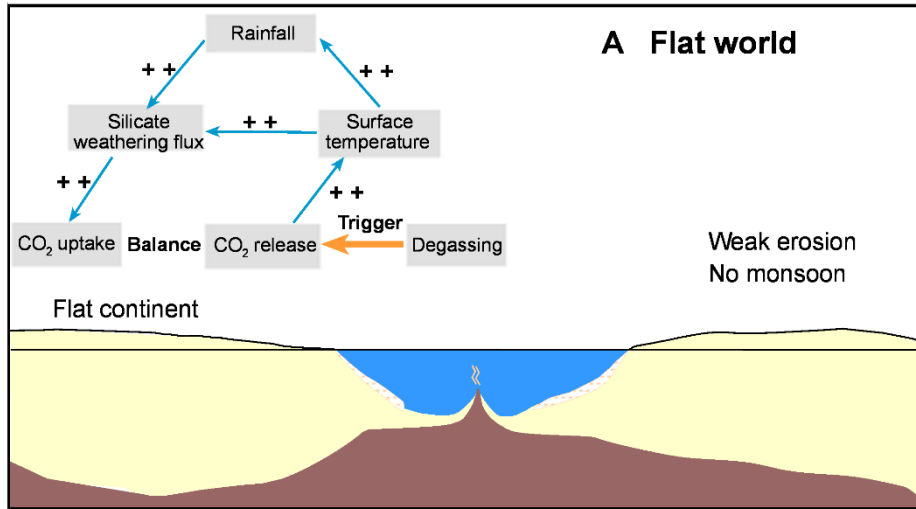
# Clay mineral records in Xining Basin



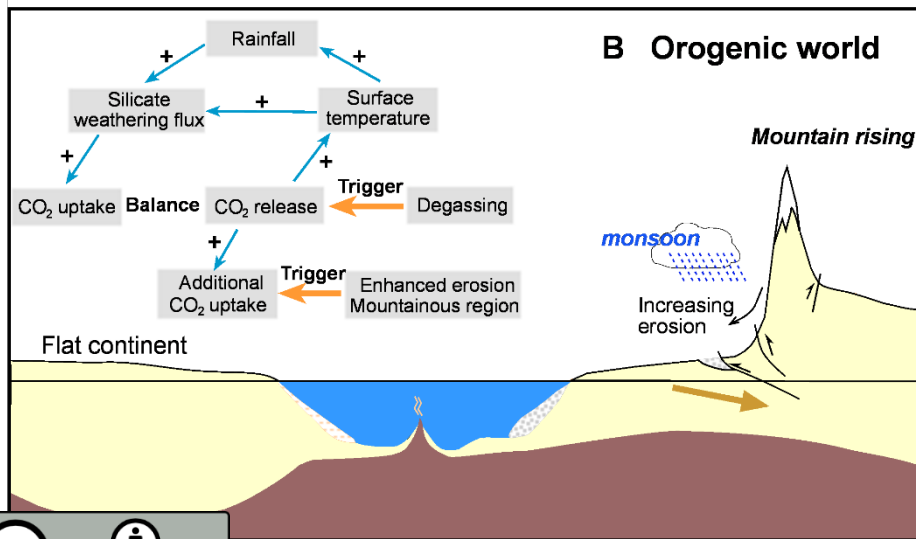
# Global cooling regulate silicate weathering intensity in the northern Tibetan Plateau



# Conceptual model



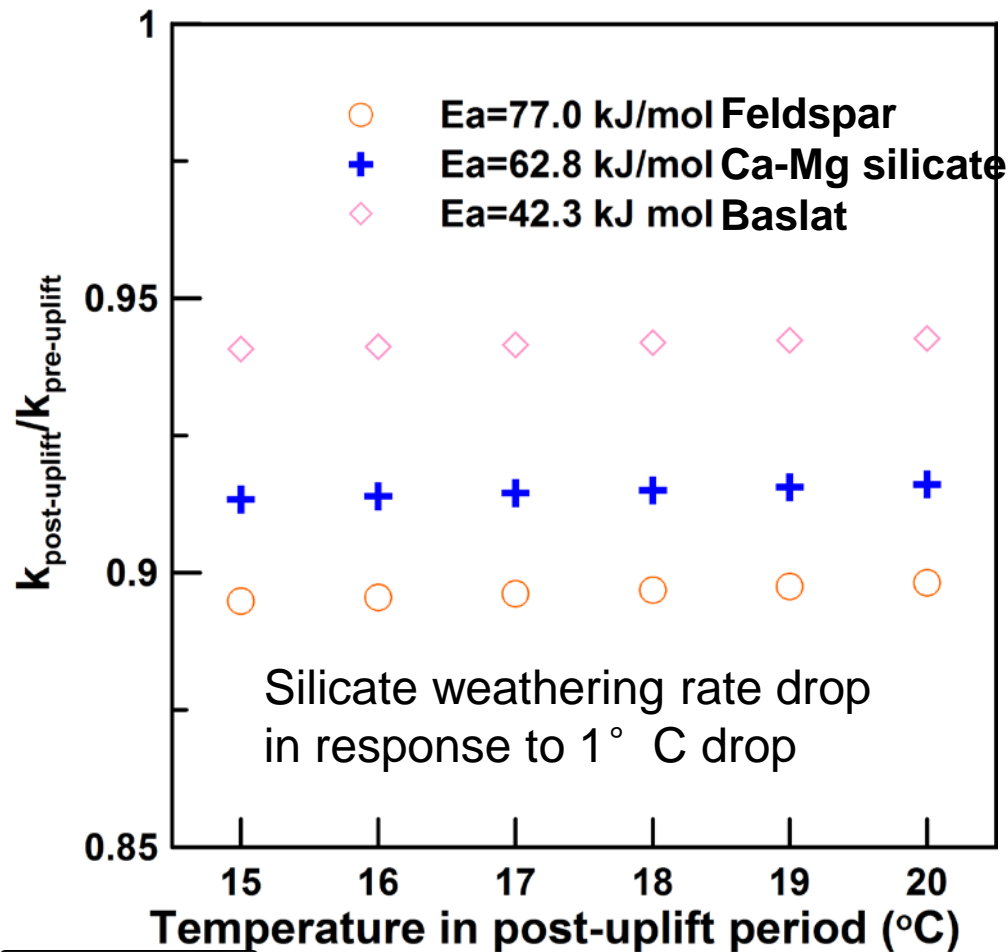
In scenario (A), any increase in degassing flux will lead to an increase in temperature and ultimately will be balanced by nearly the same amount of increase in CO<sub>2</sub> uptake flux through increases in the silicate weathering intensity and thus in the silicate weathering flux.



In scenario (B), given a constant degassing flux, any increase in erosion in orogenic belts will result in an overall increase in the CO<sub>2</sub> uptake flux in orogenic belts. This will be quasi balanced by nearly the same amount of decrease in CO<sub>2</sub> uptake flux in flat tectonically inactive regions through decreases in the silicate weathering intensity and the silicate weathering flux

# Model test

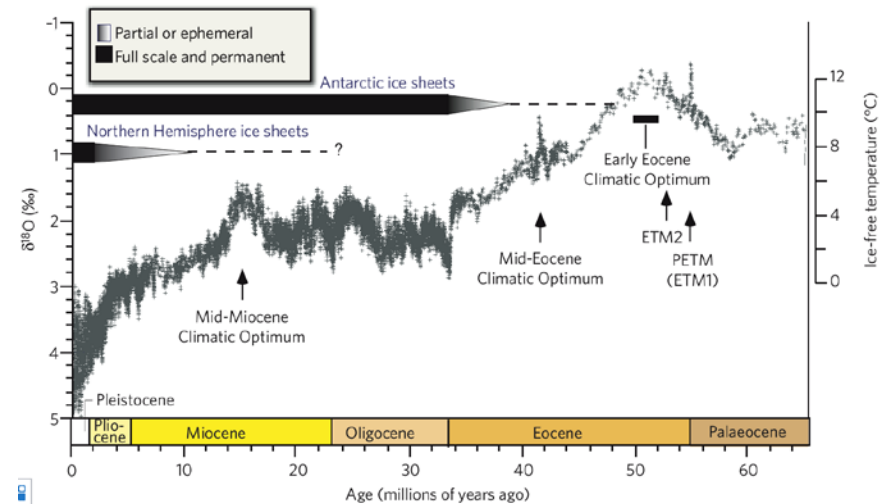
$$\ln(k_{\text{post-uplift}}/k_{\text{pre-uplift}}) = -E_a/R(1/T_{\text{post-uplift}} - 1/T_{\text{pre-uplift}})$$



52-36 Ma Temperature drop  $7-8^\circ\text{C}$ ,  
 $\sim 0.44^\circ\text{C/Myr}$

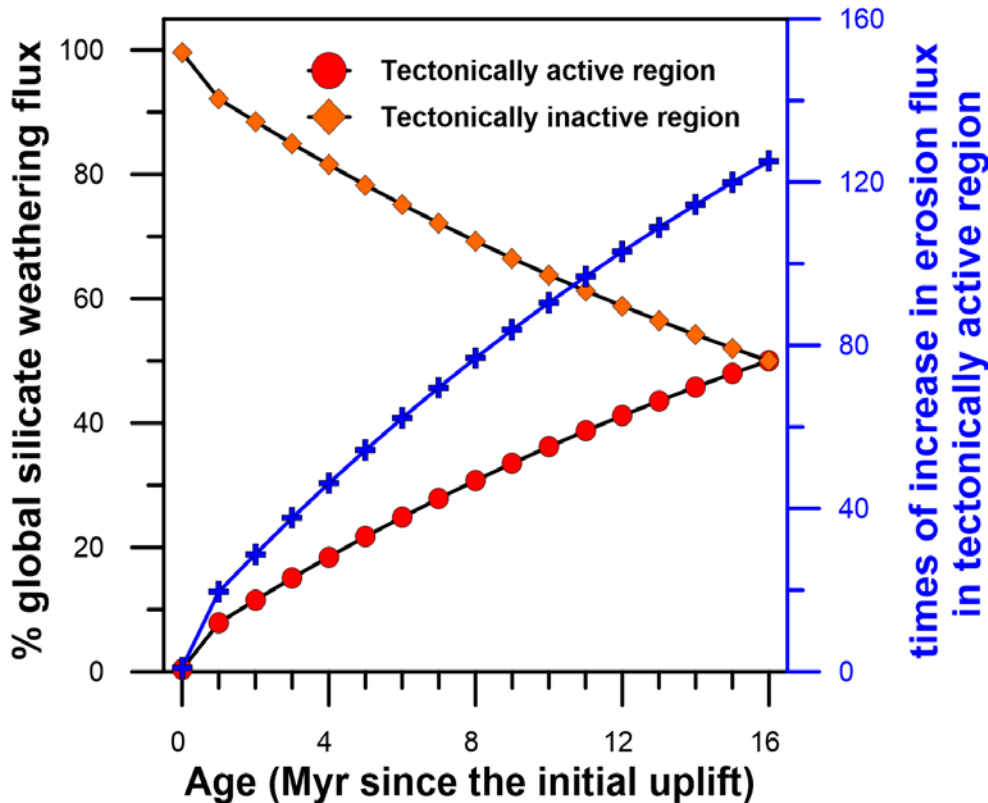


Silicate weathering rate drop  
 $4\%/Myr$



Zachos et al., 2008

# Model test



## Pre-uplift:

Area of tectonically inactive region : 99.6%

Area of tectonically active region : 0.4%

Uniform erosion flux everywhere,  
Weathering flux in proportion to area

## Post-uplift:

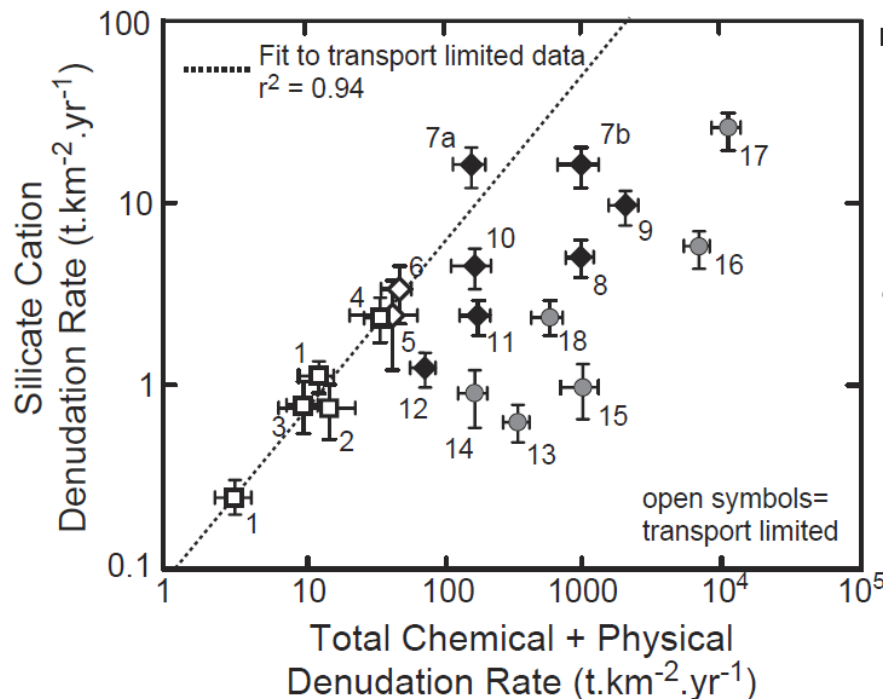
Weathering flux drop in tectonically inactive area would be balanced by erosion-induced weathering flux increase in tectonically active region

## Result:

Erosion increase in tectonically active region should be >100-fold





# Modern observations



- **Continental Cratons**
  - 1. Canadian Shield
  - 2. Siberian Shield
  - 3. African Shield
  - 4. Guyana Shield
- ◆ **Submontane Catchments**
  - 5. British Columbia
  - 6. Sabah Malaysia
  - 7. Puerto Rico
    - a. Long Term Erosion
    - b. Modern Day Erosion
  - 8. East Southern Alps
  - 9. Lesser Himalaya
  - 10. Cote d'Ivoire
  - 11. Idaho Batholith
  - 12. Appalachians
- **Alpine Catchments**
  - 13. Colorado Rockies
  - 14. Sierra Nevada
  - 15. Svalbard
  - 16. High Himalaya
  - 17. West Southern Alps
  - 18. Swiss Alps

# Main display content can be found at

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## **Paleogene global cooling–induced temperature feedback on chemical weathering, as recorded in the northern Tibetan Plateau**

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